

Overview of MIPs that have applied for CMIP6 Endorsement

Applications follow the template available on the CMIP panel website at <http://www.wcrp-climate.org/index.php/wgcm-cmip/about-cmip>

Date: 2 December 2014

Please send any feedback to these applications to the CMIP panel chair (Veronika.Eyring@dlr.de) or directly contact the individual co-chairs for questions on specific MIPs

	Short Name of MIP	Long Name of MIP
1	AerChemMIP	Aerosols and Chemistry Model Intercomparison Project
2	C4MIP	Coupled Climate Carbon Cycle Model Intercomparison Project
3	CFMIP	Cloud Feedback Model Intercomparison Project
4	DAMIP	Detection and Attribution Model Intercomparison Project
5	DCPP	Decadal Climate Prediction Project
6	ENSOMIP	ENSO Model Intercomparison Project
7	FAFMIP	Flux-Anomaly-Forced Model Intercomparison Project
8	GeoMIP	Geoengineering Model Intercomparison Project
9	GMMIP	Global Monsoons Model Intercomparison Project
10	HighResMIP	High Resolution Model Intercomparison Project
11	ISMIP6	Ice Sheet Model Intercomparison Project for CMIP6
12	LS3MIP	Land Surface, Snow and Soil Moisture
13	LUMIP	Land-Use Model Intercomparison Project
14	OCMIP6	Ocean Carbon Cycle Model Intercomparison Project, Phase 6
15	OMIP	Ocean Model Intercomparison Project
16	PDRMIP	Precipitation Driver and Response Model Intercomparison Project
17	PMIP	Palaeoclimate Modelling Intercomparison Project
18	RFMIP	Radiative Forcing Model Intercomparison Project
19	ScenarioMIP	Scenario Model Intercomparison Project
20	SolarMIP	Solar Model Intercomparison Project
21	VolMIP	Volcanic Forcings Model Intercomparison Project
	Diagnostic MIPs (i.e., no proposed experiments rather requesting that certain output is archived and/or contributing to the evaluation and analysis in a coordinated manner)	
22	CORDEX	Coordinated Regional Climate Downscaling Experiment
23	DynVar	Dynamics and Variability of the Stratosphere-Troposphere System
24	GDDEX	Global Dynamical Downscaling Experiment
25	SIMIP	Sea-Ice Model Intercomparison Project
26	VIAAB	VIA Advisory Board for CMIP6

Overview how MIPs contribute to the three CMIP6 science questions

(see Meehl et al., EOS, 2014 for details on the initial CMIP6 design)

Short Name of MIP	The experimental CMIP6 design is focused on three broad scientific questions. Please rank the three science questions in order of importance for and input from your MIP (from 1-3 with 1 being most important and 0 for not relevant at all)		
	How does the Earth System respond to forcing?	What are the origins and consequences of systematic model biases?	How can we assess future climate changes given climate variability, predictability and uncertainties in scenarios?
AerChemMIP	1	2	3
C4MIP	1	3	2
CFMIP	1	2	3
DAMIP	1	3	2
DCPP	2	3	1
ENSOMIP	1	2	3
FAFMIP	1	2	3
GeoMIP	1	2	3
GMMIP	1	2	3
HighResMIP	2	1	3
ISMIP6	1	3	2
LS3MIP	2	1	3
LUMIP	1	3	2
OCMIP6	1	1	2
OMIP	3	1	2
PDRMIP	1	2	3
PMIP	1	2	2
RFMIP	1	2	3
ScenarioMIP	2	3	1
SolarMIP	1	2	3
VolMIP	1	2	3
CORDEX	3	2	1
DynVar	2	1	2
GDDEX	1	2	2
SIMIP			
VIAAB	1	3	2

Overview how MIPs contribute to the six WCRP Grand Challenges and the theme of collaboration on biospheric forcings and feedbacks

(see <http://www.wcrp-climate.org/index.php/grand-challenges>)

Short Name of MIP	It is proposed to use as the scientific backdrop for CMIP6 the six WCRP Grand Challenges (GC), and an additional theme encapsulating questions related to biospheric forcings and feedbacks. Could you please rank the WCRP GCs and theme of collaboration in order of importance for and input from your MIP (from 1-7 with 1 being most important and 0 for not relevant at all)						
	Clouds, Circulation and Climate Sensitivity	Changes in Cryosphere	Climate Extremes	Regional Climate Information	Regional Sea-level Rise	Water Availability	Theme for collaboration: biospheric forcings and feedbacks
AerChemMIP	2	5	3	4	0	0	1
C4MIP	0	3	0	0	0	2	1
CFMIP	1	4	6	2	7	3	5
DAMIP	4	3	2	1	6	5	7
DCPP	3	3	3	1	3	2	3
ENSOMIP	1	7	3	2	6	4	5
FAFMIP	3	4	0	2	1	0	0
GeoMIP	1	3	4	2	0	5	6
GMMIP	2	0	4	1	0	3	0
HighResMIP	1	5	3	4	6	2	7
ISMIP6	5	1	6	4	2	3	7
LS3MIP	0	2	3	4	5	1	6
LUMIP	0	0	4	2	0	3	1
OCMIP6	2	7	1	3	7	7	2
OMIP	4	3	0	2	1	0	5
PDRMIP	1	0	2	4	0	3	0
PMIP	2	3	5	4	6	7	1
RFMIP	1	7	4	2	5	6	3
ScenarioMIP	7	6	3	1	4	5	2
SolarMIP	2	3	4	1	3	0	0
VolMIP	1	4	5	3	6	7	2
CORDEX	5	4	2	1	0	3	6
DynVar	1	3	2	2	0	7	3
GDDEX	5	5	1	1	3	3	5
SIMIP							
VIAAB	7	6	2	1	4	3	5

Timeline Towards MIP Endorsement

- Revised proposals sent to WGCM, WCRP GCs, biogeochemical forcing theme & projects (WGCM co-chairs), MIP co-chairs and modelling groups for review (CMIP Panel, 30 November 2014)
- Review Process Finished (15 January 2015)
- Update of interest of the modelling groups to participate in the MIPs sent to CMIP Panel (Model Groups, 15 January 2015)
- Synthesis of comments and recommendations for each MIP finished and sent to MIP co-chairs (WGCM members organized by WGCM co-chairs, 15 February 2015)
- Final MIP proposals with all information (including data request) sent to CMIP Panel and WIP co-chairs (MIP co-chairs, 31 March 2015)
- Firm Commitment from modelling groups for which MIPs they will perform all of its Tier 1 experiments and providing all the requested diagnostics needed to answer at least one of its science questions (Modelling Groups, 22 April 2015)
- For each of the MIPs, an update of the specific MIP contacts from each model group (Model Groups, 22 April 2015)
- MIP Endorsement (CMIP Panel and WGCM co-chairs, 30 April 2015)
- GMD Special Issue on the CMIP6 experimental design opens (April 2015) with envisaged submission of the April-Endorsed MIPs and the CMIP6 forcings by December 2015.

Timeline CMIP6 Data Request

- Template for CMIP data request sent to MIP co-chairs (WIP co-chairs, 15 December 2014)
- Experiment and variable list sent to WIP co-chairs (MIP co-chairs, 31 January 2015)
- Synthesized data request ready (WIP co-chairs in collaboration with CMIP Panel, 15 March 2015)
- Data request reviewed and sent to WIP co-chairs and CMIP Panel chair (Model groups and MIP co-chairs, 30 April 2015)
- Final data request published (15 July 2015)

Application for CMIP6-Endorsed MIPs

Please return to CMIP Panel Chair Veronika Eyring (email: Veronika.Eyring@dlr.de)

Date: 28 July 2014, updated on 10 November 2014

The recently proposed, revised CMIP structure (see information on the CMIP Panel website at <http://www.wcrp-climate.org/index.php/wgcm-cmip/about-cmip>) provides for a small set of experiments to be routinely performed by modeling groups whenever they develop a new model version. The output from these so-called *ongoing CMIP Diagnostic, Evaluation and Characterization of Klima (DECK)* experiments and the *CMIP6 Historical Simulation* will be distributed for community use via the ESGF infrastructure. Other Model Intercomparison Projects (MIPs) will build on the CMIP DECK experiments and the CMIP6 Historical Simulation and augment them to address a broad range of scientific questions. Additionally proposed MIP experiments together with the CMIP DECK experiments and the CMIP6 Historical Simulation will constitute the suite of simulations for the next phase of CMIP.

MIPs are invited to request endorsement for the next phase of CMIP (i.e., CMIP6). Applications from MIPs requesting status as a CMIP6-Endorsed MIP should be sent to the CMIP Panel Chair. The current set of MIP proposals is now complete and will be revised on the agreed timeline. We will review any additional proposals in a year from now at the next WGCM meeting in October 2015. A MIP may propose that a subset or even all of their experiments be included as part of the suite of simulations constituting CMIP6. The CMIP Panel will, together with the WGCM co-chairs, decide whether a MIP and its experiments meet the criteria for endorsement for CMIP6. Note that it is expected that all additional experiments proposed for CMIP6 will be scientifically analyzed and exploited by the MIP.

CMIP6-Endorsed MIPs can make full use of the ESGF infrastructure. In order to minimize the burden imposed on modeling groups wishing to participate, the MIPs seeking to be part of CMIP Phase X must agree to comply with the CMIP standards in terms of experimental design, data format and documentation. In general the WGCM encourages adhering to the standards in place for CMIP.

The main criteria for MIPs to be endorsed for CMIP6 are

1. The MIP and its experiments address at least one of the key science questions of CMIP6.
2. The MIP demonstrates connectivity to the DECK experiments and the CMIP6 Historical Simulation.
3. The MIP adopts the CMIP modeling infrastructure standards and conventions.
4. All experiments are tiered, well-defined, and useful in a multi-model context and don't overlap with other CMIP6 experiments.
5. Unless a Tier 1 experiment differs only slightly from another well-established experiment, it must already have been performed by more than one modeling group.
6. A sufficient number of modelling centers (~8) are committed to performing all of the MIP's Tier 1 experiments and providing all the requested diagnostics needed to answer at least one of its science questions.
7. The MIP presents an analysis plan describing how it will use all proposed experiments, any relevant observations, and specially requested model output to evaluate the models and address its science questions.
8. The MIP has completed the MIP template questionnaire.
9. The MIP contributes a paper on its experimental design to the CMIP6 Special Issue.
10. The MIP considers reporting on the results by co-authoring a paper with the modelling groups.

Proposals from MIPs should include the following information:

- * *Preliminary information used to determine whether a MIP should be endorsed for CMIP6 or not.*
- ** *Information that must be provided later (and before the panel can determine which experiments, if any, will be incorporated in the official CMIP6 suite).*
- Name of MIP*
- Co-chairs of MIP (including email-addresses)*
- Members of the Scientific Steering Committee*
- Link to website (if available)*
- Goal of the MIP and a brief overview*
- References (if available)*
- An overview of the proposed experiments*
- An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*
- Proposed timing*
- For each proposed experiment to be included in CMIP6**
 - the experimental design;
 - the science question and/or gap being addressed with this experiment;
 - possible synergies with other MIPs;
 - potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.
- If possible, a prioritization of the suggested experiments, including any rationale**
- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**
- List of output and process diagnostics for the CMIP DECK/CMIP6 data request**
 - whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;
 - whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);
 - whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.
 - whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;
 - the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.
- Any proposed contributions and recommendations for**
 - model diagnostics and performance metrics for model evaluation;
 - observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;
 - tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).
- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms.**
- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF.**

AerChemMIP (Aerosols and Chemistry MIP)

Application for CMIP6-Endorsed MIPs

Date: 2 December 2014

➤ Co-chairs of MIP

William Collins (UK) (W.Collins@reading.ac.uk)

Jean-François Lamarque (US) (lamar@ucar.edu)

Michael Schulz (Norway) (michael.schulz@met.no)

➤ Members of the Scientific Steering Committee

Olivier Boucher (France) (olivier.boucher@lmd.jussieu.fr)

Veronika Eyring (Germany) (veronika.eyring@dlr.de)

Arlene Fiore (US) (amfiore@ldeo.columbia.edu)

Michaela Hegglin (UK) (m.i.hegglin@reading.ac.uk)

Gunnar Myhre (Norway) (gunnar.myhre@cicero.oslo.no)

Michael Prather (US) (mprather@uci.edu)

Drew Shindell (US) (drew.shindell@duke.edu)

Steve Smith (US) (ssmith@pnnl.gov)

Darryn Waugh (US) (waugh@jhu.edu)

Goal of the MIP

Past climate change has been forced by a wide range of chemically reactive gases, aerosols, and well mixed greenhouse gases (WMGHGs), in addition to CO₂. Scientific questions and uncertainties regarding chemistry-climate interactions range from regional scales (e.g., tropospheric ozone and aerosols interacting with regional meteorology), to long-range connections (e.g., hemispheric transport of air pollution, the impacts of lower stratospheric ozone and temperatures on surface climate), to global integration (e.g., the lifetimes of CH₄ and N₂O).

AerChemMIP proposes to contribute to CMIP6 through the following: 1) diagnose forcings and feedbacks involving NTCFs, (namely tropospheric aerosols, tropospheric O₃ precursors, and CH₄) and the chemically reactive WMGHGs (e.g., N₂O, also CH₄, and some halocarbons** including impacts on stratospheric O₃), 2) document and understand past and future changes in the chemical composition of the atmosphere, and 3) estimate the global-to-regional climate response from these changes.

The AerChemMIP Tier 1 simulations focus primarily on understanding atmospheric composition changes (from NTCFs and other chemically-active anthropogenic gases) and

their impact on climate. We have devised a series of experiments that contrast the forcing of various NTCFs with that of WMGHGs in historical and future climate change. In addition, the proposed chemistry-climate simulations will enable diagnosis of changes in regional air quality (AQ) through its coupling to large-scale changes in O_3 - CH_4 - $PM_{2.5}$. We will work in collaboration with RFMIP and DAMIP to provide a comprehensive analysis of ERF and the regionally-resolved climate forcing signature from tropospheric NTCFs. For some of the specifically attributable climate forcings, in particular those at the 10s of $mW\ m^{-2}$ level, the actual climate change will be difficult to detect in a transient simulation or even a time slice of several decades. AerChemMIP is a joint, consolidated effort for CMIP6 from two international communities -- Aerosol Comparisons between Observations and Models (AeroCom, <http://aerocom.met.no/Welcome.html>) and the IGAC/SPARC Chemistry-Climate Model Initiative (CCMI, <http://www.met.reading.ac.uk/ccmi/>). Experiments suggested for CCMI Phase 2 [Eyring *et al.*, 2013b], which are traditionally run using chemistry-climate models (CCMs) with mostly prescribed sea surface temperatures and sea ice concentrations, complement this set of AerChemMIP/CMIP6 experiments. Further experiments in AeroCom phase III aim to understand sensitivity of aerosol forcing to aerosol formation and loss processes.

We do not specifically consider the very long-lived F-gases (SF_6 , PFCs, and some HFCs) as their abundance is not affected by chemistry on a century time scale.

Overview

Aerosols and ozone were identified in IPCC AR5 (Myhre *et al.*, 2013) as the main sources of uncertainty in the radiative forcing since pre-industrial times. Uncertainties in projecting the chemically reactive WMGHGs as well as future air quality from global changes were also identified in AR5 [Kirtman *et al.*, 2013]. In addition to changing anthropogenic emissions evaluated in AR5, natural aerosols originating from biogenic sources, dust or sea-salt are a primary contributor to the uncertainty in current forcing (Carslaw *et al.* 2013). Due to the nonlinear response of clouds to the background level of aerosols, the response of the climate system to human perturbations will depend critically on the natural aerosol background (Carlton *et al.*, 2010).

Beyond aerosols, the biogeochemistry of ecosystems provides large sources of the WMGHGs CH_4 and N_2O , as well as O_3 precursors (lightning and soil nitrogen oxides, volatile organic compounds, wildfire emissions). These sources are likely to be affected by climate change, leading to a variety of feedbacks that to date have only been quantified from a limited number of studies (and models) and thus demand for a coordinated set of simulations that allows for a consistent and clean comparison between models.

Anthropogenic emissions of NTCFs have been responsible for a climate forcing that is presently nearly equal in magnitude to CO_2 -forcing. These emissions have led to a variety of global climate impacts such as regional patterns of temperature and precipitation, with a magnitude similar to the global-mean equivalent ERF of WMGHG. In addition, NTCF ERF is inherently inhomogeneous, and there is some evidence that where NTCF on a regional scale

is large, the climate response differs from the globally equivalent ERF – i.e., there is some regional response to regional ERF.

NTCF emissions are also responsible for driving regional and local air quality (AQ). This has led to the recognition that a combined strategy of mitigating climate change and air pollution together has clear economic benefits compared to separate mitigation (IPCC, 2014 – WG3 SPM). In our future world, most, if not all scenarios lead to changes in the emissions and meteorology that determine air quality and create pollution episodes. The knowledge base used to manage air pollution to date must be updated based on more comprehensive information that CMIP6 will provide on future air chemistry climatologies. The exposure risks of human health and assets (agriculture, built environment, ecosystems) will be driven by daily variations in surface ozone and particulate matter in addition to deposition of nitrate and sulfate and any land-use change interacting with atmospheric changes.

The forcing of climate by ozone changes results from tropospheric increases and lower stratospheric decreases, with interaction between those. They are the result of combined impacts from climate change and multiple emission changes. For example, one of the largest components of CH₄ emissions' ERF is that from the increase in tropospheric O₃. In addition, stratospheric O₃ depletion since the 1970s has led to significant cooling of the lower stratosphere, and through the Antarctic ozone hole is linked to changes in tropospheric circulation and rainfall patterns in the southern hemisphere, especially during summer (WMO, 2014). In the Southern Hemisphere, future summertime circulation changes are controlled by both the ozone recovery rate and the rate of GHG increases [Eyring *et al.*, 2013a], indicating the need to account for ozone changes in future climate projections.

Since some models participating in CMIP6 do not have interactive chemistry and aerosol schemes, AerChemMIP will also provide historical and future time-varying aerosol, ozone, and stratospheric water vapour concentration fields for CMIP6. The ozone database will be an update of the database provided for CMIP5 by [Cionni *et al.*, 2011]. This data will be generated from a mixture of CCMs and CTMs simulations which are not themselves part of CMIP6.

Overview of the Proposed Tier 1 Experiments

The AerChemMIP Tier 1 simulations focus on three science questions

1. How have NTCF and ODS emissions contributed to global ERF and affected regional climate over the historical period?
2. How will future policies (on climate/AQ/land use) affect the NTCFs and their climate impacts?
3. How have WMGHGs forced climate (including through their chemical impacts) over the historical period?

In the following sections, we discuss each question separately and provide for each science question the description of the simulations necessary to answer the stated question. Note that we emphasize the use of the Effective Radiative Forcing (ERF) to measure climate

forcing. We have provided at the end of this document a description of the methodology associated with this calculation.

1. How have NTCF and ODS emissions contributed to global ERF and affected regional climate over the historical period?

Anthropogenic non-CO₂ emissions (e.g., NTCFs, GHGs like halocarbons and N₂O,...) have led to a climate forcing that is commensurate to CO₂-forcing on regional scales, especially over the last few decades.

By way of their associated large uncertainty in radiative forcing since pre-industrial times, ozone and aerosols in particular are a key factor behind the large uncertainty in constraining climate sensitivity over the record of observed data. These NTCFs have an inhomogeneous spatial distribution and the degree of regional temperature and precipitation responses to such heterogeneous forcing remains an open question within the scientific community. It is further unclear whether NTCFs, which are primarily located at Northern Hemisphere mid latitude land areas have led to a larger climate response there, relative to forcing from WMGHGs.

One unambiguous regional response to inhomogeneous climate forcing concerns the Southern hemisphere summertime surface circulation changes induced by the Antarctic ozone hole as an indirect response to ozone-depleting halocarbons. These changes have been argued to lead to changes in rainfall patterns, ocean circulation, and sea-ice cover. The relative role of these ozone-induced changes compared to other anthropogenic forcings and natural variability is not fully resolved by the scientific community (with some studies reaching contradictory conclusions). Hence there is a need for multi-model ensemble of simulations, especially with models resolving stratospheric chemistry that isolate the role of stratospheric ozone depletion.

Experiment 1.1: Transient historical coupled ocean climate impacts of NTCFs and of ozone depleting halocarbons (note: this builds on CMIP6-historical-simulation, which is used as the reference simulation, and requires AerChemMIP diagnostics therein)

- 1.1.1 Perturbation: Historical WMGHG (including halocarbon) concentrations, 1850 NTCF emissions. 165 years, 1-3 ensemble members
- 1.1.2 Perturbation: Historical WMGHG concentrations and NTCF emissions, 1950 halocarbons. 65 years (branched from CMIP6 historical in 1950), 1 up to the number of ensemble members performed for the CMIP6 historical

Experiment 1.2: Estimating ERFs through specified transient historical SST simulations (see note on ERFs below).

Perform 1850-2014 (1 ensemble member only) simulation with all forcings as in CMIP6 historical **but** with

- 1.2.1 1850 tropospheric ozone precursor emissions (including biomass burning) 165 years

- 1.2.2 1850 all NTCF emissions (including biomass burning). 165 years
- 1.2.3 1950 ODSs. 65 years (1950-2014)

Experiment 1.3. Time-slice simulations based on the 1850 control SSTs to compute the ERF for 1850 and 2014 for all NTCF and natural aerosols (e.g. AR5 fig 8.15). This requires four simulations

- 1.3.1 Control: 1850 WMGHG concentrations and 1850 NTCF emissions. 20 years
- 1.3.2 Perturbation: 1850 WMGHG concentrations, 2014 NTCF emissions. 20 years
- 1.3.3. Perturbation: Doubled dust emissions. 20 years
- 1.3.4. Perturbation: Doubled sea salt emissions. 20 years

2. How will future policies (on climate/AQ/land use) affect the NTCFs and their climate impact? What are the patterns of associated climate forcing, and how do these patterns translate into temperature and precipitation changes?

For the upcoming decades policy makers will be making choices in 3 broadly defined areas 1) climate change policies (targeting mostly WMGHGs), 2) air quality policies (targeting mostly NTCF emissions including CH₄ that are precursors of tropospheric aerosols and tropospheric ozone) and 3) land-use policies. AerChemMIP aims to identify the patterns of chemical change at the global and regional levels as well as the ERF associated with NTCF mitigation efforts (focusing on policy choices in areas 1 and 2 above), and their climate (surface temperature and precipitation) and environmental (health, ecosystem, visibility, ...) impact between 2015 and 2055 (this is the time frame over which aerosol and precursor emissions are expected to be significant). The impact analysis will be performed by contrasting the following simulations: a) a reference experiment with high aerosol emissions (such as SSP3-7, but the final decision will be made with ScenarioMIP) imulations (with sufficient diagnostics) and b) perturbation experiments replacing NTCF emissions in reference experiment with much reduced NTCF emissions. These perturbations will be designed in collaboration with ScenarioMIP to ensure that perturbations are consistent with the underlying story line of the scenario in consideration.

Experiment 2.1: Transient coupled ocean climate impacts

- 2.1.1 Reference: SSP3-7 (to be performed under ScenarioMIP)
- 2.1.2 Perturbation: SSP3 with reduced NTCF (aerosol and tropospheric ozone precursors, including methane) 40 years, 1-3 ensemble members

Experiment 2.2: Estimating ERFs through fixed-SST simulations (SSTs from 2.1.1)

- 2.2.1 Control: as Experiment 2.1.1 using archived SSTs from 2.1.1
40 years, one ensemble
- 2.2.2 Perturbation: Only black carbon emissions as in Experiment 2.1.2 (this is to isolate the specific role of black carbon in near-term policy decisions)
40 years, one ensemble
- 2.2.3 Perturbation: All aerosol precursor emissions (but not NO_x) as in 2.1.2,
40 years, one ensemble

- 2.2.4 Perturbation: All ozone precursors except methane kept the same as in 2.1.2, 40 years, one ensemble
- 2.2.5 Perturbation: Methane kept the same as in 2.1.2, 40 years, one ensemble

3. How have chemically reactive WMGHGs affected the forcing over the historical period?

Under this question, we focus on estimating the forcing from changes in methane and nitrous oxide on ozone (tropospheric and stratospheric), aerosol oxidation, and emissions of natural aerosols, including the climate impacts associated with those changes. Note that only ERF estimates are calculated, while the associated transient coupled simulations are in Tier 2.

Experiment 3.1: Estimating ERFs through specified SST simulations (SSTs taken from CMIP6 historical simulation)

Perform 1850-2015 (1 ensemble member only) simulation with all forcings (and including chemistry feedbacks on tropospheric and stratospheric ozone) as in transient historical **but** with

- 3.1.1 1850 CH₄. 165 years
- 3.1.2 1850 N₂O. 165 years

Total amount of simulation years (Tier 1)

Experiments 1.x.x: 705y - 1035y (overlap w DAMIP ca 330y-990y) (overlap w RFMIP ca 80y)
Experiments 2.x.x: 240y - 320y (excluding 2.1.1, run under ScenarioMIP)
Experiments 3.x.x: 330y

Synergy with other MIPs – Model diagnostics

Experiment 1.1.1/1.1.2 parallels similar forcing attribution simulations in DAMIP but include chemistry responses and diagnostics.

Experiments 1.2.4/1.2.5/3.2.1/3.2.2: These parallel similar ERF calculations in RFMIP, but start from emission changes rather than concentration changes

Experiments 2.1.1/2.1.2 extend the ScenarioMIP simulations to separate out the impact of AQ policies and NTCFs

Model diagnostics specific to AerChemMIP Tier 1 experiments need to be implemented also in the DECK and CMIP6-historical-simulation. The diagnostics will be contributed to the CMIP6 data request by AerChemMIP. If models have not all components to compute dynamic aerosols, tropospheric or stratospheric chemistry, models are requested to consider using the forcing fields of chemical compounds provided by AerChemMIP when performing AerChemMIP Tier 1 experiments.

Overview of the Proposed Tier 2 and 3 Experiments

AerChemMIP will also initiate additional experiments to document with an eventually more limited set of models complementing science questions, which are based on tier 1 experiments, and make efficient use of the general set-up of CMIP6. The Tier 2 and 3 experiments are still being discussed and are described here for completeness. We will finalize the design of the Tier 2 and 3 experiments by the end of January 2015.

Table 1: AerChemMIP experiments tier 2 and 3: [See Excel spreadsheet.](#)

Model Diagnostics and Performance Metrics for Model Evaluation

AerChemMIP will contribute to the CMIP6 data request by suggesting aerosol and chemistry related output that is required for model evaluation (including the characterization of air quality extremes) and for diagnosing radiative forcings from NTCFs. In addition, AerChemMIP will contribute to the development of the Earth System Model Evaluation Tools (ESMValTool, [Righi *et al.*, 2014]), the documentation of aerosol parameters via the AeroCom tools and will include important chemistry-related diagnostics and performance metrics for CMIP6 model evaluation.

Design of Effective Radiative Forcing simulations.

The proposed simulations combine analysis of the effective radiative forcing (ERF) and the consequent climate impacts of NTCFs. The RF from WMGHGs will be provided by RFMIP. The ERFs are calculated by comparing the net TOA radiation fluxes between two runs with the same SSTs but with perturbed NTCF emissions (see below). Internal variability (mainly clouds) generates considerable noise therefore 20 years of simulation are needed to characterize the present day ERF from NTCFs. Alternatively, models that can nudge their simulated model winds (only, towards meteorological analyses or previously generated meteorological fields) should be able to identify a statistically-significant signal with as little as 3 years of simulation. In a similar way a pair of runs driven by evolving SSTs but with and without evolving NTCF emissions will provide the time evolution of the NTCF ERF. For the temperature and precipitation impacts, simulations with a coupled ocean are needed. Again, this requires a pair with and without evolving NTCF emissions in order to compute the impacts. The internal variability in the coupled ocean models is larger than with fixed SSTs, so at least 3 ensemble members will be needed.

The effective radiative forcing (ERF) was introduced in IPCC AR5 [Boucher *et al.*, 2013; Myhre *et al.*, 2013]. The definition is given as follows: '*ERF is the change in net TOA downward radiative flux after allowing for atmospheric temperatures, water vapour and clouds to adjust, but with surface temperature or a portion of surface conditions unchanged*'. This is different from the traditional radiative forcing (RF) concept where surface and tropospheric temperature and other variables such as water vapour and clouds must be kept fixed. Quantification of a climate driver by ERF and RF provides different results for some aerosol effects where the latter concept allows quantification of semi-direct effect and second

indirect aerosol effect (ERF of aerosol-radiation interaction and aerosol-cloud interaction, respectively). For greenhouse gases RF and ERF are more similar in magnitude, but the latter has larger uncertainty.

Two ways to simulate ERF is currently used, namely; i) net TOA fluxes from fixed-sea surface temperature (SST) simulations and ii) regression of transient temperature response with the initial radiative perturbation [Gregory *et al.*, 2004]. The two methods for simulating ERF are illustrated in [Boucher *et al.*, 2013; Sherwood *et al.*, 2014]. Both ERF methods have their advantages and disadvantages [Boucher *et al.*, 2013; Myhre *et al.*, 2013]. The regression method can be applied to many of the typical CMIP runs, but require long runs (at least 20 years) with a significant radiative perturbation. The fixed-SST method can be applied to relatively small radiative perturbations, but not all modelling groups have access to fixed-SST type simulations.

The fixed-SSTs approach can further be applied with additional radiation calls to diagnose the various aerosol effects [Ghan *et al.*, 2012]. Separate diagnostics for shortwave and longwave changes are applied. To diagnose the indirect aerosol effect and semi-direct effect the scattering and absorption by aerosols are neglected by setting refractive indexes of anthropogenic aerosol to zero, see [Ghan *et al.*, 2012] for further details.

References

- Boucher, O., et al. (2013), Clouds and Aerosols, in *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by T. F. Stocker, D. Qin, G.-K. Plattner, M. Tignor, S. K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P. M. Midgley, pp. 571-657, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA.
- Cionni, I., V. Eyring, J. F. Lamarque, W. J. Randel, D. S. Stevenson, F. Wu, G. E. Bodeker, T. G. Shepherd, D. T. Shindell, and D. W. Waugh (2011), Ozone database in support of CMIP5 simulations: results and corresponding radiative forcing, *Atmos. Chem. Phys. Discuss.*, 11(4), 10875-10933.
- Eyring, V., et al. (2013a), Long-term ozone changes and associated climate impacts in CMIP5 simulations, *J Geophys Res-Atmos*, 118(10), 5029-5060.
- Eyring, V., et al. (2013b), Overview of IGAC/SPARC Chemistry-Climate Model Initiative (CCMI) Community Simulations in Support of Upcoming Ozone and Climate Assessments, *SPARC Newsletter*, 40, 48-66.
- Ghan, S. J., X. Liu, R. C. Easter, R. Zaveri, P. J. Rasch, J.-H. Yoon, and B. Eaton (2012), Toward a minimal representation of aerosols in climate models: Comparative decomposition of aerosol direct, semi-direct and indirect radiative forcing, *J. Climate*, doi: 10.1175/JCLI-D-1111-00650.00651, in press.
- Gregory, J. M., W. J. Ingram, M. A. Palmer, G. S. Jones, P. A. Stott, R. B. Thorpe, J. A. Lowe, T. C. Johns, and K. D. Williams (2004), A new method for diagnosing radiative forcing and climate sensitivity, *Geophysical Research Letters*, 31(3), L03205.
- Kirtman, B., et al. (2013), Chapter 11. Near-term Climate Change: Projections and Predictability in *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, Cambridge University Press Cambridge, United Kingdom.

- Myhre, G., et al. (2013), Anthropogenic and Natural Radiative Forcing, in *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by T. F. Stocker, D. Qin, G.-K. Plattner, M. Tignor, S. K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P. M. Midgley, pp. 659-740, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA.
- Righi, M., V. Eyring, K.-D. Gottschaldt, C. Klinger, F. Frank, P. Jöckel, and I. Cionni (2014), Quantitative evaluation of ozone and selected climate parameters in a set of EMAC simulations, *Geosci. Model Dev. Discuss.*, 7, 6549-6627.
- Sherwood, S. C., S. Bony, and J. L. Dufresne (2014), Spread in model climate sensitivity traced to atmospheric convective mixing, *Nature*, 505(7481), 37-+.

Coupled Climate Carbon Cycle MIP (C4MIP)

Application for CMIP6-Endorsed MIPs

Date: 1 December 2014

The Coupled Climate Carbon Cycle MIP, C⁴MIP, requests formal endorsement by WGCM for the next phase of CMIP (CMIP6).

Background and motivation

The carbon cycle is the key addition to physical climate models that makes them “Earth System Models” (ESMs). CMIP5 was the first CMIP phase to include ESMs as the standard climate change modeling tool and carbon cycle results featured strongly in the IPCC 5th Assessment Report (see for example WG1 SPM, TS, chapters 6, 9, 12 and WG2 chapter 4). WG1 SPM highlighted the direct link from anthropogenic emissions to global climate change through the policy relevant Transient Response to Cumulative Emissions (TCRE). This is a key advance over AR4 – an advance only possible due to the inclusion of the carbon cycle in physical climate models.

C⁴MIP has been central in this development work, and the first C⁴MIP intercomparison paper (Friedlingstein et al., J. Clim., 2006) now has more than 1000 citations. Extensive use was made of the carbon cycle simulations in the CMIP5 database which led to a *Journal of Climate* Special Issue titled “Climate–Carbon Interactions in the CMIP5 Earth System Models” (<http://journals.ametsoc.org/page/C4MIP>)

Further development, evaluation and assessment of carbon cycle processes are one of the key focus areas for global climate modelling centres and we fully expect carbon cycle processes to be of fundamental importance in CMIP6 (and beyond).

Proposals from MIPs should include the following information:

- * *Preliminary information used to determine whether a MIP should be endorsed for CMIP6 or not.*
- ** *Information that must be provided later (and before the panel can determine which experiments, if any, will be incorporated in the official CMIP6 suite).*

➤ **Name of MIP**

Coupled Climate Carbon Cycle MIP (C⁴MIP)

➤ **Co-chairs of MIP (including email-addresses) (alphabetical order)**

Vivek Arora, Canadian Centre for Climate Modelling and Analysis, Canada, vivek.arora@ec.gc.ca

Pierre Friedlingstein, University of Exeter, UK, p.friedlingstein@exeter.ac.uk

Chris Jones, Met Office Hadley Centre, UK, chris.d.jones@metoffice.gov.uk

➤ **Members of the Scientific Steering Committee**

Co-chairs plus a steering committee providing additional expertise:

- Ocean biogeochemistry: Laurent Bopp (IPSL, France) and Tatiana Ilyina (MPI, Germany)
- Nitrogen cycle: Sonke Zaehle (MIP, Germany)
- Permafrost: Charles Koven (LBL, USA)
- Observations: Heather Graven (Imperial College, UK), Martin Jung (MPI, Germany)
- Evaluation/iLAMB: Forrest Hoffman (ORNL, USA), Jim Randerson (UC Irvine, USA)
- Land use change/LUMIP: Julia Pongratz (MPI, Germany), Victor Brovkin (MPI, Germany)

Additional experts on related activities might be invited to attend SSC meetings (eg. land use forcing (LUMIP), methane emissions, offline analysis (OCMIP, TRENDY), Remote sensing data, TCRE, ...)

Several of the SSC members are also members of other MIPs : ScenarioMIP (Friedlingstein), LUMIP (Jones, Brovkin, Pongratz), OCMIP (Bopp)

➤ **Link to website (if available)**

TBC

➤ **Goal of the MIP and a brief overview**

- The primary focus of C4MIP is to understand and quantify future (century-scale) changes in land and ocean carbon storage and fluxes
- Idealized experiments will be used to separate and quantify the sensitivity of land and ocean carbon cycle to changes in climate and changes in atmospheric CO₂ concentration
- Historical experiments will be used to evaluate model performance and investigate potential for future constraints
- Future scenario experiments will be used to quantify future changes in carbon storage and hence quantify the atmospheric CO₂ concentration and related climate change for given CO₂ emissions, or diagnose the emissions compatible with a prescribed atmospheric CO₂ concentration pathway

➤ **References (if available)**

- Friedlingstein, P. and coauthors (2006): Climate–carbon cycle feedback analysis: Results from the C4MIP model intercomparison, *Journal of Climate*, 19(14), 3337-3353
- Arora, V.K. and coauthors (2013) Carbon–Concentration and Carbon–Climate Feedbacks in CMIP5 Earth System Models. *Journal of Climate*, Vol. 26, Iss. 15, pp. 5289-5314.
- Friedlingstein, P. and coauthors (2013) Uncertainties in CMIP5 Climate Projections due to Carbon Cycle Feedbacks. *Journal of Climate*, Vol. 27, Iss. 2, pp. 511-526.
- Jones, C. D. and coauthors (2013) Twenty-First-Century Compatible CO₂ Emissions and Airborne Fraction Simulated by CMIP5 Earth System Models under Four Representative Concentration Pathways. *Journal of Climate*, Vol. 26, Iss. 13, pp. 4398-4413.
- Gillett, N. P., et al. (2013) Constraining the Ratio of Global Warming to Cumulative CO₂ Emissions Using CMIP5 Simulations. *Journal of Climate*, Vol. 26, Iss. 18, pp. 6844-6858.
- Schwinger J., et al. (2014) Non-linearity of ocean carbon cycle feedbacks in CMIP5 earth system models. *Journal of Climate*, Vol. 27, Iss. 11, pp. 3869–3888.
- IPCC AR5 WG1 SPM, TS and Chapters 6, 9 and 12
- Anav, A., et al. "Evaluating the land and ocean components of the global carbon cycle in the CMIP5 Earth System Models." *Journal of Climate* 26.18 (2013): 6801-6843.

- Todd-Brown, K. E. O., et al. "Causes of variation in soil carbon simulations from CMIP5 Earth system models and comparison with observations." *Biogeosciences* 10.3 (2013): 1717-1736.
- Hoffman, Forrest M., et al. "Causes and implications of persistent atmospheric carbon dioxide biases in Earth System Models." *Journal of Geophysical Research: Biogeosciences* 119.2 (2014): 141-162.

➤ **An overview of the proposed experiments**

C4MIP will build on the DECK and the CMIP6 Historical Simulation. The following simulations are pre-requisite for C4MIP participation:

DECK simulations

Control simulation

Requested for diagnosis of model drift (drift in land and ocean carbon pools)

CMIP 1% per year increasing CO₂ up to 4xCO₂ simulation

Requested as baseline for the C4MIP climate-carbon cycle feedback analysis; also requested for assessment of TCRE.

CMIP6 Historical Simulation

CMIP6 Historical Simulation with prescribed CO₂ emissions

Requested for model evaluation and continuity with scenario simulations from Tier1

C4MIP SIMULATIONS

C4MIP- Tier 1

Tier1.1: 1%BGC: biogeochemically-coupled version of 1% per year increasing CO₂ up to 4xCO₂ simulation. CO₂ increase only affects carbon cycle models, radiative code sees pre-industrial CO₂
Requested for the C4MIP climate-carbon cycle feedback analysis.

Tier1.2: Emission-driven future scenario (SSP-based RCP SSP5-8.5) up to 2100

Requested for analysis of impact of carbon cycle feedbacks on climate projections over the 21st century. Also requested for assessment of cumulative emissions compatible with climate targets.

C4MIP- Tier 2

Tier 2.1: 1%RAD: radiatively-coupled version of the 1% per year increasing CO₂ up to 4xCO₂ simulation. CO₂ increase only affects the radiative code, carbon cycle models see pre-industrial CO₂
Requested for further C4MIP climate-carbon cycle feedback analysis (non-linearities/synergies)

Tier 2.2: Emission-driven future scenario (SSP-based RCP SSP5-8.5) extension to 2300

Requested for analysis of impact of carbon cycle feedbacks on climate projections for slow components (vegetation, permafrost, oceanic circulation)

Tier 2.3: Emission-driven CMIP6 Historical simulation and SSP5-8.5, BGC mode

Requested for assessment of CO₂-carbon cycle feedbacks over the 21st century; also for assessment of CO₂ induced warming. Extension to 2300 recommended for groups doing Tier2-2. simulation

Possible additional simulation under Tier 2 (to be decided with ScenarioMIP and/or GEOMIP):

Overshoot on an SSP-2.6 scenario (emission-driven or concentration driven to be decided). This will require further discussion with ScenarioMIP and GEOMIP to agree on a potential scenario to be used, which science questions it would primarily address, and hence in which MIP it would better fit.

CURRENT C4MIP PROPOSAL

Category	Type of Scenario	Emission or concentration driven	Coupling mode	Simulation years	Short name	Use by other MIPs
Tier 1						
1%BGC	Idealised 1% per year CO ₂ only, BGC mode	C-driven	CO ₂ affects BGC	140	esm1pcbgc	OCMIP, LS3MIP
SSP5-8.5	SSP5-8.5 up to 2100	E-driven	Fully coupled	85	esmssp5-85	ScenarioMIP, LUMIP, OCMIP, LS3MIP
Tier 2						
1%RAD	Idealised 1% per year CO ₂ only, RAD mode	C-driven	CO ₂ affects RAD	140	esm1pcrad	OCMIP, LS3MIP
SSP5-8.5	SSP5-8.5 extension to 2300	E-driven	Fully coupled	200	esmssp5-85ext	ScenarioMIP, LUMIP, OCMIP, LS3MIP, ISMIP
SSP5-8.5	Historical+SSP5-8.5 up to 2100 or 2300, BGC mode	E-driven	CO ₂ affects BGC	155 + 85 or 285	esmhstbgc, esmssp5-85bgc and esmssp5-85extbgc	ScenarioMIP, OCMIP, LS3MIP, DAMIP

To be discussed

All 1% CO₂ simulations (DECK and C4MIP). Need to define Nitrogen deposition forcing

C4MIP Tier 2.3 Biogeochemically-coupled version of Emission-driven CMIP6 Historical simulation and SSP5-8.5. Need to decide whether all radiative forcings or only CO₂ forcing are kept at pre-industrial level in the BGC runs.

➤ An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments

C4MIP for CMIP6 will have a strengthened focus on model evaluation against observation-based estimates of carbon quantities.

Using the emission-driven historical simulation from DECK:

- Model evaluation: coordinated top-down and bottom-up metrics of performance for key land and ocean quantities (as in Anav et al., 2013, Hoffman et al., 2014), in emission driven historical simulations: analysis of simulated atmospheric CO₂ and evaluation against long-term observations (eg. Mauna Loa)
- Use of emerging constraints based on carbon cycle interannual variability to constrain future projections (as in Cox et al., Nature 2013, Wenzel et al., JGR 2014) using the historical simulation and the CMIP1% simulations
- Quantify and explain changes since CMIP5 (show “demonstrable progress”)
- Link with WGCM/WGNE metric panel for essential carbon cycle variables and with Obs4MIP for observation datasets.

Using the CMIP 1% DECK and the 1% BGC, RAD simulations from C4MIP:

- Quantification of the strength of carbon-concentration and carbon-climate feedbacks terms of Friedlingstein et al. (JCLim. 2006) and their non-linearities (as in Gregory et al. JCLim. 2009), assessment of magnitude, uncertainty, and changes since CMIP5
- Assessment of the transient climate response to cumulative carbon emissions (TCRE), its magnitude, uncertainty and changes since CMIP5 as well as the impact of climate change on TCRE (using COU and BGC simulations).
- Quantification of response of natural CH₄ and N₂O emissions to climate and CO₂ changes

Using the SSP5-8.5 simulations

- In the emission driven case, these simulations will allow to quantify the effect of climate change (and land use change) on the global carbon cycle and atmospheric CO₂ and hence on the climate response (when compared to the SSP5-8.5 concentration driven simulation from ScenarioMIP)
- Quantification the uncertainty in simulated atmospheric CO₂ concentration (emissions-driven simulations).
- Analysis of TCRE based on SSP scenarios and its comparison with TCRE estimated from the idealized CMIP 1% simulations. Characterization of cumulative emissions allowed to likely stay below a given climate target (as in IPCC AR5).
- Analysis of changes in land and ocean carbon pools for future scenarios as a result CO₂ increase, climate change and of anthropogenic LUC (in coordination with LUMIP).

- Assessment of risk of longer term carbon release from permafrost, vegetation dieback, change in oceanic circulation and impact on ocean carbon sink for the extension up to 2300.
- For the SSP5-8.5, BGC mode: analysis of CO₂-carbon cycle feedbacks over the 21st century in a scenario world (as opposed to the idealized 1% world); also for assessment of CO₂ induced warming (by comparison with the fully coupled scenario run).

Potential interest for an overshoot scenario:

- Analysis of the feasibility of overshooting and returning to lower forcing in terms of capacity of the land and ocean to continue to act as carbon sinks, commitment/ irreversibility following the overshoot period.
- Assessment of ability of simple models (e.g. MAGICC) used in IAMs, to recreate ESM carbon-sink behavior in non-monotonic scenarios.

➤ Proposed timing*

We propose to ask modeling groups to provide results by end of 2016 from their COU and BGC 1% per year CO₂ simulations. The analyses of historical and scenarios will depend on when their results become available (to be coordinated with Historical CMIP6 and ScenarioMIP).

➤ **Synergies with other MIPS**

ScenarioMIP

ScenarioMIP will coordinate the concentration-driven scenario simulations. C4MIP will coordinate the emission-driven scenario simulations. This will allow investigating the impact of carbon cycle feedbacks on climate projections. It will hence confirm (or infirm) the CO₂ concentration pathways used in ScenarioMIP and provided by the IAM models.

LUMIP

Scenario simulations will include land-use change as a forcing. The analysis of its impact on land carbon cycle and climate system

OCMIP

Analysis of oceanic response in 1% and SSP scenarios will be done in collaboration with OCMIP.

LS3MIP

Analysis of land response in 1% and SSP scenarios will be done in collaboration with LS3MIP.

DAMIP

Analysis of the emission-driven historical run in fully coupled mode and BGC only mode will help detection and attribution of historical role of CO₂ emissions on climate and carbon cycle.

➤ **Data request**

Work in Progress (in collaboration with LUMIP and OCMIP)

Output request will certainly evolve from CMIP5, but we're not yet at a stage to offer a complete, revised list. Land-based MIPs (C4MIP, LUMIP, LS3MIP, TRENDY) should coordinate on changes/specifications needed for CMIP6; likewise for C4MIP, OCMIP and OMIP on marine variables.

- Note that requirements last time were not defined precisely enough and some groups reported wrongly/subtly different variables. Unlike some physical model outputs, BGC variables can be very model specific (PFTs, soil C pools etc), so we will need much better/clearer/more detailed explanation of how to process outputs to a common requirement. Suggest a GMD paper or similar to augment a CMIP overview document.

➤ **Model diagnostics and performance metrics for model evaluation**

- We propose to use iLAMB/ESMVal and other software evaluation packages for evaluation of the carbon cycle in the historical emissions forced simulations

- For each proposed experiment to be included in CMIP6**
 - the experimental design;
 - the science question and/or gap being addressed with this experiment;
 - possible synergies with other MIPs;
 - potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.
- If possible, a prioritization of the suggested experiments, including any rationale**
- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**
- List of output and process diagnostics for the CMIP DECK/CMIP6 data request**
 - whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;
 - whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);
 - whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.
 - whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;
 - the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.
- Any proposed contributions and recommendations for**
 - model diagnostics and performance metrics for model evaluation;
 - observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;
 - tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).
- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms.**
- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF.**

Cloud Feedback Model Intercomparison Project (CFMIP)

Application for CMIP6-Endorsed MIPs

*Mark Webb, Chris Bretherton, Sandrine Bony, Jen Kay, Steve Klein, Pier Siebesma,
Bjorn Stevens, George Tselioudis, Masahiro Watanabe,
Peter Good, Timothy Andrews, Roger Marchand, Robin Chadwick and Hervé Douville*

Updated 27th November 2014

The primary goal of CFMIP is to inform improved assessments of climate change cloud feedbacks. However, the CFMIP approach is increasingly also being used to understand other aspects of climate response, such as regional-scale precipitation and non-linear changes.

CFMIP started in 2003 and its first phase (CFMIP-1) organised an intercomparison based on perpetual July SST forced Cess style +2K experiments and 2xCO₂ equilibrium mixed-layer model experiments containing ISCCP simulator in parallel with CMIP3 (McAvaney and Le Treut, 2003). Results from CFMIP-1 had a substantial impact on the evaluation of clouds in models and in the identification of low level cloud feedbacks as the primary cause of inter-model spread in cloud feedback, and featured prominently in the fourth and fifth IPCC assessments.

The subsequent objective of CFMIP-2 was to inform improved assessments of climate change cloud feedbacks by providing better tools to support evaluation of clouds simulated by climate models and to understand cloud-climate feedback processes. CFMIP-2 organized further experiments as part of CMIP5, introducing seasonally varying SST perturbation experiments for the first time, as well as fixed SST CO₂ forcing experiments to examine cloud adjustments, and idealized ‘aquaplanet’ experiments to establish the contributions of land and zonally asymmetric circulations to cloud feedback uncertainties (Bony et al., 2011). CFMIP-2 also introduced satellite simulators to CMIP via the CFMIP Observation Simulator Package (COSP), not only the ISCCP simulator, but additional simulators to facilitate the quantitative evaluation clouds using a new generation of active RADARs and LIDARs in space. Additionally CFMIP-2 introduced into CMIP5 process diagnostics such as temperature and humidity budget tendency terms and high frequency ‘cfSites’ outputs at 120 locations around the globe. CFMIP also organized a joint project with the GEWEX Global Atmospheric System Study (GASS) called CGILS (the CFMIP-GASS Intercomparison of LES and SCMs) to develop cloud feedback intercomparison cases to assess the physical credibility of cloud feedbacks in climate models by comparing Single Column Models (SCM) versions of GCMs with high resolution Large Eddy Simulations (LES) models. Additionally CFMIP-2 developed the CFMIP-OBS data portal and the CFMIP diagnostic codes repository (see <http://www.cfmip.net> for more details).

Early studies arising from CFMIP-2 include numerous model evaluation studies using COSP, studies attributing cloud feedbacks and cloud adjustments to different cloud types, and the finding that idealized ‘aquaplanet’ experiments without land or Walker circulations are able to capture the essential differences between models’ global cloud feedbacks and cloud adjustments. Process outputs from CFMIP have also been used to develop and test physical mechanisms proposed to explain and constrain inter-model spread in cloud feedbacks in the CMIP5 models. CGILS has demonstrated a consensus in the responses of LES models to climate forcings and identified a number of shortcomings in the physical representations of cloud feedbacks in climate models. Additionally the CFMIP experiments have, due to their idealized nature, proven useful in a number of studies not directly related to clouds, but instead analyzing the responses of regional precipitation and circulation patterns to CO₂ forcing and climate change. Studies using CFMIP-2 outputs from CMIP5 remain ongoing and many further results are expected to feed into future assessments of the representation of clouds and cloud feedbacks in climate

models. For a list of publications arising from CFMIP-2, please refer to the CFMIP publications page at <http://www.cfmip.net>.

Given the previous record of CFMIP activities and the case outlined below we would like to request that CFMIP be endorsed as a CMIP6 project to continue support for community activities in this important area of research. We provide information on our plans for CFMIP-3 structured according to the provided criteria below.

Name of MIP: The Cloud Feedback Model Intercomparison Project (CFMIP)

Co-chairs: Mark Webb mark.webb@metoffice.gov.uk, Chris Bretherton breth@washington.edu

Members of the Scientific Steering Committee: Mark Webb (Met Office), Chris Bretherton (U. Washington), Sandrine Bony (IPSL), Jen Kay (CIRES), Steve Klein (PCMDI), Pier Siebesma (KNMI), Bjorn Stevens (MPG), George Tselioudis (NASA GISS), Masahiro Watanabe (U. Tokyo)

Link to website: <http://www.cfmip.net>

Goal of the MIP and a brief overview: The primary goal of CFMIP is to inform improved assessments of climate change cloud feedbacks. However, the CFMIP approach is increasingly being used to understand other aspects of climate response, such as circulation, regional-scale precipitation and non-linear changes. This involves bringing climate modelling, observational and process modelling communities closer together and providing better tools and community support for evaluation of clouds and cloud feedbacks simulated by climate models and for understanding of the mechanisms underlying them. This is to be achieved by:

- Ongoing organized coordinated model inter-comparison activities which include experimental design as well as specification of model output diagnostics to support quantitative evaluation of modelled clouds with observations (e.g. COSP) and in-situ measurements (e.g. cfSites) as well as process-based investigation of cloud maintenance and feedback mechanisms (e.g. cfSites, budget tendency terms, etc.)
- Ongoing development and improvement of COSP and CFMIP-OBS infrastructure.
- Ongoing collaboration with the cloud process modelling community (via GASS collaboration) on CGILS and via new efforts to develop a hierarchy of experiments linking GCMs with cloud resolving models (CRMs) and Large Eddy Simulation (LES) models run on large domains (e.g. via the IMPULSE project consortium).
- Organising annual meetings to provide a focus for community activities relevant to CFMIP and also to the broader community working to understand changes in clouds, circulation and precipitation which impact regional projections of climate change. (These two communities are increasingly becoming connected because the experiments designed for CFMIP are also useful in addressing a broader range of questions not directly related to clouds.)

References:

- Andrews, T., (2014), Using an AGCM to diagnose historical effective radiative forcing and mechanisms of recent decadal climate change. *J. Climate*, 27, 1193–1209, doi:10.1175/JCLI-D-13-00336.1.
- Bony, S., Webb, M., Bretherton, C. S., Klein, S. A., Siebesma, P., Tselioudis, G., & Zhang, M. (2011). CFMIP: Towards a better evaluation and understanding of clouds and cloud feedbacks in CMIP5 models. [Clivar Exchanges](#), 56(2), 20-22.
- Good, P., Andrews, T., Bouttes, N., Chadwick, R., Gregory, J. M., Lowe, J. A. (2014). The nonlinMIP intercomparison project: physical basis, experimental design and analysis principles. In preparation; (attached)
- McAvaney BJ, Le Treut H (2003) The cloud feedback intercomparison project: (CFMIP). In: CLIVAR Exchanges—supplementary contributions. 26: March 2003.

- Skinner, C.B., M. Ashfaq, and N.S. Diffenbaugh (2012). Influence of twenty-first-century atmospheric and sea surface temperature forcing on West African climate. *J. Climate*, 25, 527-542.
- Stevens B., Bony S., Frierson, D.M, Jakob, C., Kageyama, M., Pincus, R, Shepherd, T., Sherwood, S., Siebesma, A. P., Sobel, A., Watanabe, M., Webb, M.J. (2014). Clouds, Circulation and Climate Sensitivity: A Grand Science Challenge. [World Climate Research Programme Report No. 8/2014](#)

We argue below the CFMIP and its proposed experiments meet the requirements laid out by the CMIP panel, as outlined below.

1. CFMIP and its experiments directly address the key science questions of CMIP6. The question that CFMIP most directly addresses is 'How does the Earth system respond to forcing?' The CFMIP emphasis on understanding cloud feedbacks makes CFMIP highly relevant to this question. The next most relevant question is 'What are the origins and consequences of systematic model biases?' CFMIP has a strong model evaluation component via the use of satellite simulators, process diagnosis and comparisons with LES, and a proven track record in investigating the link between errors in cloud processes and cloud feedbacks. CFMIP is also relevant to the question 'How can we assess future climate changes given climate variability, climate predictability, and uncertainties in scenarios?' CFMIP will continue to supplement fully coupled CMIP experiments with idealised experiments that focus on basic understanding of the dominant uncertainties associated with cloud feedbacks. This will continue to support work which relates variability on observable timescales (e. g. seasonal to decadal) to longer term climate change responses (e.g. via 'emergent constraints'). For example the amipPiForcing experiment proposed below will support studies relating cloud variability on observable timescales to long term cloud feedbacks (Andrews, 2014).

Note also that the WCRP Grand Challenge on Clouds, Circulation and Climate is led by two CFMIP committee members (Bony and Stevens), and has three additional CFMIP committee members on its steering committee (Webb, Siebesma, Watanabe), including one of the CFMIP co-chairs. This puts CFMIP in an excellent position to directly address the questions arising from the WCRP Grand Challenge.

2. CFMIP builds on and connects to the shared CMIP DECK and CMIP6 historical experiments. The AMIP experiment is the control simulation for the CFMIP amip4K, amip4xCO2 and amipFuture experiments which were proposed by CFMIP for CMIP5 and which we would like to see continued in CMIP6 as Tier I experiments. The proposed Tier II experiments also connect to the AMIP DECK experiment; the AMIP preindustrial forcing experiment and amip minus 4K experiments also use the DECK AMIP experiment as a control. The abrupt +/- 4% solar constant experiments build on and contrast with the DECK abrupt4xCO2 experiment, as do the abrupt4xCO2 and abrupt0.5CO2 experiments. Additionally the atmosphere-only timeslice experiments build on the abrupt4xCO2 experiment, decomposing the regional response of each model's abrupt4xCO2 run into separate responses to each aspect of forcing and warming. Additionally CFMIP will propose additional process diagnostics and simulator outputs for the CMIP6 historical experiment, which will allow process based comparisons with the AMIP experiments to assess the impact of coupled SST errors on the simulation of clouds and regional precipitation patterns in the CMIP6 models.

3. CFMIP will continue to follow the CMIP modeling infrastructure standards and conventions, in terms of experimental design, data format and documentation. CFMIP-2 experiments were organized as part of CMIP5 and the CFMIP co-chairs have demonstrated the ability to follow all of the relevant standards in experimental protocols, in specification of diagnostic output requests, data formats and documentation. We commit to continuing in this spirit for CFMIP experiments which are coordinated through CMIP6.

4. All experiments are tiered, well-defined, and useful in a multi-model context and don't overlap with other CMIP6 experiments.

These are outlined below, and detailed specifications are provided in the accompanying spreadsheet. They are tiered into Tiers I and II. Additionally we give guidance on other experiments currently under development which we may propose as additional Tier II experiments in the future. Alternatively these additional experiments may be coordinated outside of CMIP.

These experiments are we believe useful in the multi-model context because the common purpose that they share is a focus on understanding the inter-model uncertainty/spread in cloud adjustments and cloud feedbacks as well as that in regional precipitation and circulation change and non-linear change. Investigation of inter-model requires multi-model analysis and hence all of these experiments are useful (and in fact require) a multi-model context. The usefulness of the Tier I experiments to a number of climate researchers has already been demonstrated by the large number of publications produced using CFMIP-2 experiments.

We have checked for overlaps with other CMIP6 experiments and are confident that links with other MIPS (e.g. nonLinMIP, GeoMIP, SolarMIP, ENSOMIP, RFMIP and PMIP) are based on complementary but non-overlapping experiments.

Summary of proposed experiments

Tier I Science questions, activities and experiments

1.1 Continuation of CFMIP-2 experiments - Lead coordinator: Mark Webb (Met Office)

Science Question: What are the physical mechanisms underlying the range of cloud feedbacks and cloud adjustments predicted by climate models, and which models have the most credible cloud feedbacks?

The CMIP5/CFMIP-2 experiments and diagnostic outputs have enabled considerable progress on these questions but participation by a larger fraction of modelling groups is required in CMIP6 for a more comprehensive assessment of the uncertainties across the full multi-model ensemble. Our proposal is essentially to retain the CFMIP-2/CMIP5 experiments in Tier I for CMIP6. The experiments to be retained are amip4K, amip4xCO₂, amipFuture, aquaControl, aqua4xCO₂ and aqua4K. These build on the amip DECK experiment. As the output requirements for the DECK and for CFMIP are not yet agreed, it is possible that the DECK AMIP experiment will not contain all of the output diagnostics required for CFMIP. In this event we may request an additional AMIP ensemble member including additional CFMIP diagnostics, both for model evaluation and for interpretation of feedbacks and adjustments in conjunction with other Tier I CFMIP experiments.

Tier II Science questions, activities and experiments

2.1 Abrupt +/-4% Solar Forced AOGCM experiments - Lead coordinators: Chris Bretherton (UW), Roger Marchand (UW), Bjorn Stevens(MPI)

Science Question: How do responses in the climate system due to changes in solar forcing differ from changes due to CO₂, and is the response sensitive to the sign of the solar forcing?

Rapid adjustments in clouds and precipitation are now recognized as significant components of models' responses to CO₂ forcing. While they can easily be separated from conventional feedbacks in SST forced experiments, such a separation in coupled models is complicated by various issues, including the response of the ocean on decadal timescales. A number of studies have examined cloud feedbacks in coupled models subject to a solar forcing, which is generally associated with much smaller cloud and precipitation adjustment, due to a smaller atmospheric absorption for a given top of atmosphere forcing.

Solar forcing also has a weaker impact on the stratosphere than CO₂, potentially resulting in different upper tropospheric meridional temperature gradients and storm track responses.

A +4% solar experiment would be equivalent to the abrupt4xCO₂ experiment but would increase the solar constant abruptly by 4 percent, resulting in a radiative forcing of a similar magnitude to that due to CO₂ quadrupling. This would provide a useful complement to the DECK abrupt4xCO₂ experiment, and would support our understanding of regional responses of the coupled system with and without CO₂ adjustments. A complementary -4% abrupt solar forcing experiment would allow the examination of feedback asymmetry under climate cooling, and would also help with the interpretation of model responses to geo-engineering scenarios and volcanic forcing, and relate to past climates.

2.2 Abrupt2xCO₂ and abrupt0.5xCO₂ Experiments (nonLinMIP) - Lead Coordinator Peter Good (Met Office Hadley Centre)

Science Question: To what extent is regional-scale climate change per CO₂ doubling state-dependent (nonlinear), and why? How does the balance of mechanisms differ for high-forcing compared to low-forcing scenarios or paleoclimate simulations?

To address this question we propose two new experiments for Tier II, abrupt2xCO₂ and abrupt0.5xCO₂, to explore global and regional-scale nonlinear responses, highlighting different behavior under business-as-usual scenarios, mitigation scenarios and paleoclimate simulations. Additional experiments may be proposed for Tier II in the future, or coordinated via CFMIP outside of CMIP6. These include 100-year extensions to abrupt4xCO₂ and abrupt2xCO₂; a 1% ramp-down from the end of the 1pctCO₂ experiment; an abrupt step-down to 1xCO₂ from year 100 of the abrupt4xCO₂. These would be used to explore longer-timescale responses, quantify nonlinear mechanisms more precisely and understand the reversibility of climate change.

2.3 amipMinus4K Experiment: Lead Coordinator: Mark Webb (Met Office)

Science Question: Are cloud feedbacks symmetric when subject to climate cooling rather than warming, and if not, why not?

An amipMinus4K experiment would take a similar form to the amip4K experiment, except that the sea surface temperatures would be uniformly reduced by 4K. This will be used to investigate asymmetric responses of clouds to a cooling climate in an idealized experiment, providing a link to PMIP. This experiment also complements the abrupt0.5xCO₂ and the -4% solar experiments in that one can identify asymmetries in the warming/cooling response with and without interactions with the ocean. This experiment has been proposed for CFMIP following discussions with PMIP representatives (Pacale Braconnot, Masa Kageyama, and Masakazu Yoshimori).

2.4 Feedbacks in AMIP experiments: Lead Coordinator: Tim Andrews (Met Office)

Science question: Are climate feedbacks during the 20th century different to those acting on long term climate change and climate sensitivity?

Experiment and rationale: The previous CFMIP design was unable to diagnose time-dependent feedbacks that potentially undermine the simple linear forcing-feedback paradigm and which may be relevant to the gap between observed and modeled estimates of climate sensitivity. To address this we propose an additional experiment called 'amipPiForcing' (amip pre-industrial forcing), which is exactly the same as the standard amip run (i.e. SSTs and sea-ice) but with constant pre-industrial forcings (i.e. all anthropogenic and natural forcing boundary conditions identical to the piControl run). Since the forcing constituents do not change in this experiment it readily allows a simple diagnosis of the

simulated atmospheric feedbacks to observed SST changes, which can then be compared to feedbacks representative of long term change and climate sensitivity (e.g. from abrupt4xCO₂ or amip4K). This has an advantage over the alternative approach of first estimating the forcing and adjustments (e.g. from RFMIP) and removing them from the amip experiment since the approach here only requires a single experiment (rather than pairs) which reduces the noise. The experiment has the additional benefit, by differencing with the standard amip run, of providing detailed information on the transient effective radiative forcing and adjustments in models relative to pre-industrial. The inclusion of CFMIP process diagnostics not available in the RFMIP experiments will also enable a deeper understanding of the factors underlying forcing and feedback differences in the present and future climate.

2.5 Timeslice experiments for understanding regional climate responses to CO₂ forcing. Co-ordinators: Rob Chadwick (Met Office) and Hervé Douville (CNRM)

Science questions:

- How do regional climate responses (of e.g. precipitation) in a coupled model arise from the combination of responses to different aspects of CO₂ forcing and warming (uniform SST warming, pattern SST warming, direct CO₂ effect, plant physiological effect)?
- Which aspects of forcing/warming are most important for causing inter-model uncertainty in regional climate projections?
- Can inter-model differences in regional projections be related to underlying structural or resolution differences between models through improved process understanding, and could this help us to constrain the range of regional projections?
- What impact do coupled model SST biases have on regional climate projections?

We propose a set of 6 20-year atmosphere-only timeslice experiments to decompose the regional responses of each model's abrupt4xCO₂ run into separate responses to each aspect of forcing and warming (uniform SST warming, pattern SST change, increased CO₂, plant physiological effect). As well as allowing regional responses in each individual model to be better understood, this set of experiments should prove especially useful for understanding the causes of model uncertainty in regional climate change.

The experiments are: 1) sstPi – the same as amip but with monthly-varying SSTs and sea-ice from years 101-120 of each model's own control run rather than observed fields; 2) sstPi4K – the same as sstPi but with SSTs uniformly increased by 4K; 3) sstPi4xCO₂ – the same as sstPi but CO₂ as seen by the radiation scheme is quadrupled; 4) sstPi4xCO₂Veg – the same as sstPi4xCO₂ but with the plant physiological response also able to respond to the increased CO₂; 5) sstPiFuture – the same as sstPi but a seasonally varying monthly mean climatology of the SST pattern anomaly taken from years 91-140 of each model's own abrupt4xCO₂ minus piControl is scaled to have a global mean increase of 4K and applied; 6) sstPiTot – the same as sstPiFuture but also with 4xCO₂ including the plant effect. sstPiTot is used to establish whether a timeslice experiment can adequately recreate the coupled abrupt4xCO₂ response in each model, and then forms the basis for a decomposition using the other experiments.

We also propose an additional amip based experiment, amipTot: the same as sstPiTot but with the SST pattern anomaly climatology from sstPiFuture added instead to the observed background SSTs and sea-ice (as for other amip experiments). Comparison of amipTot and sstPiTot should help to illuminate the impact of SST biases on regional climate responses in each model, and how this contributes to inter-model uncertainty. These experiments complement the amipFuture4xCO₂ experiment of ENSOmip (where a composite SST pattern is applied), and will allow the influence of different mean SST pattern change and background SSTs on the ENSO precipitation and circulation response to forcing to be examined.

2.6 Atmosphere-only experiments for understanding the role of cloud-radiative effects in the large-scale atmospheric circulation in current and perturbed climates. Co-ordinators: Sandrine Bony (IPSL) and Bjorn Stevens (MPI).

Science questions:

- How do cloud-radiative effects impact the structure, the strength and the variability of the general atmospheric circulation in the present-day climate?
- How much do cloud-radiative feedbacks contribute to the spread of circulation and precipitation responses in climate change?
- Can we identify robust aspects of the climate response to global warming that do not depend on cloud-radiative feedbacks?

It is increasingly recognized that clouds, and cloud-radiative effects in particular, play a critical role in the general circulation of the atmosphere (ITCZ, MJO, storm tracks, hurricanes) and its response to global warming. A better assessment of this role would greatly help interpret model biases (how much do biases in cloud-radiative properties contribute to biases in the structure of the ITCZ, in the position and strength of the storm tracks, in the lack of intra-seasonal variability, etc) and to inter-model differences in simulations of the current climate and in climate change projections (especially changes in regional precipitation and extreme events). More generally, a better understanding of how clouds couple to circulation is expected to improve our ability to answer two of the four science questions raised by the WCRP Grand Challenge on Clouds, Circulation and Climate Sensitivity: what controls the position, the strength, and the variability of the storm tracks and of the tropical rainbelts?

These questions provided the scientific motivation for the Clouds On/Off Klima Intercomparison Experiment (COOKIE) project proposed by the European consortium EUCLIPSE and CFMIP in 2012. The COOKIE experiments, which have been run by 4 to 8 climate models (depending on the experiment), consisted in switching off the cloud-radiative effects (clouds seen by the radiation code - and the radiation code only- were artificially made transparent) in an atmospheric model forced by prescribed SSTs. By doing so, the atmospheric circulation could feel the lack of cloud-radiative heating within the atmosphere, but the land surface could also feel the lack of cloud shading, which led to changes in land-sea contrasts. The change in circulation between On and Off experiments was resulting from both effects, obscuring a bit the mechanisms through which the atmospheric cloud-radiative effects interact with the circulation for given surface boundary conditions. As the LW cloud-radiative effects are felt mostly within the troposphere (and represent most of the LW+SW cloud-radiative heating) while the SW effects are felt mostly at the surface, we could better isolate the role of tropospheric cloud-radiative effects on the circulation by running atmosphere-only experiments in which clouds are made transparent to radiation only in the LW.

We propose in Tier II a set of simple experiments similar to the amip, amip4K, aquaControl and aqua4K experiments of CMIP5/CFMIP2 (and Tier 1 of CMIP6) but in which cloud-radiative effects are switched off in the LW part of the radiation code. These experiments will be referred to as offlwamip, offlwamip4K, offlwaquaControl and offlwaqua4K. The analysis of idealized (aqua-planet) experiments will allow us to assess the robustness of the impacts found in more realistic (AMIP) configurations. It will also facilitate the interpretation of the results using simple dynamical models or theories, in collaboration with large-scale dynamicists (e.g. DynVar). The comparison of the inter-model spread of simulations between AMIP and offlwAMIP experiments for present-day and globally warmer climates will help identify which aspects of the spread depend on the representation of cloud-radiative effects, and which aspects do not, thus better highlighting other sources of spread.

Additional CFMIP experiments under consideration for the future

We also propose to use these CMIP6 experiments as the foundation for further experiments planned in the context of the Grand Challenge on Clouds, Circulation and Climate Sensitivity. These will include for example sensitivity experiments to assess the impacts of different physical processes on cloud feedbacks and regional circulation/precipitation responses, and others designed to test specifically proposed cloud feedback mechanisms. Additional experiments further idealizing the aquaplanet

framework to a non-rotating rotationally symmetric case are also under development. These will be proposed as additional Tier II experiments at a future time, or coordinated by CFMIP outside of CMIP6.

5. Unless a Tier I experiment differs only slightly from another well-established experiment, it must already have been performed by more than one modeling group. All of the proposed Tier I experiments were previously included in CMIP5 and so are well established and already performed by multiple groups.

6. A sufficient number of modelling centers (~8) are committed to performing all of CFMIP's Tier I experiments and providing all the requested diagnostics needed to answer at least one of its science questions. Fourteen modeling groups have so far agreed to participate in CFMIP as part of CMIP6, implying that they are prepared to perform the Tier I experiments. These are ACCESS (Australia), BCC (China), CanESM (Canada), CESM (USA), CNRM (France), FGOALS (China), GFDL (USA), IPSL-ESM (France), MIROC6-GCM (Japan) NICAM (Japan), MPI-ESM (Germany), MRI (Japan) and UKESM (United Kingdom).

7. The MIP presents an analysis plan describing how it will use all proposed experiments, any relevant observations, and specially requested model output to evaluate the models and address its science questions. Our analysis plan is outlined below.

We commit to contributing to the creation of the CMIP6 data request and to analyzing the data, as we did for CMIP5. This will include making proposals for an updated COSP request in CMIP6 (see the proposal from the COSP PMC), and also additional improvements to the CFMIP diagnostic specifications relating to temperature and humidity budget increments, 3D radiative fluxes, inclusion of aerosol diagnostics across CFMIP experiments, and the introduction of a more representative land locations in the cfSites specification (details to follow).

We also commit to identifying observations needed for model evaluation and improved process understanding, and to contributing directly to making such datasets available as part of obs4MIPs. For example the CFMIP community has up to now played a central role in providing versions of CloudSat and CALIPSO datasets designed for direct comparison with CMIP5 data through the CFMIP-OBS website (see <http://climserv.ipsl.polytechnique.fr/cfmip-obs/>) and part of this work has recently involved publishing this data via the ESG and linking into obs4MIPS (see for example references to CFMIP-OBS on the obs4MIPS website at <https://www.earthsystemcog.org/projects/obs4mips/aboutus>). This work will continue.

CFMIP analysis activities are ongoing and the CFMIP community is ready to analyse CMIP6 data at any time. We would like modelling groups to perform the proposed CFMIP/CMIP6 experiments at the same time or shortly after their DECK experiments. Subsequent CFMIP experiments which are not included in CMIP6 will build on the proposed DECK and CMIP6/CFMIP experiments and some will start as soon as CMIP6 DECK experiments start to become available. We envisage a succession of CFMIP related intercomparisons addressing different questions arising from the GC spanning the duration of CMIP6.

We commit to scientifically analyze, evaluate and exploit the proposed experiments, and have identified leads within CFMIP for different aspects of this activity. An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments follows:

- CFMIP will continue to exploit the CMIP DECK and CMIP6 experiments to understand and evaluate cloud processes and cloud feedbacks in climate models. The wide range of analysis activities described above in the context of CFMIP-2 will be continued in CFMIP-3 using the CMIP DECK and CMIP6 experiments, allowing the techniques developed in CFMIP-2 to applied to an expanding number of models, including the new generation of models currently

under development. These activities will include evaluation of clouds using additional simulators (see attached proposal regarding COSP), investigation of cloud processes and cloud feedback/adjustment mechanisms using process outputs (cfSites, budget tendency terms, etc). The inclusion of COSP and budget tendency terms in additional DECK experiments (e.g. abrupt4xCO2 and some scenario experiments, also see attached proposal for COSP) will enable the CFMIP approach to be applied to a wider range of experimental configurations. (Lead coordinator Mark Webb).

- Analysis of the +/-4% solar model runs would include an evaluation of both rapid adjustments and longer-term responses on global and regional top-of-atmosphere radiative fluxes, cloud types (using ISCCP and other COSP simulators) and precipitation characteristics, as well as comparison of these responses with responses in DECK abrupt4xCO2 experiments. GeoMIP and SolarMIP have expressed a strong interest in these CFMIP experiments and joint analysis of these CFMIP experiments with GeoMIP and SolarMIP experiments is anticipated, specifically with the goal of determining to what degree results from abrupt solar forcing ONLY experiments and abrupt CO2 ONLY experiments can be used to predict what happens when both forcing are applied simultaneously, as done in the GeoMIP experiments (Lead coordinator Chris Bretherton).
- Analysis of nonlinear climate processes will primarily involve comparing the abrupt4xCO2, abrupt2xCO2 and abrupt0.5xCO2 experiments over the same timescale (Good et al., 2014). (Lead coordinator Peter Good).
- Analysis of amipPiForcing has already been done in detail for a single model in Andrews (2014). We propose to use this as a starting point for a multi-model analysis. (Lead coordinator Timothy Andrews).
- An overview analysis of regional responses and model uncertainty in the timeslice and amipTot experiments will be carried out by the co-ordinators, in collaboration with members of contributing modeling groups. We anticipate that further detailed analysis on the processes at work in different regions will be carried out by a variety of research groups with interest and expertise in a particular region: for example a set of similar experiments has previously been used to examine the climate response of the West African monsoon in CCSM3 (Skinner et al. 2012). The timeslice and amipTot experiments have already been successfully run with HadGEM2 (Met Office), and are currently in the planning stage for CNRM. (Lead coordinator Robin Chadwick).
- When analyzed together with the amip4K experiment, the amipMinus4K experiment allows one to exploit the CFMIP process diagnostics to understand for asymmetries in the climate response to warming and cooling which have been noted in PMIP experiments. These might arise from cloud phase responses in middle- and high-latitude clouds or from the adiabatic cloud liquid water path response feedback which is important over land regions and which would be expected to be weaker with cooling because of the non-linearity in the Clausius-Clapeyron relation. (Lead coordinator Mark Webb).

8. The MIP has completed the MIP template questionnaire. We have done this.

9. The MIP contributes a paper on its experimental design to the CMIP6 Special Issue. We agree to do this.

10. The MIP considers reporting on the results by co-authoring a paper with the modelling groups. We agree to do this. Separate papers will be prepared for each of the experiment groups proposed.

Answers to other questions in the MIP template questionnaire

All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale. We have no objection to this.

List of output and process diagnostics for the CMIP DECK/CMIP6 data request. To be provided at a later date.

Any proposed contributions and recommendations for model diagnostics and performance metrics, observations/reanalysis data products, tools, code or scripts. To be provided at a later date.

Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms. None expected.

Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF. None expected.

The nonlinMIP intercomparison project: physical basis, experimental design and analysis principles

P. Good¹, T. Andrews¹, N. Bouttes², R. Chadwick¹, J. M. Gregory^{2,1}, J. A. Lowe¹

[1]{ Met Office Hadley Centre, Exeter, United Kingdom}

[2]{ NCAS-Climate, University of Reading, Reading, United Kingdom}

Correspondence to: P. Good (peter.good@metoffice.gov.uk)

Abstract

nonlinMIP aims to quantify and understand, at regional scales, climate responses that are non-linear under CO₂ forcing (mechanisms for which doubling the CO₂ forcing does not double the response). Non-linear responses can be large at regional scales, with important implications for understanding mechanisms and for GCM emulation techniques (e.g. energy balance models and pattern-scaling methods). However, these processes are hard to explore using traditional experiments, explaining why they have had little attention in previous studies. Some single model studies have established novel analysis principles and some physical mechanisms. There is now a need to explore robustness and uncertainty in such mechanisms across a range of models.

nonlinMIP addresses this using a simple, small set of CO₂-forced experiments that are able to separate linear and non-linear mechanisms cleanly, with good signal/noise – while being demonstrably traceable to realistic transient scenarios. The design builds on the CMIP5 and CMIP6 DECK protocols, and is centred around a suite of abruptCO₂ experiments, with a ramp-up-ramp-down experiment to test traceability to gradual forcing scenarios. The understanding gained will help interpret the spread in policy-relevant scenario projections.

Here we outline the basic physical principles behind nonlinMIP, and the method of establishing traceability from abruptCO₂ to gradual forcing experiments, before detailing the experimental design and finally some analysis principles. The discussion on traceability of abruptCO₂ to transient experiments is also relevant to the abrupt4xCO₂ experiment in the CMIP5 and CMIP6 DECK protocols.

1 Introduction

Climate impacts assessments require, at regional scales, understanding of physical mechanisms of climate change in GCM projections. Also required is the ability to emulate (using fast simplified climate models) GCM behaviour for a much larger range of policy-relevant scenarios than may be evaluated using GCMs directly. These two requirements may be combined into a single question: what is the simplest conceptual framework that has quantitative predictive power and captures the key mechanisms behind GCM scenario projections?

Often, a pragmatic choice has been to assume some form of linearity. In studies of the global energy balance, linearity is often assumed in the form of a constant climate feedback parameter. This parameter may be used to quantify feedbacks in different models (e.g. Zelinka et al., 2013) or, in emulation methods, to parameterise global energy balance models (e.g. Huntingford and Cox, 2000). In understanding or emulating regional patterns of climate change, it is often assumed that regional climate change is roughly proportional to global mean warming. In emulation work, this is termed 'pattern scaling' (Mitchell, 2003; Santer et al., 1990; Tebaldi and Arblaster, 2014), but this assumption may also be applied either explicitly or implicitly in understanding mechanisms. Sometimes, patterns of change per K of global warming are quantified; often, physical mechanisms are studied for a single period of a single forcing scenario (implicitly assuming that the understanding is relevant for other periods or scenarios).

While these approximations appear to work well under some circumstances, significant limitations are increasingly being revealed in such assumptions. These are of two types: different timescales of response, and non-linear responses. In discussing this, a complication arises in that different linearity assumptions exist. Henceforth we define 'linear' as meaning 'consistent with linear systems theory' - i.e. responses that are linear in model forcing (i.e. where doubling the forcing doubles the response; this is different from assuming that pairs of responses are linearly related to each other – as in pattern scaling).

Even in a linear system (where responses are linear in forcing), the relationship between two system outputs (e.g. between global-mean temperature and regional sea surface temperature - SST) will in general be non-linear. This is due to different timescales of response in different locations and/or variables. Examples include lagged surface ocean warming due to a connection with the deeper ocean (Chadwick et al., 2013; Held et al., 2010; Williams et al., 2008; Manabe et al., 1990; Andrews and Ringer, 2014) or the direct response of precipitation to forcings (Andrews et al., 2010; Allen and Ingram, 2002; Mitchell et al., 1987). One (generally false) assumption of pattern scaling, then, is that regional climate responds over the same

timescale as global-mean temperature. Different timescales of response are especially important in understanding and predicting behaviour under mitigation and geoengineering scenarios (or over very long timescales).

Non-linear system responses (e.g. Schaller et al., 2013) are more complex to quantify, understand and predict than those of linear systems. Some examples have been known for some time, such as changing feedbacks through retreating snow/sea-ice (Colman and McAvaney, 2009; Jonko et al., 2013), or the Atlantic Meridional Overturning Circulation. More recently, substantial non-linear precipitation responses have been demonstrated in spatial patterns of regional precipitation change in two Hadley Centre climate models with different atmospheric formulations (Good et al., 2012; Chadwick and Good, 2013). This is largely due to simultaneous changes in pairs of known robust pseudo-linear mechanisms (Chadwick and Good, 2013). Non-linearity has also been demonstrated in the response under idealised geoengineering scenarios, of ocean heat uptake, sea-level rise, and regional climate patterns, with different behaviour found when forcings are decreasing than when they are increasing (Bouttes et al., 2013; Schaller et al., 2014).

Investigation of these mechanisms at regional scales has been constrained by the type of GCM experiment typically analysed. Most previous analyses (e.g. Solomon et al., 2007) have used results from transient forcing experiments, where forcing changes steadily through the experiment. There are three main problems with this approach. First, information about different timescales of response is masked. This is because the GCM response at any given time in a transient forcing experiment is a mixture of different timescales of response (Good et al., 2013; Held et al., 2010; Li and Jarvis, 2009), including short-timescale responses (e.g. ocean mixed layer response from forcing change over the previous few years) through long-timescale behaviour (including deeper ocean responses from forcing changes multiple decades to centuries earlier). Secondly, in transient forcing experiments, non-linear behaviour is hard to separate from linear mechanisms. For example, in an experiment where CO₂ is increased by 1% per year for 140 years ('1pctCO₂'), we might find different spatial patterns at year 70 (at 2xCO₂) than at year 140 (at 4xCO₂). This could be due to nonlinear mechanisms (due to the different forcing level and associated different climate state). However, it could also be due to linear mechanisms: year 140 follows 140 years of forcing increase, so includes responses over longer response timescales than at year 70 (only 70 years of forcing increase). Thirdly, signal/noise ratios of regional climate change can be relatively poor in such experiments.

These three issues may be addressed by the use of idealised abruptCO₂ GCM experiments (Forster et al., 2012; Zelinka et al., 2013; Jonko et al., 2013; Good et al., 2013; Good et al., 2012; Chadwick and Good, 2013; Chadwick et al., 2013; Bouttes et al., 2013; Gregory et al., 2004): an experiment where CO₂ forcing is abruptly changed, then held constant. In abrupt CO₂ experiments, responses over different timescales are separated from each other. Further, responses at different forcing levels may be directly compared, e.g. by comparing the response in abrupt2xCO₂ and abrupt4xCO₂ experiments over the same timescale - both have identical forcing time histories, apart from the larger forcing magnitude in abrupt4xCO₂. Thirdly, high signal/noise is possible: averages may be taken over periods of 100 years or more (after the initial ocean mixed layer adjustment, change is gradual in such experiments). Recent work (Good et al., 2011; Good et al., 2012; Good et al., 2013; Zelinka et al., 2013) has established that these experiments contain global and regional-scale information quantitatively traceable to more policy-relevant transient experiments - and equivalently, that they form the basis for fast simple climate model projections traceable to the GCMs.

The CMIP5 abrupt4xCO₂ experiments have thus been used widely: including quantifying GCM forcing and feedback behaviour (Gregory et al., 2004; Zelinka et al., 2013), and for traceable emulation of GCM projections of global-mean temperature and heat uptake (Good et al., 2013; Stott et al., 2013). Abrupt4xCO₂ is also part of the CMIP6 DECK protocol (Meehl et al., 2014).

NonlinMIP extends the CMIP5 and CMIP6 DECK designs to explore non-linear responses (via additional abruptCO₂ experiments at different forcing levels. It also explores responses over slightly longer timescales (extending the CMIP5 abrupt4xCO₂ experiment by 100 years).

2 Relating abruptCO₂ to gradual forcing scenarios: the step-response model

In using the highly-idealised abruptCO₂ experiments, it is essential that their physical relevance (traceability) to more realistic gradual forcing experiments is determined. Some GCMs could respond unrealistically to the abrupt forcing change. A key tool here is the step-response model (described below). This response-function method aims to predict the GCM response to any given transient-forcing experiment, using the GCM response to an abruptCO₂ experiment. Such a prediction may be compared with the GCM transient-forcing simulation, as part of a traceability assessment (discussed in detail in section 5).

Once some confidence is established in traceability of the abruptCO₂ experiments to transient-forcing scenarios, the step-response model has other roles: to explore the implications, for different forcing scenarios, of physical understanding gleaned from abruptCO₂ experiments; to help separate linear and nonlinear mechanisms (section 5); and potentially as a basis for GCM emulation. The method description below also serves to illustrate the assumptions of linear system theory.

The step-response model represents the evolution of radiative forcing in a scenario experiment by a series of step changes in radiative forcing (with one step taken at the beginning of each year). The method makes two linear assumptions. First, the response to each annual forcing step is estimated by linearly scaling the response in a CO₂ step experiment according to the

magnitude of radiative forcing change. Second, the response y_i at year i of a scenario experiment is estimated as a sum of responses to all previous annual forcing changes (see Figure 1 of Good et al., 2013 for an illustration):

$$y_i = \sum_{j=0}^i w_{i-j} x_j \quad (1a)$$

where x_j is the response of the same variable in year j of the CO₂ step experiment. w_{i-j} scales down the response from the step experiment (x_j) to match the annual step change in radiative forcing from year i to year j of the scenario (denoted ΔF_{i-j}):

$$w_{i-j} = \frac{\Delta F_{i-j}}{\Delta F_s} \quad (1b)$$

where ΔF_s is the radiative forcing change in the CO₂ step experiment. All quantities are expressed as anomalies with respect to a constant-forcing control experiment.

This approach can in principle be applied at any spatial scale for any variable for which the assumptions are plausible (e.g. Chadwick et al., 2013).

3 Linear and non-linear mechanisms, and the relevance of abruptCO2 experiments

Here we discuss further, with examples, the distinction between linear and nonlinear mechanisms, when they are important, and the relevance of abruptCO2 experiments.

3.1 Linear mechanisms: different timescales of response

Even in a linear system, regional climate change per K of global warming will evolve during a scenario simulation. This happens because different parts of the climate system have different timescales of response to forcing change.

This may be due to different effective heat capacities. For example, the ocean mixed layer responds much faster than the deeper ocean, simply due to a thinner column of water (Li and Jarvis, 2009). However, some areas of the ocean surface (e.g. the Southern Ocean and south-east subtropical Pacific) show lagged warming, due to a greater connection (via upwelling or mixing) with the deeper ocean (e.g. Manabe et al., 1990; Williams et al., 2008). The dynamics of the ocean circulation and vegetation may also have their own inherent timescales (e.g. vegetation change may lag global warming by years to hundreds of years, Jones et al., 2009). At the other extreme, some responses to CO₂ forcing are much faster than global warming: such as the direct response of global mean precipitation to forcings (Allen and Ingram, 2002; e.g. Andrews et al., 2010; Mitchell et al., 1987) and the physiological response of vegetation to CO₂ (Field et al., 1995).

In a linear system, patterns of change per K of global warming are sensitive to the forcing history. For example in Figure 1, a scenario is illustrated where forcing is ramped up, then stabilized. Three periods are highlighted, which may have different patterns of change per K of global warming, due to different forcing histories: at the leftmost point, faster responses will be relatively more important, whereas at the right, the slower responses have had some time to catch up. This is illustrated in Figure 2 for sea-level rise. The blue curves show that for RCP2.6, global-mean warming ceases after 2050, while sea-level rise continues at roughly the same rate throughout the century. This is largely because deep ocean heat uptake is much slower than ocean mixed-layer warming.

By design, abruptCO2 experiments separate different timescales of GCM response to forcing change. This is used, for example, (Gregory et al., 2004) to estimate radiative forcing and feedback parameters for GCMs: plotting radiative flux anomalies against global mean warming can separate 'fast' and 'slow' responses (see e.g. Figure 3).

3.2 Non-linear responses

Nonlinear mechanisms arise for a variety of reasons. Often, however, it is useful to describe them as state-dependent feedbacks. For example, the snow-albedo feedback becomes small at high or low snow depth. Sometimes, nonlinear mechanisms may be better viewed as simultaneous changes in pairs of properties. For example, convective precipitation is broadly a product of moisture content and dynamics (Chadwick and Good, 2013; Chadwick et al., 2012). Both moisture content and atmospheric dynamics respond to CO₂ forcing, so in general we might expect convective precipitation to have a nonlinear response to CO₂ forcing. Of course, more complex nonlinear responses exist, such as for the Atlantic Meridional Overturning Circulation.

In contrast to linear mechanisms, nonlinear mechanisms are sensitive to the magnitude of forcing. For example, the two points highlighted in Figure 4 may have different patterns of change per K of global warming, due to nonlinear mechanisms.

An example is given in Figure 5, which shows the albedo feedback declining with increased global temperature, due to declining snow and ice cover, and the remaining snow and ice being in areas of lower solar insolation (Colman and McAvaney, 2009).

AbruptCO₂ experiments may be used to separate nonlinear from linear mechanisms. This can be done by comparing the responses at the same timescale in different different abruptCO₂ experiments. Figure 6 compares abrupt2xCO₂ and abrupt4xCO₂ experiments over years 50-149. A 'doubling difference' is defined, measuring the difference in response to the first and second CO₂ doublings. In most current simple climate models (e.g. Meinshausen et al., 2011), the radiative forcing from each successive CO₂ doubling is assumed identical (because forcing is approximately linear in log[CO₂], Myhre et al., 1998). With this assumption, a linear system would have zero doubling difference everywhere. Therefore, the doubling difference is used as a measure of nonlinearity. The question of which abruptCO₂ experiments to compare, and over which timescale, is discussed in section 5.

In some GCMs, the forcing per CO₂ doubling has been shown to vary with CO₂ (Colman and McAvaney, 2009; Jonko et al., 2013). However, this variation depends on the specific definition of forcing used (Jonko et al., 2013). Currently this is folded into our definition of nonlinearity. If a robust definition of this forcing variation becomes available in future, it could be used to scale out any difference in forcing between pairs of abruptCO₂ experiments, to calculate an 'adjusted doubling difference'.

As an example, Figure 7 maps the response to abrupt2xCO₂ and abrupt4xCO₂, and the doubling difference, for precipitation in HadGEM2-ES over the ocean (taken from Chadwick and Good). The nonlinearities are large - comparable in magnitude to the responses to abrupt2xCO₂, albeit with a different spatial pattern.

4 Experimental design

nonlinMIP is composed of a set of abruptCO₂ experiments (the primary tools), plus a CO₂-forced transient experiment. These build on the CMIP5 and CMIP6 DECK protocols (the required runs from these are detailed in Table 1). The additional nonlinMIP runs (Table 2) are assigned three priority levels. Three options for participation are: 1) only the 'essential' simulation; 2) all 'high priority' plus the 'essential' simulations; or, preferably, 3) all simulations. The experiments in Table 1 are required in all cases. All experiments must be initialized from the same year of a pre-industrial control experiment, except for abrupt4xto1x (see Table 2). A typical analysis procedure is outlined in section 5.

The nonlinMIP design is presently limited to CO₂ forcing, although the same principles could be applied to other forcings.

5 Basic analysis principles

This section outlines the general principles behind analysis of nonlinMIP results. The primary idea is to find where the step-response model (section 2) breaks: since the step-response model is based on a linear assumption, this amounts to detecting non-linear responses.

The aim is to focus subsequent analysis. If non-linearities in a quantity of interest are found to be small, then analysis may focus on understanding different timescales of response from a single abruptCO₂ experiment: linearity means that the physical response (over a useful range of CO₂ concentrations) is captured by a single abruptCO₂ experiment. This represents a considerable simplification. If, on the other hand, non-linearities are found to be important, the focus shifts to understanding the different responses in different abruptCO₂ experiments. The choice of which abruptCO₂ experiments to focus on, and over which timescales, is discussed below.

5.1 First step: check basic traceability of abrupt4xCO₂ to the transient-forced response near 4xCO₂

This is to confirm that the abruptCO₂ experiments contain realistic physical responses in the variables of interest (as previously done for global-mean temperature and heat uptake for a range of CMIP5 models (Good et al., 2013), and for other global-mean quantities for HadCM3 (Good et al., 2011). This also, rules out the most pathological non-linearities (e.g. if the response to an abrupt CO₂ change in a given GCM was unrealistic).

The linear step-response model should first be used with the abrupt4xCO₂ response, to predict the response near year 140 of the 1pctCO₂ experiment (i.e. near 4xCO₂). This prediction is then compared with the actual GCM 1pctCO₂ result. This should first be done for global mean temperature: this assessment has been performed for a range of CMIP5 models (Good et al., 2013; see Figure 8), giving an idea of the level of accuracy expected. If the abruptCO₂ response is fundamentally

unrealistic, it is likely to show up in the global temperature change. This approach may then be repeated for spatial patterns of warming, and then for the quantities of interest. Abrupt4xCO₂ is used here as it has larger signal/noise than abrupt2xCO₂, yet is representative of forcing levels in a business-as-usual scenario by 2100. However, the tests may also be repeated using abrupt2xCO₂ – but compared with year 70 of the 1pctCO₂ experiment (i.e. at 2xCO₂).

The step-response model emulation under these conditions should perform well for most cases: the state at year 140 of the 1pctCO₂ experiment is very similar to that of abrupt4xCO₂ (same forcing, similar global-mean temperature), so errors from non-linear mechanisms should be minimal. If large errors are found, this may imply caution about the use of abruptCO₂ experiments for these variables, or perhaps point to novel non-linear mechanisms that may be understood by further analysis.

5.2 Second step: detecting nonlinear responses

Having established some level of confidence in the abruptCO₂ physical response, the second step is to look for nonlinear responses. This first involves repeating the tests from step 1 above, but for different parts of the 1pctCO₂ and 1pctCO₂ ramp-down experiments, and using different abruptCO₂ experiments for the step-response model.

An example is given in Figure 9 (but for different transient-forcing experiments). This shows results for global-mean precipitation in the HadCM3 GCM (Good et al., 2012). Here, the step-response model prediction using abrupt4xCO₂ (red curves) only works where a transient-forced experiment is near to 4xCO₂. Similarly, the prediction using abrupt2xCO₂ (blue curves) works only near 2xCO₂. Otherwise, quite large errors are seen, and the predictions with abrupt2xCO₂ and abrupt4xCO₂ are quite different from each other. This implies that there are large non-linearities in the precipitation response in this GCM, and that they may be studied by comparing the responses in the abrupt2xCO₂ and abrupt4xCO₂ experiments.

Having identified some non-linear response, and highlighted two or more abruptCO₂ experiments to compare (in the previous example, abrupt2xCO₂ and abrupt4xCO₂), the non-linear mechanisms may be studied in detail by comparing the responses in the different abruptCO₂ experiments over the same timescale (e.g. via the doubling difference, as in Figures 6,7). This allows (Good et al., 2012; Chadwick and Good, 2013) non-linear mechanisms to be separated from linear mechanisms (not possible in a transient-forcing experiment).

6 Conclusions

This paper outlines the basic physical principles behind the nonlinMIP design, and the method of establishing traceability from abruptCO₂ to gradual forcing experiments, before detailing the experimental design and finally some general analysis principles that should apply to most studies based on this dataset.

Acknowledgements

This work was supported by the Joint UK DECC/Defra Met Office Hadley Centre Climate Programme (GA01101). Nathaëlle Bouttes received funding from the European Research Council under the European Community's Seventh Framework Programme (FP7/2007-2013), ERC Grant Agreement 247220, project "Seachange."

References

- Allen, M. R., and Ingram, W. J.: Constraints on future changes in climate and the hydrologic cycle, *Nature*, 419, 224-+, 10.1038/nature01092, 2002.
- Andrews, T., Forster, P. M., Boucher, O., Bellouin, N., and Jones, A.: Precipitation, radiative forcing and global temperature change, *Geophysical Research Letters*, 37, Artn L14701
Doi 10.1029/2010gl043991, 2010.
- Andrews, T., and Ringer, M. A.: Cloud feedbacks, rapid adjustments, and the forcing-response relationship in a transient co2 reversibility scenario, *Journal of Climate*, 27, 1799-1818, Doi 10.1175/Jcli-D-13-00421.1, 2014.
- Bouttes, N., Gregory, J. M., and Lowe, J. A.: The reversibility of sea level rise, *Journal of Climate*, 26, 2502-2513, Doi 10.1175/Jcli-D-12-00285.1, 2013.
- Chadwick, R., Boutle, I., and Martin, G.: Spatial patterns of precipitation change in cmip5: Why the rich don't get richer., *Journal of Climate*, accepted, 2012.
- Chadwick, R., and Good, P.: Understanding non-linear tropical precipitation responses to co2 forcing, *Geophysical Research Letters*, 40, 10.1002/grl.50932, 2013.
- Chadwick, R., Wu, P. L., Good, P., and Andrews, T.: Asymmetries in tropical rainfall and circulation patterns in idealised co2 removal experiments, *Climate Dynamics*, 40, 295-316, DOI 10.1007/s00382-012-1287-2, 2013.
- Colman, R., and McAvaney, B.: Climate feedbacks under a very broad range of forcing, *Geophysical Research Letters*, 36, L01702
10.1029/2008gl036268, 2009.
- Field, C. B., Jackson, R. B., and Mooney, H. A.: Stomatal responses to increased co2 - implications from the plant to the global-scale, *Plant Cell Environ*, 18, 1214-1225, DOI 10.1111/j.1365-3040.1995.tb00630.x, 1995.
- Forster, P. M., Andrews, T., Good, P., Gregory, J. M., Jackson, L., and Zelinka, M. D.: Evaluating adjusted forcing and model spread for historical and future scenarios in the cmip5 generation of climate models, *Journal of Geophysical Research- Atmospheres* (accepted pending minor revisions), 2012.
- Good, P., Gregory, J. M., and Lowe, J. A.: A step-response simple climate model to reconstruct and interpret aogcm projections, *Geophysical Research Letters*, 38, Artn L01703
Doi 10.1029/2010gl045208, 2011.
- Good, P., Ingram, W., Lambert, F. H., Lowe, J. A., Gregory, J. M., Webb, M. J., Ringer, M. A., and Wu, P. L.: A step-response approach for predicting and understanding non-linear precipitation changes, *Climate Dynamics*, 39, 2789-2803, DOI 10.1007/s00382-012-1571-1, 2012.
- Good, P., Gregory, J. M., Lowe, J. A., and Andrews, T.: Abrupt co2 experiments as tools for predicting and understanding cmip5 representative concentration pathway projections, *Climate Dynamics*, 40, 1041-1053, DOI 10.1007/s00382-012-1410-4, 2013.
- Gregory, J. M., Ingram, W. J., Palmer, M. A., Jones, G. S., Stott, P. A., Thorpe, R. B., Lowe, J. A., Johns, T. C., and Williams, K. D.: A new method for diagnosing radiative forcing and climate sensitivity, *Geophysical Research Letters*, 31, L03205
10.1029/2003gl018747, 2004.
- Held, I. M., Winton, M., Takahashi, K., Delworth, T., Zeng, F. R., and Vallis, G. K.: Probing the fast and slow components of global warming by returning abruptly to preindustrial forcing, *Journal of Climate*, 23, 2418-2427, Doi 10.1175/2009jcli3466.1, 2010.
- Huntingford, C., and Cox, P. M.: An analogue model to derive additional climate change scenarios from existing gcm simulations, *Climate Dynamics*, 16, 575-586, 2000.
- IPCC: Summary for policymakers, in: *Climate change 2013: The physical science basis. Contribution of working group i to the fifth assessment report of the intergovernmental panel on climate change*, edited by: Stocker, T. F., Qin, D., Plattner, G. K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M., Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Jones, C., Lowe, J., Liddicoat, S., and Betts, R.: Committed terrestrial ecosystem changes due to climate change, *Nat Geosci*, 2, 484-487, Doi 10.1038/Ngeo555, 2009.
- Jonko, A. K., Shell, K. M., Sanderson, B. M., and Danabasoglu, G.: Climate feedbacks in cesm3 under changing co2 forcing. Part ii: Variation of climate feedbacks and sensitivity with forcing, *Journal of Climate*, 26, 2784-2795, Doi 10.1175/Jcli-D-12-00479.1, 2013.
- Li, S., and Jarvis, A.: Long run surface temperature dynamics of an a-ogcm: The hadcm3 4xco(2) forcing experiment revisited, *Climate Dynamics*, 33, 817-825, 10.1007/s00382-009-0581-0, 2009.
- Manabe, S., Bryan, K., and Spelman, M. J.: Transient-response of a global ocean atmosphere model to a doubling of atmospheric carbon-dioxide, *J Phys Oceanogr*, 20, 722-749, Doi 10.1175/1520-0485(1990)020<0722:Troago>2.0.Co;2, 1990.
- Meehl, G. A., Moss, R., Taylor, K. E., Eyring, V., Stouffer, R. J., Bony, S., and Stevens, B.: Climate model intercomparisons: Preparing for the next phase, *Eos Trans. AGU*, 95, 77, 2014.
- Meinshausen, M., Raper, S. C. B., and Wigley, T. M. L.: Emulating coupled atmosphere-ocean and carbon cycle models with a simpler model, *magicc6-part 1: Model description and calibration*, *Atmos Chem Phys*, 11, 1417-1456, DOI 10.5194/acp-11-1417-2011, 2011.
- Mitchell, J. F. B., Wilson, C. A., and Cunningham, W. M.: On co2 climate sensitivity and model dependence of results, *Q J Roy Meteor Soc*, 113, 293-322, 1987.
- Mitchell, T. D.: Pattern scaling - an examination of the accuracy of the technique for describing future climates, *Climatic Change*, 60, 217-242, 2003.

Myhre, G., Highwood, E. J., Shine, K. P., and Stordal, F.: New estimates of radiative forcing due to well mixed greenhouse gases, *Geophysical Research Letters*, 25, 2715-2718, 1998.

Santer, B., Wigley, T., Schlesinger, M., and Mitchell, J. F. B.: Developing climate scenarios from equilibrium gcm results, Report No. 47, Max Planck Institute for Meteorology, Hamburg, 1990.

Schaller, N., Cermak, J., Wild, M., and Knutti, R.: The sensitivity of the modeled energy budget and hydrological cycle to co2 and solar forcing, *Earth Syst Dynam*, 4, 253-266, DOI 10.5194/esd-4-253-2013, 2013.

Schaller, N., Sedláček, N. J., and Knutti, R.: The asymmetry of the climate system's response to solar forcing changes and its implications for geoengineering scenarios, *Journal of Geophysical Research: Atmospheres*, 10, 5171–5184, 2014.

Solomon, S., Qin, D., Manning, M., Chen, Z., Marquis, M., Averyt, K. B., Tignor, M., and Miller, H. L.: Contribution of working group i to the fourth assessment report of the intergovernmental panel on climate change, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2007.

Stott, P., Good, P., Jones, G., Gillett, N., and Hawkins, E.: The upper end of climate model temperature projections is inconsistent with past warming, *Environ Res Lett*, 8, Artn 014024
Doi 10.1088/1748-9326/8/1/014024, 2013.

Tebaldi, C., and Arblaster, J. M.: Pattern scaling: Its strengths and limitations, and an update on the latest model simulations, *Climatic Change*, 122, 459-471, DOI 10.1007/s10584-013-1032-9, 2014.

Williams, K. D., Ingram, W. J., and Gregory, J. M.: Time variation of effective climate sensitivity in gcms, *Journal of Climate*, 21, 5076-5090, Doi 10.1175/2008jcli2371.1, 2008.

Zelinka, M. D., Klein, S. A., Taylor, K. E., Andrews, T., Webb, M. J., Gregory, J. M., and Forster, P. M.: Contributions of different cloud types to feedbacks and rapid adjustments in cmip5, *Journal of Climate*, 26, 5007-5027, Doi 10.1175/Jcli-D-12-00555.1, 2013.

Table 1. List of CMIP5/CMIP6 DECK experiments required by nonlinMIP.

Experiment	Description	Role
piControl	Pre-industrial control experiment	
Abrupt4xCO2	CO2 abruptly quadrupled, then held constant for 150 years.	Separate different timescales of response.
1pctCO2	CO2 increased at 1% per year for 140 years (i.e. as CMIP5 1pctCO2 experiment), then decreased by 1% per year for 140 years (i.e. returning to pre-industrial conditions).	To test traceability of the abruptCO2 experiments to more realistic transient-forcing conditions. Adding the ramp-down phase explores physics relevant to mitigation and geo-engineering scenarios.

Table 2. NonlinMIP experimental design. Three options are: only the ‘essential’ simulation; all ‘high priority’ plus the ‘essential’ simulations; or, preferably, all simulations. The experiments in Table 1 are required in all cases.

Experiment (priority)	Description	Role
Abrupt2xCO2 (essential)	As abrupt4xCO2 (see Table 1), but at double pre-industrial CO2 concentration.	To diagnose non-linear responses (in combination with abrupt4xCO2). Assess climate response and (if appropriate) make climate projections with the step-response model at forcing levels more relevant to mid- or low-forcing scenarios.
1pctCO2 ramp-down (high priority)	Initialised from the end of 1pctCO2. CO2 is decreased by 1% per year for 140 years (i.e. returning to pre-industrial conditions).	To test traceability of the abruptCO2 experiments to more realistic transient-forcing conditions. Adding the ramp-down phase explores a much wider range of physical responses, providing a sterner test of traceability. Relevant also to mitigation and geo-engineering scenarios, and offers a sterner test of.
Extend both abrupt2xCO2 and abrupt4xCO2 by 100 years (high priority)		Allow traceability tests (via the step-response model) against most of the 1pctCO2 ramp-up-ramp-down experiment. Explore longer timescale responses than in CMIP5 experiment. Permit improved signal/noise in diagnosing some regional-scale non-linear responses Provide a baseline control for the abrupt4xto1x experiment.
Abrupt4xto1x (medium priority)	Initialised from year 100 of abrupt4xCO2, CO2 is abruptly returned to pre-industrial levels, then held constant for 150 years.	Quantify non-linearities over a larger range of CO2 (quantifies responses at 1xCO2). Assess non-linearities that may be associated with the direction of forcing change.
Abrupt8xCO2 (medium priority)	As abrupt4xCO2, but at 8x pre-industrial CO2 concentration. Only 150 years required here.	Quantify non-linearities over a larger range of CO2.

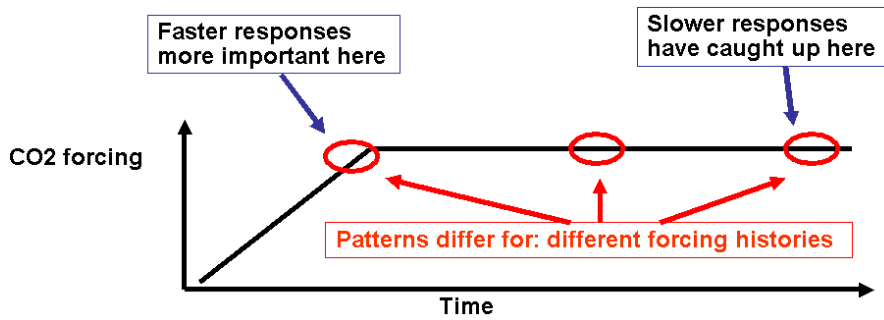


Figure 1. Schematic illustrating a situation where linear mechanisms can cause climate patterns to evolve. This represents a scenario where forcing (black line) is ramped up, then stabilised.

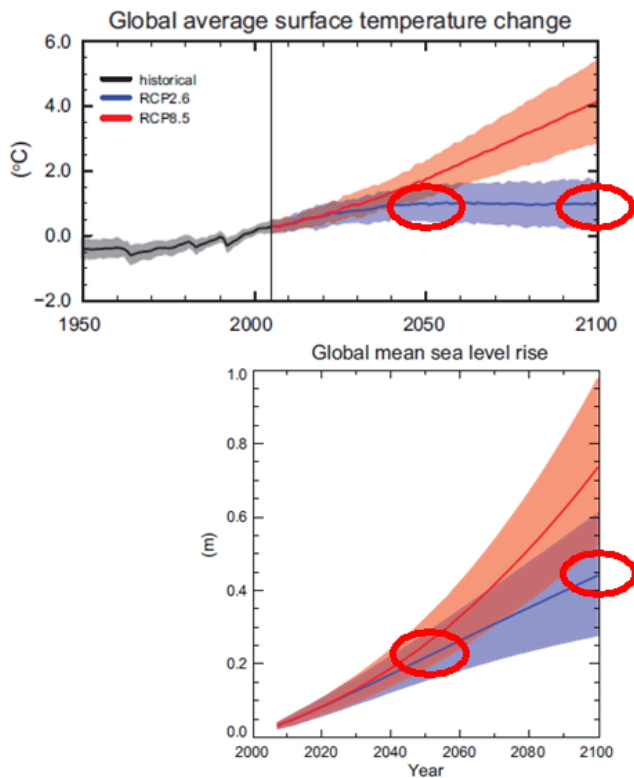


Figure 2. Adapted (red ovals overlaid) from the IPCC Fifth Assessment Report (IPCC, 2013), Figures SPM.7 and SPM.9. Global mean warming (top) and global mean sea level rise (bottom), relative to 1986-2005, for rcp8.5 (red) and rcp2.6 (blue).

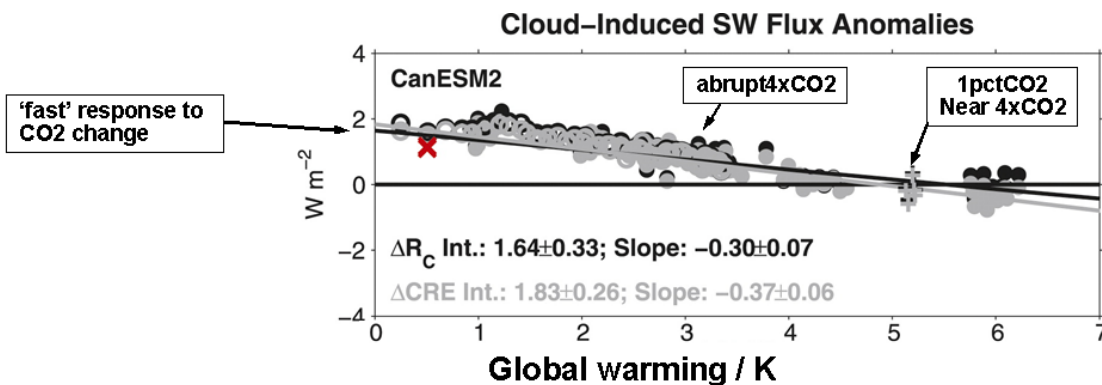


Figure 3. Illustrating a method (Gregory et al., 2004) for separating ‘fast’ and ‘slow’ responses to radiative forcing change. Figure adapted (labels in rectangles overlaid) from Zelinka et al. (2013). Global-mean cloud-induced SW flux anomalies against global warming, for the CanESM2 model (black & grey represent two methods of calculating cloud-induced fluxes). This also illustrates one test of traceability of abrupt4xCO2 to 1pctCO2 responses: the linear fit to the abrupt4xCO2 response (straight lines) passes through the 1pctCO2 response near 4xCO2 (i.e. near year 140 of that experiment).

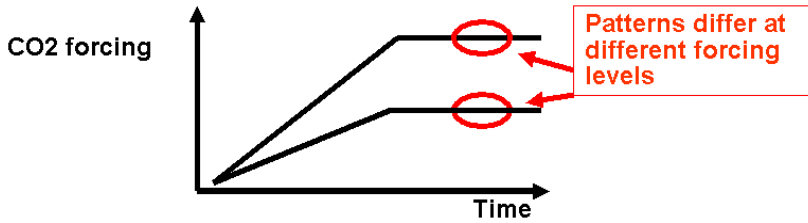


Figure 4. Schematic illustrating the point that nonlinear mechanisms can cause climate patterns to differ at different forcing (and hence global temperature) levels.

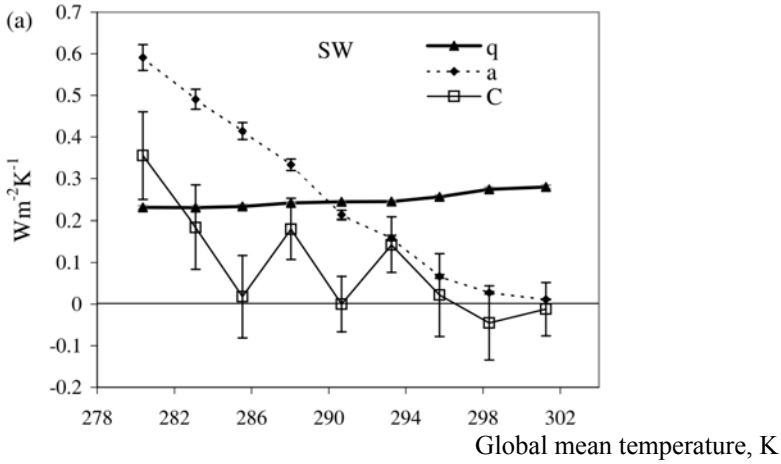


Figure 5. Albedo feedback (dotted line) strength (y-axis) decreasing with global mean temperature (x-axis, K) in a climate model (figure from Colman and McAvaney, 2009).

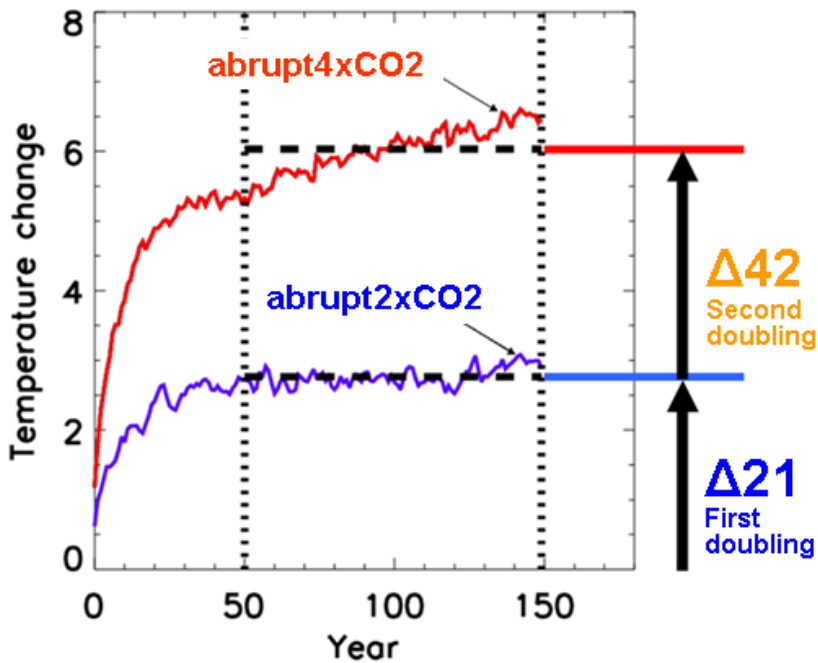


Figure 6. Defining the ‘doubling difference’. Doubling difference = $\Delta 42 - \Delta 21$ (the difference in response between the first and second CO₂ doublings. This is defined for a specific timescale after the abrupt CO₂ change – in this example, it is the mean over years 50-149.

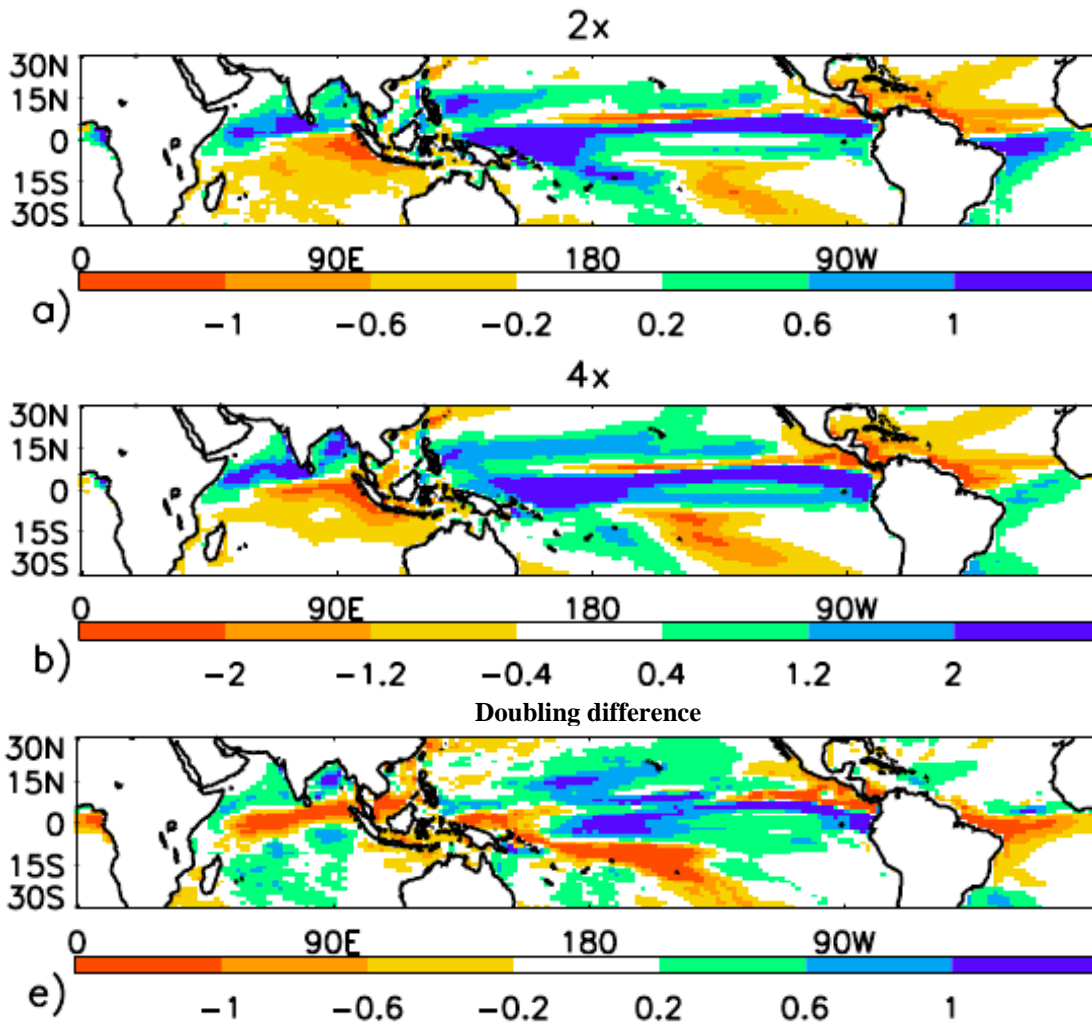


Figure 7. Non-linear regional precipitation responses over the ocean in HadGEM2-ES (figure from Chadwick and Good, 2013). Precipitation change (mm/day) averaged over years 50-149 for (top) abrupt2xCO₂ and (middle) abrupt4xCO₂, and the doubling difference (bottom). Note that the top and bottom panels have the same scale.

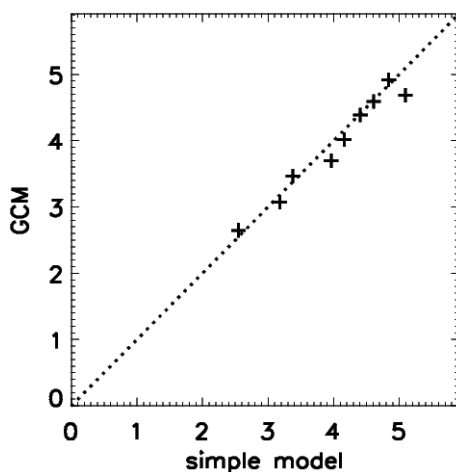


Figure 8. Checking basic traceability of abrupt4xCO₂ to a transient forcing experiment (1pctCO₂) (figure from Good et al., 2013). Global-mean warming (K) averaged over years 120-139 of 1pctCO₂ for (y-axis) the GCM simulation and (x-axis) the reconstruction from abrupt4xCO₂ using the step-response method.

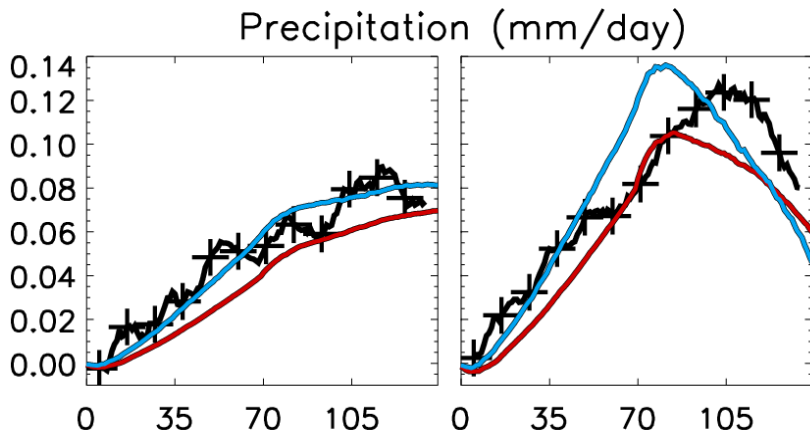


Figure 9. Finding nonlinear responses in transient forcing experiments. (figure from Good et al., 2012). Left: where CO₂ is increased by 1% per year, then stabilised at 2x pre-industrial levels. Right: where CO₂ is increased by 2% per year for 70 years, then decreased by 2% per year for 70 years. Black: GCM. Red: step-response model using the abrupt4xCO₂ response. Blue: the abrupt2xCO₂ response.

Proposal of request of COSP diagnostics for CMIP/DECK experiments

COSP Project Management Committee

August, 2014

1 Introduction

The initial design for CMIP6 has recently been published [Meehl et al., 2014]. It includes a small set of experiments to be run by modelling groups whenever they develop a new model version:

- AMIP (1979-end)
- Pre-industrial control
- 1% yr⁻¹ CO₂ increase up to 4xCO₂
- Abrupt 4xCO₂
- RCP8.5

These experiments are called the CMIP Diagnostic, Evaluation and Characterization of Klima (DECK) experiments.

In this document, we present the proposal of the list of COSP diagnostics to be requested for the DECK experiments. This proposal is the outcome of initial discussions by the COSP Project Management Committee (PMC), and it is now open for comments. The COSP diagnostic request for CMIP5/CFMIP2 is summarised and motivated in the CFMIP-2 proposal document [Bony et al., 2009], and documented in detail in the CMIP5 Standard Output documentation at http://cmip-pcmdi.llnl.gov/cmip5/output_req.html in excel spreadsheet format (Worksheet 'CFMIP output' indicates which tables appear in which experiments and for which periods, which other worksheets such as cfMon, cfDay etc indicate the variables in each table). Our view is that the CFMIP-2 diagnostics set fundamentally sound and form a suitable basis for the COSP request for the DECK, subject to some modifications. Thus, we present this proposal as changes with respect to the CFMIP-2 protocol in the spreadsheet attached. We have highlighted the changes with the following colour code:

- Red: to be deleted (or moved to a different table)
- Orange: additions or modifications

We have tried to address the concerns raised in the CMIP5 survey by simplifying the technical difficulty of the requests (sometimes at the expense of extra data) and basing the requests upon a frozen well-tested and already-released version of COSP (v1.4).

In the sections below we present and motivate the specific requested changes.

2 Description of proposed changes

2.1 Change #1: Replacement of curtain data by full 3D fields, and deletion of cfOff table (proposed by Alejandro Bodas-Salcedo)

In CFMIP-2, the production of data along the A-train track ("curtain" data, table cf3hr offline) involved a substantial amount of post-processing. A second post-processing step required the gridding and time-averaging of these data to produce the monthly means

requested in the cfOff table. This proved quite difficult for many modelling centres. Although not from the ESG archive, this type of data has been used in several model evaluation papers [Bony et al., 2009; Bodas-Salcedo et al., 2008; Field et al., 2011; Williams et al., 2013] involving case-study comparison of models with along-track observations from CloudSat and Calipso. We believe that by simplifying the request, the modelling centres will find easier to contribute these data. Hence, we propose to drop the orbital sampling, i.e. to request globally-complete fields on a standard lat/lon grid. Given this change, the calculation of monthly-averages from gridded 3-hourly data is trivial, and therefore we propose to delete the cfOff table.

2.2 Change #2: New table cfMonExtra. Add CloudSat and CALIPSO CFADs to cfMonExtra (proposed by Alejandro Bodas-Salcedo and Mark Webb)

Optimisations to the code in COSP v1.4 mean that it is now practical to run the CloudSat simulator inline in models and so for longer periods. We propose the introduction of a new table cfMonExtra for the inclusion of monthly mean COSP diagnostics used for model evaluation in present day experiments such as AMIP, but which we don't consider appropriate for climate change experiments. In this new table we include Cloud Frequency/Altitude Diagram (CFAD) diagnostics for CloudSat and CALIPSO for the entire AMIP integration. CFADs for CloudSat and CALIPSO have appeared in a number of published studies [e.g. Nam et al., 2014; Franklin et al., 2013; Bodas-Salcedo et al., 2011; Bodas-Salcedo et al., 2012; Nam and Quaas, 2012; Nam and Quaas, 2013; Kay et al., 2012; Kodama et al., 2012; Marchand et al., 2009; Abel and Boutle, 2012] and their inclusion as monthly means in the AMIP DECK experiment will make them available for analysts in a more convenient form than the higher frequency outputs currently requested in CMIP5.

2.3 Change #3: Standard monthly COSP and daily COSP 2D outputs in all of the DECK experiments (proposed by Mark Webb and Steve Klein)

Many of the standard monthly COSP and daily COSP 2D have been shown to be valuable in the CMIP5 experiments, not only for cloud evaluation [e.g. Franklin et al., 2013; Bodas-Salcedo et al., 2012; Nam and Quaas, 2013; Lacagnina and Selten, 2014; Bodas-Salcedo et al., 2014; Klein et al., 2013; Cesana and Chepfer, 2012; Tsushima et al., 2013] but also in quantifying the contributions of different cloud types to cloud feedbacks and forcing adjustments in climate change experiments [e.g. Tsushima et al., 2013; Zelinka et al., 2012a; Zelinka et al., 2012a; Zelinka et al., 2013; Zelinka et al., 2014]. We propose to include these in all DECK experiments as standard, so support evaluation of cloud, cloud feedbacks and cloud adjustments in a wider range of scenarios, for example relating trends in the observational record to changes in more realistic future scenarios compared to the idealised scenarios to which they have been applied thus far.

2.4 Change #4: Move PARASOL reflectance to cfMonExtra (proposed by Robert Pincus)

Top-of-atmosphere reflectance measurements from PARASOL were part of the standard request for CMIP5. They have been used in some applications [e.g. Nam et al. 2012] but

have not been widely exploited. The proposal is to move them from the cfMon to cfMonExtra tables to reduce the number of integrations for which they are requested and to focus on model evaluation applications.

2.5 Change #5: Add MISR CTH-COD to cfMonExtra. Add MISR CTH-COD and ISCCP CTP-OD histograms to cf3hr (proposed by Roger Marchand)

Histograms of cloud-top-height (or cloud-top-pressure) and optical-depth produced by ISCCP have been widely used in the evaluation of climate models, often in combination with the ISCCP-simulator now part of COSP. Because top-of-atmosphere outgoing longwave fluxes are related to cloud-top-height and outgoing shortwave fluxes are related to cloud-optical-depth this framework provides a way to evaluate the distribution of model clouds in a way that is closely related to their radiative impact. Similar histograms of cloud-top-height and optical-depth are being produced from observations by the Multiangle Imaging Spectro-Radiometer (MISR). While similar, the cloud-top-height in the MISR dataset is obtained using a stereo-imaging technique that is purely geometric and insensitive to the calibration of the MISR cameras. This technique provides more accurate retrievals of cloud-top-height for low-level and mid-level clouds, and more reliable discrimination of mid-level clouds from other clouds, while ISCCP provides greater sensitivity to optically-thin high-level clouds. In addition, ISCCP and MISR histograms can be combined to separate optically-thin high-level clouds into multi-layer and single-layer categories [Marchand et al. 2010]. We therefore recommend using both ISCCP and MISR observations and instrument-simulators in the evaluation of climate model, and such an analysis is underway using a few CFMIP5 models that have run the MISR simulator [Hillman et al. 2014]. While monthly data are useful for the broad evaluation of models on monthly or longer time scales, the acquisition of high frequency (Three hourly) data will enable analysis of events that are not well resolved with monthly data, including the diurnal cycle, the Madden-Julian Oscillation (MJO) and various synoptic states or weather patterns, such as frontal passages. We recognize that this represents a large increase in data-volume compared with monthly averages and propose collection of this three hourly data only for a period of about 1 year.

2.6 Change #6: Add MODIS cloud fractions (total, liquid, ice) to cfMonExtra (proposed by Robert Pincus)

The partitioning between liquid and ice phase has significant impacts on the energy and hydrologic impacts of clouds. As models move towards predicting more details of the aerosol distributions, including the ice nucleation ability, evaluation of the phase partitioning on the global scale will become more important. Evaluation to date has been based primarily on polarization measurements from active and passive sensors [e.g. Doutriaux-Boucher and Quaas, 2004; Komurcu et al., 2014] and height-resolved partitioning estimates from the CALIPSO sensor are requested below. Cloud phase estimates from the MODIS simulator were not available in CFMIP2 but may prove a useful complement by virtue of greater geographic sampling and longer time records.

2.7 Change #7: MODIS COT-particle size histograms by phase in cfMonExtra, cfDayExtra, cf3hr (proposed by Robert Pincus)

The joint distribution of optical thickness and particle size provides a window on the microphysical processes within clouds [Nakajima et al., 1991] and is influenced by direct and some indirect effects of aerosols on cloud optical properties [Han et al. 2002]. As models move towards predicting more details of the aerosol properties and cloud-aerosol interactions the assessment of these processes becomes more pressing.

Estimate of particle size from MODIS have been difficult to use for model evaluation to date because of observational artefacts not treated by the MODIS simulator. These artefacts are reduced by the use of observations at wavelengths with greater absorption by condensed water (e.g. by exploiting reflectance at 3.7 μm instead of 2.1 μm). The MODIS simulator and accompanying data for CFMIP3 will use measurements at 3.7 μm to infer particle size. This will also act to make output from the MODIS simulator roughly consistent with the PATMOS-X observations in the same way that distributions of optical thickness from the MODIS, MISR, and ISCCP simulators are nearly equivalent.

2.8 Change #8: add CALIPSO ice and liquid 3D cloud fractions to cfMonExtra (proposed by H el ene Chepfer)

Changes in cloud optical depth associated with cloud phase feedbacks can dominate the changes in high-latitude clouds in future climate projections [e.g. Senior and Mitchell, 1993]. Cloud phase identification capabilities have been recently added to the CALIPSO simulator in COSP, and a compatible observational dataset has been produced [Cesana and Chepfer, 2013]. We propose to include these in the AMIP DECK experiment to support the evaluation of the simulation of cloud phase.

2.9 Change #9: CALIPSO total cloud fraction and PARASOL reflectance to cfDayExtra (proposed by H el ene Chepfer and Dimitra Konsta)

The multi-sensor A-train observations (CALIPSO-GOCCP and MODIS, PARASOL) allow to make the correlations between the different cloud variables at the instantaneous time scale, and at high resolution. The use of the high-frequency relationships between different variables allows for process-oriented model evaluation. These diagnostics will help test the realism of the co-variation of key cloud properties that control cloud feedbacks in models. Konsta et al. (2014) have used these diagnostics in a pilot analysis.

3 References Using Satellite Simulators for the Evaluation of Clouds in Models (other refs below)

- Bodas-Salcedo, A. et al., 2008: Evaluating cloud systems in the Met Office global forecast model using simulated CloudSat radar reflectivities, *J. Geophys. Res.*, 113, D00A13. DOI: 10.1029/2007JD009620.
- Bodas-Salcedo, A., et al., 2011: Satellite simulation software for model assessment. *Bull. Am. Meteorol. Soc.*, 92. DOI: 10.1175/2011BAMS2856.1.
- Bodas-Salcedo, A., et al., 2012: The surface downwelling solar radiation surplus over the Southern Ocean in the Met Office model: the role of midlatitude cyclone clouds, *J. Climate*, 25. DOI: 10.1175/JCLI-D-11-00702.1.

- Bodas-Salcedo, A., et al., 2014: Origins of the Solar Radiation Biases over the Southern Ocean in CFMIP2 Models, *J. Climate*, 27. DOI: 10.1175/JCLI-D-13-00169.1.
- Cesana, G., and Chepfer, H., 2012: How well do climate models simulate cloud vertical structure? A comparison between CALIPSO-GOCCP satellite observations and CMIP5 models, *Geophys. Res. Lett.* DOI: 10.1029/2012GL053153.
- Cesana, G., and Chepfer, H., 2013: Evaluation of the cloud thermodynamic phase in a climate model using CALIPSO-GOCCP, *J. Geophys. Res.*, 118, 7922–7937. DOI: 10.1002/jgrd.50376.
- Doutriaux-Boucher, M., and J. Quaas, 2004: Evaluation of cloud thermodynamic phase parametrizations in the LMDZ GCM by using POLDER satellite data, *Geophys. Res. Lett.*, 31, L06126. DOI: 10.1029/2003GL019095.
- Field, P. R., et al., 2011: Using model analysis and satellite data to assess cloud and precipitation in midlatitude cyclones, *Q. J. R. Meteorol. Soc.*, 137, 1501-1515. DOI: 10.1002/qj.858.
- Franklin, C. N., et al., 2013: Evaluation of clouds in ACCESS using the satellite simulator package COSP: Global, seasonal, and regional cloud properties, *J. Geophys. Res.* DOI: 10.1029/2012JD018469.
- Hillman, B., R. Marchand, T. P. Ackerman, A. Bodas-Salcedo, J. Cole, J.-C. Golaz, J. E. Kay, 2014: Comparing Cloud Biases in CMIP5: Insights Using MISR and ISCCP Observations and Satellite Simulators, *in preparation*.
- Kay, J. E., et al., Exposing Global Cloud Biases in the Community Atmosphere Model (CAM) Using Satellite Observations and Their Corresponding Instrument Simulators, *J. Climate*, 25, 2012. DOI: 10.1175/JCLI-D-11-00469.1.
- Klein, S. A. et al., 2013: Are climate model simulations of clouds improving? An evaluation using the ISCCP simulator, *J. Geophys. Res.*, 118. DOI: 10.1002/jgrd.50141.
- Kodama, C., et al., 2012: An assessment of the cloud signals simulated by NICAM using ISCCP, CALIPSO, and CloudSat satellite simulators, *J. Geophys. Res.*, 117. DOI: 10.1029/2011JD017317.
- Komurcu, M., T. Storelvmo, I. Tan, U. Lohmann, Y. Yun, J. E. Penner, Y. Wang, X. Liu, and T. Takemura, 2014: Intercomparison of the cloud water phase among global climate models, *J. Geophys. Res.*, 119, 3372–3400. DOI:10.1002/2013JD021119.
- Konsta, D., J.-L. Dufresne, H. Chepfer, A. Idelkadi and G. Cesana, 2014: Evaluation of clouds simulated by the LMDZ5 GCM using A-train satellite observations (CALIPSO, PARASOL, CERES). *Climate Dynamics*, under review.
- Lacagnina, C., and Selten, F., 2014: Evaluation of clouds and radiative fluxes in the EC-Earth general circulation model, *Clim. Dyn.* DOI: 10.1007/s00382-014-2093-9.
- Marchand, R., et al., 2009: A comparison of simulated cloud radar output from the multiscale modeling framework global climate model with CloudSat cloud radar observations, *J. Geophys. Res.*, 114, D00A20. DOI: 10.1029/2008JD009790.
- Marchand, R., T. Ackerman, M. Smyth, and W. B. Rossow, 2010: A review of cloud top height and optical depth histograms from MISR, ISCCP, and MODIS, *J. Geophys. Res.*, 115, D16206. DOI:10.1029/2009JD013422.
- Nam, C., S. Bony, J.-L. Dufresne, and H. Chepfer, 2012: The "too few, too bright" tropical low-cloud problem in CMIP5 models, *Geophys. Res. Lett.*, 39. DOI:10.1029/2012GL053421.

- Nam, C. C. W., and Quaas, J., 2012: Evaluation of Clouds and Precipitation in the ECHAM5 General Circulation Model Using CALIPSO and CloudSat Satellite Data. I, *J. Climate*, 25, 4975-4992. DOI:10.1175/JCLI-D-11-00347.1.
- Nam, C. C. W. and Quaas, J., 2013: Geographically versus dynamically defined boundary layer cloud regimes and their use to evaluate general circulation model cloud parameterizations, *Geophys. Res. Lett.* DOI: 10.1002/grl.50945.
- Nam, C. W. W., et al., 2014: Evaluation of boundary layer cloud parametrizations in the ECHAM5 general circulation model using CALIPSO and CloudSat satellite data. *JAMES*. DOI: 10.1002/2013MS000277.
- Tsushima, Y. et al., 2013: Quantitative evaluation of the seasonal variations in climate model cloud regimes, *Clim. Dyn.*, 41. DOI: 10.1007/s00382-012-1609-4.
- Williams, K. D., et al., 2013: The Transpose-AMIP II experiment and its application to the understanding of Southern Ocean cloud biases in climate models, *J. Climate*, 26, 3258-3274. DOI: 10.1175/JCLI-D-12-00429.1.
- Zelinka, M. D., et al., 2012a: Computing and Partitioning Cloud Feedbacks Using Cloud Property Histograms. Part I: Cloud Radiative Kernels. *J. Climate*. DOI: 10.1175/JCLI-D-11-00248.1.
- Zelinka, M. D., et al., 2012b: Computing and Partitioning Cloud Feedbacks Using Cloud Property Histograms. Part II: Attribution to Changes in Cloud Amount, Altitude, and Optical Depth. *J. Climate*. DOI: 10.1175/JCLI-D-11-00249.
- Zelinka, M. D., et al., 2013: Contributions of Different Cloud Types to Feedbacks and Rapid Adjustments in CMIP5, *J. Climate*. DOI: 10.1175/JCLI-D-12-00555.1.
- Zelinka, M. D., et al., 2014: Quantifying Components of Aerosol-Cloud-Radiation Interactions in Climate Models. *J. Geophys. Res.* DOI: 10.1002/2014JD021710.

4 Other References

- Abel, S. J., and Boutle, I. A., 2012: An improved representation of the raindrop size distribution for single-moment microphysics schemes, *Q. J. R. Meteorol. Soc.*, 138, 2151-2162. DOI: 10.1002/qj.1949.
- Bony, S. et al., 2009: The Cloud Feedback Model Intercomparison Project : Summary of Activities and Recommendations for Advancing Assessments of Cloud-Climate Feedbacks, (<http://www.cfmip.net> -> CFMIP Strategy and Plans -> CFMIP2_experiments_March20th2009.pdf).
- Han, Q, W B. Rossow, J Zeng, R Welch, 2002: Three Different Behaviors of Liquid Water Path of Water Clouds in Aerosol-Cloud Interactions. *J. Atmos. Sci.*, 59, 726-735. DOI: 10.1175/1520-0469(2002)059<0726:TDBOLW>2.0.CO;2.
- Meehl, G. et al., 2014: Climate Model Intercomparisons: Preparing for the Next Phase, *Eos Trans. AGU*, 95(9), 77. DOI: 10.1002/2014EO090001.
- Nakajima, T, M D. King, J D. Spinhirne, L F. Radke, 1991: Determination of the Optical Thickness and Effective Particle Radius of Clouds from Reflected Solar Radiation Measurements. Part II: Marine Stratocumulus Observations, *J. Atmos. Sci.*, 48, 728-751. DOI: 10.1175/1520-0469(1991)048<0728:DOTOTA>2.0.CO;2.
- Senior, C. A., and Mitchell, J. F. B., 1993: Carbon Dioxide and Climate. The Impact of Cloud Parameterization. *J. Climate*, 6, 393-418. DOI: 10.1175/1520-0442(1993)006<0393:CDACTI>2.0.CO;2.

Acknowledgements

Thanks to D. Konsta for providing text for Section 2.9.

Detection and Attribution Model Intercomparison Project (DAMIP)

Application for CMIP6-Endorsed MIPs

Date: 26th November 2014

➤ Name of MIP*

Detection and Attribution Model Intercomparison Project (DAMIP)

➤ Co-chairs of MIP as contacts*

Nathan Gillett, Canadian Centre for Climate Modelling and Analysis (nathan.gillett@ec.gc.ca)

Hideo Shiogama, National Institute for Environmental Studies (shiogama.hideo@nies.go.jp)

➤ Members of the Scientific Steering Committee*

Gabriele Hegerl, University of Edinburgh (gabi.hegerl@ed.ac.uk)

Dáithí Stone, Lawrence Berkeley National Laboratory (dstone@lbl.gov)

Claudia Tebaldi, National Center for Atmospheric Research (claudia.tebaldi@gmail.com)

Reto Knutti, ETH Zurich (reto.knutti@env.ethz.ch)

➤ Link to website (if available)*

Not yet available.

➤ References (if available)*

Bindoff, N.L., et al., 2013: Detection and Attribution of Climate Change: from Global to Regional. In: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change [Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley (eds.)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA.

Jones, G. S., P. A. Stott, and N. Christidis, 2013: Attribution of observed historical near surface temperature variations to anthropogenic and natural causes using CMIP5 simulations. *J. Geophys. Res. Atmos.*, doi:10.1002/jgrd.50239.

Gillett, N. P., V. K. Arora, D. Matthews, P. A. Stott, and M. R. Allen, 2013: Constraining the ratio of global warming to cumulative CO₂ emissions using CMIP5 simulations. *J. Clim.*, doi:10.1175/JCLI-D-12-00476.1.

➤ Goal of the MIP and a brief overview*

The primary goals of DAMIP are to facilitate improved estimation of the contribution of anthropogenic and natural forcing changes to observed global warming; to facilitate improved estimation of the contribution of those forcings to observed global and regional changes in other climate variables; to contribute to the estimation of how historical emissions have altered and are altering contemporary climate risk; and to facilitate and improve observationally-constrained projections of future climate change. Detection and attribution studies typically require unforced control simulations and historical simulations including all major anthropogenic and natural

forcings. Such simulations will be carried out as part of the DECK and the *CMIP6 historical simulation* (hereafter we referred to the *CMIP6 historical simulation* as *histALL*). In addition such studies require additional simulations with individual forcings or subsets of forcings. We propose some such separated forcing experiments as Detection and Attribution MIP (DAMIP) for CMIP6. Combinations of *histALL* and separated forcing experiments from models participating in CMIP6 will be useful for model evaluation, better understanding of historical climate changes, and for deriving observational constraints on future climate change projections.

➤ An overview of the proposed experiments*

We propose some historical experiments using individual forcings or subsets of forcings. These experiments are CO₂-concentration driven for ESMs. These simulations should start at the same time as the *histALL* simulations and continue to at least 2020. Forcings identical to those in the *histALL* simulations should be used up to the end of those simulations, followed by forcings from the SSP5-8.5 simulation (*future emission scenarios are subject to change based on further consultation with other MIPs, e.g. ScenarioMIP, DCP and RFMIP*). Multi-member ensembles are vital for the separation of forced responses and internal variability. We recommend at least 3 ensemble members with different initial conditions for each experiment, and recommend that modeling groups which cannot afford to do this for all requested runs start by carrying out at least 3-member ensembles of the 1st priority simulations. We also request three extension experiments with individual forcings up to 2100 under SSP5-8.5: well-mixed GHG changes only; ozone changes only; and anthropogenic aerosol changes only. The minimum ensemble size of these is one. We also recommend modelling groups to perform a 500-year or longer piControl run to allow robust estimates of internal variability.

We propose four Tier 1 experiments for DAMIP/CMIP6. The first one is the enlargement of the ensemble size of **histALL** (the CMIP6 historical simulation) to at least three members. The other three Tier 1 experiments are Natural-only (**histNAT**), GHG-only (**histGHG**) and Aerosols-only (**histAER**). Here, “XXX-only” mean that the agent XXX changes as in the *histALL* runs, but the other conditions are imposed and kept constant as in the piControl experiments. We require that forcing agents are perturbed exactly as in the *histALL* simulations: For example in the *histGHG* simulations the same well-mixed GHG concentrations are prescribed as in the *histALL* simulations. We request modelling groups to report what sets of emissions and boundary conditions are used in each run.

One Tier 2 experiment is proposed: the extension of GHG-only up to 2100 (**ssp585GHG**). Tier 3 experiments are Ozone-only (**histOZ**), the extension of Ozone-only to 2100 (**ssp585OZ**), Volcanic-only (**histVLC**), and the extension of Aerosol-only up to 2100 (**ssp585AER**).

Both DAMIP and DCP propose an initial condition ensemble of *histALL* simulations. *histALL*, *histNAT*, *histGHG*, *histAER*, *ssp585GHG* of DAMIP correspond closely to transient AGCM experiments for estimates of radiative forcing proposed in RFMIP-ERF. Combinations of DAMIP and RFMIP-Historical simulations will allow uncertainties in the aerosol response to be separated into those associated with simulating the climate response to a given distribution of aerosols, and the full uncertainty based on specified aerosol precursor emissions. DAMIP also co-sponsors "ALL minus land-use (LND_noLULCC_hist)" simulations of LUMIP. Some experiments of AerChemMIP are based on *histGHG* of DAMIP. *histALL* and *histNAT* runs from DAMIP will be used in diagnoses of GMMIP. Solar-only runs of SolarMIP are complementary with the *histNAT* ensemble of DAMIP. *histVLC* runs are useful for GeoMIP and VolMIP. Combinations of SSP5-8.5 runs of ScenarioMIP and *ssp585GHG*, *ssp585OZ* and *ssp585AER* of DAMIP will allow the investigation of future climate responses to different forcing agents and observational constraints on future projections.

➤ An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*

A number of detection and attribution analyses of anthropogenic and natural forcing influences on historical climate changes are anticipated. Those analyses will likely address historical changes in temperature, the hydrological cycle, atmospheric circulation, ocean properties, cryospheric variables, extreme indices and other variables, from global to regional scales. The extension of DAMIP experiments from 2005 in CMIP5 to 2020 is essential to understand reasons for the recent *hiatus* of climate warming and improve signal-to-noise ratio for detection and attribution of changes in high-noise variables such as precipitation. The DAMIP experiments are also important for observational constraints on future climate change projections, climate sensitivity, TCR and TCRE. Using combinations of experiments from DAMIP, RFMIP and AerChemMIP, we can compare transient climate responses per unit radiative forcing across different forcing factors.

It is anticipated that analyses of DAMIP simulations will form the basis of the assessment of the detection and attribution of climate change in the next IPCC assessment report. Observationally-constrained estimates based on DAMIP simulations may also provide a major contribution to projections of future climate in this report.

➤ Proposed timing *

After modelling centers perform piControl (we recommend 500-year or longer simulations), histALL (the CMIP6 historical simulation) and SSP5-8.5 (ScenarioMIP Tier 1) experiments.

➤ For each proposed experiment for CMIP Phase 6**

Tier 1 experiments

(1.0) Enlarging ensemble size of the CMIP6 historical simulations to at least three members (histALL)

○ the experimental design

- All forcing historical simulations
- Enlarging ensemble size of histALL to at least three members with different initial conditions. Please use forcings from SSP5-8.5 during 2015-2020.
- Please provide outputs of experiments under the name of the CMIP6 historical simulation, not histALL.
- DCPD proposes a 10-member ensemble of histALL.

○ the science question and/or gap being addressed with this experiment

Combinations of histALL, histNAT and histGHG will allow us to attribute observed climate changes to contributions from GHG, the other anthropogenic factors and natural forcing. Because better signal to noise ratio is vital to D&A analyses, we request at least 3 members for all historical experiments. Larger numbers of simulations also provide much larger samples of extreme events for climate risk analysis.

○ possible synergies with other MIPs

- This experiment will benefit all researchers who analyze historical climate changes, the present climatology and future changes from the present climate.
- DCPD: DCPD proposes a 10 member ensemble of histALL.

- RFMIP-ERF: Combining radiative forcing estimated from RFMIP-ERF and transient climate responses from DAMIP (histALL, histNAT, histGHG, histAER, ssp585GHG and ssp585AER), we can investigate how feedbacks and adjustments vary with forcing factors.
 - RFMIP-Historical: Combinations of DAMIP (histALL, histNAT, histAER) and RFMIP-Historical will allow us to separate uncertainties in climate response based on specified aerosol evolution from the overall uncertainties in climate response to specified aerosol precursor emissions.
 - GMMIP: histALL and histNAT runs from DAMIP will be used in diagnoses of GMMIP.
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.
Larger ensemble sizes of histALL should benefit (A) and (C) due to better signal to noise ratios of climate change signals and information about uncertainties associated with internal variability.

(1.1) Natural-only run (histNAT)

- the experimental design
 - Historical simulations forced by natural forcing agents only (i.e., solar irradiance change and volcanic activity), exactly as in histALL.
- the science question and/or gap being addressed with this experiment
histALL and histNAT simulations will allow us to attribute observed changes to anthropogenic and natural influences. histALL, histNAT and histGHG simulations will allow us to attribute observed climate changes to contributions from GHG, the other anthropogenic factors and natural forcing.
- possible synergies with other MIPs
 - C20C+ Detection and Attribution Project: The event attribution project of C20C+ will make use of the histNAT and histALL simulations to estimate boundary SST conditions for their AGCM simulations of the hypothetical counterfactual world without human influences.
 - RFMIP-ERF: Combining radiative forcing estimated from RFMIP-ERF and transient climate responses from DAMIP (histALL, histNAT, histGHG, histAER, ssp585GHG and ssp585AER), we can investigate how feedbacks and adjustments vary with forcing factors.
 - RFMIP-Historical: Combinations of DAMIP (histALL, histNAT, histAER) and RFMIP-Historical will allow us to separate uncertainties in climate response based on specified aerosol evolution from the overall uncertainties in climate response to specified aerosol precursor emissions.
 - VolMIP: VolMIP proposes both historical and mechanism based simulations with a focus on volcanic eruptions.
 - SolarMIP: histNAT and histVLC of DAMIP and Solaronly of SolarMIP allow us to investigate volcanic and solar influences on climate and to check additivity.
 - GMMIP: histALL and histNAT runs from DAMIP will be used in diagnoses of GMMIP.

- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

histALL, histGHG and histNAT in CMIP5 were vital for IPCC AR5 to conclude “more than half of the observed increase in global mean surface temperature from 1951 to 2010 is very likely due to the observed anthropogenic increase in greenhouse gas concentrations”. The updated forcings, longer simulations, and larger ensemble sizes of these experiments using new models in DAMIP/CMIP6 will facilitate even more robust attribution assessments and a better understanding of observed climate changes. histALL and histNAT will be used for event attribution analyses of recent extreme weather and climate events, and can be used for D&A analyses of impact assessments. Those attribution studies will provide essential information for discussion of mitigation and adaptation policies.

(1.2) well-mixed GHG-only run (histGHG)

- the experimental design
 - Historical simulations forced by well mixed greenhouse gas changes only, as in the histALL simulations. Models with interactive chemistry schemes should either turn off the chemistry or use a preindustrial climatology of stratospheric and tropospheric ozone in their radiation schemes. This will ensure that simulated responses in models with and without coupled chemistry are comparable.
- the science question and/or gap being addressed with this experiment
 - Combinations of histALL, histNAT and histGHG will allow the quantification of climate change attributable to GHG changes, other anthropogenic forcings and natural forcings.
 - Allows observationally-constrained TCR and TCRE to be estimated together the 1PCTCO2 simulation in the DECK.
- possible synergies with other MIPs
 - C4MIP: Carbon flux changes related to the GHG concentration changes.
 - RFMIP-ERF: Combining radiative forcing estimated from RFMIP-ERF and transient climate responses from DAMIP (histALL, histNAT, histGHG, histAER, ssp585GHG and ssp585AER), we can investigate how feedbacks and adjustments vary with forcing factors.
 - AerChemMIP: The following two experiments of AerChemMIP rely on the histGHG simulation of DAMIP: “GHG-only with all NTCF precursors but both aerosol and ozone interacting with radiation” and “GHG-only with all NTCF precursors but only aerosol interacting with radiation”.
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

histALL, histGHG and histNAT in DAMIP/CMIP6 will facilitate more robust attribution assessments and a better understanding of climate changes than those based on CMIP5. Furthermore, histGHG and ssp585GHG may be used to derive observationally constrained future climate projections, climate sensitivity, TCR and TCRE. Observationally-constrained projections will provide useful information to inform discussion of mitigation and adaptation policies.

(1.3) Anthropogenic-Aerosols-only runs (histAER)

Two experimental designs are proposed for histAER - Please select one of them.

If you like to perform both this experiment and the corresponding simulation in RFMIP-ERF, please apply the same setup.

(1.3a) Anthropogenic-Aerosols-only runs (histAER)

- the experimental design
 - Historical simulations forced by anthropogenic aerosol concentrations only or aerosol and aerosol precursor emissions only as in the histALL simulation (sulfate, black carbon, organic carbon, ammonia, NO_x and VOCs).
 - (1.3a) is only for models in which changes in GHG concentrations do not affect aerosols and changes in aerosol precursors do not affect ozone. In addition models in which these interactions do occur, but for which 1.3b cannot be implemented, should carry out this experiment.

(1.3b) Anthropogenic-Aerosols-only runs (histAERchem)

- the experimental design
 - Historical simulations forced by aerosol and aerosol precursor emissions only as in the histALL simulation (sulfate, black carbon, organic carbon, ammonia, NO_x and VOCs).
 - Changes in well-mixed-GHGs, aerosol precursors and ozone precursors are prescribed as in histALL runs. However, in the radiation scheme, the concentrations of well-mixed-GHGs and the ozone climatology from the piControl runs are used. This procedure will allow the simulation of aerosol burdens consistent with histALL runs, and the simulation of their influences on climate.
 - (1.3b) is only for models in which changes in GHG concentrations affect aerosols or changes in aerosol precursors affect ozone.

- the science question and/or gap being addressed with this experiment

Aerosols are a key source of uncertainty in historical and future climate simulations and the prime reason for high uncertainty in TCR and ECS constraints. Together with the histNAT and histALL simulations, these simulations will allow us to attribute observed climate changes to contributions from natural forcings, aerosols and “GHG+ozone+land use change”. This approach will likely result in more tightly constrained estimates of attributable warming since the aerosol response, which is more uncertain, will be directly simulated.

- possible synergies with other MIPs

- RFMIP-ERF: Combining radiative forcing estimated from RFMIP-ERF and transient climate responses from DAMIP (histALL, histNAT, histGHG, histAER, ssp585GHG and ssp585AER), we can investigate how feedbacks and adjustments vary with forcing factors.
- RFMIP-Historical: Combinations of DAMIP (histALL, histNAT, histAER) and RFMIP-Historical will allow us to separate uncertainties in climate response based on specified aerosol evolution from the overall uncertainties in climate response to specified aerosol precursor emissions.
- AerChemMIP: histAER and the experiments of AerChemMIP are useful to understand climate impacts from NTCF.

- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

Aerosols are large uncertainty sources for historical and future climate simulations. The histAER and ssp585AER experiments will allow the climate modelling community to further understand aerosol impacts on climate. Because the hydrological cycle and shortwave radiation are sensitive to aerosols, histAER and ssp585AER may be useful for impact studies regarding water availability and shortwave radiation inputs. A better understanding of aerosol influence on climate is also important for policies controlling aerosol emissions.

Tier 2 experiments

(2.1) Extension of well-mixed GHG-only run (ssp585GHG)

- the experimental design
 - extensions of histGHG runs up to 2100 using the SSP5-8.5 concentrations. As in histGHG, models with interactive chemistry schemes should either run with the chemistry scheme turned off or use a preindustrial climatology of ozone in the radiation scheme.
- the science question and/or gap being addressed with this experiment

Combinations of histALL, histGHG, histNAT, ssp585GHG and SSP5-8.5 (ScenarioMIP) will allow us to make estimates of future temperature changes that are constrained by observed historical changes. Simulated future responses to aerosols and GHG are scaled based on the scaling factors by which the historical simulated responses to aerosols and GHG must be multiplied to best fit observations.
- possible synergies with other MIPs
 - RFMIP-ERF: Combining radiative forcing estimated from RFMIP-ERF and transient climate responses from DAMIP (histALL, histNAT, histGHG, histAER, ssp585GHG and ssp585AER), we can investigate how feedbacks and adjustments vary with forcing factors.
 - ScenarioMIP: Allows the separation of future climate change signals of GHG and the other anthropogenic forcing factors.
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

The histGHG and ssp585GHG simulations can be used to obtain observationally constrained future climate projections, climate sensitivity, TCR and TCRE that wide communities of CM, IAM, IAV and policy makers are interested in.

Tier 3 experiments

(3.1) Ozone-only (histOZ)

- the experimental design

Historical simulations forced by changes in stratospheric and tropospheric ozone concentrations. In models with coupled chemistry, the simulated tropospheric and stratospheric ozone concentrations from the histALL simulations should be prescribed. In models without coupled chemistry the same ozone prescribed in histALL should be prescribed.

- the science question and/or gap being addressed with this experiment

Ozone changes have driven large changes in stratospheric temperature, and are also responsible for driving circulation changes in the Southern Hemisphere, with associated climate impacts. Only a few CMIP5 models carried out histOZ simulations. A larger multi-model ensemble will allow us to more robustly identify in models and perhaps also in observations the influence of ozone on the stratosphere, the tropospheric circulation and climate, and the Southern Ocean and associated carbon cycle aspects.

- possible synergies with other MIPs

- AerChemMIP: Comparison with the ODS-only simulation of AerChemMIP will allow the net climate effects of ODSs to be compared with the net climate effects of ozone changes, a key issue of concern for the WMO/UNEP Ozone Assessment.

- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, (D) policy makers, and (E) Ozone research and policy community.

These simulations will be of most relevance to (E) since they will allow the climate impacts of past ozone changes to be directly assessed.

(3.2) Extension of Ozone-only run (ssp585OZ)

- the experimental design

- extensions of histOZ runs up to 2100 using the ozone concentrations prescribed in the SSP5-8.5 simulation of ScenarioMIP.

- the science question and/or gap being addressed with this experiment

These simulations will allow the contribution of future ozone changes to future climate change to be evaluated, including for example contributions to future Southern Hemisphere atmospheric circulation change, oceanic circulation changes, and carbon cycle impacts. These simulations will be relevant to future WMO/UNEP Ozone Assessments.

- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, (D) policy makers, and (E) Ozone research and policy community.

These simulations will be of most relevance to (E) since they will allow the climate impacts of future ozone changes to be directly assessed.

(3.3) Volcanic-only run (histVLC)

- the experimental design

Historical simulations forced by volcanic forcing as in histALL

- the science question and/or gap being addressed with this experiment

The combination of the histNAT and histVLC simulations will allow us to separate contributions from volcanic (VLC) and solar (NAT minus VLC) forcings to historical climate change. The CMIP5 ensemble tended to overestimate the historical volcanic

response. histVLC will be used for better understanding errors in the volcanic forcing and responses. histNAT and histVLC of DAMIP and Solaronly of SolarMIP allow us to investigate volcanic and solar influences on climate and to check additivity.

- possible synergies with other MIPs
 - GeoMIP & VolMIP: The volcanic response of models can be validated against observations using histVLC, whereas GeoMIP experiments cannot. Thus histVLC experiments will provide useful context for interpreting simulated responses to stratospheric aerosol across models in the GeoMIP experiment. While VolMIP includes simulations of individual eruptions it does not include simulations of the transient response to historical eruptions, allowing better validation of long-term transient effects against observations, and its focus is on 19th century eruptions.
 - SolarMIP: histNAT and histVLC of DAMIP and Solaronly of SolarMIP allow us to investigate volcanic and solar influences on climate and to check additivity.
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

Solar radiation management has recently attracted interest from the wide communities of climate model, IAM, IAV and policy makers. The histVLC experiment can be used for the validation of model responses to stratospheric aerosol injections against observations.

(3.4) Extension of anthropogenic Aerosol-only run (ssp585AER)

- the experimental design
 - extensions of histAER runs up to 2100 using the SSP5-8.5 scenario. Please use the same setup as the histAER runs but the SSP5-8.5 forcing.
- the science question and/or gap being addressed with this experiment
 - Combinations of histALL, histAER, histNAT, ssp585AER and SSP5-8.5 (ScenarioMIP) will allow us to make estimates of future temperature changes that are constrained by observed historical changes. Combining radiative forcing estimated from RFMIP-ERF and transient climate responses from DAMIP (histALL, histNAT, histGHG, histAER, ssp585GHG and ssp585AER), we can investigate how feedbacks and adjustments vary with forcing factors.
- possible synergies with other MIPs
 - RFMIP-ERF: Combining radiative forcing estimated from RFMIP-ERF and transient climate responses from DAMIP (histALL, histNAT, histGHG, histAER, ssp585GHG and ssp585AER), we can investigate how feedbacks and adjustments vary with forcing factors.
 - ScenarioMIP and AerChemMIP: Allows the separation of future climate change signals of aerosols and the other anthropogenic forcing factors.
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

Aerosols are large uncertainty sources for historical and future climate simulations. The histAER and ssp585AER experiments will allow the climate modelling community to further understand aerosol impacts on climate. Because the hydrological cycle and

shortwave radiation are sensitive to aerosols, histAER and ssp585AER may be useful for impact studies regarding water availability and shortwave radiation inputs. A better understating of aerosol influence on climate is also important for policies controlling aerosol emissions.

- If possible, a prioritization of the suggested experiments, including any rationale **

Tier 1

- (1.0) Enlarging ensemble size of histALL in DECK to at least three members (histALL)
- (1.1) Natural-only run (histNAT)
- (1.2) well-mixed GHG-only run (histGHG)
- (1.3) Anthropogenic-Aerosols-only run (histAER)

Tier 2

- (2.1) Extension of well-mixed GHG-only run (ssp585GHG)

Tier 3

- (3.1) Ozone-only (histOZ)
- (3.2) Extension of Ozone-only run (ssp585OZ)
- (3.3) Volcanic-only run (histVLC)
- (3.4) Extension of Anthropogenic-Aerosols-only run (ssp585AER)

To keep consistency between CMIP6 and the previous MIPs, we have histALL, histGHG and histNAT in Tier 1. The analysis of the CMIP5 ensemble has highlighted that aerosols remain the largest source of uncertainty in D&A analyses. Therefore we also propose histAER in the high priority, which will allow the aerosol response to be directly estimated and reduce uncertainties in regression coefficients. histALL, histNAT, histGHG and histAER correspond closely to some experiments of RFMIP and AerChemMIP.

The ssp585GHG simulations will be used for observational constraints of future projections, which will attract interest from wide communities. Therefore we also place ssp585GHG in Tier 2.

- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale. **

No objection to open access.

- List of output and process diagnostics for the CMIP DECK/CMIP6 data request**
 - whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;
 - whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);

- whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.
- whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;
- the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.

Standard outputs as in CMIP5, but subject to change based on further consultation with climate modeling communities and other MIPs (e.g., RFMIP, AerChemMIP, ScenarioMIP, CFMIP, DCP, GMMIP, SolarMIP and ISI-MIP).

- Any proposed contributions and recommendations for**
 - model diagnostics and performance metrics for model evaluation;
 - observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;
 - tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).

Specification of the exact estimate of external forcing data used.

- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms.**

None.

- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF.**

NA

The Decadal Climate Prediction Project (DCPP)

Name of MIP: The Decadal Climate Prediction Project (DCPP)

Co-chairs of MIP: George Boer (george.boer@ec.gc.ca), Doug Smith (doug.smith@metoffice.gov.uk)

Members of the Scientific Steering Committee: The DCPP Panel consists of:

George Boer <George.Boer@ec.gc.ca>,
Christophe Cassou <cassou@cerfacs.fr>,
Francisco Doblas-Reyes <francisco.doblas-reyes@ic3.cat>,
Gokhan Danabasoglu <gokhan@ucar.edu>,
Ben Kirtman <bkirtman@rsmas.miami.edu>,
Yochanan Kushnir <kushnir@ldeo.columbia.edu>
Kimoto Masahide <kimoto@aori.u-tokyo.ac.jp>,
Jerry Meehl <meehl@ucar.edu>,
Rym Msadek <rym.msadek@noaa.gov>,
Wolfgang Mueller <wolfgang.mueller@mpimet.mpg.de >,
Doug Smith <doug.smith@metoffice.gov.uk>,
Karl Taylor <taylor13@llnl.gov>,
Francis Zwiers <fwzwers@uvic.ca>

Link to website. DCPP material is available at (<http://www.wcrp-climate.org/dcp-overview>) and at (<http://dcpp.pacificclimate.org/>).

Goals and overview:

The decadal hindcast component of CMIP will follow the example of other coordinated experiments as a protocol-driven multi-model multi-national project with data production and data sharing as integral components.

The Goals of the decadal prediction component of CMIP include:

- the promotion of the science and practice of decadal prediction (forecasts on timescales up to and including 10 years)
- the provision of information potentially useful for the IPCC WG1 AR6 assessment report and other studies and reports on climate prediction and evolution
- the production and retention of a multi-year multi-model collection of decadal hindcast data in support of climate science and of use to the Global Framework for Climate Services

Scientific aspects of the DCPP to which CMIP-decadal can contribute include:

- a system view (data; analyses; initial conditions; ensemble generation; models and forecast production; post processing and assessment) of decadal prediction
- investigation of broad questions (e.g. sources and limits of predictability, current abilities with respect to decadal prediction, potential applications, ...)
- provision of benchmarks against which to compare improvements in models and prediction quality
- information on processes and mechanisms of interest, e.g., the hiatus, climate shifts, AMOC etc., in a collection of hindcasts

Practical aspects include:

- the coordination of efforts based on agreed experimental structures and timelines in order to promote research, intercomparison, multimodel approaches, applications, and to provide justification for research directions
- a contribution to the development of infrastructure, in particular a multi-purpose data archive of decadal hindcasts useful for a broad range of scientific and application questions and of benefit to national and international climate prediction and climate services organizations

References:

Many decadal prediction papers have been published referring to the decadal component of CMIP5 and to other decadal prediction results. Chapter 11 of the AR5 also gives information and pertinent references.

Overview of the proposed experiments: The DCPP is organized into three Components:

- *Hindcasts:* the design and organization of a coordinated decadal prediction (hindcast) component of CMIP6 in conjunction with the seasonal prediction and climate modelling communities
- *Forecasts:* the ongoing production of experimental quasi-operational decadal climate predictions in support of multi-model annual to decadal forecasting and the application of the forecasts
- *Predictability, mechanisms, and case studies:* the organization and coordination of decadal climate predictability studies and of case studies of particular climate shifts and variations including the study of the mechanisms that determine these behaviours

Overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments:

- basic means and variability statistics pertaining to the forecasts
- bias adjustment information
- individual and multi-model predictability and skill measures

The basic analysis consists of first and second order climate statistics characterizing the hindcasts including as a primary output the geographic distribution of individual and multi-model prediction skills of annual and multi-annual means of climate variables of both practical and theoretical interest.

Proposed timing:

The climate prediction and modelling communities will be surveyed for comments on the proposed experiments and for suggestions as to timing. Timing will depend on community advice as well as on the timing of other aspects of CMIP6.

Experimental design:

- See the Goals and Overview section above for the scientific and practical questions addressed by CMIP-decadal.
- See the attached file (*DCPP27Nov14.pdf*) which includes some of the above material together with
 - A listing of the proposed experiments
 - A prioritized listing of the proposed data archive
- See also the attached file (*DCPP_MIPconnectionTable27Nov14.pdf*) which lists potential connections between the DCPP and other MIPs

Model diagnostics, performance metrics:

For the hindcast experiments, standard forecast skill metrics will be a major output. Other aspects of model behavior of interest include model drift and variability which directly affect forecast performance and predictability as noted above. Many other analyses will be performed by interested participants.

Changes from CMIP5 in documentation and data treatment:

The DCPP Panel's Data Subgroup comprised of K. Taylor (a member of the WIP), F. Doblas-Reyes, Rym Msadek and W. Mueller have been, and are, in the process of developing detailed submissions to the WIP and to the CMIP Panel.

As noted also in the Experimental Design material, the hope is that, in conjunction with the WIP, a coordinated set of “basic” or “common” tiered data tables can be developed across MIPs together with “MIP specific” tables associated with individual MIPs.

The Decadal Climate Prediction Project (DCPP)

The term “decadal prediction” encompasses predictions on annual, multi-annual to decadal timescales. The possibility of making skilful forecasts on these timescales and the ability to do so is investigated by means of predictability studies and retrospective predictions (hindcasts) made using the current generation of climate models and by empirical methods. Skilful decadal prediction of relevant climate parameters is a Key Deliverable of the WCRP’s Grand Challenge of providing Regional Climate Information (<http://www.wcrp-climate.org/index.php/gc-regionalclimate>).

The DCPP envisions three Components and invites groups to participate in any and/or all of them:

- *Hindcasts*: the design and organization of a coordinated decadal prediction (hindcast) component of CMIP6 in conjunction with the seasonal prediction and climate modelling communities
- *Forecasts*: the ongoing production of experimental quasi-operational decadal climate predictions in support of multi-model annual to decadal forecasting and the application of the forecasts
- *Predictability, mechanisms, and case studies*: the organization and coordination of decadal climate predictability studies and of case studies of particular climate shifts and variations including the study of the mechanisms that determine these behaviours

Many scientific and practical questions are involved. The understanding of the physical processes that govern the long timescale predictability of the climate system is vital to improving decadal predictions and these are explored using observations, climate model studies and the results of decadal hindcasts. The analysis of available observations for initializing forecasts, the improvement of the models used in the production of the forecasts, post processing of forecasts including bias adjustment, calibration and multi-model combination, together with the production and application of probabilistic decadal forecasts, are all involved in the research and development efforts contributing to the DCPP. As has been the case for weather forecasting, continued improvement in each of the components of a decadal forecasting system is expected to yield improvement in decadal prediction skill.

The Decadal Climate Prediction Panel in conjunction with the Working Group on Seasonal to Interannual Prediction ([WGSIP](#)) and the Working Group on Coupled Modelling ([WGCM](#)) is a focus for the coordination of the scientific and practical aspects of the DCPP. The description of the first two of the DCPP Components follow. The remaining DCPP Component is under development. A Survey will seek community input on the nature, design and timing of the DCPP.

DCPP Component A: CMIP-decadal

A multi-year multi-model decadal hindcast experiment

The decadal hindcast component of CMIP will follow the example of other coordinated experiments as a protocol-driven multi-model multi-national project with data production and data sharing as integral components.

The Goals of the decadal prediction component of CMIP include:

- the promotion of the science and practice of decadal prediction (forecasts on timescales up to and including 10 years)
- the provision of information potentially useful for the IPCC WG1 AR6 assessment report and other studies and reports on climate prediction and evolution
- the production and retention of a multi-year multi-model collection of decadal hindcast data in support of climate science and of use to the Global Framework for Climate Services ([GFCS](#))

Scientific aspects of the DCPP to which CMIP-decadal can contribute include:

- a system view (data; analyses; initial conditions; ensemble generation; models and forecast production; post processing and assessment) of decadal prediction
- investigation of broad questions (e.g. sources and limits of predictability, current abilities with respect to decadal prediction, potential applications, ...)
- provision of benchmarks against which to compare improvements in models and prediction quality
- information on processes and mechanisms of interest, e.g., the hiatus, climate shifts, AMOC etc., in a collection of hindcasts

Practical aspects of CMIP-decadal include:

- the coordination of efforts based on agreed experimental structures and timelines in order to promote research, intercomparison, multimodel approaches, applications, and to provide justification for research directions
- a contribution to the development of infrastructure, in particular a multi-purpose data archive of decadal hindcasts useful for a broad range of scientific and application questions, and of benefit to national and international climate prediction and climate services organizations

The basic elements of CMIP-decadal are:

- a coordinated set of multi-model multi-member ensembles of retrospective forecasts made each year from 1960 to the present.
- an associated hierarchy of data sets of results generally and readily available to the scientific and applications communities

Consultation and timing for CMIP-decadal:

- the climate prediction and modelling communities will be surveyed for comments on the experimental design
- the timing will depend on community advice and on the timing of other aspects of CMIP6

Details of the proposed CMIP-decadal prediction component are listed below.

CMIP-decadal hindcast protocols

The approach parallels that of the core “Near-term Decadal” component of CMIP5 (Taylor et al., 2009, the version dated 22 January, 2011, together with the Experiment Design Addendum at (http://cmippcmdi.llnl.gov/cmip5/experiment_design.html). An important addendum is the call for hindcasts to be produced every year, rather than every 5 years, over the hindcast period

Table 1. Basic CMIP-decadal experiments

#	Experiment	Notes	# of years
PRIORITY 1: Hindcast/forecast information			
1.	Ensembles of at least 5-year, but much preferably 10-year, <i>hindcasts</i> and <i>forecasts</i>	<p>Coupled models with initialization based on observations</p> <p>Start date <i>every year</i> from 1960 to the present</p> <p>Start date on or before 31 Dec of the year preceding the forecast period (start dates on or before Nov 30 allow DJF seasonal forecast results and are recommended)</p> <p>10 ensemble members preferred (more if possible)</p> <p>Prescribed historical values of atmospheric composition and/or emissions (and other conditions including volcanic aerosols) and a suitable scenario for future years</p>	$60 \times 10 \times (5-10) = 3000-6000$ yrs of integration
PRIORITY 2: To quantify the effects of initialization			
2.	Ensembles of historical and near-future climate <i>simulations</i>	<p>Made with same model as used for hindcasts</p> <p>1850 to present plus 10 forecast years with initial conditions from a preindustrial control simulation</p> <p>10 ensemble members preferred (more if possible)</p> <p>Prescribed historical and future forcing as for Experiment 1</p>	$170 \times 10 = 1700$ yrs of integration

Table 2. Other CMIP-decadal experiments of interest if resources are available

#	Experiment	Notes	# of years
Effects of increased ensemble size			
3.	Increase ensemble size for Experiment 1	m additional ensemble members to improve skill and examine dependence of skill on ensemble size	$60 \times 10 \times m = 600m$ yrs of integration
Improved estimates of hindcast skill			
4.	Ensembles of at least 5-year, but much preferably 10-year, hindcasts and forecasts	As Experiment 1 but with no information from the future with respect to the forecast Radiative and other forcing information (e.g. greenhouse gas concentrations, aerosols etc.) maintained at initial state value or projected in a simple way. No inclusion of volcano or other short term forcing unless available at initial time.	$60 \times 10 \times (5-10) = 3000-6000$ yrs of integration
Improved estimates of the effects of initialization			
5	Ensembles of at least 5-year, but much preferably 10-year, hindcasts and forecasts	Historical climate simulations up to the start dates of corresponding forecast with prescribed forcing Simulations continued from forecast start date but with the same forcing as in Experiment 4, i.e. with NO information from the future with respect to the start date. These are uninitialized versions of Experiment 4 hindcasts.	$60 \times 10 \times (5-10) = 3000-6000$ yrs of integration

Table 3. The CMIP-decadal data archive

Data to be served via the Earth System Grid (ESG) with protocols paralleling CMIP5 although with modifications as specified by the WGCM Infrastructure Panel (WIP).

Priority	Description	Notes
Priority 1 - basic time invariant fields	- as CMIP5 “fx” table	
Priority 1 (new) - monthly means - basic variables - single level files	- surface air temperature, precipitation, mean sea level pressure, sea-ice, snow - fluxes of energy, moisture and momentum (wind stress components) at the TOA and surface - vertically integrated amounts of energy, salt in the ocean - Atlantic MOC	- basic data sets for many investigations
Priority 1 - daily mean data	- as CMIP5 “day” Table	- daily data for applications
Priority 2 - monthly means 2D fields on atmospheric grid	- as CMIP5 “Xmon” Tables	- basic monthly data
Priority 3 - monthly means of atmospheric 3D fields on pressure levels	- as CMIP5 “Xmon” Tables	

The CMIP5 tables referred to are those in the CMIP5 “List of Requested Model Output” at http://cmip-pcmdi.llnl.gov/cmip5/data_description.html. These will be updated

Explanatory comments

Table 1 lists the main CMIP-decadal experiments. Experiment 1 parallels the corresponding CMIP5 decadal prediction experiment in using the same specified forcing during the forecasts as is used for the historical climate simulations of Experiment 2. The specification of historical and scenario forcing introduces some information from the future with respect to the forecast and may lead to slightly overestimated historical forecast skill measures. The main effect is expected to be due to the specification of short term radiative forcings such as volcanoes which occur during a forecast. Other forcings, such as those associated with greenhouse gas and aerosol emissions and/or concentrations, vary comparatively slowly over the five or ten year period of a forecast so affect the results very little. The benefits of using specified forcings include the use of common values across models, the ease of treatment within models, the possibility of documenting improvements with respect to CMIP5 hindcasts, the ability to estimate the effects of initialization by comparing forecasts and simulations which use the same forcings, and the estimation of drift corrections from hindcasts which include the forcings and so are more suitable for the purpose of future decadal forecasts.

Table 2 lists additional experiments which are of interest if resources are available. Experiment 3 increases the ensemble size in order to quantify the expected benefits and as a guide to future forecast applications. Experiments 4 and 5 are lower priority since they demand a large commitment of resources although they are of interest in order to quantify the effects of specifying forcing during the forecast period together with the corresponding effects of initialization.

Table 3 lists the components of the CMIP-decadal data archive. The data are to be served via the Earth System Grid (ESG) and to parallel CMIP5 although with changes to protocols as specified by the WGCM Infrastructure Panel (WIP). At this time, 6-hourly decadal prediction data for dynamical downscaling is not considered a priority although a restricted “test case” may be proposed after further discussion.

The hope is that, in conjunction with the WIP, a coordinated set of “basic” or “common” tiered data tables can be developed across MIPs together with “MIP specific” tables associated with individual MIPs.

DCPP Component B: Experimental real-time multi-model decadal predictions

The real-time decadal prediction component of CMIP will also follow the example of other coordinated experiments as a protocol-driven multi-model multi-national project with data production and data sharing as integral components. It will build on the WMO structure already in place for [seasonal forecasts](#). Forecasts and verification statistics will be made available via the web at WMO designated “Lead Centres” and mirrored via the ESGF. Lead Centres will collect forecast and verification data from designated “Contributing Centres”. Lead Centres will produce a multi-model forecast together with uncertainties, and maintain an archive of previous real-time forecasts from Contributing Centres along with an assessment of its performance as verifying observations become available.

Goals

- the promotion of the science and practice of decadal prediction by generating *real-time* multi-model decadal predictions
- the production and retention of ongoing multi-year multi-model decadal forecast data in support of the Global Framework for Climate Services ([GFCS](#))
- the provision of information potentially useful for the IPCC WG1 AR6 assessment report and other studies and reports on climate prediction and evolution

Scientific aspects

- assess decadal predictions of key variables including temperature, precipitation, mean sea level pressure, the AMO, PDO, Arctic sea ice, the NAO, and tropical storms
- assess uncertainties and generate a consensus forecast
- assess decadal predictions of associated climate impacts of societal relevance

Practical aspects

- the coordination of efforts based on agreed experimental structures and timelines as specified in the protocol below
- a contribution to the development of infrastructure, in particular a multi-purpose data archive of ongoing decadal forecasts useful for a broad range of scientific and application questions and of benefit to national and international climate prediction and climate services organizations

The basic elements

- an ongoing coordinated set of multi-model multi-member ensembles of real-time forecasts updated each year.
- an associated hierarchy of data sets of results generally and readily available to the scientific and applications communities

Details of the proposed CMIP-decadal prediction component are listed below.

CMIP/WMO real-time decadal forecast protocols

Table 1. Basic CMIP-decadal experiments

#	Experiment	Notes	# of years
PRIORITY 1: Real-time forecasts			
1.	Ensembles of ongoing real-time 5-year forecasts	<p>Coupled models with initialization based on observations</p> <p>Start date <i>every year</i> ongoing</p> <p>Start date on or before 31 Dec (start dates on or before Nov 31 allow DJF seasonal forecast results and are recommended)</p> <p>10 ensemble members preferred (more if possible)</p> <p>Atmospheric composition and/or emissions (and other conditions including volcanic aerosols) to follow the RCP4.5 scenario (or its replacement)</p>	10x5=50 yrs of integration for 5-year forecasts
PRIORITY 2: Increased ensemble size and duration			
2a	Increase ensemble size	m additional ensemble members to reduce noise and improve skill	$5m$ yrs of integration
2b	Extend forecast duration to 10 years	To provide forecast information for the period 5 to 10 years ahead	10x5=50 yrs of integration

Table 2. Data

Data to be served via WMO Lead Centres and mirrored on the Earth System Grid (ESG) with protocols paralleling CMIP5 although with modifications as specified by the WGCM Infrastructure Panel (WIP). Data to be archived by March 31st each year.

Priority	Description	Notes
Priority 1 - monthly means - basic variables - single level files	- surface air temperature, precipitation, mean sea level pressure, sea-ice, snow, 500hPa geopotential height, 850hPa temperature - vertically integrated amounts of energy, salt in the ocean - Atlantic MOC - fluxes of energy and moisture at the TOA and surface	- basic data sets for many investigations
Priority 2a - hindcast data for skill assessment and forecast calibration	- surface air temperature, precipitation, mean sea level pressure	Hindcast data for models which have contributed to the multi-model prediction exercise since CMIP5
Priority 2b - hindcast data for skill assessment and forecast calibration	Same variables as Priority 1	Data as provided by new models and participants as part of DCP component A
Priority 3 - monthly means - multi-level data	- as CMIP5 “Xmon” Tables	
Priority 4 - daily mean data	- as CMIP5 “day” Table	- daily data for applications

The CMIP5 tables referred to are those in the CMIP5 “List of Requested Model Output” at http://cmippcmdi.llnl.gov/cmip5/data_description.html. These will be updated.

Once again the hope is that, in conjunction with the WIP, a coordinated set of “basic” or “common” tiered data tables can be developed across MIPs together with “MIP specific” tables associated with individual MIPs

DCPP Component C: Predictability, Mechanisms and Case Studies

The climate system varies on multiple timescales which may be studied using physically based and statistical models. Diagnostic studies investigate climate system behaviour inferred indirectly from a long series of observations and/or model simulations. Prognostic studies investigate the behaviour of models when initial conditions or model features such as physical parameterizations, numerics or forcings are perturbed. The mechanisms involved are of great interest as they underpin the inherent predictability of the system and as they govern forecast skill.

Predictability studies based on perturbations to models may be referred to as “perfect model” studies in the sense that one has perfect knowledge of the modelled climate system in terms of the computer code. They represent “attainable predictability” only to the extent that the model is sufficiently similar to the real system and it is important also to study their applicability. Predictability studies are intended to give an indication of the regions and timescales for which skilful forecasts may be possible and may also be used to study aspects of the physical mechanisms and processes involved.

Case studies are hindcasts which focus on a particular climatic event and the mechanisms and impacts involved. These are typically hindcast studies of an observed event although they can include particular kinds of events in model integrations (variations of AMOC and the associated variation of N Atlantic SSTs in models are an example). Studies of the skill with which a particular event (e.g. the hiatus, climate shift, extreme year, etc.) can be forecast and the mechanisms which support (or perhaps make difficult) a skilful prediction are all of interest.

The DCPP and CLIVAR are together developing coordinated multi-model investigations of a restricted number of mechanism/predictability/case studies believed to be of broad interest to the community. Two research foci currently being developed are:

- Hiatus+: an investigation of the origin, mechanisms and predictability of long timescale variations in global mean temperature (and other variables) including periods of both enhanced warming and cooling with a focus on the current “hiatus”
- Volcanoes and prediction: an investigation of the influence and consequences of volcanic eruptions on decadal prediction and predictability

Both of these areas are being independently investigated by different groups and should benefit from a coordinated approach.

The proposed experiments for these research foci will be reviewed in conjunction with the forthcoming MiKlip-SPECS workshop in February and a detailed description of Component C experiments produced at that time. A subsequent AGCI Workshop in Aspen in the summer of 2015 will provide the opportunity to review Component C and to suggest modifications and/or extensions guided by available results to that time.

Connections between the DCPD and other MIPs

We attempt to distinguish between MIPs with a direct connection to the DCPD, with MIPs which have the potential of an indirect connection, and MIPs for which no connection is envisioned, largely because they are forcing/response experiments. This is summarized in the tables below.

MIPs with Potentially Direct Connections with the DCPD	
DAMIP	<ul style="list-style-type: none"> • both DCPD and DAMIP propose <i>historical</i> climate simulations and projections which extend beyond the “historical” period for a number of years • we will choose a “most probable” <i>common scenario</i> for the period beyond that for which historical CMIP forcing is available
GMMIP	<ul style="list-style-type: none"> • GMMIP also proposes an ensemble of <i>historical</i> simulations as for the DCPD and DAMIP above • the DCPD hindcast results can be analyzed for skill and predictability for aspects of the monsoons • some of the proposed pacemaker integrations could potentially align with DCPD Component C integrations
RFMIP and ScenarioMIP	<ul style="list-style-type: none"> • decadal predictions depend on forcings as they affect system initial conditions and forecasts over the historical-to-near-future period • the DCPD (as well as DAMIP and GMMIP) envision an ensemble of simulations for the historical period and therefore <i>depend on</i> the clear and timely specification of CMIP6 historical forcings as informed by RFMIP and ScenarioMIP • for forecasts and simulations beyond the historical period, the DCPD (and DAMIP) also depend on scenario information, again as informed by RFMIP and ScenarioMIP • for near-future simulations and forecasts a “most probable” common scenario should be chosen across the DCPD and DAMIP
VolMIP	<ul style="list-style-type: none"> • the results of the idealized volcano forcing/response VolMIP experiments are of direct interest to the DCPD Component C volcano studies • the proposed DCPD Component C experiments investigating the effects of historical volcanoes in the context of initialized decadal forecasts is directly related to VolMIP interests • DCPD and VolMIP will propose a number of joint experiments

MIPs with Potentially Indirect Connections with the DCP	
AerChemMIP	<ul style="list-style-type: none"> • The effects of aerosols in particular are of interest as they affect prediction and predictability. They will be specified as external forcing for the DCP so it is important that there be agreement on these forcings across CMIP6 • Component C of the DCP will consider the effect of volcanic aerosols on prediction and predictability but we do not see a more intimate connection with AerChemMIP at this time
C4MIP	<ul style="list-style-type: none"> • it is of interest to consider the prediction and predictability of biogeochemical quantities if ESMs are used for the DCP and this would benefit from a connection with C4MIP <ul style="list-style-type: none"> ◦ some coordination of the data retention tables between C4MIP and the DCP would be necessary ◦ the availability of biogeochemical verification data over the DCP prediction period is a key connection
ENSOMIP	<ul style="list-style-type: none"> • the DCP hindcast results can be analyzed for ENSO skill and predictability
FAFMIP	<ul style="list-style-type: none"> • DCP results can provide hindcasts and estimates of predictability for sea level in those models for which sea level is a prognostic variable • DCP results can provide hindcast information on the steric component of sea level for other models without prognostic sea level • DCP results can provide forecast surface flux anomalies if these are of interest
HighResMIP	<ul style="list-style-type: none"> • some DCP hindcast experiments may be performed with high resolution models
LS3MIP	<ul style="list-style-type: none"> • the DCP hindcasts can be analyzed wrt to skill (to the extent that verifying data is available) and predictability for land surface variables
SIMIP	<ul style="list-style-type: none"> • the DCP hind/forecasts, while not specifically directed toward predictions of sea ice, can be analyzed in terms of cryospheric variables if suitable quantities are retained
SolarMIP	<ul style="list-style-type: none"> • the DCP hindcasts can nominally be analyzed wrt to skill and predictability as they relate to solar variability during the hindcast period

MIPs where a Connection with the DCP is not foreseen
CFMIP
GeoMIP
GDDEX
ISMIP6
LUMIP
nonlinMIP
OCMIP6
PDRMIP
PMIP
SensMIP1
CORDEX <ul style="list-style-type: none"> • the difficulties of data retention, bias adjustment etc. make the downscaling of decadal hind/forecasts a difficult and expensive task • consultations with CORDEX co-chairs and others suggests that the time is not ripe
VIAAB

El Niño Response and Teleconnections under Climate Change (ENSOMIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

Co-chairs: Mat Collins (M.Collins@exeter.ac.uk),
Scott Power (S.Power@bom.gov.au)

Scientific Steering Committee: Wenju Cai, Eric Guilyardi, Matthieu Lengaigne, Masahiro Watanabe, Andrew Wittenberg, Sang-Wook Yeh.

Website: TBA

Goal of ENSOMIP

The basic pattern of climate change in the tropical Pacific in climate models shows an increase in rainfall towards the central and eastern equatorial Pacific, anchored to a broad equatorial local maximum in SST warming (Xie et al., 2010). This leads to an intensification and eastward shift of rainfall anomalies associated with canonical El Niño events (Chung et al., 2014; Power et al., 2013) potentially leading to changes in both tropical and mid-latitude teleconnections (Schneider et al., 2009; Weare, 2013; Zhou et al., 2014). When defining El Niño in terms of rainfall anomalies, this intensification leads to an increase in the frequency of 'extreme' events, in which significant convective rainfall moves from its climatological west Pacific position into the central and eastern Pacific, is found (Cai et al., 2014).

While these changes seem relatively robust in coupled models, the models themselves exhibit persistent systematic biases, in particular a cold tongue that is generally too cold and extends too far into the west Pacific with associated rainfall and trade wind biases. SST anomalies associated with ENSO in coupled models also show a wide diversity in terms of evolution, spatial pattern, amplitude and frequency (Bellenger et al., 2014). In addition, the mechanisms responsible for changes in El Niño teleconnections, in particular those associated with ENSO events that have different spatial patterns, amplitude and evolution, are not well understood.

The aim of this MIP is to characterize and understand the response of the atmosphere to ENSO SST anomalies under enhanced CO₂ conditions under controlled conditions of identical ENSO SST anomalies, superimposed on a mean pattern of SST change that is the same in each atmosphere-only model simulation. Thus we can 'control out' some of the uncertainty associated with mean SST biases and with ENSO SST diversity.

References:

- Bellenger, H., Guilyardi, E., Leloup, J., Lengaigne, M. and Vialard, J., 2014. ENSO representation in climate models: from CMIP3 to CMIP5. *Climate Dynamics*, 42(7-8): 1999-2018.
- Cai, W., Borlace, S., Lengaigne, M., van Rensch, P., Collins, M., Vecchi, G., Timmermann, A., Santoso, A., McPhaden, M.J., Wu, L., England, M.H., Wang, G., Guilyardi, E. and Jin, F.-F., 2014. Increasing frequency of extreme

- El Nino events due to greenhouse warming. *Nature Climate Change*, 4(2): 111-116.
- Chung, C., Power, S., Arblaster, J., Rashid, H. and Roff, G., 2014. Nonlinear precipitation response to El Nino and global warming in the Indo-Pacific. *Climate Dynamics*, 42(7-8): 1837-1856.
- Power, S., Delage, F., Chung, C., Kociuba, G. and Keay, K., 2013. Robust twenty-first-century projections of El Nino and related precipitation variability. *Nature*, 502(7472): 541-+.
- Schneider, E., Fennessy, M. and Kinter, J., 2009. A Statistical-Dynamical Estimate of Winter ENSO Teleconnections in a Future Climate. *Journal of Climate*, 22(24): 6624-6638.
- Weare, B., 2013. El Nio teleconnections in CMIP5 models. *Climate Dynamics*, 41(7-8): 2165-2177.
- Xie, S., Deser, C., Vecchi, G., Ma, J., Teng, H. and Wittenberg, A., 2010. Global Warming Pattern Formation: Sea Surface Temperature and Rainfall. *Journal of Climate*, 23(4): 966-986.
- Zhou, Z-Q, Xie, S-P, Zheng, X-T, Liu, Q, Wang, H. 2014. Global warming-induced changes in El Niño teleconnections over 1 the North Pacific and North America. In press.

Overview of Experiments:

This goal may be achieved by performing a new experiment that builds on the list of experiments proposed by CFMIP. Following their naming convention, we call this experiment 'ampFuture4xCO2'. It is an AMIP experiment where SSTs are subject to a composite SST warming pattern derived from coupled models, scaled to a global mean of 4K and in which the CO₂ is quadrupled from preindustrial values. Changing both the SSTs and the CO₂ is important in order to capture the impact of both forcings on precipitation.

While the experiment might be subsumed into the CFMIP suite of runs, we have decided to keep them separate. The scientific focus of ENSOMIP is slightly different from that of CFMIP. We aim to produce the most realistic simulation of future ENSO events that may be used for process understanding and for impacts studies. The CFMIP project is clear focused on understanding clouds feedbacks in models. In addition, by having a clearly identified MIP it will allow us to plan a series of diagnostic studies within the ENSO community. A further extension to ENSOMIP could include experiments with different patterns of SST change specified to the sensitivity to this aspect of climate change.

Number of ensemble members: Ideally more ensemble members would be required to increase signal to noise, particularly when separating the impacts of ENSO events with a different structure

Overview of Analysis:

A single paper, giving the headline results, would be coordinated by the steering group. Thereafter it would be up to the community to focus on more detailed analysis of processes. We expect these experiments could be useful in other areas of climate projection science.

Timing:

The experiments would be reliant on the CFMIP group to define the composite pattern of mean SST in the future. It would be sensible to perform the ENSOMIP experiment at the same time as the CFMIP atmosphere-only experiments.

Flux-anomaly-forced model intercomparison experiment (FAFMIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

Steering committee

Jonathan Gregory (chair, j.m.gregory@reading.ac.uk), Detlef Stammer (detlef.stammer@zmaw.de), Stephen Griffies (stephen.griffies@noaa.gov)

Goals and overview of experiments

FAFMIP is proposed in support of the WCRP Grand Challenge on sea level rise and regional impacts. Projections of regional sea level change by CMIP5 AOGCMs, like earlier AOGCM generations, show a substantial spread due to the different models' differing simulations of regional ocean density and circulation changes, especially in high latitudes and the North Atlantic (Yin, 2012, 10.1029/2012GL052947; Bouttes et al., 2012, 10.1029/2012GL054207; IPCC AR5 WG1 chapter 13, Church et al., 2013; Slangen et al, 2014, 10.1007/s10584-014-1080-9). By applying flux perturbations from a range of CMIP5 models to the same AOGCM, previous analyses have shown that a substantial fraction, but not all, of the diversity of sea level projections arises from the spread in AOGCM projections of changes in surface fluxes of momentum (windstress), heat and freshwater (Bouttes et al., 2012, cited above; Bouttes et al., 2014, 10.1007/s00382-013-1973-8; Bouttes and Gregory, 2014, 10.1088/1748-9326/9/3/034004).

In the FAFMIP experiments, a prescribed set of surface flux perturbations will be applied to the ocean. These perturbations will be obtained from the ensemble-mean changes simulated at time of doubled CO₂ by CMIP5 AOGCMs under the 1pctCO₂ scenario, so they are representative of projected anthropogenic climate change. The aims of the experiments are:

- to quantify the difference in the geographical patterns of sea level change due to ocean density and circulation change simulated by the models, when given common surface flux perturbations.
- to provide information about the efficiency and interior distribution of ocean heat uptake in response to climate change; the AOGCM spread in these phenomena contributes to their spread in transient climate response and global mean sea level rise due to thermal expansion.
- to provide information about the sensitivity of the Atlantic meridional overturning circulation (AMOC) to prescribed buoyancy forcing of the character expected for CO₂ forcing, rather than idealised freshwater forcing such as has been used in previous AMOC intercomparisons; change in the AMOC is of relevance to both regional and global sea level rise, as well as to regional climate change.

The FAFMIP experiments are aimed at increased physical understanding. They are not themselves policy-relevant scenarios, but obviously the uncertainties in projection of global and regional sea level and AMOC change are of great policy relevance.

The steering committee undertakes to ensure that a paper on the FAFMIP design will be prepared, and all participants will be encouraged to collaborate in producing a paper on the results. At the time of writing (25 Nov 2014) there are nine groups who plan to run FAFMIP experiments (ACCESS, CanESM, CNRM/CERFACS, GISS, GFDL, MIROC6, MPI, MRI, UKESM) and one other possibility (IPSL).

Design of experiments (see <http://www.met.reading.ac.uk/~jonathan/FAFMIP>)

All the experiments will add anomalies to the surface fluxes computed by the AOGCM (like a flux adjustment). The fluxes themselves will not be replaced because this would typically cause a very large climate drift and possible instability, and is technically more complicated than adding an anomalous flux. The surface flux anomalies are a function of (longitude, latitude, time of year) and constant throughout the experiments, which are proposed to be 70 years long (but shorter experiments

would still be useful if 70 years cannot be afforded). The experiments will branch from and be analysed by comparison with the standard CMIP DECK pre-industrial control. All the experiments have pre-industrial atmospheric conditions.

There are three tier-1 experiments, most important first. The **bold** word is the name of the experiment.

wind: Impose a perturbation in surface zonal and meridional windstress. We propose this experiment first because the windstress change appears to have the largest effect on sea level in CMIP5 scenario experiments. In addition to its relevance to sea level, this experiment will also be of interest regarding the phenomenon of eddy saturation (relative insensitivity of the circumpolar circulation to windstress change), especially in eddy-resolving models, and to study the influence of windstress change on advecting circumpolar deep water towards the Antarctic continental shelf, where it could affect ice-shelf melting and hence sea-level rise through the effect on ice-sheet dynamics (a different aspect of the Grand Challenge). The perturbation is made to windstress, rather than to wind speed in the atmosphere, because windstress is the flux experienced by the ocean. AOGCMs typically use other diagnostics of wind speed to supply turbulent mixing energy to the ocean in addition to windstress. Perturbing these quantities is not included in the proposed design at present.

heat: Impose a perturbation in surface heat flux, which is second in importance in its influence on patterns of sea level change. It has also been found in a previous analyses to be the main influence on AMOC change. In an AOGCM, imposing a heat flux perturbation is not straightforward, because it alters the SST, which affects the surface heat flux calculated by the atmosphere model and tends to cancel out the perturbation. In this experiment, we propose to use a passive tracer to avoid this feedback (see documents on website), and also the tier-2 experiments (see below).

water: Impose a perturbation in the surface freshwater flux (including the contribution from runoff change). This is the least influential surface flux.

There are three tier-2 experiments.

- ◆ There is a pair of experiments which implement an alternative method to impose the heat flux anomaly, in which the sea surface temperature used by the atmosphere model is prescribed from the monthly control climatology, rather than the present state of the ocean, thus suppressing the feedback due to the heat flux perturbation. Because this interference with the coupling will introduce a climate drift, this method requires its own control (**heataltcontrol**, with no perturbative flux) as well as the anomaly experiment (**heataltanomaly**). Comparison of **heat–piControl** with **heataltanomaly–heataltcontrol** will allow the effect of ocean advection on surface heat flux feedback to be assessed (cf. Winton et al., 2013, 10.1175/JCLI-D-12-00296.1).
- ◆ In the **all** experiment, the anomalous fluxes of wind, heat and water are simultaneously applied, using the passive-tracer method for heat.

Diagnostics

No changes to the standard CMIP set of diagnostics or CF, CMOR or ESG are required. The analysis of sea level change will mainly use zos, zostoga, thetaso and so. Analyses of ocean heat uptake efficiency will use thetaso. Analyses of the AMOC will use msftmyz, msftyzyz, uo and vo. It is strongly recommended that 3D ocean diagnostics should be implemented for monthly-mean temperature and salinity tendencies ($\partial T/\partial t$ and $\partial S/\partial t$) due to the various physical processes which modify the state (advection, diffusion, etc.). These diagnostics have been included in Table 2.9 of the recommendations from the CLIVAR Ocean Model Development Panel committee on CMIP6 ocean model output for use in all CMIP6 experiments, including DECK for instance, but their usefulness for FAFMIP is particularly noted there. If the $\partial/\partial t$ diagnostics are not submitted for all experiments, for FAFMIP they are particularly requested for the DECK piControl as well as for the FAFMIP experiments, since piControl is the control for **wind**, **heat**, **water** and **all**.

Proposed timing

The required input fields will be prepared and tested by the end of 2014 and experiments can be done thereafter by any interested groups.

Geoengineering Model Intercomparison Project (GeoMIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

Name of MIP: The Geoengineering Model Intercomparison Project (GeoMIP)

Co-chairs of MIP: Ben Kravitz (ben.kravitz@pnnl.gov) and Alan Robock (robock@envsci.rutgers.edu)

Members of the Scientific Steering Committee:

- Ben Kravitz (Pacific Northwest National Laboratory, ben.kravitz@pnnl.gov)
- Alan Robock (Rutgers University, robock@envsci.rutgers.edu)
- Olivier Boucher (Laboratoire de Météorologie Dynamique, IPSL, CNRS/UMPC, olivier.boucher@lmd.jussieu.fr)
- Mark G. Lawrence (Institute for Advanced Sustainability Studies, mark.lawrence@iass-potsdam.de)
- John C. Moore (Beijing Normal University, john.moore.bnu@gmail.com)
- Ulrike Niemeier (Max Planck Institute for Meteorology, ulrike.niemeier@mpimet.mpg.de)
- Trude Storelvmo (Yale University, trude.storelvmo@yale.edu)
- Simone Tilmes (National Center for Atmospheric Research, tilmes@ucar.edu)
- Robert Wood (University of Washington, robwood2@uw.edu)

Website: <http://climate.envsci.rutgers.edu/GeoMIP/>

Goal of the MIP and a brief overview: As anthropogenic climate change continues unabated, society is exploring research into options for addressing the effects of greenhouse gas emissions. One of these options could be geoengineering. Therefore, research on the climate effects and impacts of geoengineering is crucial. The goal of GeoMIP is to understand robust climate model response to geoengineering.

GeoMIP directly addresses the key CMIP6 theme of geoengineering. Moreover, the study of geoengineering, particularly through climate model simulations under GeoMIP, have proven to address multiple key CMIP6 focus areas, including clouds/circulation, chemistry/aerosols, characterizing forcing, carbon cycle, regional climate/extremes, scenarios, and ocean/sea ice. The effects and impacts of global scale interventions in the climate system are broad, and GeoMIP is well poised to address such breadth with its wide variety of participants and interests.

References:

1. S. Tilmes, M. J. Mills, U. Niemeier, H. Schmidt, A. Robock, B. Kravitz, J.-F. Lamarque, G. Pitari, and J. M. English (2014), A new Geoengineering Model Intercomparison Project (GeoMIP) experiment designed for climate and

- chemistry models, *Geoscientific Model Development Discussions*, 7, 5447-5464.
2. Xia, L., A. Robock, J. N. S. Cole, C. L. Curry, D. Ji, A. Jones, B. Kravitz, J. C. Moore, H. Muri, U. Niemeier, B. Singh, S. Tilmes, S. Watanabe, and J.-H. Yoon (2014), Solar Radiation Management impacts on agriculture in China: A case study in the Geoengineering Model Intercomparison Project (GeoMIP), *Journal of Geophysical Research*, 119, 8695-8711, doi:10.1002/2013JD020630.
 3. Kravitz, B., D. G. MacMartin, A. Robock, P. J. Rasch, K. L. Ricke, J. N. S. Cole, C. L. Curry, P. J. Irvine, D. Ji, D. W. Keith, J. E. Kristjánsson, J. C. Moore, H. Muri, B. Singh, S. Tilmes, S. Watanabe, S. Yang, and J.-H. Yoon (2014), A multi-model assessment of regional climate disparities caused by solar geoengineering, *Environmental Research Letters*, 9, 074013, doi:10.1088/1748-9326/9/7/074013.
 4. Irvine, P. J., O. Boucher, B. Kravitz, K. Alterskjær, J. N. S. Cole, D. Ji, A. Jones, D. J. Lunt, J. C. Moore, H. Muri, U. Niemeier, A. Robock, B. Singh, S. Tilmes, S. Watanabe, and J.-H. Yoon (2014), Key factors governing uncertainty in the response to sunshade geoengineering from a comparison of the GeoMIP ensemble and a perturbed parameter ensemble, *Journal of Geophysical Research*, 119, 7946-7962, doi:10.1002/2013JD020716.
 5. Huneus, N., O. Boucher, K. Alterskjær, J. N. S. Cole, C. L. Curry, D. Ji, A. Jones, B. Kravitz, J. E. Kristjánsson, J. C. Moore, H. Muri, U. Niemeier, P. J. Rasch, A. Robock, B. Singh, H. Schmidt, M. Schulz, S. Tilmes, S. Watanabe, and J.-H. Yoon (2014), Forcings and feedbacks in the GeoMIP ensemble for a reduction in solar irradiance and increase in CO₂, *Journal of Geophysical Research*, 119, 5226-5239, doi:10.1002/2013JD021110.
 6. Curry, C. L., J. Sillmann, D. Bronaugh, K. Alterskjær, J. N. S. Cole, B. Kravitz, J. E. Kristjánsson, H. Muri, U. Niemeier, A. Robock, and S. Tilmes, A multi-model examination of climate extremes in an idealized geoengineering experiment, *Journal of Geophysical Research*, 119, 3900-3923, doi:10.1002/2013JD020648.
 7. Pitari, G., V. Aquila, B. Kravitz, A. Robock, S. Watanabe, N. De Luca, G. Di Genova, E. Mancini, S. Tilmes, and I. Cionni (2014), Stratospheric ozone response to sulfate geoengineering: Results from the Geoengineering Model Intercomparison Project (GeoMIP), *Journal of Geophysical Research*, 119, 2629-2653, doi:10.1002/2013JD020566.
 8. Berdahl, M., A. Robock, D. Ji, A. Jones, B. Kravitz, J. C. Moore, and S. Watanabe (2014), Arctic cryosphere response in the Geoengineering Model Intercomparison Project (GeoMIP) G3 and G4 scenarios, *Journal of Geophysical Research*, 119, 1308-1321, doi:10.1002/2013JD020627.
 9. Moore, J. C., A. Rinke, X. Yu, D. Ji, X. Cui, Y. Li, K. Alterskjær, J. E. Kristjánsson, O. Boucher, N. Huneus, B. Kravitz, A. Robock, U. Niemeier, H. Schmidt, M. Schulz, S. Tilmes, and S. Watanabe (2014), Arctic sea ice and atmospheric circulation under the GeoMIP G1 scenario, *Journal of Geophysical Research*, 119, 567-583, doi:10.1002/2013JD021060.
 10. Kravitz, B., P. J. Rasch, P. M. Forster, T. Andrews, J. N. S. Cole, P. J. Irvine, D. Ji, J. E. Kristjánsson, J. C. Moore, H. Muri, U. Niemeier, A. Robock, B. Singh, S. Tilmes, S. Watanabe, and J.-H. Yoon (2013), An energetic perspective on

- hydrological cycle changes in the Geoengineering Model Intercomparison Project (GeoMIP), *Journal of Geophysical Research*, *118*, 13087-13102, doi:10.1002/2013JD020502.
11. K. Alterskjær, J. E. Kristjánsson, O. Boucher, H. Muri, U. Niemeier, H. Schmidt, M. Schulz, and C. Timmreck (2013), Sea salt injections into the low-latitude marine boundary layer: The transient response in three Earth System Models, *Journal of Geophysical Research*, *118*, 12195-12206, doi:10.1002/2013JD020432.
 12. Niemeier, U., H. Schmidt, K. Alterskjær, and J. E. Kristjánsson (2013), Solar irradiance reduction via climate engineering--impact of different techniques on the energy balance and the hydrological cycle, *Journal of Geophysical Research*, *118*, 11905-11917, doi:10.1002/2013JD020445.
 13. Kravitz, B., A. Robock, P. M. Forster, J. M. Haywood, M. G. Lawrence, and H. Schmidt (2013), An overview of the Geoengineering Model Intercomparison Project (GeoMIP), *Journal of Geophysical Research*, *118*, 13103-13107, doi:10.1002/2013JD020569.
 14. Tilmes, S., J. Fasullo, J.-F. Lamarque, D. R. Marsch, M. Mills, K. Alterskjær, O. Boucher, J. N. S. Cole, C. L. Curry, J. M. Haywood, P. J. Irvine, D. Ji, A. Jones, D. B. Karam, B. Kravitz, J. E. Kristjánsson, J. C. Moore, H. O. Muri, U. Niemeier, P. J. Rasch, A. Robock, H. Schmidt, M. Schulz, B. Singh, S. Watanabe, S. Yang, and J.-H. Yoon (2013), The hydrological impact of geoengineering in the Geoengineering Model Intercomparison Project (GeoMIP), *Journal of Geophysical Research*, *118*(19), 11036-11058, doi:10.1002/jgrd.50868.
 15. Kravitz, B., P. M. Forster, A. Jones, A. Robock, K. Alterskjær, O. Boucher, A. K. L. Jenkins, H. Korhonen, J. E. Kristjánsson, H. Muri, U. Niemeier, A.-I. Partanen, P. J. Rasch, H. Wang, and S. Watanabe (2013), Sea spray geoengineering experiments in the Geoengineering Model Intercomparison Project (GeoMIP): Experimental design and preliminary results, *Journal of Geophysical Research*, *118*(19), 11175-11186, doi:10.1002/jgrd.50856.
 16. Jones, A., J. M. Haywood, K. Alterskjær, O. Boucher, J. N. S. Cole, C. L. Curry, P. J. Irvine, D. Ji, B. Kravitz, J. E. Kristjánsson, J. C. Moore, U. Niemeier, A. Robock, H. Schmidt, B. Singh, S. Tilmes, S. Watanabe, and J.-H. Yoon (2013), The impact of abrupt suspension of solar radiation management (termination effect) in experiment G2 of the Geoengineering Model Intercomparison Project (GeoMIP), *Journal of Geophysical Research*, *118*(17), 9743-9752, doi:10.1002/jgrd.50762.
 17. Kravitz, B., K. Caldeira, O. Boucher, A. Robock, P. J. Rasch, K. Alterskjær, D. Bou Karam, J. N. S. Cole, C. L. Curry, J. M. Haywood, P. J. Irvine, D. Ji, A. Jones, J. E. Kristjánsson, D. J. Lunt, J. Moore, U. Niemeier, H. Schmidt, M. Schulz, B. Singh, S. Tilmes, S. Watanabe, S. Yang, and J.-H. Yoon (2013), Climate model response from the Geoengineering Model Intercomparison Project (GeoMIP), *Journal of Geophysical Research*, *118*(15), 8320-8332, doi:10.1002/jgrd.50646.
 18. Haywood, J. M., A. Jones, N. Bellouin, and D. Stephenson (2013), Asymmetric forcing from stratospheric aerosols impacts Sahelian rainfall, *Nature Climate Change*, *3*, 660-665, doi:10.1038/nclimate1857.

19. Schmidt, H., K. Alterskjær, D. Bou Karam, O. Boucher, A. Jones, J. E. Kristjánsson, U. Niemeier, M. Schulz, A. Aaheim, F. Benduhn, M. Lawrence, and C. Timmreck (2012), Solar irradiance reduction to counteract radiative forcing from a quadrupling of CO₂: Climate responses simulated by four Earth system models, *Earth System Dynamics*, 3, 63-78, doi:10.5194/esd-3-63-2012.
20. Kravitz, B., A. Robock, O. Boucher, H. Schmidt, K. E. Taylor, G. Stenchikov, and M. Schulz (2011), The geoengineering model intercomparison project (GeoMIP), *Atmospheric Science Letters*, 12, 162-167, doi:10.1002/asl.316.

An overview of the proposed experiments: Please see the attached document detailing the experiment design and the scientific questions/motivation behind each experiment. We are proposing four Tier 1 experiments to be included in this phase of GeoMIP. This document will form the basis of a manuscript to be submitted to *Geoscientific Model Development*.

All experiments are built on the CMIP DECK or the ScenarioMIP Tier 1 simulations¹ and will be useful complements to the standard CMIP analyses. We anticipate the model output from GeoMIP6 simulations will be of use to the climate modeling community (comparisons of future trajectories of climate change and understanding climate feedbacks), the Impacts Adaptation and Vulnerability (IAV) community (understanding how geoengineering can alleviate, modify, or exacerbate the impacts of climate change), and policy makers (determining which geoengineering technologies may work, the possible effects of geoengineering, and which climate impacts geoengineering can alter).

An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments: GeoMIP currently has a special joint issue of *Atmospheric Chemistry and Physics* and *Geoscientific Model Development*, which will include the following analyses:

1. Special issue introduction – Ben Kravitz, Alan Robock, others
2. G1ocean-albedo overview – Ben Kravitz
3. G4cdnc overview - Hannele Korhonen
4. G4sea-salt overview – Jon Egill Kristjánsson
5. Vegetation response in G1 - Susanne Glienke and Pete Irvine
6. Carbon cycle feedbacks – Andrew Lenton
7. Sea level rise in G4 – John Moore
8. Extreme events in G4 – John Moore
9. Agricultural impacts – Lili Xia and Alan Robock
10. Stratospheric dynamics in G3 and G4 – Hauke Schmidt
11. Comparison of G3 and G3solar – Simone Tilmes
12. Effects on ENSO – Corey Gabriel and Alan Robock

¹ As of the drafting of this document, the ScenarioMIP Tier 1 scenarios are not finalized. Both Tier 1 and Tier 2 will have high, medium, and low forcing scenarios (referring to the level of anthropogenic forcing exhibited during the scenarios). The GeoMIP experiments will be based upon the Tier 1 ScenarioMIP experiments, regardless of what their final form may be.

13. Southern Hemisphere Circulation in G3 and G3solar – Steven Phipps
14. Effects on the QBO – Ulrike Niemeier
15. Effects on the Indian Monsoon – Saroj Kanta Mishra
16. Ocean circulation in G1 – Phil Rasch
17. Scavenging processes in G4sea-salt – Hailong Wang
18. Stratospheric sulfate aerosol microphysics – Jason English

Each potential paper is listed with a responsible first author who will be leading the analysis. All of these papers will involve analysis of currently existing model output. As output from the GeoMIP6 experiments becomes available, any resulting papers will also be included in this special issue.

There is also an experiment design paper (attached) that will be submitted to the CMIP6 special issue. We will request that this manuscript be cross-linked to the aforementioned GeoMIP special issue.

Proposed timing: The GeoMIP community is currently engaged in analyzing output from the already existing model output from previous experiments; results from this will be included in the previously described special joint issue, which will be open for submission for two years. Simulations of the new experiments will be conducted in concordance with the CMIP6 timetable; they will begin as soon as the final MIP proposals are submitted to the CMIP Panel (31 March 2015). Analyses will begin shortly thereafter, with presentations on preliminary work to be given at the next GeoMIP meeting 20-24 July 2015 at NCAR. The community is constantly generating new ideas for experiments (for examples, please see the GeoMIP Testbed in the attached document), and we anticipate that GeoMIP is likely to continue in a rolling fashion, designing new experiments as ideas emerge.

A prioritization of the suggested experiments, including any rationale: Please see the attached document describing the experiments.

All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.: We are quite happy with open access and strongly encourage any interested party to contact us for guidance on how to download GeoMIP output and an up-to-date list on the current areas of analysis that are already being explored.

List of output and process diagnostics for the CMIP DECK/CMIP6 data request: GeoMIP has proven to have broad appeal to a wide variety of communities, so we do not wish to restrict the output to any narrow set of variables that might be requested by any particular focus of analysis. We are requesting that all participating models output the standard set of variables requested by CMIP6. We have been quite happy with the wide variety of output that has been saved by modeling groups in the past.

We anticipate that our output will be used for both understanding of the climate system and for downstream users, such as the impacts assessment community. We have begun preliminary discussions with broader communities, such as the impacts assessment community, as well as other social science communities, to determine how our output can best be used to promote understanding of geoengineering.

Any proposed contributions and recommendations for

- **model diagnostics and performance metrics for model evaluation;**
- **observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;**
- **tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).**

GeoMIP itself is unlikely to contribute any of these materials, although individual researchers may provide contributions.

Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms. No requested changes.

Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF. No requested changes.

The main criteria for MIPs to be endorsed for CMIP6 are

- **The MIP and its experiments address at least one of the key science questions of CMIP6;** Geoengineering is one of the key themes in CMIP6, which GeoMIP directly addresses. GeoMIP also directly addresses all three broad scientific questions of CMIP6. Understanding the Earth System's response to forcing is at the core of all geoengineering studies, and GeoMIP seeks to understand the commonalities and differences of model response to those forcings. GeoMIP studies have also revealed that geoengineering simulations can be a useful method of constraining model feedbacks, thus providing a unique way of revealing the sources of systematic model biases. Geoengineering is a method of directly addressing future climate change, and there have been several studies published showing how geoengineering can manage some of the uncertainties in the climate system. All of the proposed GeoMIP6 experiments (see attached document) are based upon future climate change scenarios, and two of them directly address the question of using geoengineering as part of a portfolio of approaches to manage future climate change.
- **The MIP demonstrates connectivity to the DECK experiments and the CMIP6 Historical Simulation;** The proposed GeoMIP6 experiments are entirely based upon the DECK experiments or the ScenarioMIP Tier 1 experiments¹, which are extensions of the CMIP6 Historical Simulation.
- **The MIP adopts the CMIP modeling infrastructure standards and conventions;** All of the output from our experiments will be processed and

documented in accordance with CMOR standards. The same teams that are preparing the DECK and CMIP6 Historical Simulations will also be conducting simulations for GeoMIP6.

- **All experiments are tiered, well-defined, and useful in a multi-model context and don't overlap with other CMIP6 experiments;** The attached document provides a detailed description of all of the proposed experiments, which are divided into two tiers. Justification for completing the experiments in a multi-model context is provided for each experiment. We have identified synergies between our proposed experiments and other participating MIPs, but we are unaware of any other group proposing geoengineering experiments.
- **Unless a Tier 1 experiment differs only slightly from another well-established experiment, it must already have been performed by more than one modeling group;** The proposed experiment G1ext is quite similar to experiment G1, which has been simulated by 13 modeling groups. Experiments G6sulfur and G6solar are similar to past experiment G3, which has been performed by 5 modeling groups. These two experiments are proposed here because they are better defined than G3 was, easier to simulate than G3 (thus hopefully garnering greater participation from modeling groups), and more relevant to future scenario experiments. Test simulations of G7cirrus have been performed by at least three models; preliminary results and citations are included in the attached document.
- **A sufficient number of modeling centers are committed to performing all of the MIP's Tier 1 experiments and providing all the requested diagnostics needed to answer at least one of its science questions;** We currently have 16 models participating in GeoMIP. Commitment to performing the proposed GeoMIP6 experiments has been pledged by modeling teams representing current or projected new versions of the models BNU-ESM, CanESM2, CCSM4, CESM-CAM5, CSIRO-Mk3L, EC-Earth, GISS-E2-R, HadGEM2-ES, IPSL-CM5A-LR, MIROC-ESM, MPI-ESM-LR, and NorESM1-M (12 models total). In addition, diagnostic simulations will be performed by CESM-WACCM, and there will be substantial participation in GeoMIP by chemistry climate models through the proposed experiment G4-SSA.
- **The MIP presents an analysis plan describing how it will use all proposed experiments, any relevant observations, and specially requested model output to evaluate the models and address its science questions;** The attached document provides a detailed description of the science questions central to GeoMIP6 and how each experiment will address those questions. We have also provided above a list of papers that have been pledged for the ACP/GMD special issue of GeoMIP, and we have provided a list of references detailing past GeoMIP studies. All of these analyses will remain relevant for the new proposed GeoMIP6 experiments, and we expect additional proposed papers on these topics, as well as new topics, to emerge as time progresses.
- **The MIP has completed the MIP template questionnaire;** These have been completed, and we have no additional updates to them.
- **The MIP contributes a paper on its experimental design to the CMIP6 Special Issue;** The attached experiment description will be this paper.

- **The MIP considers reporting on the results by co-authoring a paper with the modeling groups;** Papers of this type are the key outputs of GeoMIP. Producing these papers is standard practice for participating in GeoMIP, and all participants have repeatedly shown they are quite eager to analyze and publish results. The “official” GeoMIP policy is that if a paper is written using GeoMIP output produced within the past 12 months (approximately), all modeling groups that produced that output should be invited to contribute to the paper as co-authors. Specific dates applying to all GeoMIP experiments are posted on the GeoMIP website.

Potential synergies with other MIPs

- Several of the new proposed experiments are based upon the Tier 1 experiments in ScenarioMIP¹; it is common practice for new GeoMIP experiments to be based on core experiments describing future projections of climate change, so we expect this synergy to continue long into the future.
- As is outlined in the attached document, there are many uncertainties in the accurate representation of sulfate aerosol microphysics associated with stratospheric sulfate aerosol geoengineering. A key step in narrowing these uncertainties is understanding the microphysical evolution of stratospheric sulfate aerosols from volcanic eruptions, especially large eruptions in which the coagulation processes strongly affect aerosol lifetime. This is at the core of VolMIP and is also a priority for DAMIP. The results obtained from these two MIPs will be important for informing proposed and future simulations in GeoMIP. In particular, the DAMIP experiment histVLC and all of the Type I and Type II experiments in VolMIP are highly relevant to GeoMIP.
- The questions in nonlinMIP are central to the experimental design of GeoMIP: How do abrupt changes in radiative forcing impact the climate? As was described in the nonlinMIP proposal, GeoMIP experiments can be used to inform the analyses central to nonlinMIP, especially considering that geoengineering simulations can alter the strengths of temperature-related feedbacks that are the source of many climate system nonlinearities.
- Changes in the hydrological cycle are some of the key motivations behind GeoMIP: what are the effects of CO₂ on the hydrological cycle, and how are these changes mitigated or enhanced by geoengineering? Addressing these concerns is at the center of PDRMIP. Simulation G1ext in GeoMIP consists of a combination of the single forcing simulations that are proposed to be included in PDRMIP, so we anticipate that results from GeoMIP can inform the analyses of PDRMIP, and vice versa.
- The goals of GeoMIP and RFMIP are quite concordant; GeoMIP at its core seeks to understand the relationship between radiative forcing and climate response. It has been shown that some GeoMIP simulations (particularly G1ext) can provide a novel way of constraining the climate response to better isolate the quantities that can aid in the mission of RFMIP to quantify radiative forcing. The GeoMIP6 Tier 2 experiments involving fixed sea surface temperatures can also be analyzed in

concert with the RFMIP leads, providing additional information about effective radiative forcing.

- It has long been known that land use change is a substantial climate driver; this is at the core of LUMIP. A new proposed experiment in the GeoMIP Testbed is aimed at idealized simulations of land use change as a method of geoengineering. The findings of LUMIP will inform the boundaries of land use modification as a method of geoengineering, providing information about the feasibility of this newly proposed Testbed experiment.
- GeoMIP has a natural connection to the Chemistry Climate Model Initiative (CCMI), which is a key partner in AerChemMIP. Many of the interests in aerosol-chemistry and climate interactions overlap between AerChemMIP and GeoMIP. Moreover, G4-SSA is a proposed experiment to CCMI.
- CFMIP has proposed two experiments involving a 4% increase or decrease in total solar irradiance. This simulation is quite complimentary to the DECK experiment abrupt4xCO₂ and to our experiment G1ext (and the past experiment G1). G1ext is a simulation of two combined forcings: CO₂ and solar irradiance changes. Applying each of those forcings individually is a crucial step in understanding the climate response in G1ext. Similarly, analyses of G1ext can provide information about rapid adjustments and feedback strengths, especially related to clouds, that arise in both abrupt4xCO₂ and the proposed CFMIP solar experiments. The coordinators of CFMIP have invited GeoMIP to cosponsor and coordinate their analysis of these experiments dealing with solar irradiance changes.

1 The Geoengineering Model Intercomparison Project Phase 6 (GeoMIP6):
2 Simulation Design and Preliminary Results

3
4
5 Ben Kravitz,^{1*} Alan Robock,² Simone Tilmes,³ Olivier Boucher⁴, Jason M. English⁵,
6 Peter J. Irvine⁶, Andy Jones⁷, Mark G. Lawrence⁶, Michael MacCracken⁸, Helene
7 Muri⁹, Ulrike Niemeier¹⁰, Steven J. Phipps¹¹, Jana Sillmann¹², Trude Storelvmo¹³,
8 Hailong Wang¹, Shingo Watanabe¹⁴

9
10
11 In preparation for *Geoscientific Model Development*

12
13
14 ¹Atmospheric Sciences and Global Change Division, Pacific Northwest National
15 Laboratory, Richland, WA, USA

16 ²Department of Environmental Sciences, Rutgers University, New Brunswick, NJ,
17 USA

18 ³National Center for Atmospheric Research, Boulder, CO, USA

19 ⁴Laboratoire de Météorologie Dynamique, IPSL, CNRS/UPMC, Paris, France

20 ⁵Laboratory for Atmospheric and Space Physics, University of Colorado, Boulder, CO,
21 USA

22 ⁶Institute for Advanced Sustainability Studies, Potsdam, Germany

23 ⁷Met Office Hadley Centre, Exeter, United Kingdom

24 ⁸The Climate Institute, Washington, D.C., USA

25 ⁹Department of Geosciences, University of Oslo, Oslo, Norway

26 ¹⁰Max Planck Institute for Meteorology, Hamburg, Germany

27 ¹¹University of New South Wales, Sydney, Australia

28 ¹²Center for International Climate and Environmental Research, Oslo, Norway

29 ¹³Department of Geology & Geophysics, Yale University, New Haven, CT, USA

30 ¹⁴Japan Agency for Marine-Earth Science and Technology, Yokohama, Japan

31
32
33
34 *To whom correspondence should be addressed : P.O. Box 999, MSIN K9-24,
35 Richland, WA 99352, USA (ben.kravitz@pnnl.gov)

40 **Abstract.** We present a suite of new climate model experiment designs for the
41 Geoengineering Model Intercomparison Project (GeoMIP). This set of experiments,
42 named GeoMIP6 (to be consistent with the Coupled Model Intercomparison Project
43 Phase 6), is designed to study several important topics, including key uncertainties
44 in extreme events, use of geoengineering as part of a portfolio of responses to
45 climate change, and the relatively new idea of cirrus cloud thinning to allow more
46 longwave radiation to escape to space. We discuss experiment designs, as well as
47 the rationale for those designs, showing preliminary results from individual models
48 when available. We introduce a new feature, called the GeoMIP Testbed, which
49 provides a platform for simulations that will be performed with single models and
50 then subsequently assessed to determine whether the simulation designs will be
51 adopted as core (Tier 1) GeoMIP experiments. The GeoMIP Testbed is meant to
52 encourage various stakeholders to propose targeted experiments that address their
53 key open problems, with the goal of making GeoMIP more relevant to a broader set
54 of communities.

1. Introduction

As anthropogenic climate change continues largely unabated, society is exploring research into options for addressing the effects of greenhouse gas emissions. Along with mitigation and adaptation, a further option that is under consideration is solar radiation management (SRM). A form of geoengineering, SRM involves deliberate modification of the climate system to offset the radiative effects of increasing anthropogenic greenhouse gases by either increasing the reflection of solar radiation back to space or increasing the outgoing flux of terrestrial radiation. Better understanding the potential role that SRM might have in addressing climate change requires research on the climate effects and impacts, as well as the underlying processes involved and their uncertainties.

The goal of the Geoengineering Model Intercomparison Project (GeoMIP) is to understand the robust climate model responses to geoengineering (Kravitz et al., 2011). GeoMIP has achieved a number of successes. So far, there have been seven core climate model experiments designed for analyzing the effects of solar irradiance reduction, stratospheric sulfate aerosols, and marine cloud (or sky) brightening (Kravitz et al., 2011, 2013a), as well as several additional experiments proposed by various groups. Table 1 lists all of the designed experiments to date. As of the writing of this paper, GeoMIP has resulted in 20 peer-reviewed publications, and results from GeoMIP were featured in the Fifth Assessment Report of the Intergovernmental Panel on Climate Change (Boucher et al., 2013).

These past efforts were designed to target specific areas but were not designed to answer all questions about the climate effects of geoengineering; there are still many unanswered questions in geoengineering research. The Coupled Model Intercomparison Project is beginning its sixth phase (CMIP6), and one of their focus areas is geoengineering (Meehl et al., 2014). Now is an opportune moment to address some of the key uncertainties in geoengineering by introducing designs for a new suite of climate modeling experiments. Some of the more pressing questions we would like to address are

1. How would geoengineering affect changes in less easily detectable fields, such as extreme events, modes of climate variability, regional climate impacts, or long timescale climate processes?
2. What are common responses in climate model simulations of cirrus cloud thinning as a relatively new proposed geoengineering method?
3. What are common responses in climate models if geoengineering were to be used in concert with mitigation and adaptation? That is, what if geoengineering is used to only partially offset climate change?
4. What are robust differences in the climate model response between stratospheric sulfate aerosol injection and solar irradiance reduction?

In this paper, we outline four Tier 1 experiments for the next phase of GeoMIP. To be consistent with the numbering convention of CMIP, we call this next phase GeoMIP6. The experiment design for GeoMIP6 is based on discussions held at the

101 Fourth GeoMIP Workshop (Paris, April 2014; Kravitz et al., 2014a), the SCRiM All
102 Hands Meeting (State College, May 2014), the Exploring the Potential and Side
103 Effects of Climate Engineering (EXPECT) workshop (Oslo, June 2014), and an
104 experiment proposed for inclusion in the Chemistry Climate Model Initiative (CCMI;
105 Tilmes et al., 2014). All of the proposed experiments are listed in Table 1 along with
106 all previous GeoMIP and GeoMIP-affiliated experiments.
107

108 **2. Tier 1 experiments in GeoMIP6**

109

110 In this section, we outline the four Tier 1 experiments that are proposed for
111 GeoMIP6. These same experiments have also been proposed for inclusion in CMIP6,
112 with GeoMIP serving as an officially endorsed model intercomparison project.
113

114 The general experiment protocol is somewhat different from that of the previous
115 experiments (Kravitz et al., 2011; Kravitz et al., 2013a; also see Table 1). There has
116 recently been interest in conducting geoengineering studies that examine
117 phenomena for which there is a low signal-to-noise ratio: for example, extreme
118 temperature and precipitation events (Curry et al., 2014). To aid in the ability to
119 obtain more robust estimates of potential changes in extremes, we are now
120 requesting that all simulations be conducted for longer than 50 years. Cessation or
121 termination (in which the background scenario continues, but geoengineering is no
122 longer conducted) is no longer part of the experimental protocol. Many of the broad
123 messages associated with the so-called termination effect were well captured by
124 Jones et al. (2013), so additional efforts to represent termination are not currently a
125 high priority.
126

127 We request that all modeling groups produce the following at daily frequency:
128 minimum and maximum near-surface air temperature (reference height; usually
129 1.5-2 m), total surface precipitation, surface convective precipitation, and near-
130 surface (usually 10 m) wind speed, and hourly surface ozone, if available. If
131 possible, precipitation and convective precipitation should be reported as a
132 cumulative value at 6-hourly frequency, and wind speed should be reported as an
133 instantaneous value at 6-hourly frequency. Each modeling group should produce a
134 minimum of three ensemble members for each experiment; ideally, groups would
135 complete five or more ensemble members.
136

137 As before, the Tier 1 experiments will be based on core experiments in CMIP. The
138 newest version of the core CMIP6 experiments is called the CMIP Diagnostic,
139 Evaluation and Characterization of Klima (DECK) experiment portfolio (Meehl et al.,
140 2014). This will include many different simulations, but the DECK simulations that
141 are relevant for GeoMIP6 are piControl and abrupt4xCO2, both of which were also
142 included in CMIP5. Additionally, simulations involving future projections of climate
143 change scenarios will be based on the Tier 1 simulations of ScenarioMIP, which are
144 in the process of being finalized (O'Neill et al., 2014). Tier 1 of ScenarioMIP will
145 consist of a high, medium, and low forcing scenario, referring to the maximum
146 amount of anthropogenic radiative forcing exhibited in that scenario.

147

148 **2.1. G1ext**

149 This experiment will be an extended version of Experiment G1 (*Kravitz et al., 2011*).
150 G1 proposes that, beginning from a preindustrial simulation (piControl), the net top
151 of atmosphere (TOA) radiative flux imbalance due to an abrupt quadrupling of the
152 CO₂ concentration (abrupt4xCO₂) would be balanced via a reduction in total solar
153 irradiance (Figure 1). Here, “balance” is defined as the global mean value top-of-
154 atmosphere net radiative flux being within $\pm 0.1 \text{ W m}^{-2}$ of the piControl experiment
155 over an average of years 1-10 of simulation. The original G1 was conducted for 50
156 simulation years, so this will be a simple extension of the previous experiment.
157 Modeling groups that have already moved on to a new model version, or for
158 whatever reason are not able to extend their previous model run, should run
159 experiment G1ext for the full 100 years with their new version.

160

161 G1 has proven quite successful in revealing the underlying climate behavior in
162 response to solar irradiance reduction. The models have been modified since
163 CMIP5, so evaluating climate response to G1 in the new model versions could serve
164 as a useful comparison with the older model versions. A longer simulation will also
165 improve the detection of changes in extreme events and modes of climate
166 variability. Moreover, some processes of interest, such as changes in ice sheet
167 dynamics or deep ocean circulation, take longer to resolve than 50 years. Although
168 100 years is probably an insufficient length of time to assess changes in these fields,
169 it may nevertheless allow enough time for an early detection of features that emerge
170 above the noise level of the climate system; this early detection will be aided by
171 having multiple ensemble members in the simulations.

172

173 G1 is the only original experiment from Kravitz et al. (2011) that is proposed to be
174 lengthened. The climate response in G2 is very similar to that of G1, but with a lower
175 signal-to-noise ratio, so extending G2 is unlikely to provide substantial additional
176 information. A new experiment (G6sulfur, below) has been proposed that will
177 accomplish similar goals to G3, but without some of the inherent ambiguities that
178 caused difficulties in interpreting results from G3 in certain cases. G4 may be
179 extended in the future, but such a simulation is not a high priority at this time.

180

181 **2.2. G6sulfur**

182 Previous GeoMIP experiments (G3 and G4) used RCP4.5 as a background scenario.
183 To maintain relevance to the newly designed experiments in CMIP6, our bases are
184 changed to follow the ScenarioMIP Tier 1 scenarios, described above.

185

186 Under experiment G6sulfur (Figure 2), stratospheric sulfate aerosols will be injected
187 into the model with the goal of reducing the value of net anthropogenic radiative
188 forcing from the ScenarioMIP Tier 1 high forcing scenario to that of the ScenarioMIP
189 Tier 1 medium forcing scenario (within $\pm 0.1 \text{ W m}^{-2}$). The motivation for this choice
190 is to evaluate a climate in which geoengineering is used as only a partial measure of
191 offsetting climate change, leaving the remainder for mitigation and adaptation. The
192 choice of the medium forcing scenario as the target instead of the low forcing

193 scenario (as in Section 4.1) is because the required amount of sulfate aerosol
194 injection to achieve a low anthropogenic forcing is quite large, and many of the risks
195 associated with such a large amount of geoengineering are currently unknown.
196

197 For this experiment, geoengineering will be simulated over years 2020-2100. All
198 atmospheric constituents in the ScenarioMIP Tier 1 scenarios are well defined
199 through the year 2100. Modeling groups that have an internal sulfate aerosol
200 treatment should calibrate the radiative response to sulfate aerosols individually so
201 that the results will be internally consistent. This procedure will be more difficult
202 for models that have a complex microphysical treatment of the aerosols, which may
203 require more sophisticated methods of meeting the goals of G6sulfur. One method
204 to calculate the necessary amount of sulfate aerosol is a double radiation call, once
205 with and once without the stratospheric aerosols. Another potential method
206 involves using feedback methods (Jarvis and Leedal, 2012; Kravitz et al., 2014b;
207 MacMartin et al., 2014). For models that have no dynamical treatment of sulfate
208 aerosols, we will provide a data set of aerosol optical depth, as well as ozone fields
209 that are consistent with this aerosol distribution; these fields will be consistent with
210 the fields generated for G4-SSA (see Section 3.2 for further details). The amount of
211 sulfate injection needed for a given model to achieve the goals of this experiment
212 may vary, so modeling groups should scale the aerosol optical depth as necessary.
213

214 Of notable importance is that the lifecycle of stratospheric sulfate aerosols is very
215 complex. To date, there are no comprehensive simulations of stratospheric sulfate
216 aerosol geoengineering that include aerosol microphysical processes, explicit size
217 representation, interactive chemistry, clouds, and radiation. Of the more
218 comprehensive simulations conducted, some studies include aerosol microphysics
219 and explicit size representation but do not allow oxidants to evolve (e.g.,
220 Heckendorn et al., 2009) or do not allow aerosol heating to interact with radiation
221 and dynamics (e.g., English et al., 2012). Other studies include aerosol microphysics
222 and heating, but represent the aerosol size distribution in assumed lognormal
223 modes of prescribed constant width (e.g. Niemeier et al., 2011, 2013). Because
224 geoengineering has not been conducted, there are no observations to constrain
225 these particular physical processes in models. Even with large volcanic eruptions,
226 Kokkola et al. (2009) show that capturing the evolution of the aerosol size
227 distribution is more difficult with increasing amounts of SO₂ injection. Development
228 of models that can represent these processes and thus constrain the uncertainties
229 that may arise is ongoing, and we expect that substantial progress will be made by
230 the time the GeoMIP6 experiments will begin. Nevertheless, the goal of GeoMIP is to
231 use the best available models and attempt to characterize uncertainties introduced
232 by structural uncertainties in those models.
233

234 All simulations will be conducted as if the aerosols or aerosol precursors are
235 emitted in a line from 10°S to 10°N along a single longitude band. This setup differs
236 somewhat from a single point source injection in that it allows models with a strong
237 stratospheric transport barrier to achieve a reasonable global distribution of sulfate
238 aerosol rather than an aerosol optical depth maximum in the tropics. The size of the

239 injection zone can substantially alter the resulting aerosol size distribution (English
240 et al., 2012), but we do not wish to add additional complications to the simulation
241 design at this time, so our design does not strongly deviate from the design of a
242 point-source injection. Injected aerosols or aerosol precursors should be evenly
243 spread across model layers between 16 and 25 km, a similar setup to that of the
244 original sulfate aerosol experiments (Kravitz et al., 2011). Models will use their own
245 individual treatments of aerosol optical properties, as this would be too difficult to
246 specify in a consistent way across all participating models.

247

248 **2.3. G6solar**

249 Experiment G3solar was proposed as an unofficial counterpart to experiment G3
250 (Kravitz et al., 2011); instead of achieving the goals of G3 using stratospheric sulfate
251 aerosol injections, the goals would be achieved using solar irradiance reduction.
252 Comparison of these two simulations would reveal differential effects of sulfate
253 aerosols and solar irradiance reduction. Preliminary results from a limited set of
254 models show some differences in the results of the two experiments, particularly
255 related to the hydrological cycle response (Niemeier et al., 2013).

256

257 Because of the difficulties in setting up experiment G3, few groups performed either
258 Experiment G3 or G3solar. We proposed above the G6sulfur experiment, which is
259 better specified than G3 and better aligned with the core simulations of CMIP, so it
260 should garner substantially greater participation. As a parallel experiment, to
261 compare the effects of solar reduction with those of stratospheric aerosols, we
262 propose G6solar, which uses the same setup as G6sulfur, but geoengineering is
263 performed using solar irradiance reduction (Figure 2).

264

265 **2.4. G7cirrus**

266 A recent proposal in the geoengineering literature is the idea of seeding cirrus
267 clouds, thinning them and thus allowing more longwave radiation to escape to space
268 (Mitchell et al., 2009; Storelvmo et al., 2013). Although cirrus cloud microphysics
269 are very complex and in many cases poorly understood, simulations have shown
270 that simply increasing the ice crystal fall speed of cirrus clouds captures the
271 dominant intended effects of cirrus seeding as simulated by models with more
272 complex microphysical treatments (Muri et al., 2014). As such, simulating the main
273 effects of cirrus cloud seeding should be possible in many different general
274 circulation models.

275

276 The design of G7cirrus (Figure 3) is comparable to previous GeoMIP experiments.
277 Against a background of the ScenarioMIP Tier 1 high forcing scenario, cirrus seeding
278 will begin in 2020 and continue through the year 2100. The goal of this experiment
279 is to seed cirrus by a constant amount that reduces average global mean
280 temperature in the decade 2020-2029 to that of the decade 1970-1979 (as
281 calculated in a historical run). The decade 1970-1979 was chosen to avoid the
282 climate effects of the 1982 El Chichón eruption, the 1991 Mount Pinatubo eruption,
283 and the unusually large El Niño events in 1982 and 1998. Unlike G6sulfur or
284 G6solar, G7cirrus does not propose to return net radiative forcing from one

285 ScenarioMIP Tier 1 scenario to another, as it is yet unclear what levels of forcing are
286 achievable through cirrus seeding.

287

288 Representing cirrus seeding in the same way in all participating climate models has
289 proven to be challenging. The goal of cirrus seeding in the real world would be to
290 cause cirrus clouds to consist of fewer but larger ice crystals, thus increasing the fall
291 speed. A simplistic approach is to multiply cirrus cloud optical depth in the
292 radiation code by a factor $\epsilon < 1$ without modifying the actual cirrus fields.

293 Implementing this approach can be difficult in some models, as they may only
294 distinguish between liquid and ice clouds; for such models, the factor ϵ is only
295 implemented for ice clouds with temperature below -35°C and pressure below 600
296 mb. Other models formulate the effects of cirrus clouds in the infrared as a
297 modification to atmospheric emissivity, not optical depth.

298

299 Figure 4 shows preliminary results from GISS ModelE2 (Schmidt et al., 2014) for
300 various values of ϵ . Global mean surface air temperature changes appear to be
301 linear with ϵ , but the required cooling is not nearly substantial enough to achieve
302 the goal of G7cirrus. We hypothesize that these results are due to cirrus clouds
303 being very efficient absorbers of longwave radiation, even if they are optically thin.
304 To achieve substantial cooling, it appears necessary to reduce cirrus cloud coverage,
305 not just optical depth. Single model simulations of cirrus thinning that incorporate a
306 treatment of cloud microphysics show more substantial surface cooling. Storelvmo
307 and Herger (2014) found global cooling of 0.25°C with regional cooling by as much
308 as 3°C ; Muri et al. (2014) found global mean cooling of $\sim 1^{\circ}\text{C}$; and Storelvmo et al.
309 (2014) found global mean cooling of 1.4°C in coupled simulations of high latitude
310 cirrus cloud thinning. As such, we conclude that the simplistic method of decreasing
311 cirrus cloud optical depth does not capture the relevant effects necessary to
312 represent cirrus cloud thinning.

313

314 A more complicated representation of cirrus cloud thinning would be to double the
315 ice crystal fall speed. Figure 5 shows that simulations using NorESM1-ME, in which
316 the fall speed was changed, can result in substantial cooling. This representation is
317 also not ideal, as fall speed is greater for large crystals. Actually introducing ice
318 nuclei (IN) would result in large ice crystals (although not so large as to fall out
319 quickly), but increasing the fall speed causes all large crystals to fall out quickly,
320 resulting in an unrealistically small size distribution of crystals. Doubling the size of
321 the ice crystals would be a better representation of cirrus cloud seeding, but it is not
322 well defined how one doubles a size distribution. Moreover, a change in size of the
323 ice crystals would change the scattering properties of the crystals; accounting for
324 this effect in a way that is consistent across all participating models would be quite
325 complicated.

326

327 A simple approach that roughly captures the desired effect is to add a new local
328 variable that replaces (in all locations where temperature is colder than 235 K) the
329 ice mass mixing ratio in the calculation of the sedimentation velocity with a value
330 that is eight times the original ice mass mixing ratio; we recommend that all

331 simulations of G7cirrus be conducted in this way. We acknowledge that this
332 approach has many shortcomings. Increasing the sedimentation velocity may not
333 capture part of the cooling effect due to the increase in crystal size. It would also
334 artificially increase fall speed without having larger ice crystals. However, this
335 method captures many of the broad effects of cirrus thinning and avoids the very
336 difficult task of including a doubling of the ice crystal size in both radiative transfer
337 and fall speed calculations; this more complicated approach is not straightforward
338 to incorporate in all models.

339

340 Figure 6 shows results from NorESM1-ME (Tjiputra et al., 2013) for an octupling of
341 the ice crystal fall speed against a background of RCP8.5. Relative humidities in the
342 upper troposphere are reduced by over 30% in the tropical upper troposphere,
343 which is consistent with the aims of cirrus cloud thinning. These results clearly
344 show that the simplistic approach of reducing cirrus cloud optical depth would not
345 be sufficient to capture the main effects of cirrus thinning. They also show that,
346 despite the shortcomings listed previously, increasing the sedimentation velocity of
347 the ice crystals captures many of the hypothesized effects of cirrus thinning,
348 particularly upper troposphere humidity changes.

349

350 Storelvmo and Herger (2014) found that the majority of the cirrus thinning effects
351 on net cloud forcing and surface temperatures are due to cirrus seeding outside of
352 the tropics; including the tropics in the regions that are seeded caused a modest
353 additional effect. However, so as not to introduce artificial boundaries in the
354 regions where cirrus clouds are altered, cirrus clouds will be modified at all
355 latitudes.

356

357 **3. Tier 2 Experiments in GeoMIP6**

358

359 In addition to the four Tier 1 experiments, we propose another set of experiments
360 that will aid in diagnosing climate model response. These Tier 2 experiments will
361 not be included as official contributions to CMIP6, but will instead be conducted
362 independently by GeoMIP.

363

364 **3.1. Timeslice Simulations**

365

366 Separately calculating the rapid adjustments and the feedback response (also call
367 the fast and slow responses, respectively) can reveal fundamental climate behavior.
368 This has been shown to be particularly useful for geoengineering simulations
369 (Tilmes et al., 2013; Kravitz et al., 2013b; Huneus et al., 2014). As such, we are
370 requesting that all participating modeling groups conduct timeslice simulations
371 (e.g., Cubasch et al., 1995) for each of the Tier 1 experiments to aid in diagnosing
372 radiative forcing for the scenarios proposed here.

373

374 These timeslice experiments involve fixed sea surface temperature (SST)
375 simulations for a period of 10 years; these are similar to Radiative Flux Perturbation
376 simulations (Haywood et al., 2009). In these simulations, SSTs, sea ice, and all

377 boundary conditions (greenhouse gas concentrations, aerosols, and other climate
378 forcing agents) are to be prescribed at a constant climatology for the entire 10-year
379 simulation. In most of the timeslice simulations, an external forcing is applied. For
380 this forcing, the climatology is derived from the appropriate geoengineering
381 experiment. For all the other boundary conditions, the climatologies are derived
382 from the appropriate reference scenarios, in which no geoengineering is applied.
383 Each Tier 1 experiment will have two associated timeslice simulations, one at the
384 beginning of the coupled simulation and one at the end. The timeslice simulations
385 are described in more detail in Table 2.
386

387 **3.2. G4-Specified Stratospheric Aerosol experiment (G4-SSA)**

388
389 There are several issues in simulations of geoengineering with prognostic
390 stratospheric sulfate aerosols, as differences in the resulting aerosol distribution can
391 have prominent effects on the climate impacts of geoengineering and thus can
392 produce large differences in the response between the models. To remove this
393 difference between the models, Tilmes et al. (2014a) have designed an experiment
394 for chemistry climate models (CCMs) called G4-SSA. This experiment is designed so
395 all models use the same prescribed stratospheric sulfur distribution, allowing for
396 assessments of the range of climate responses for different representations of
397 aerosol-chemistry and climate interactions. This experiment is connected to the
398 other experiments in the Chemistry Climate Model Initiative (CCMI).
399

400 The experiment design takes inspiration from GeoMIP experiment G4. Against a
401 background of RCP6.0, a layer of stratospheric aerosols will be injected into the
402 model at a rate of 8 Tg SO₂ per year. Instead of allowing the models to calculate
403 their aerosol distributions, a distribution of surface area density and other aerosol
404 parameters will be provided to all models. The described distribution can also be
405 scaled so as to apply to other scenarios, such as the ScenarioMIP scenarios (this is
406 relevant for Experiment G6sulfur). We will provide time series of aerosol optical
407 depth and ozone concentration that are consistent with the aerosol distribution at
408 the website [https://www2.acd.ucar.edu/gcm/geomip-g4-specified-stratospheric-](https://www2.acd.ucar.edu/gcm/geomip-g4-specified-stratospheric-aerosol-data-set)
409 [aerosol-data-set](https://www2.acd.ucar.edu/gcm/geomip-g4-specified-stratospheric-aerosol-data-set).
410

411 Although G4-SSA was developed for CCMs, it would be useful to obtain results from
412 general circulation models (GCMs) as well, hence the inclusion in GeoMIP6. These
413 two classes of models have very different treatments of the atmosphere, including
414 stratospheric chemistry, aerosol microphysics, and representation of the quasi-
415 biennial oscillation. Comparing results from these two groups would reveal some of
416 the mechanisms behind climate model response to stratospheric aerosol
417 geoengineering, as well as provide a guideline for which processes are important to
418 improve in models.
419

420 **4. The GeoMIP Testbed**

421

422 A new feature of GeoMIP is termed the *GeoMIP Testbed*. This is a set of experiments
423 that are potentially useful geoengineering studies that have been proposed by
424 individual groups. The idea is that each group understands the key problems in its
425 own sector and is thus uniquely posed to design a simulation that will best address
426 those problems. That simulation design will then be vetted by individual models
427 before a decision can be made as to whether they should be implemented in the full
428 model suite.

429

430 **4.1. G6sulfur_low**

431

432 Experiment G6sulfur is designed to reduce radiative forcing in a high emissions
433 scenario to that of a moderate emissions scenario via simulating stratospheric
434 sulfate aerosol injection. This experiment is useful in assessing the effectiveness of
435 geoengineering as part of a portfolio of responses to climate change. However, this
436 experiment does not address feasibility or limits of stratospheric sulfate aerosol
437 injection. As was stated in Section 2.2, increasing amounts of stratospheric SO₂
438 injection would cause particles to coagulate and fall out more rapidly. Therefore,
439 the relationship between the amount of injection and the resulting radiative forcing
440 is projected to be sublinear. This problem prompts a natural question: What is the
441 limit of achievable radiative forcing from stratospheric sulfate aerosol injection?
442

443

444 A natural first step in addressing this problem involves a similar setup to that of
445 G6sulfur. Against a background of the ScenarioMIP Tier 1 high forcing scenario,
446 sulfate aerosol precursors will be injected into the stratosphere in sufficient
447 amounts to reduce anthropogenic radiative forcing from the levels in the high
448 forcing scenario to levels in the low forcing scenario. Because the low forcing
449 scenario is a ScenarioMIP Tier 1 experiment, so if it is deemed that this simulation
450 would be worthwhile to conduct among the full suite of GeoMIP participants,
451 simulations can be done with relatively little preparation.

452

453 Figure 7 shows the required amount of stratospheric aerosol injection to achieve
454 given amounts of radiative forcing; these simulations were performed in ECHAM-
455 HAM (Stier et al., 2005), a general circulation model coupled to an aerosol
456 microphysical model that simulates the physical evolution and particle growth of
457 sulfate aerosols. The sublinear relationship between injection amount and radiative
458 forcing is clearly illustrated. The difference between RCP8.5 and RCP2.6 in the year
459 2100 is 5.9 W m⁻², or the approximate radiative forcing of a tripling of the
460 preindustrial CO₂ concentration; this difference is similar to the expected difference
461 in forcing between the ScenarioMIP Tier 1 high forcing scenario and the Tier 1 low
462 forcing scenario, when those scenarios are finalized. Extrapolating from the results
463 of Figure 7, achieving this radiative forcing would require an injection of 40-50 Tg S
464 (80-100 Tg SO₂) per year. This injection rate is equivalent to 4-5 1991 Mount
465 Pinatubo eruptions per year. Some efforts to evaluate the climate effects of such a
466 scenario are already underway (Niemeier et al., in preparation).

4.2. GeoFixed10, GeoFixed20, GeoFixed50

A different way of quantifying the effects of stratospheric aerosol geoengineering is to perform a series of experiments in which the hypothetical rate of injection of stratospheric sulfate aerosols is constrained. Such a simulation would be well suited to ascertain the range of model responses to a fixed amount of SO₂ injection, highlighting model diversity. Against a background of the ScenarioMIP Tier 1 high forcing scenario, the modeling groups will inject 10, 20, or 50 Tg of sulfur dioxide per year into the lower stratosphere, in a similar setup to Experiment G4 (Kravitz et al., 2011).

4.3. GeoLandAlbedo

Experiment G1ocean-albedo has simulated the effects of marine cloud brightening by imposing a uniform increase in the ocean albedo (Kravitz et al., 2013). However, GeoMIP has not yet explored terrestrial-based approaches towards solar radiation management. Such approaches could readily be implemented on the regional scale, as human activities already control the albedo of much of the land surface. We therefore propose an alternative experiment in which the land surface albedo is increased, against a background of the CMIP5 abrupt4xCO₂ experiment.

Under experiment GeoLandAlbedo, the land surface albedo would be increased by a uniform amount of 0.1 across all urban and agricultural areas. Such an increment represents a reasonable estimate of the maximum large-scale albedo increase that could be achieved in practice (Lobell et al., 2006; Lenton and Vaughan, 2009; Davin et al., 2014). The aim of experiment GeoLandAlbedo would not be to achieve global energy balance, but rather to determine the extent to which land surface albedo changes could offset the effects of increasing greenhouse gases on a regional basis.

To some degree, different aspects of this problem have been explored. Irvine et al. (2011) determined that different types of surface albedo geoengineering were incapable of offsetting the radiative forcing from a doubling of the CO₂ concentration, and the adverse side effects of such attempts could be large. Focusing only on bio-engineering crops to increase crop canopy albedo (Ridgwell et al., 2009) could cause local cooling effects (Doughty et al., 2011) but would likely have a small global impact (Singarayer et al., 2009; Singarayer and Davies-Barnard, 2012).

All of the previous studies on terrestrial-based albedo increases were conducted with single models, so the robustness of the effectiveness of this particular method of geoengineering, as well as the side effects, have not yet been tested. Assessing the range of responses to terrestrial-based geoengineering is especially important, given the wide range of structural and parametric uncertainties associated with modeling land surface processes.

512 **5. Conclusions**

513

514 The climate model experiment designs presented here mark the beginning of a
515 concerted effort to include broader perspectives within GeoMIP. The extension of
516 all experiments to at least 80 years is recommended to obtain more robust
517 estimates of changes in extremes and modes of variability; it will be particularly
518 interesting to compare what results can be obtained from G1ext that were not
519 obtainable through analyses of Experiment G1, particularly related to extreme
520 events (Curry et al., 2014) and modes of climate variability. The two G6
521 experiments were designed to open the door toward possible conversations with
522 designers of climate change scenarios. We would eventually like to explore
523 potential synergies with ScenarioMIP, on which our core simulations are based.
524 Experiment G7cirrus is the first model intercomparison of the new idea of cirrus
525 thinning and is designed to open avenues of investigation in both geoengineering
526 and cirrus cloud microphysical representations. G4-SSA was designed to explore
527 commonalities and differences between general circulation models and CCMs,
528 potentially highlighting important processes in representing aerosol on chemistry,
529 but also on dynamics and climate.

530

531 Geoengineering has the potential to impact climate systems at all scales, so by
532 incorporating requirements from communities studying these different systems, we
533 can broaden the usefulness of GeoMIP to a wider variety of scientists, policy makers,
534 and other stakeholders. The GeoMIP Testbed is a key part of this effort. Under this
535 new framework, individual communities can propose and test experiments that are
536 designed to address problems in their sectors, providing invaluable information as
537 to whether simulations by the full GeoMIP community are warranted.

538

539 Nevertheless, there remain some key gaps in GeoMIP; these can provide a roadmap
540 for future experiment design. One notable area is in impacts assessment. GeoMIP is
541 quite adept at calculating expected climate effects from particular geoengineering
542 scenarios, but translating those effects into impacts on people has only been
543 explored in a limited set of studies (e.g., Xia et al., 2014). Interaction with the
544 impacts assessment communities is one of the highest priorities for future
545 directions of GeoMIP. This is particularly applicable for effects on developing
546 countries, many of which will be most affected by climate change, and thus might
547 also be most affected by geoengineering.

548

549 Although we expect that this new suite of climate model experiments will be useful
550 in addressing many uncertainties in geoengineering research, there will remain
551 many key questions. These experiment designs are idealized and are not
552 representative of how geoengineering may be done in the real world, if society were
553 to decide to deploy it. These designs also do not include studies of feasibility; some
554 of the designed experiments may be more easily implemented than others.
555 Moreover, while physical science studies are necessary for gaining information
556 about the effects and impacts of geoengineering, they are not sufficient to stand as
557 the sole basis for decision making. A multitude of concerns are crucial for making

558 informed decisions about geoengineering, including natural science, social science,
559 humanities, and the humanitarian sector (e.g., Robock, 2014).

560

561 **Acknowledgments.** Ben Kravitz and Hailong Wang are supported by the Fund for
562 Innovative Climate and Energy Research (FICER). The Pacific Northwest National
563 Laboratory is operated for the U.S. Department of Energy by Battelle Memorial
564 Institute under contract DE-AC05-76RL01830. Simulations performed by Ben
565 Kravitz were supported by the NASA High-End Computing (HEC) Program through
566 the NASA Center for Climate Simulation (NCCS) at Goddard Space Flight Center.
567 Alan Robock is supported by NSF grants AGS-1157525 and GEO-1240507. The
568 National Center for Atmospheric Research is funded by the National Science
569 Foundation. Andy Jones was supported by the Joint UK DECC/Defra Met Office
570 Hadley Centre Climate Programme (GA01101). Helene Muri is supported by the
571 Norwegian Research Council project EXPECT (grant no. 229760/E10) and
572 computing time was provided by NOTUR. Jana Sillmann is supported by the
573 Norwegian Research Council project NAPEX (229778). Shingo Watanabe is
574 supported by the SOUSEI program, MEXT, Japan.

575 **References**

576

577 Boucher, O., et al. (2013), Clouds and Aerosols. In: *Climate Change 2013: The Physical*
578 *Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the*
579 *Intergovernmental Panel on Climate Change* [Stocker, T.F., D. Qin, G.-K. Plattner, M.
580 Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley (eds.)].
581 Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA,
582 pp. 571–658, doi:10.1017/CBO9781107415324.016.

583

584 Cubasch, U., J. Waszkewitz, G. Hegerl, and J. Perlwitz (1995), Regional climate
585 changes as simulated in time-slice experiments, *Climatic Change*, 31, 372-304,
586 doi :10.1007/BF01095150.

587

588 Curry, C. L., et al. (2014), A multi-model examination of climate extremes in an
589 idealized geoengineering experiment, *J. Geophys. Res.*, 119, 5226-5239,
590 doi:10.1002/2013JD020648.

591

592 Davin, E.L., S.I. Seneviratne, P. Ciais, A. Oliso, and T. Wang, 2014: Preferential
593 cooling of hot extremes from cropland albedo management. *Proc. Natl Acad. Sci.*
594 Published online, doi:10.1073/pnas.1317323111.

595

596 Doughty, C. E., C. B. Field, and A. M. S. McMillan (2011), Can crop albedo be increased
597 through the modification of leaf trichomes, and could this cool regional climate?
598 *Climatic Change*, 104, 379-387.

599

600 English, J. M., O. B. Toon, and M. J. Mills (2012), Microphysical simulations of sulfur
601 burdens from stratospheric sulfur geoengineering, *Atmos. Chem. Phys.*, 12, 4775-
602 4793, doi:10.5194/acp-12-4775-2012.

603

604 Haywood, J., L. J. Donner, A. Jones, and J.-C. Golaz (2009), Global indirect radiative
605 forcing caused by aerosols: IPCC (2007) and beyond, *Clouds in the Perturbed*
606 *Climate System* (J. Heintzenberg and R. J. Charlson, eds.), 451-467, MIT Press,
607 Cambridge.

608

609 Heckendorn, P., D. Weisenstein, S. Fueglistaler, B. P. Luo, E. Rozanov, M. Schraner, L.
610 W. Thomason, and T. Peter (2009), The impact of geoengineering aerosols on
611 stratospheric temperature and ozone, *Environ. Res. Let.*, 4, 045108,
612 doi:10.1088/1748-9326/4/4/045108.

613

614 Huneus, N. et al. (2014), Forcings and feedbacks in the GeoMIP ensemble for a
615 reduction in solar irradiance and increase in CO₂, *J. Geophys. Res.*, 119, 5226-5239,
616 doi:10.1002/2013JD021110.

617

618 Irvine, P. J., A. Ridgwell, and D. J. Lunt (2011), Climatic effects of surface albedo
619 geoengineering, *J. Geophys. Res.*, 116, D24112, doi:10.1029/2011JD016281.

620

621 Jarvis, A. and D. Leedal (2012), The Geoengineering Model Intercomparison Project
622 (GeoMIP): A control perspective, *Atmos. Sci. Lett.*, *13*, 157-163, doi:10.1002/asl.387.
623

624 Jones, A., et al. (2013), The impact of abrupt suspension of solar radiation
625 management (termination effect) in experiment G2 of the Geoengineering Model
626 Intercomparison Project (GeoMIP), *Journal of Geophysical Research*, *118*(17), 9743-
627 9752, doi:10.1002/jgrd.50762.
628

629 Kokkola, H., R. Hommel, J. Kazil, U. Niemeier, A. -I. Partanen, J. Feichter, and C.
630 Timmreck (2009), Aerosol microphysics modules in the framework of the ECHAM5
631 climate model – Intercomparison under stratospheric conditions, *Geosci. Model Dev.*,
632 *2*, 97-112, doi:10.5194/gmd-2-97-2009.
633

634 Kravitz, B., A. Robock, O. Boucher, H. Schmidt, K. E. Taylor, G. Stenchikov, and M.
635 Schulz (2011), The Geoengineering Model Intercomparison Project (GeoMIP),
636 *Atmos. Sci. Lett.*, *12*, 162-167, doi:10.1002/asl.316.
637

638 Kravitz, B., et al. (2013a), Sea spray geoengineering experiments in the
639 Geoengineering Model Intercomparison Project (GeoMIP): Experimental design and
640 preliminary results, *J. Geophys. Res.*, *118*(19), 11175-11186,
641 doi:10.1002/jgrd.50856.
642

643 Kravitz, B., et al. (2013b), An energetic perspective on hydrological cycle changes in
644 the Geoengineering Model Intercomparison Project, *J. Geophys. Res.*, *118*, 13087-
645 13102, doi:10.1002/2013JD020502.
646

647 Kravitz, B., A. Robock, and O. Boucher (2014a), Future directions in simulating solar
648 geoengineering, *Eos Trans. Amer. Geophys. Union*, *95*, 280,
649 doi:10.1002/2014EO310010.
650

651 Kravitz, B., D. G. MacMartin, D. T. Leedal, P. J. Rasch, and A. J. Jarvis (2014b), Explicit
652 feedback and the management of uncertainty in meeting climate objectives with
653 solar geoengineering, *Environ. Res. Lett.*, *9*, 044006, doi:10.1088/1748-
654 9326/9/4/044006.
655

656 Lenton, T. M., and N. E. Vaughan (2009), The radiative forcing potential of different
657 climate geoengineering options, *Atmos. Chem. Phys.*, *9*, 5539-5561.
658

659 Lobell, D.B., et al. (2006), Biogeophysical impacts of cropland management changes
660 on climate, *Geophys. Res. Lett.*, *33*, L06708, doi:10.1029/2005GL025492.
661

662 MacMartin, D. G., B. Kravitz, D. W. Keith, and A. Jarvis (2014), Dynamics of the
663 coupled human-climate system resulting from closed-loop control of solar
664 geoengineering, *Climate Dynam.*, *43*, 243-258, doi:10.1007/s00382-013-1822-9.
665

666 Meehl, G. A., R. Moss, K. E. Taylor, V. Eyring, R. J. Stouffer, S. Bony and B. Stevens
667 (2014), Climate Model Intercomparisons: Preparing for the Next Phase, *Eos Trans.*
668 *AGU*, 95(9), 77, doi:10.1002/2014E0090001.
669

670 Mitchell, D. L. and W. Finnegan (2009), Modification of cirrus clouds to reduce global
671 warming, *Environ. Res. Lett.*, 4, 045102, doi:10.1088/1748-9326/4/4/045102.
672

673 Muri, H., J. E. Kristjánsson, T. Storelvmo, and M. A. Pfeffer (2014), The climatic effects
674 of modifying cirrus clouds in a climate engineering framework, *J. Geophys. Res.*, 119,
675 4174-4191, doi:10.1002/2013JD021063.
676

677 Niemeier, U., H. Schmidt, and C. Timmreck (2011), The dependency of
678 geoengineered sulfate aerosol on the emission strategy, *Atmos. Sci. Lett.*, 12(2), 189-
679 194, doi:10.1002/asl.304.
680

681 Niemeier, U., H. Schmidt, K. Alterskjær, and J. E. Kristjánsson (2013), Solar
682 irradiance reduction via climate engineering—impact of different techniques on the
683 energy balance and the hydrological cycle, *J. Geophys. Res.*, 118, 11905-11917,
684 doi:10.1002/2013JD020445.
685

686 Niemeier, U., et al., Is there a limit for sulfate injections?, in preparation.
687

688 O'Neill, B. C., et al. (2014), A new scenario framework for climate change research:
689 The concept of shared socioeconomic pathways, *Climatic Change*, 122, 387-400,
690 doi:10.1007/s10584-013-0905-2.
691

692 Ridgwell, A., J. S. Singarayer, A. M. Hetherington, and P. J. Valdes (2009), Tackling
693 regional climate change by leaf albedo bio-geoengineering, *Current Biology*, 19, 146-
694 150, doi:10.1016/j.cub.2008.12.025.
695

696 Robock, A. (2014) Stratospheric aerosol geoengineering, *Issues Env. Sci. Tech.*
697 (special issue “Geoengineering of the Climate System”), 38, 162-185.
698

699 Schmidt, G. A., et al. (2014), Configuration and assessment of the GISS ModelE2
700 contributions to the CMIP5 archive, *J. Adv. Model. Earth Syst.*, 6, 141-184,
701 doi:10.1002/2013MS000265.
702

703 Singarayer, J. S., A. Ridgwell, and P. Irvine (2009), Assessing the benefits of crop
704 albedo bio-geoengineering, *Environ. Res. Lett.*, 4, 045110, doi:10.1088/1748-
705 9326/4/4/045110.
706

707 Singarayer, J. S. and T. Davies-Barnard (2012), Regional climate change mitigation
708 with crops: Context and assessment, *Phil. Trans. Roy. Soc. A*, 370, 4301-4316, doi:
709 10.1098/rsta.2012.0010.
710

711 Stier, P., et al. (2005), The aerosol-climate model ECHAM5-HAM, *Atmos. Chem. Phys.*,
712 5, 1125-1156, doi:10.5194/acp-5-1125-2005.
713
714 Storelvmo, T., J. E. Kristjansson, H. Muri, M. Pfeffer, D. Barahona and A. Nenes
715 (2013), Cirrus cloud seeding has potential to cool climate, *Geophys. Res. Lett.*, 40,
716 178–182, doi:10.1029/2012GL054201.
717
718 Storelvmo, T., and N. Herger (2014), Cirrus cloud susceptibility to the injection of ice
719 nuclei in the upper troposphere, *J. Geophys. Res. Atmos.*, 119, 2375–2389,
720 doi:10.1002/2013JD020816.
721
722 Storelvmo, T., W. R. Boos, and N. Herger (2014), Cirrus cloud seeding: A climate
723 engineering mechanism with reduced side effects? *Phil. Trans. Roy. Soc. A.*,
724 20140116, doi:10.1098/rsta.2014.0116.
725
726 Tilmes, S., et al. (2014), A new Geoengineering Model Intercomparison Project
727 (GeoMIP) experiment designed for climate and chemistry models, *Geosci. Model Dev.*
728 *Discuss.*, 7, 5447-5464, doi:10.5194/gmdd-7-5447-2014.
729
730 Tjiputra, J. F., Roelandt, C., Bentsen, M., Lawrence, D. M., Lorentzen, T., Schwinger, J.,
731 Seland, Ø., and Heinze, C.: Evaluation of the carbon cycle components in the
732 Norwegian Earth System Model (NorESM), *Geosci. Model Dev.*, 6, 301-325,
733 doi:10.5194/gmd-6-301-2013, 2013.
734
735 Xia, L., et al. (2014), Solar radiation management impacts on agriculture in China: A
736 case study in the Geoengineering Model Intercomparison Project (GeoMIP), *J.*
737 *Geophys. Res.*, 119, 8695-8711, doi:10.1002/2013JD020630.

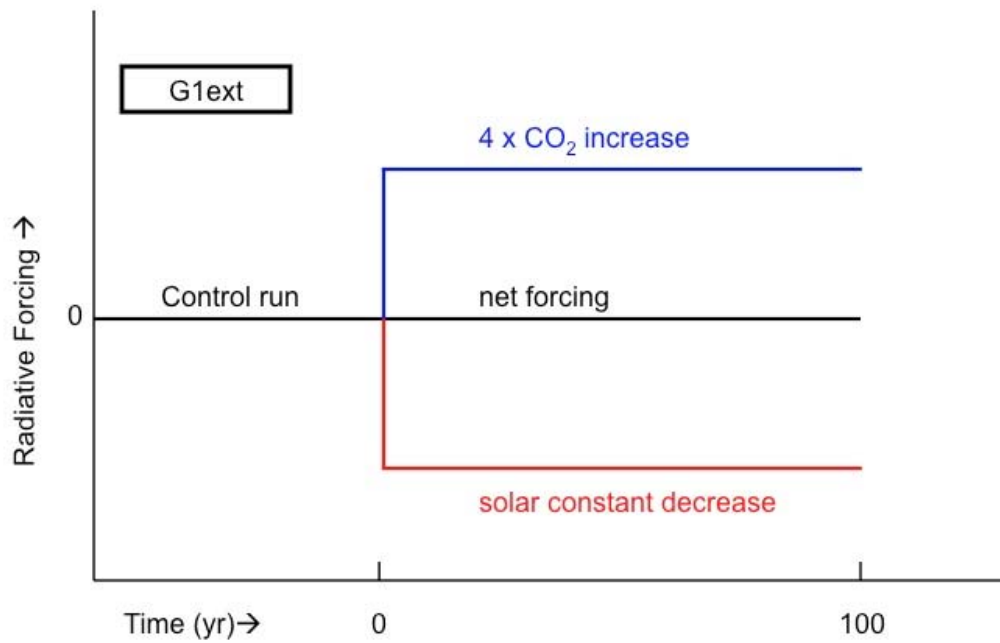
738 **Table 1.** All core GeoMIP experiments up to this point, including the additional
739 proposed Tier 1 GeoMIP6 experiments. Proposed Tier 2 experiments are listed in
740 Table 2. For each experiment, the name is given, along with a short description and
741 reference. Newly proposed experiments are printed in boldface. G5 is not a core
742 GeoMIP experiment but is included for completeness.
743

Experiment name	Description	Reference
G1	Balance 4xCO ₂ via solar irradiance reduction	Kravitz et al. (2011)
G1ext	Same as G1 but extended an extra 50 years	This document
G1ocean-albedo	Balance 4xCO ₂ via ocean albedo increase	Kravitz et al. (2013)
G2	Balance 1% CO ₂ increase per year via solar irradiance reduction	Kravitz et al. (2011)
G3	Keep TOA radiative flux at 2020 levels against RCP4.5 via stratospheric sulfate aerosols	Kravitz et al. (2011)
G4	Injection of 5 Tg SO ₂ into lower stratosphere per year	Kravitz et al. (2011)
G4cdnc	Increase CDNC in marine low clouds by 50% against a background of RCP4.5	Kravitz et al. (2013)
G4sea-salt	Inject sea salt aerosols into tropical marine boundary layer to achieve ERF of -2.0 W m ⁻² against a background of RCP4.5	Kravitz et al. (2013)
G5	Identical setup as G3 but using sea salt injection into marine low clouds (IMPLICC experiment; named SALT in Niemeier et al., 2013)	Alterskjær et al. (2013); Niemeier et al. (2013)
G6sulfur	Reduce forcing from ScenarioMIP Tier 1 high forcing scenario to the medium forcing scenario with stratospheric sulfate aerosols	This document
G6solar	Reduce forcing from ScenarioMIP Tier 1 high forcing scenario to the medium forcing scenario with solar irradiance reduction	This document
G7cirrus	Reduce forcing by constant amount via increasing cirrus ice crystal fall speed	This document

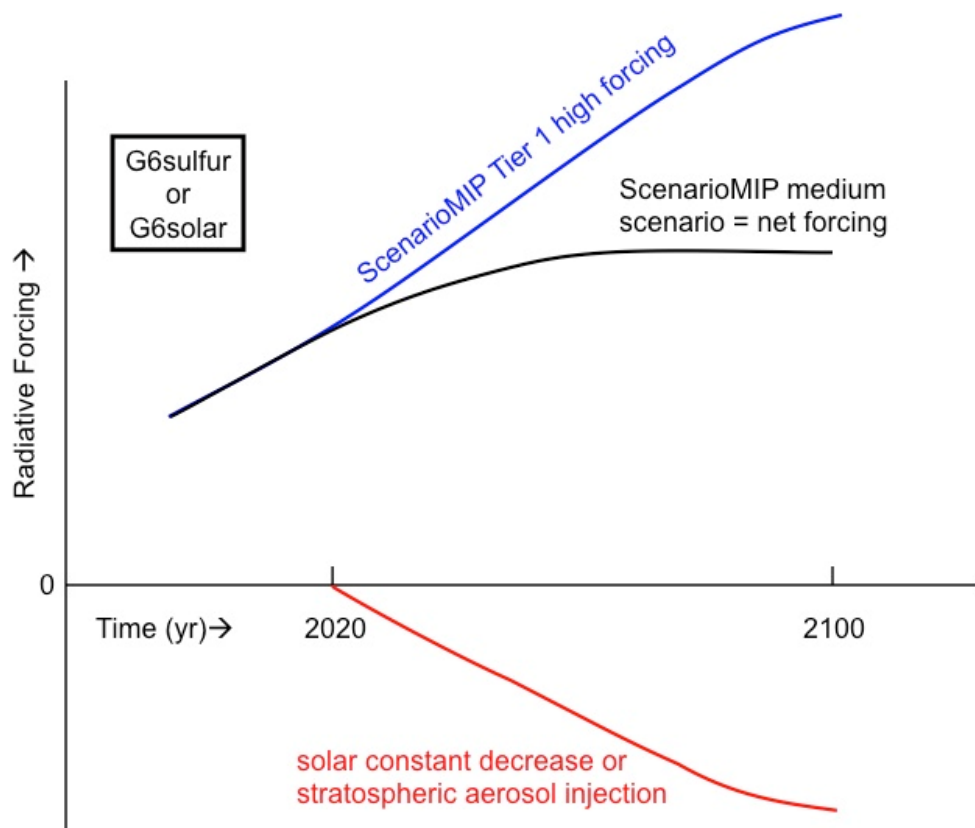
744 **Table 2.** Timeslice simulations associated with each of the four Tier 1 experiments.
 745 Further description of the timeslice simulations is given in Section 3.1. Each tier 1
 746 has two associated timeslice simulations: one for the beginning of the coupled
 747 simulation and one at the end of the coupled simulation. Note that the first timeslice
 748 simulations for G6sulfur and G6solar is identical, as no geoengineering has been
 749 applied yet. As such, this simulation is simply called G6Slice1.
 750

Experiment Name	Applied forcing	Boundary conditions
G1extSlice1	4xCO2	piControl
G1extSlice2	4xCO2	abrupt4xCO2 after 100 years
G6Slice1	None	ScenarioMIP Tier 1 high forcing scenario in year 2020
G6sulfurSlice2	G6sulfur in year 2100	ScenarioMIP Tier 1 high forcing scenario in year 2100
G6solarSlice2	G6solar in year 2100	ScenarioMIP Tier 1 high forcing scenario in year 2100
G7cirrusSlice1	G7cirrus in year 2020	ScenarioMIP Tier 1 high forcing scenario in year 2020
G7cirrusSlice2	G7cirrus in year 2100	ScenarioMIP Tier 1 high forcing scenario in year 2100

751

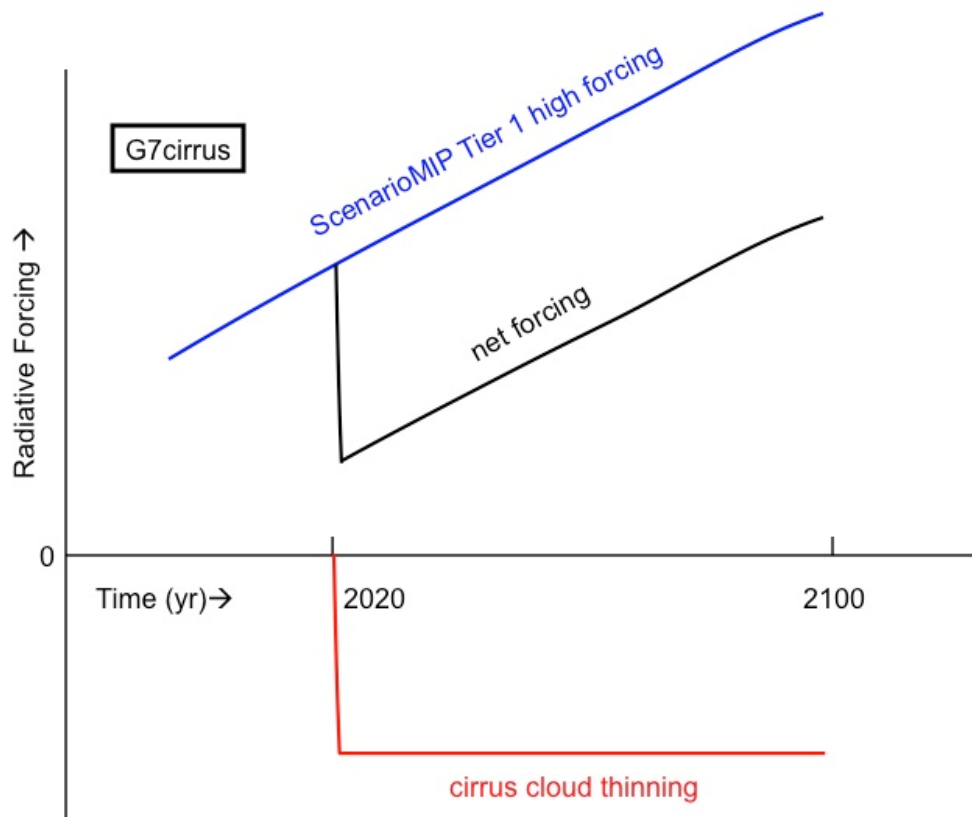


752
 753 **Figure 1.** Schematic of experiment G1ext. The experiment is started from a
 754 preindustrial control run. The instantaneous quadrupling of the CO₂ concentration
 755 from its preindustrial value is balanced by a reduction in solar irradiance for 100
 756 years.
 757



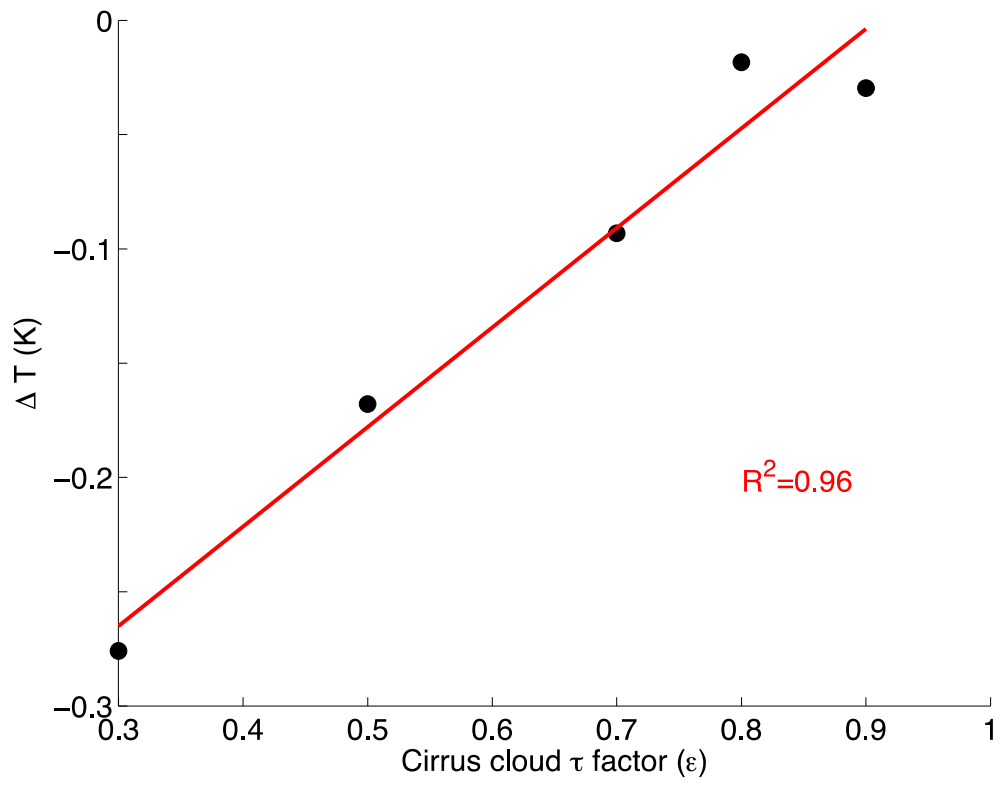
758
 759
 760
 761
 762
 763
 764
 765

Figure 2. Schematic of experiments G6sulfur and G6solar. Against a background of the ScenarioMIP Tier 1 high forcing scenario, geoengineering will be conducted at time-varying amounts to return net anthropogenic radiative forcing to the levels of the ScenarioMIP Tier 1 medium forcing scenario. Geoengineering will be accomplished by stratospheric aerosol injection (G6sulfur) or solar irradiance reduction (G6solar).



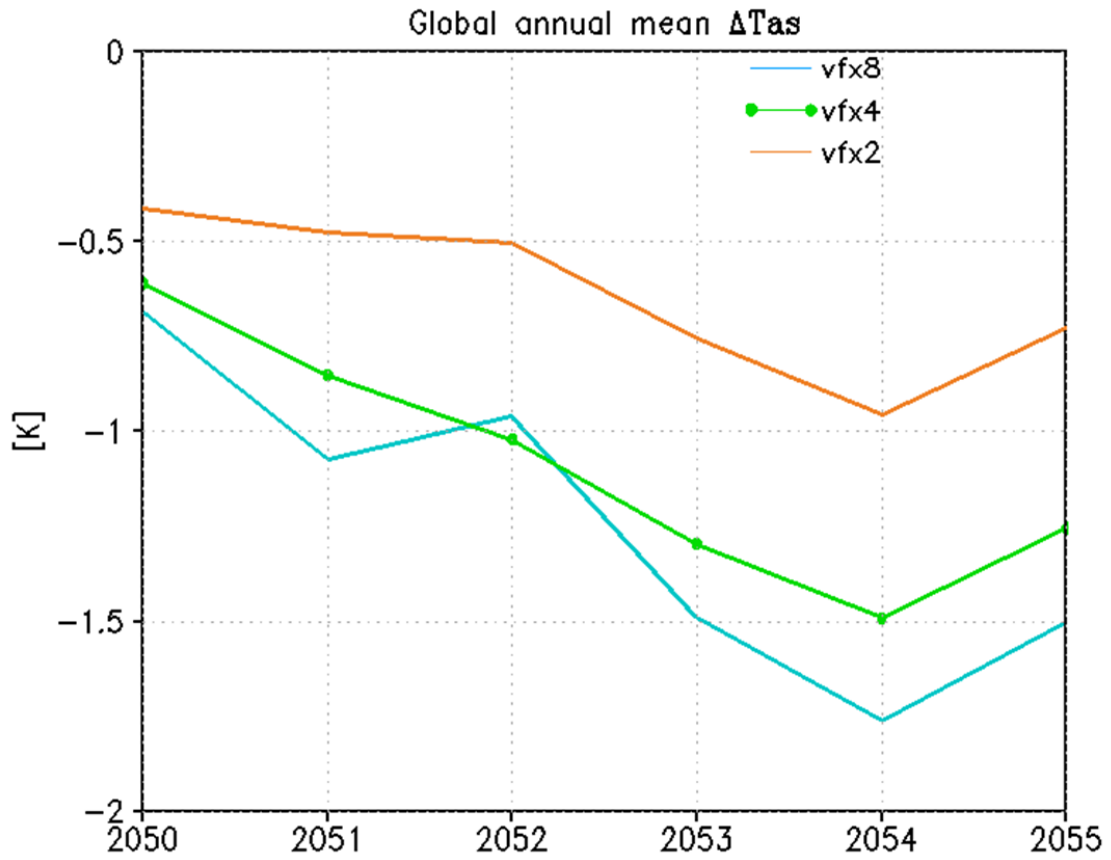
766
 767
 768
 769
 770
 771

Figure 3. Schematic of experiment G7cirrus. Against a background scenario of the ScenarioMIP Tier 1 high forcing scenario, a representation of cirrus cloud seeding will reduce net forcing by a constant amount. This simulation will begin in 2020 and will be conducted for 100 years.



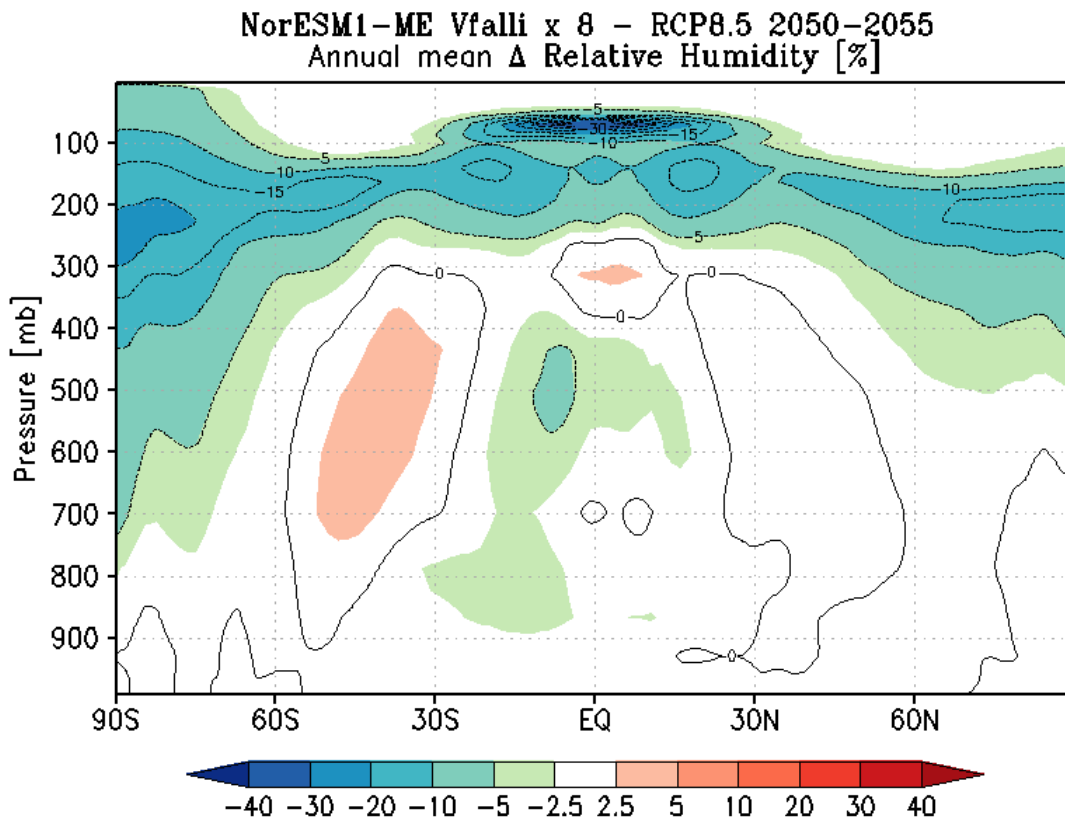
772
 773
 774
 775
 776
 777
 778
 779

Figure 4. Test simulations of reducing cirrus cloud optical depth (τ) as described in Section 2.4. τ was scaled by a factor $\epsilon < 1$ (x-axis). The amount of surface air temperature change due to this scaling (y-axis) was measured over a 4 year average; 0 indicates the global mean surface air temperature over years 2020-2023 in an RCP8.5 simulation. All simulations were performed using GISS ModelE2 (Schmidt et al., 2014).



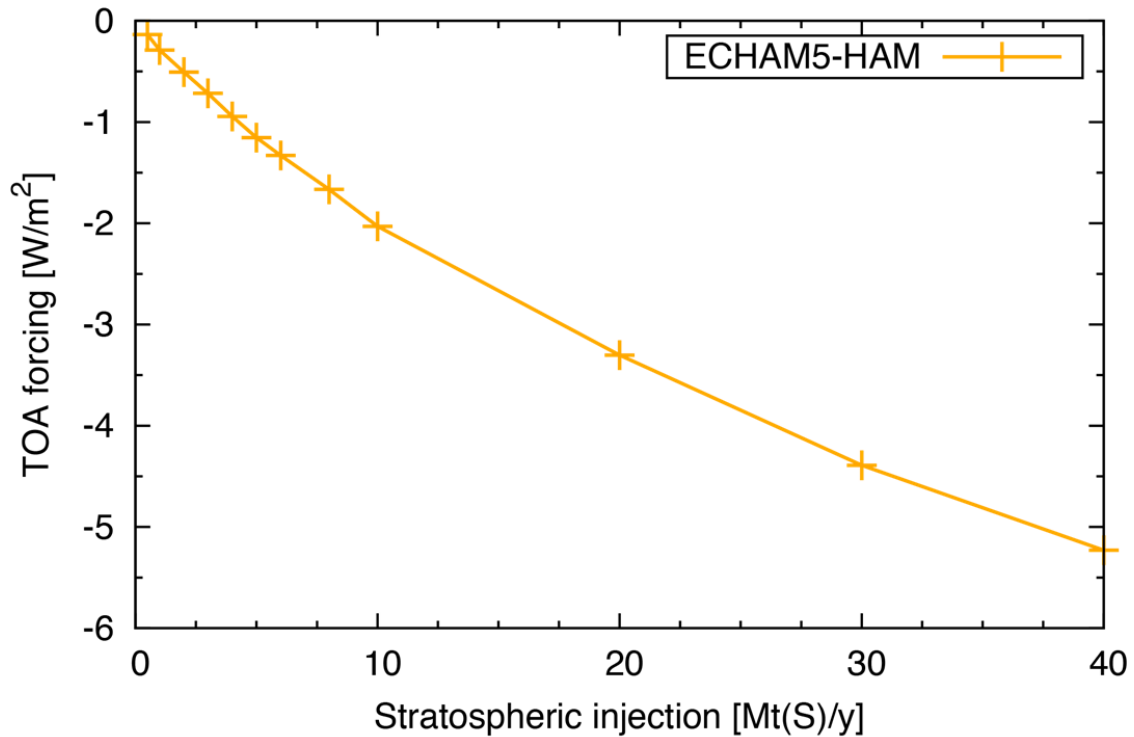
780
781
782
783
784
785
786

Figure 5. A sensitivity study of the effects of changing cirrus ice crystal sedimentation velocity in NorESM1-ME. vfx2, vfx4, and vfx8 indicate an increase in the sedimentation velocity by 2, 4, and 8 times, respectively. y-axis shows the global mean temperature change as a function of year (x-axis); differences are calculated with respect to an average over years 2050-2055 under an RCP8.5 scenario.



787
 788
 789
 790
 791

Figure 6. Zonally averaged annual mean of the difference in relative humidity (%) from NorESM1-ME for an octupling of the cirrus ice crystal fall speed. Differences are calculated as an average over years 2050-2055 against a background of RCP8.5.



792
793

794 **Figure 7.** This figure shows the amount of annual stratospheric injection (x-axis)
795 required to offset a given level of TOA net radiative flux imbalance (y-axis) in
796 ECHAM5-HAM, an atmospheric general circulation model with a treatment of the
797 microphysical evolution of sulfate aerosols. Maintaining 2020 values of net TOA
798 radiative flux imbalance against a background of RCP8.5 requires an injection of
799 approximately 70 Tg(S)/year in 2100. All values were calculated for injection of SO₂
800 into one grid box over the equator; other injection strategies would likely require a
801 different injection rate to achieve the same radiative forcing.

Global Monsoons Modeling Inter-comparison Project (GMMIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

Name of MIP:

Global Monsoons Modeling Inter-comparison Project (GMMIP)

Chairs:

Tianjun Zhou, Institute of Atmospheric Physics, Chinese Academy of Sciences, China
(zhoutj@lasg.iap.ac.cn)

Andy Turner, University of Reading, UK (a.g.turner@reading.ac.uk)

James Kinter, COLA, George Mason University (ikinter@gmu.edu)

Suggested Members of the Scientific Steering Committee (*new members from African and American monsoon communities will be included*)

Bin Wang, University of Hawaii, USA (wangbin@hawaii.edu)

Yun Qian, Atmospheric Sciences & Global Change Division, Pacific Northwest National Laboratory, USA (yun.qian@pnnl.gov)

Bin Wang, Institute of Atmospheric Physics, Chinese Academy of Sciences, China
(wab@lasg.iap.ac.cn)

Website:

<http://www.lasg.ac.cn/gmmip>

Proposed by:

CLIVAR AAMP, CLIVAR-GEWEX MP, CLIVAR/C20C+, in collaboration with LASG/IAP, China and PNNL, USA

Goal of GMMIP:

Changes in the precipitation and atmospheric circulation in the global monsoons are of

great scientific and societal importance owing to their impacts on more than two-thirds of the world's population. Monsoons occur in various regions around the world. Prediction of the monsoon rainfall change in the coming decades is of deep societal concern and vital for infrastructural planning, water resource management, and sustainable economic development.

The dominant monsoon systems in the world include the Asian-Australian, African, and the American monsoons. Each monsoon system generally has its own unique and specific characteristics in terms of variability. At the same time, the connections in the global divergent circulation necessitated by mass conservation link the various regional monsoons as they evolve through the season. On interannual-to-multidecadal time scales, there is evidence that monsoon precipitation in the Northern Hemisphere (NH) and Southern Hemisphere (SH) varies coherently, driven by ENSO and other global modes of climate variability at the lower boundary of the atmosphere.

The combination of changes in monsoon area and rainfall intensity has led to an overall weakening trend of global land monsoon rainfall accumulation since the 1950s. This decreasing tendency is dominated by the African and South Asian monsoons, due to the significant decreasing tendencies of both rainfall intensity and monsoon coverage. Beginning in the 1980s, however, the NH global monsoon precipitation has shown an upward trend. Understanding the mechanisms of precipitation changes in the global monsoons and identifying the roles of natural and anthropogenic forcing agents have been foci of the monsoon research community.

While all monsoons are large-scale cross-equatorial overturning circulations, major differences between characteristics of the different regional monsoons arise because of the different orography. This is most apparent for the Asia region, due to the TIP/Himalaya.

Climate models are useful tools in climate variability and climate change studies. However, the performance of the current state-of-the-art climate models is very poor and needs to be greatly improved over the monsoon domains. The Global Monsoons Model Inter-comparison Project (hereafter **GMMIP**) aims to improve our understanding of physical processes in global monsoon systems and to better simulate the mean state, interannual variability and long-term change of global monsoons by performing multi-model inter-comparisons. The contributions of internal variability (IPO-Interdecadal Pacific

Oscillation, AMO-Atlantic Multidecadal Oscillation) and external anthropogenic forcing to the historical evolution of global monsoons in the 20th and 21st century will be addressed.

Primary Science Questions:

- 1) What are the relative contributions of internal processes and external forcing that are driving the 20th century historical evolution of global monsoons?
- 2) To what extent and how does the atmosphere-ocean interaction contribute to the interannual variability and predictability?
- 3) What are the effects of Eurasian orography, in particular the Himalaya/Tibetan Plateau, on the regional/global monsoons?
- 4) How well can developing high-resolution models and improving model dynamics and physics help to reliably simulate monsoon precipitation and its variability and change?

By focusing on addressing these four questions we expect to deepen our understanding of models' capability in reproducing the monsoon mean state and its natural variability as well as the forced response to natural and anthropogenic forcing, which ultimately will help to reduce model uncertainty and improve the credibility of models in projecting future changes in the monsoon. The coordinated experiments will also help advance our physical understanding and prediction of monsoon changes.

Due to the uncertainties in the physical parameterizations in current models, the best way to address these questions is through a multi-model framework. CMIP6 provides a good opportunity for advancement of monsoon modeling and understanding. GMMIP will contribute to four of the five grand challenges of the WCRP, viz. Regional Climate Information, Water Availability, Climate Extremes, and Clouds, Circulation and Climate Sensitivity.

Proposed Experiments:

The main experiments of GMMIP will be divided into Tier 1 and Tier 2, with further optional ideas in Tier 3. The total experiments of GMMIP are summarized in Table 1. The **Tier-1** experiments will be extended AMIP runs. This is the *entry card for GMMIP*.

Table 1: Experiment list of GMMIP

	EXP name	Integration time	Short description and purpose of the EXP design	Model type
Tier-1	AMIP20C	1870-2013	Extended AMIP run that covers 1850-2014. All natural and anthropogenic historical forcings as used in <i>CMIP6 Historical Simulation</i> will be included. AGCM resolution as <i>CMIP6 Historical Simulation</i> . The HadISST data will be used. Minimum number of integrations is 1.	AGCM
Tier-2	HIST-IPO	1870-2013	Pacemaker 20 th century historical run that includes all forcing as used in <i>CMIP6 Historical Simulation</i> , and the observational historical SST is restored in the tropical lobe of the IPO domain (20°S-20°N, 175°E-75°W); to understand the forcing of IPO-related tropical SST to global monsoon changes. Models resolutions as <i>CMIP6 Historical Simulation</i> . The HadISST data will be used. Minimum number of integrations is 1.	CGCM with SST restored to the model climatology plus observational historical anomaly in the tropical lobe of IPO domain
Tier-2	HIST-AMO	1870-2013	Pacemaker 20 th century historical run that includes all forcing as used in <i>CMIP6 Historical Simulation</i> , and the observational historical SST is restored in the AMO domain (0°-70°N, 70°W-0°); to understand the forcing of AMO-related SST to global monsoon changes. Models resolutions as <i>CMIP6 Historical Simulation</i> . The HadISST data will be used. Minimum number of integrations is 1.	CGCM with SST restored to the model climatology plus observational historical anomaly in the AMO domain
Tier-3	DTIP	1979-2013	The topography of the TIP is modified by setting surface elevations to 500m; to understand the combined thermal and mechanical forcing of the TIP. Same model as DECK. Minimum number of integrations is 1.	AGCM
Tier-3	DTIP-DSH	1979-2013	Surface sensible heat released at the elevation above 500m over the TIP is not allowed to heat the atmosphere; to compare of impact of removing thermal effects. Same model as DECK. Minimum	AGCM

			number of integrations is 1.	
Tier-3	DHLD	1979-2013	The topography of the highlands in Africa, N. America and S. America TP is modified by setting surface elevations to a certain height (500m). Same model as DECK. Minimum number of integrations is 1.	AGCM

The **Tier-2 HIST-IPO run** is Pacemaker 20th century historical climate simulation that includes all forcing, and the sea surface temperature (SST) restored to the model climatology plus observational historical anomaly in the tropical lobe of the Interdecadal Pacific Oscillation (IPO; Power et al. 1999; Folland et al. 2002) domain (20°S-20°N, 175°E-75°W): the weight=1 in the inner box (15°S-15°N, 180°-80°W), linearly reduced to zero in the buffer zone (zonal and meridional ranges are both 5°) from the inner to outer box.

The **Tier-2 HIST-AMO run** is Pacemaker 20th century historical climate simulation that includes all forcing, and the SST restored to the model climatology plus observational historical anomaly in the Atlantic Multidecadal Oscillation (AMO; Enfield et al. 2001; Trenberth and Shea 2006) domain (0°-70°N, 70°W-0°): the weight=1 in the inner box (5°N-65°N, 65°W-5°W), linearly reduced to zero in the buffer zone (zonal and meridional ranges are both 5°) from the inner to outer box.

In **Tier-3 DTIP run**, following Boos and Kuang (2011, 2013) and Wu et al. (2007, 2012), the topography of the Tibetan Plateau(hereafter TIP) (20-60°N, 25-120°E) in the model is modified by leveling off the TIP to a certain height (e.g. 500m), with the surface properties unchanged. Other settings of the integration are same as the standard DECK AMIP run. This experiment represents perturbations to both thermal and mechanical forcing of the TIP with respect to the standard DECK AMIP run.

In **Tier-3 DTIP-DSH run**, the surface sensible heat flux at elevations above 500m over the TIP is not allowed to heat the atmosphere, i.e., the vertical diffusive heating term in the atmospheric thermodynamic equation is set to zero (Wu et al. 2012). Other settings of the integration are same as the standard DECK AMIP run. The differences between the standard DECK AMIP run and the DTIP-DSH are considered to represent the removal of TIP thermal forcing only and thus the circulation pattern of DTIP-DSH reflects the impacts of mechanical forcing.

Description of the analysis of GMMIP experiments:

There are four tasks in the analysis of GMMIP:

- 1) **Task-1:** Understanding 20th century changes of global monsoons
- 2) **Task-2:** The role of Eurasian orography on the regional/global monsoons (Himalaya/Tibetan Plateau experiment)
- 3) **Task-3:** Interannual variability of global monsoon precipitations
- 4) **Task-4:** High resolution modeling of global monsoons

The analysis of four tasks will use the outputs of GMMIP experiments, DAMIP (Detection and Attribution MIP) experiments, HighResMIP experiments, the CMIP6 Historical Simulation, and the AMIP experiments of DECK.

Connection with DECK and CMIP6 Historical Simulation

The DECK simulations will serve as an entry card for the CMIP6-Endorsed MIPS. The

DECK experiments are:

- AMIP simulations
- Pre-industrial control simulations
- 1%/yr increase in CO₂ concentration
- Switch-on 4XCO₂

The **CMIP6 Historical Simulation** experiment is:

- Historical simulation of fully coupled models (1850-2014)

The AMIP DECK simulation with the standard CMIP6 resolution will be used in the analysis of GMMIP. The Tier-1 AGCM experiment of GMMIP will specify the specific forcings which are consistent with the historical simulation from 1850-2014, viz. the CMIP6 Historical Simulation.

Connection with other MIPS

DAMIP (Detection and Attribution MIP):

The histALL (enlarging ensemble size of historical ALL forcing runs in DECK), histNAT (Historical natural-only run), histGHG (Historical well-mixed GHG-only run), histAER experiments (Historical anthropogenic-Aerosols-only run) of DAMIP will be used in the analysis of Task-1 of GMMIP.

Combinations of histALL, histNAT and histGHG will allow us to understand the observed 20th century global monsoon precipitation and circulation changes in the context of contributions from GHG, the other anthropogenic factors and natural forcing. The contributions of these external forcings will be compared to those from internal variability modes such as IPO and AMO.

HighResMIP:

The Tier-1 experiments of HighResMIP, which are AMIP runs but with minimum 25-50 km at mid-latitudes for high resolution + a standard resolution configuration (1950-2014), will be used in the analysis of Task-4 of GMMIP, which aims to examine the performance of high-resolution models in reproducing both the mean state and year-to-year variability of global monsoons.

The Tier-2 experiments of HighResMIP, which are coupled runs consisting of pairs of both historic runs and control runs using fixed 1950s forcing, will be used in the analysis of Task-3 of GMMIP, which aims to understand the role of air-sea interaction process in the improvement of monsoon mean state and year-by-year variability.

WCRP Grand Challenges:

GMMIP will address the grand challenges of the WCRP in the following way:

Regional Climate Information (rank 1)

GMMIP will improve our understanding of the 20th climate changes in global monsoon domains. The contributions of external anthropogenic forcings (GHG, aerosol), natural forcing, and internal variability modes (IPO, AMO) will be identified. These would provide useful information to climate prediction/projections in the highly populated global monsoon domains.

Water Availability (rank 2)

The water resources in global monsoon domains are greatly affected by the anomalous activities of monsoons. Understanding the mechanisms of monsoon variability as posed by GMMIP will lead to improvement of monsoon prediction/projection and provide useful information to policymakers in water availability-related decision making.

Climate Extremes (rank 2)

Extreme events such as mega-droughts and flooding have been frequently occurred in global monsoon domains. GMMIP is hopefully to identify the useful ways of improving the simulation/prediction of climate extremes in global monsoon domains.

Clouds, Circulation and Climate Sensitivity (rank 2).

A reasonable simulation of monsoon circulation and clouds is a prerequisite for a successful simulation of monsoon precipitation. By comparing the performances of climate models with high and normal resolutions, model simulations with/without air-sea interaction processes, the implementation of GMMIP will link the monsoon circulations to monsoon precipitation in the context of reducing model bias and improving model performances.

GEWEX and CLIVAR

Monsoon has been a research focus of GEWEX and CLIVAR. The scientific questions listed in GMMIP were originally identified by the CLIVAR Asian-Australian monsoon panel, the GEWEX/CLIVAR Monsoons Panel, and CLIVAR/C20C+ project. The questions have also been highlighted by the reports of CLIVAR Research Opportunities Tiger Team on “*Decadal Variability in the Climate System and its Predictability*”, and CLIVAR Research Opportunities Tiger Team on “*Intra-seasonal, Seasonal and Interannual Variability and Predictability of Monsoon Systems*”.

Participation:

Participation in GMMIP is voluntary and open. GMMIP will be coordinated by a small working group composed of engaged representatives from climate diagnosis, climate change attribution and climate modeling communities. This working group will engage the broadest degree of input and involvement from members of the scientific community.

The Scientific Steering Committee (SSC) of GMMIP will be composed of

representatives from CLIVAR & GEWEX monsoon panels, relevant projects and the global monsoon community. The SSC will provide comments and instructions for the analysis of GMMIP with focus on the scientific questions listed in the proposal.

The following modeling centers have expressed their interests in participating in GMMIP:

- BCC, China
- BNU, China
- CanESM, Canada (as many as possible)
- CESM, USA (will be run by PNNL scientists)
- CFS- IITM-ESM, India
- IAP, China
- IPSL, France
- FIO, China
- GISS, USA
- MPI-ESM, Germany (preference for prioritized Tier 1 experiments)
- MRI, Japan
- Nor-ESM, Norway
- UKESM, UK
- KMA model, South Korea

Proposed timing

Start of the experiments: Beginning of 2016

End of the experiments: No fixed date.

References

Boos, W. R. & Z. M. Kuang 2010: Dominant control of the South Asian monsoon by orographic insulation versus plateau heating. *Nature* 463, 218-223.

Boos, W.R. & Z. M. Kuang 2013: Sensitivity of the South Asian monsoon to elevated and non-elevated heating. *Sci. Rep.* 3, 1192; DOI:10.1038/srep01192.

- Cook, B.I., and R. Seager, 2013: The response of the North American Monsoon to increased greenhouse gas forcing. *J. Geophys. Res.*, 118, 1690-1699
- Cook, K. H., G. A. Meehl, and J. M. Arblaster, 2012: Monsoon regimes and processes in CCSM4. Part II: African and American monsoon systems. *J. Climate*, 25, 2609-2621
- Enfield, D., A. Mestas-Nuñez, and P. Trimble (2001), The Atlantic Multidecadal Oscillation and its relation to rainfall and river flows in the continental U.S., *Geophys. Res. Lett.*, 28, 2077–2080, doi:10.1029/2000GL012745.
- Folland, C. K., J. A. Renwick, M. J. Salinger, and A. B. Mullan (2002), Relative influences of the Interdecadal Pacific Oscillation and ENSO on the South Pacific Convergence Zone, *Geophys. Res. Lett.*, 29(13), 1643, doi:10.1029/2001GL014201.
- Kitoh, A., H. Endo, K. Krishna Kumar, I.F.A. Cavalcanti, P. Goswami, T. Zhou, 2013: Monsoons in a changing world: a regional perspective in a global context. *J. Geophys. Res.*, 118, 3053-3065, doi:10.1002/jgrd.50258.
- Lin R., T. Zhou, Y. Qian, 2014: Evaluation of Global Monsoon Precipitation Changes based on Five Reanalysis Datasets, *Journal of Climate*, 27(3), 1271-1289
- Liu, J., B. Wang, M.A. Cane, S.-Y. Yim, and J.-Y. Lee, 2013: Divergent global precipitation changes induced by natural versus anthropogenic forcing. *Nature*, 493, 656-659
- Power, S., T. Casey, C. Folland, A. Colman, and V. Mehta (1999), Interdecadal modulation of the impact of ENSO on Australia, *Clim. Dyn.*, 15, 319–324.
- Qiu J. 2013: Monsoon Melee. *Science*, 340 (6139), 1400-1401; DOI:10.1126/science.340.6139.1400.
- Song, F., T. Zhou, and Y. Qian (2014), Responses of East Asian summer monsoon to natural and anthropogenic forcings in the 17 latest CMIP5 models, *Geophys. Res. Lett.*, 41, doi:10.1002/2013GL058705
- Song, F., T. Zhou, 2014: Interannual Variability of East Asian Summer Monsoon Simulated by CMIP3 and CMIP5 AGCMs: Skill Dependence on Indian Ocean–Western Pacific Anticyclone Teleconnection. *J. Climate*, 27, 1679-1697
- Song F., T. Zhou, 2014: The climatology and inter-annual variability of East Asian summer monsoon in CMIP5 coupled models: Does air-sea coupling improve the simulations ? *Journal of Climate*, doi:10.1175/JCLI-D-14-00396.1

- Sperber, K.R., H. Annamalai, I.-S. Kang, A. Kitoh, A. Moise, A. Turner, B. Wang and T. Zhou, 2013: The Asian summer monsoon: an intercomparison of CMIP5 vs. CMIP3 simulations of the late 20th century. *Clim. Dynam.*, 41, 2771-2744, doi:10.1007/s00382-012-1607-6.
- Trenberth, K. E., and D. J. Shea (2006), Atlantic hurricanes and natural variability in 2005, *Geophys. Res. Lett.*, 33, L12704, doi:10.1029/2006GL026894.
- Wang B., Q. H. Ding, X. H. Fu, I.-S. Kang, K. Jin, J. Shukla, and F. Doblas-Reyes, 2005: Fundamental challenge in simulation and prediction of summer monsoon rainfall. *Geophys. Res. Lett.*, 32, L15711, doi:10.1029/2005GL022734.
- Wang, B., J. Liu, H.-J. Kim, P.J. Webster, S.-Y. Yim, and B. Xiang, 2013: Northern Hemisphere summer monsoon intensified by mega-El Niño/southern oscillation and Atlantic multidecadal oscillation. *PNAS*, 110 (14), 5347-5352, doi:10.1073/pnas.1219405110.
- Wang, B., J. Liu, H.-J. Kim, P.J. Webster, and S.-Y. Yim, 2012: Recent Change of the Global Monsoon Precipitation (1979-2008). *Clim. Dyn.*, 39 (5), 1123-1135
- Wu, G., Y. Liu, T. Wang, et al., 2007: The Influence of the Mechanical and Thermal Forcing of the Tibetan Plateau on the Asian Climate. *J. Hydrometeorology* 8: 770-789.
- Wu, G., Y. Liu, B. He, Q. Bao, A. Duan & F.-F. Jin, 2012: Thermal Controls on the Asian Summer Monsoon. *Sci. Rep.* 2, 404; DOI:10.1038/srep00404.
- Zhou T., R. Yu, H. Li, and B. Wang, 2008, Ocean Forcing to Changes in Global Monsoon Precipitation over the Recent Half-Century, *Journal of Climate*, 21(15), 3833–3852
- Zhou T., B. Wu, B. Wang, 2009, How Well Do Atmospheric General Circulation Models Capture the Leading Modes of the Interannual Variability of the Asian-Australian Monsoon? *Journal of Climate*, 22, 1159-1173

Appendix: Description of the scientific objectives of four tasks of GMMIP

TASK-1: Understanding 20th century changes of global monsoons

The global monsoons have shown multi-decadal changes in the 20th century. Understanding the mechanisms of global monsoon changes and identifying the contributions of natural and anthropogenic forcing agents have been foci of the monsoon research community. **TASK-1** aims to reveal the role of forcing from the global oceans on monsoon precipitation change, and identify the relative contributions of natural and anthropogenic forcing (greenhouse gases and aerosols) by performing coupled, uncoupled, and partly coupled runs that cover the period from 1870 to 2013.

TASK-2: The role of Eurasian orography on the regional/global monsoons (Himalaya/Tibetan Plateau experiment)

Although monsoons are generally large-scale overturning circulations, apparent differences between characteristics of regional monsoons arise because of the different orography. This is most apparent for the Asia region, due to the existence of Tibetan – Iranian Plateau (TIP). The influence of the large-scale orography on the Asian summer monsoon includes both mechanical and thermal forcing. Various mechanisms have been suggested concerning the topographic effects; however, an overarching paradigm delineating the dominant factors determining these effects and the strength of impacts remains debated. **The goals of TASK-2** are to provide a benchmark of current model behavior in simulating the relationship of the monsoon to the Tibetan-Iranian Plateau (TIP, the highlands in 20-60°N, 25-120°E) so as to stimulate further research on the thermodynamical and dynamical effects of the TIP on the monsoon system. In particular the relative contributions of thermal and orographic mechanical forcing by the TIP to the Asian monsoon will be addressed. The task extends the studies from the TIP to other highlands including highlands in Africa, N. America and S. America.

TASK-3: Interannual variability of global monsoon precipitations

AGCM simulations with specified SST generally have low skill in simulating the summer precipitation over global monsoon domains, especially the Asian-western Pacific summer monsoon domain. This can be partly attributed to the exclusion of air-sea coupled processes. It is argued that in the real world the air-sea interaction in monsoon domains appears as “monsoon-driving-ocean”, but in an AMIP simulation, the interaction mechanism is “ocean-driving-monsoon” by construction (Wang et al. 2005). **The TASK-3 aims** to understand the air-sea interaction process in driving the interannual variability of global monsoons.

TASK-4: High resolution modeling of global monsoons

The monsoon rainbands are usually at a maximum width of 200 km. Climate models with low or moderate resolutions are generally unable to realistically reproduce the mean state and variability of monsoon precipitation. This is partly due to the model resolution. **The TASK-4 aims** to examine the performance of high-resolution models in reproducing both the mean state and year-to-year variability of global monsoons.

High Resolution Model Intercomparison Project (HighResMIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

Name of MIP

HighResMIP

Chairs

Rein Haarsma, KNMI, The Netherlands. (haarsma@knmi.nl)

Malcolm Roberts, Met. Office, UK. (malcolm.roberts@metoffice.gov.uk)

Suggested Members of the Scientific Steering Committee

Graeme Stephens, JPL, USA (graeme.stephens@jpl.nasa.gov)

Masahide Kimoto, Tokyo, University (kimoto@aori.u-tokyo.ac.jp)

Christiane Jablonowski, Univ. Michigan, USA (cjablono@umich.edu)

Lai-Yung (Ruby) Leung, PNNL, USA, Leung, (Ruby.Leung@pnnl.gov)

Websites

<http://www.wcrp-climate.org/index.php/modelling-wgcm-mip-catalogue/modelling-wgcm-mips/429-wgcm-hiresmip>

<https://dev.knmi.nl/projects/highresmip/wiki>

Goal of HighResMIP

For the first time, we want to assess the robustness of improvements in the representation of important climate processes with “weather-resolving” global model resolutions (~25km or finer), within a simplified framework using the physical climate system with constrained aerosol forcing.

Recent simulations with global high-resolution climate models have demonstrated the added value of enhanced resolution compared to the output from models in the CMIP3 and CMIP5 archive. They showed significant improvement in the simulation of aspects of the large scale circulation such as such as El Niño Southern Oscillation (ENSO) (Shaffrey et al 2009), Tropical Instability Waves (Roberts et al 2009), the Gulf Stream and its influence on the atmosphere (Chassignet and Marshall 2008; Kuwano-Yoshida et al 2010), the global water cycle (Demory et al. 2014), extra-tropical cyclones and storm tracks (Hodges et al. 2011) and Euro-Atlantic blocking (Jung et al 2012). In addition, the increased resolution enables more realistic simulation of small scale phenomena with potentially severe impacts such as tropical cyclones (Zhao et al. 2009), tropical-extratropical interactions (Haarsma et al. 2013) and polar lows. Other phenomena that are sensitive to increasing resolution are ocean mixing, sea-ice dynamics and monsoons. The improved simulation of climate also results in better representation of extreme events such as heat waves, droughts and floods.

The requirement for a multitude of multi-centennial simulations, including poorly constrained Earth System processes and feedbacks, has meant that model resolution within CMIP has progressed very slowly. In CMIP3 the typical resolution was 250km in the atmosphere and 1.5° in the ocean, while more than seven years later in CMIP5 this had only increased to 150km and 1° respectively. Until now high-resolution simulations have been performed at only a few research centers without overall coordination. Due to the large computer resources needed for these simulations, synergy will be gained

if these runs are done in a coordinated way, which enables the construction of a multi-model ensemble (since ensemble size for each model will be limited) with common integration periods, forcing and boundary conditions. The CMIP3 and CMIP5 data bases provide outstanding examples of the success of this approach. The multi-model mean has proven often to be superior to individual models in seasonal and decadal forecasting. Moreover, significant scientific understanding has been gained from analyzing the inter-model spread and attempting to attribute to model formulation.

HighResMIP will coordinate the efforts in the high-resolution modeling community. Joint analysis, based on process-based assessment and seeking to attribute model performance to emerging physical climate processes (without the complications of Earth System feedbacks) and sensitivity of model physics to model resolution, will further highlight the impact of enhanced resolution on the simulated climate. As a result of the widespread impact of resolution on the simulation of the climate, HighResMIP will contribute to all of the five grand challenges of the WCRP, and hence such analysis may begin to reveal at what resolution particular processes can be robustly represented.

The European institutes in Table I have submitted in September 2014 a proposal (PRIMAVERA) to the European Commission that coordinates the simulation and analyses of high-resolution runs. If HighResMIP is endorsed they will follow that protocol. During the WGCM meeting in October 2014 the following modelling centres have expressed their interest in participating in HighResMIP

- EC-Earth consortium, (KNMI, IC3, SMHI)
- Met. Office, UK
- NCAR, USA
- CMCC, Italy
- GFDL, USA
- CNRM, France
- MRI, Japan
- MPI, Germany
- CPTEC, Brazil
- IAP, China

In addition institutes that have expressed their interest in carrying out the HighResMIP experiments are Lawrence Berkeley National Laboratory (USA), JAMSTEC (Japan), PNNL (USA). Also institutes that are not able to undertake the HighResMIP simulations currently due to limited computer resources, such as the ARC Centre of Excellence for Climate System Science (Australia) have expressed their strong interest in analyzing the HighResMIP simulations.

Institution	MO/NCAS/ NOCS	KNMI/SMHI/ IC3/CNR	CERFACS	MPI	CMCC	ECMWF	AWI
Model names	UM / NEMO	ECEarth / NEMO	Arpege / NEMO	ECHAM / MPIOM	CCESM / NEMO	IFS / NEMO	ECHAM/ FESOM
Atmospheric resolution	60-25	T239-T799	T359	T255	25km	T239- T799	T255
Oceanic resolution	1/4-1/12°	1/4°-1/12°	1/4-1/12°	1/4-1/10°	1/4	1/4	1/4 - 1/12 spatially variable

Table I: European institutes, together with the models and the resolutions that are committed to HighResMIP (note that the eddy resolving 1/10-1/12° ocean may be used for a small subset of

simulations).

Proposed Experiments

The main experiments will be divided between Tier 1 and Tier 2, with further optional ideas in Tier 3 .

The Tier 1 experiments will be AMIP runs. A few institutes have already performed high resolution AMIP runs and published their results. These runs will not impose such prohibitively large technical difficulties and it is feasible for a considerable number of institutes to deliver a coordinated and coherent set of experiments.

For the coupled experiments the situation is somewhat different. Although a few institutes already have carried out high resolution coupled simulations, there still remain issues with for instance biases and spin-up. Due to these issues and the large amount of computer resources needed, only a limited number of institutes will be able to afford these coupled simulations, and hence they will be done in Tier 2.

Standard CMIP6 resolution experiments

To evaluate the impact of increased resolution the experiments in Tier 1 and Tier 2 will be repeated with the standard CMIP6 resolution. The experimental set-up and design of the standard resolution experiments will be exactly the same as for the high-resolution runs. This enables the use of HighResMIP simulations for sensitivity studies investigating the impact of resolution.

- **Tier 1**

AMIP runs

Resolution: *minimum 25-50 km at mid-latitudes for high resolution + a standard resolution configuration.*

This resolution is significantly higher than used in CMIP5. Century integrations for this resolution are now feasible.

Periods of integration: *Mid term 1950-2050*

The mid term period is relevant for decision makers, whereas prominent changes in climate and variability will only become more visible at the end of the 21st century. The start year of the integrations is 1950 to cover significant historical changes, and to allow a longer period of assessment than is found in standard AMIP-type simulations (typically 1979-2008).

Forcing: *CMIP6 scenario's*

CMIP6 scenarios that span the range from middle to high end scenarios. For the historical period all forcings natural and anthropogenic will be included.

For optimal comparison between the models aerosol concentrations should be used and not emissions – we plan coordination with HistMIP/RFMIP in order to secure historical aerosol concentrations (or else enough information to allow us to calculate concentrations from aerosol optical properties and number

concentrations).

The full details of the forcing datasets and strategy proposed can be found from the WCRP website (<http://www.wcrp-climate.org/index.php/modelling-wgcm-mip-catalogue/429-wgcm-hiresmip>). In summary, SST and sea-ice from 1950-present will likely be much the same as for standard AMIP-II integrations (based on HadISST since this is the only dataset that is long enough), while the future period will use methodologies such as Mizuta et al (2008) but adapted to make a continuous timeseries.

Minimum number of integrations: 1

Any manageable number is too low for a rigorous estimate of the internal variability. However, because the aim of the high-res protocol is to perform simulations at the highest possible resolution, the ensemble size has to be kept low. By using a strictly common protocol that is followed by many institutes, the effective multi-model ensemble will be much larger, enabling a much wider sampling than previously of the multi-model robustness of resolution impacts. In addition, if models can be proven to be portable, the ensemble size could be increased if other computer resources are available (discussions are already underway with the European PRACE supercomputing infrastructure). Some centers may be able to produce much larger ensembles, enabling a more robust estimate of internal variability.

- **Tier 2**

Coupled runs

The coupled runs will consist of pairs of both scenario (historic for the past) runs and, for comparison, control runs using fixed 1950s forcings. This will allow an evaluation of the model drift in addition to the climate change signal. It may be possible to use the ocean initial condition from pre-existing or already planned spin-up from historic or similar integrations.

Resolution: Atmosphere same as AMIP-runs. Ocean ~0.25 degree.

This enables the ocean to have some variability (compared to non-eddy permitting models), particularly in the tropics, and has been shown to change the strength of atmosphere-ocean interactions (Kirtman et al, 2012).

Period of integration:

- Scenario runs: Same as for AMIP runs
- Control runs: Minimal length as AMIP runs.

Forcing: Same as for AMIP runs

Minimum number of integrations: 1 for each of control and historic forcings

Ideally the ensemble number would be of order 3 simulations for each forcing, to help in evaluating model drift and enabling an improved sampling of internal variability, but this will quickly become very onerous on computing.

Coupling: Minimal daily coupling between ocean and atmosphere. Preferably more frequent, 3hr or 1hr.

Ocean-atmosphere interaction occurs on all time scales. With 3hr or 1hr the diurnal time scale can be resolved.

Initial state: Due to limited computer resources an equilibrated initial ocean state is not feasible. Possible solutions to circumvent this are bias correction or the interpolation of an initial state of the low resolution DECK runs. For the latter a prerequisite is that the dynamics of the low- and high resolution ocean model are sufficient similar.

- **Tier 3**

Optional additional simulations to be discussed by interested parties

These could include

1. Extension of the AMIP simulations to 2100 with agreed forcings, to give a stronger signal to noise ratio
2. Additional ensemble members for both AMIP and coupled simulations. Even if these are primarily at the standard resolution, it would enable a better understanding of internal variability, and hence be able to say if the high resolution differs significantly from that distribution.
3. Aqua planet simulations. These idealized simulations facilitate a more straight forward interpretation of the impact of resolution on model physics and dynamical behavior.
4. Switch-on 4xCO₂ in coupled models. This will enable assessment of possible changes in climate extremes and in climate sensitivity due to improved resolution which cannot be well simulated by the DECK-counterpart.

Connection with DECK

The DECK simulations will serve as an entry card for the CMIP6-Endorsed MIPS. The DECK experiments are

- AMIP simulations
- Pre-industrial control simulations
- 1%/yr increase in CO₂ concentration
- Switch-on 4XCO₂

The AMIP DECK simulation with the standard CMIP6 resolution will serve as the entry card for the Tier 1 HighResMIP simulations. For Tier 2 the other three coupled simulations of DECK with the standard CMIP6 simulations will serve as an entry card. This applies also to the CMIP6 Historical simulation which consists of a historical simulation from 1850-2014 using specific forcings consistent with CMIP6.

For the high-resolution simulations the DECK is too expensive in computer resources, but the comparison between the standard resolution simulations within HighResMIP and the DECK simulations will be informative in themselves.

Connection with other MIPS

GMMIP for global monsoons.

There is known sensitivity to monsoon flow and rainfall with model resolution in the West African monsoon, Indian monsoon (particularly via monsoon depressions) and possibly East Asian monsoon. As stated in GMMIP the monsoon rainbands are usually at a maximum width of 200 km. Climate

models with low or moderate resolutions are generally unable to realistically reproduce the mean state and variability of monsoon precipitation for the right reasons. This is partly due to the model resolution. The Tier 1 AMIP runs of HighResMIP will be used in Task-4 of GMMIP to examine the performance of high-resolution models in reproducing both the mean state and year-to-year variability of global monsoons.

SensMIP for parameter sensitivity

It is unclear how much the experimental design in SensMIP and HighResMIP overlap or complement each other. The multi-model high resolution ensemble could give one axis of uncertainty/variability from models, while a corresponding parameter sensitivity study would explore a different axis, but the limited number of parameters proposed to change in SensMIP may limit its use here.

Grand Challenges

HighResMIP will address the grand challenges of the WCRP in the following way

Clouds, Circulation and Climate Sensitivity (Rank 1)

HighResMIP will address this Grand Challenge in many different ways. The sensitivity of increasing resolution on water vapour loading, cloud formation, circulation characteristics and climate sensitivity will be investigated.

To improve the robustness of our understanding, the multi-model ensemble at different resolutions, together with the longer period AMIP integrations, will allow us to:

- (i) link tropospheric circulation to changing patterns of SSTs, land-surface properties, and understanding the role of cloud processes in natural variability
- (ii) examine the extent and limits of our understanding of patterns of precipitation
- (iii) examine changes in model biases (such as humidity) with resolution, since there are some indications that these may be linked to climate sensitivity

Increasing resolution affects in particular small scale process such as the formation of clouds. Although the formation of clouds has still to be parameterized in the resolution of HighResMIP the dynamical constraints for the formation of clouds, such as the location and magnitude of upwards and downwards motion, as well as moisture availability, are sensitive to resolution. This also applies to the response of the circulation to cloud formation.

Cryosphere in a Changing Climate (Rank 5)

Because in the Tier 1 experiments the sea-ice distribution is prescribed the contribution to this grand challenge is limited. Its main impact will be on the distribution of snow fall and subsequent accumulation and melting of snowpack that affect land surface hydrology. For instance the occurrence of intense polar systems, such as deep polar-lows that are accompanied by abundant snowfall will be better represented with increasing resolution.

In the Tier 2 coupled simulations the historic simulation will affect the growth of sea-ice and the air-sea heat flux, processes that are strongly affected by small scale processes. Here we can study the effect of model resolution on Arctic sea-ice variability, and possible influences on mid-latitude circulation.

Understanding and Predicting Weather and Climate Extremes (Rank 3)

HighResMIP is strongly related to this grand challenge. Increasing resolution of climate models will bring us closer to the ultimate goal of seamless prediction of weather and climate. Extremes mostly

occur and are driven by processes on small temporal and spatial scales that are not well resolved by standard CMIP6 climate models. Dynamical down scaling only partially resolves this due to the non-linear interaction between large and small spatial scales and the importance of representing global teleconnection patterns. We aim to improve our understanding of the interaction between global modes of variability (e.g. ENSO, NAO, PDO) and regional climate inter-decadal variability and extremes.

Regional Climate Information (Rank 4)

Regional climate information focuses on smaller scales and extreme events, which are relevant for stakeholders and adaptation strategies. This requires high resolution modeling to provide reliable information. Recent high resolution modeling studies (Di Luca et al. 2012; Bacmeister et al. 2013) and comparisons of CMIP3 and CMIP5 results (Watterson et al. 2014) have demonstrated the added value of increased resolution for regional climate information. Model outputs from HighResMIP could also be used by the regional climate modeling community for comparison of dynamical downscaling and global high resolution approaches and for further downscaling by cloud resolving regional models.

Sea-Level Rise and Regional Impacts (Rank 6)

For Tier 1 simulations there is no contribution to this grand challenge. For Tier 2 the contribution is limited although there is the potential for large contribution. If for instance the deep water formation and MOC response appears to be highly sensitive to resolution than there is a considerable impact on regional sea level rise. In addition resolving the topographic effect at high-resolution should have profound impacts on regional details about the sea level rise that are relevant for policy making and planning.

Changes in Water Availability (Rank 2)

HighResMIP is very relevant to this grand challenge. Resolution affects the hydrological cycle by modifying the land/sea partitioning of precipitation. Increasing resolution in general increases the moisture convergence over land (Demory et al. 2014) although regionally this can be reversed such for instance in Europe during the winter due to changes in the position of the storm track (Van Haren et al. 2014). In addition simulation of extreme precipitation events are highly sensitive to increasing resolution. How robust are these results across the multi-model ensemble? Can higher resolution models help to give insight into inconsistencies between global precipitation and energy balance datasets?

Biospheric forcings and feedbacks (Rank 7)

There is no direct link to this collaboration theme as the biosphere is not explicitly modelled. Because the response of the biosphere depends critically on the accurate simulation of the physical environment there is potential for spin-off studies, for instance by interpreting diagnostic information about vegetation production. Recycling of water is an important aspect of biospheric forcings and feedbacks, and the way that vegetation responds to drying depends on their role in recycling water - given the small scales of the involved processes this is strongly affected by model resolution.

GEWEX

HighResMIP fits in the GEWEX research focus of “Develop accurate global model formulation of the energy and water budget and demonstrate predictability of their variability and response to climate forcing”. Accurate modeling of the energy and water budget is sensitive to the adequate simulation of the energy conversions and phase transitions as well as the transport that occur on small spatial scales.

Overview of the proposed evaluation and analysis

The analysis will focus on the impact of increasing resolution on the simulation of the climate. The robustness of the impact of increasing resolution on the simulation of these phenomena among the different HighResMIP models will be investigated and their response to global warming assessed. One of the primary strengths of the simple experimental design for HighResMIP is that it enables a wide range of process-based analysis –simulation campaigns which included 1-2 models such as UPSCALE (Mizielinski et al. 2014) and Athena (Kinter et al. 2014) already have an extremely active number of analysis projects associated with them and insightful papers.

The results of the analysis of HighResMIP will be compared with the CMIP6 DECK experiments. Their experimental design, data format and documentation will follow the DECK experiments as far as possible.

The storage and distribution of the high resolution model data is a challenging issue that requires further discussion within HighResMIP. In the EU-proposal PRIMAVERA the JASMIN platform will be used for data exchange and as a common analysis platform.

Proposed timing

Start of the experiments: Beginning of 2016

End of the experiments: No fixed date.

References

Bacmeister, J.T., M.F. Wehner, R.B. Neale, A. Gettelman, C.E. Hannay, P.H. Lauritzen, J.M. Caron, and J.E. Truesdale, 2014: Exploratory high-resolution climate simulations using the Community Atmosphere Model (CAM). *Journal of Climate*, 27, 3073–3099, DOI: 10.1175/JCLI-D-13-00387.1.

Bell, R.J., J. Strachan, P.L. Vidale, K.I. Hodges and M. Roberts, 2013: Response of tropical cyclones to idealized climate change experiments in a global high resolution coupled general circulation model. *J. Climate*, 26 (20), 7966-7980.

Chassignet, E.P., and D. P. Marshall, 2008: Gulf stream separation in numerical ocean models. In: Hecht, M., Hasumi, H. (eds.), *Eddy-Resolving Ocean Modeling*, AGU Monog. Ser., 39-62.

Demory, M.-E., et al. (2014). The role of horizontal resolution in simulating drivers of the global hydrological cycle. *Clim. Dyn.*, 42, 2201-2225.

A Di Luca, R de Elfa, R Laprise, 2014: Potential for added value in precipitation simulated by high-resolution nested Regional Climate Models and observations. *Clim. Dyn.*, 38, 1229–1247. DOI 10.1007/s00382-011-1068-3

Haarsma, R.J., W. Hazeleger, C. Severijns, H. de Vries, A. Sterl, R. Bintanja, G.J. van Oldenborgh and H.W. van den Brink, 2013: More hurricanes to hit Western Europe due to global warming. *Geophys.*

Res. Lett. doi:10.1002/grl.50360.

Hertwig, E., J.-S. von Storch, D. Handorf, K. Dethloff, I. Fast and T. Krismer, 2014: Effect of horizontal resolution on ECHAM6-AMIP performance. *Clim. Dyn.* (in press).doi:10.1007/s00382-014-2396-x

Hodges, K. I., et al., 2011: A comparison of extratropical cyclones in recent re-analyses ERA-Interim, NASA MERRA, NCEP CFSR, JRA-25. *J. Clim.*, 24, 4888-4906.

Jung, T., et al., 2012: High-resolution global climate simulation with the ECMWF model in project Athena: Experimental design, model climate, and seasonal forecast skill. *J. Clim.*, 25, 3155-3172.

Kinter, III, J. L., B. Cash, D. Achuthavarier, J. Adams, E. Altshuler, P. Dirmeyer, B. Doty, B. Huang, E. K. Jin, L. Marx, J. Manganello, C. Stan, T. Wakefield, T. Palmer, M. Hamrud, T. Jung, M. Miller, P. Towers, N. Wedi, M. Satoh, H. Tomita, C. Kodama, T. Nasuno, K. Oouchi, Y. Yamada, H. Taniguchi, P. Andrews, T. Baer, M. Ezell, C. Halloy, D. John, B. Loftis, R. Mohr, and K. Wong, 2013: Revolutionizing Climate Modeling with Project Athena: A Multi-Institutional, International Collaboration. *Bull. Amer. Meteor. Soc.*, 94, 231–245. doi: <http://dx.doi.org/10.1175/BAMS-D-11-00043.1>

Kirtman, B. P., et al, 2012: Impact of ocean model resolution on CCSM climate simulations. 39, 1303-1328. <http://rd.springer.com/article/10.1007/s00382-012-1500-3>.

Kuwano-Yoshida A, Minobe S, Xie S-P, 2010: Precipitation response to the Gulf Stream in an Atmospheric GCM. *J. Clim* 23, 3676–3698.

Mizielinski, M. S., et al, 2014: High resolution global climate modelling; the UPSCALE project, a large simulation campaign. *Geosci. Model Dev.*, 7, 563-591.

Mizuta, R., Y. Adachi, S. Yukimoto, and S. Kusunoki, 2008: Estimation of the future distribution of sea surface temperature and sea ice using the CMIP3 multi-model ensemble mean, Tech. Rep. 56, 28 pp., Meteorol. Res. Inst., Tsukuba, Japan.

Roberts, M. J., et al., 2009: Impact of Resolution on the Tropical Pacific Circulation in a Matrix of Coupled Models. *J. Clim.*, 22, 2541-2556.

Shaffrey, L. C., et al., 2009: U.K. HiGEM: the new U.K. High-Resolution Global Environment Model-model description and basic evaluation. *J. Clim.*, 22,1861-1896.

Van Haren, R., R.J. Haarsma, G.J. van Oldenborgh and W. Hazeleger, 2014: Resolution dependence of European precipitation in a state-of-the-art atmospheric general circulation model. *J. Clim.* In review.

Watterson, I.G., J. Bathols, and C. Heady, 2014: What Influences the Skill of Climate Models over the Continents?. *Bull. Amer. Meteor. Soc.*, 95, 689–700. doi: <http://dx.doi.org/10.1175/BAMS-D-12-00136.1>

Zhao, M., et al., 2009: Simulations of Global Hurricane Climatology, Interannual Variability, and Response to Global Warming Using a 50km Resolution GCM. *J. Climate*, 33, 6653-6678.

Ice Sheet Model Intercomparison Project for CMIP6 (ISMIP6)

Application for CMIP6-Endorsed MIPs

Date: 23 November 2014

Proposals from MIPs should include the following information:

- * *Preliminary information used to determine whether a MIP should be endorsed for CMIP6 or not.*
- ** *Information that must be provided later (and before the panel can determine which experiments, if any, will be incorporated in the official CMIP6 suite).*

➤ Name of MIP*

ISMIP6: Ice Sheet Model Intercomparison Project for CMIP6

➤ Co-chairs of MIP (including email-addresses)*

Eric Larour, NASA Jet Propulsion Laboratory, USA, eric.larour@jpl.nasa.gov

Sophie Nowicki, NASA Goddard Space Flight Center, USA, sophie.nowicki@nasa.gov

Tony Payne, University of Bristol, UK, a.j.payne@bristol.ac.uk

➤ Members of the Scientific Steering Committee*

Helene Seroussi, NASA Jet Propulsion Laboratory, USA,

Heiko Goelzer, Vrije Universiteit Brussel, BE,

Andrew Shepherd, University of Leeds, UK,

William Lipscomb, Los Alamos National Laboratory, USA,

Jonathan Gregory, University of Reading and Met Office Hadley Center, UK,

Ayako Abe Ouchi, The University of Tokyo, JP

➤ Link to website (if available)*

➤ Goal of the MIP and a brief overview*

The primary goal of ISMIP6 is to improve projections of sea level rise via improved projections of the evolution of the Greenland and Antarctic ice sheets under a changing climate, along with a quantification of associated uncertainties (associated with both uncertainty in climate forcing and in the response of the ice sheets). As depicted in Figure 1, this goal requires an evaluation of AOGCM climate over and surrounding the ice sheets; analysis of simulated ice-sheet response from standalone models forced “offline” with CMIP AOGCM outputs and, where possible, with coupled ice sheet-AOGCM models; and experiments with standalone ice sheet models targeted at exploring the uncertainty associated with ice sheets physics, dynamics and numerical implementation. A secondary goal is to investigate the role of feedbacks between ice sheets and climate in order to gain insight into the impact of increased mass loss from the ice sheets on regional and global sea level, and of the implied ocean freshening on the coupled ocean-atmosphere circulation.

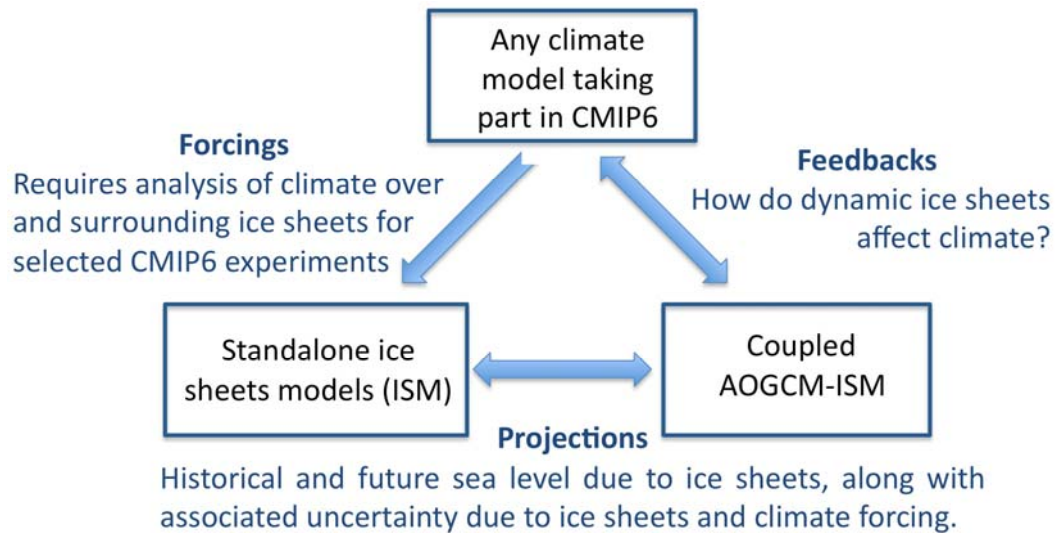


Figure 1: Overview of the ISMIP6 effort.

ISMIP6 is directly related to the WCRP Grand Challenges on ‘Changes in the Cryosphere’ and ‘Regional Sea-level Rise’. A white paper on the former identifies the need for “a focused effort on developing ice sheet models, with specific emphasis on the role of ice sheet dynamics on the rate of the sea-level rise”, which ISMIP6 is ideally placed to deliver by linking the improved process-based understanding delivered within CiC (and elsewhere) to projections of future ice-sheet mass budget. While a white paper on the latter identifies several open issues that strongly relate to our proposed activity, including the need understand the ocean’s response to high latitude freshwater forcing and the impact of ice sheet dynamics. ISMIP6 is primarily focused on the CMIP6 scientific question “How does the Earth System respond to forcing?” and offers the exciting opportunity of widening the current CMIP definition of Earth System to include (for the first time) the ice sheets. The emphasis on standalone, ensemble modelling will also shed light on the question “How can we assess future climate changes given climate variability, predictability and uncertainties in scenario” for the mass budget of the ice sheets and its impact of global sea level.

➤ References (if available)*

ISMIP6 is based on a long history of Ice Sheet Model Intercomparison Projects (ISMIP <http://homepages.vub.ac.be/~phuybrec/ismip.html>) and the more recent ice2sea (www.ice2sea.eu), Sea level Response to Ice Sheet Evolution (SeaRISE http://websrv.cs.umt.edu/isis/index.php/SeaRISE_Assessment), and COMBINE (<https://www.combine-project.eu/>) efforts. ISMIP6 brings together for the first time a consortium of international ice sheet models and coupled ice sheet-climate models to fully explore the sea level rise contribution from the Greenland and Antarctic ice sheets. Papers generated by these recent activities, that involved the ice-sheet modeling community, include:

- Bindshadler, R. et al. (2013) Ice-Sheet Model Sensitivity to Environmental Forcing and Their Use in Projecting Future Sea levels (The SeaRISE Project). *Journal of Glaciology*, 59 (214), 195-224.
- Edwards, T.L., et al. (2014) Effect of uncertainty in surface mass balance elevation feedback on projections of the future sea level contribution of the Greenland ice sheet. *The Cryosphere* 8(1), 195-208.
- Favier, L., et al. (2014) Retreat of Pine Island Glacier controlled by marine ice-sheet instability. *Nature Clim. Change* 4(2), 117-121.
- Nowicki et al. (2013) Insights into spatial sensitivities of ice mass response to environmental change from the SeaRISE ice sheet modeling project I: Antarctica. *Journal of Geophysical Research- Earth Surface*, 118 (2), 1002-1024.

- Nowicki et al. (2013) Insights into spatial sensitivities of ice mass response to environmental change from the SeaRISE ice sheet modeling project II: Greenland. *Journal of Geophysical Research- Earth Surface*, 118(2), 1025-1044.
- Pattyn, F., et al. (2013) Grounding-line migration in plan-view marine ice-sheet models: results of the ice2sea MISIP3d intercomparison. *Journal of Glaciology* 59(215), 410-422.
- Rae, J., et al. (2012) Greenland ice sheet surface mass balance: evaluating simulations and making projections with regional climate models. *The Cryosphere* 9(6), 1275-1294.
- Shannon, S.R., et al. (2013) Enhanced basal lubrication and the contribution of the Greenland ice sheet to future sea level rise. *Proceedings of the National Academy of Sciences* 110(35), 14156-14161.
- Shepherd, A., et al. (2012) A Reconciled Estimate of Ice-Sheet Mass Balance. *Science* 338(6111), 1183-1189.

➤ An overview of the proposed experiments*

The overall framework for ISMIP6 is designed to deliver projections of the ice sheet contribution to sea level rise. Together with the proposed glacier CliC (Climate and Cryosphere) targeted activity and projections of thermal expansion (that already sit within the CMIP framework), this will allow sea level to become part of the family of variables for which CMIP can provide routine IPCC-style projections. The proposed experiments will both use and augment the CMIP6-DECK, Historical and ScenarioMIP experiments, as summarized in Table 1. ISMIP6 will use the standard CMIP AGCM and AOGCM experiments for analysis of the climate over and surrounding the ice sheets, and as forcing for the standalone ice sheet models (ISM) projections. Additional sensitivity experiments will be performed with the ISM to investigate the uncertainty associated with these projections arising from ice sheet models. The key output will be an ensemble of historical and future estimates of ice sheet contribution to sea level. To address the feedbacks introduced by interactive ice sheets, we propose that a small number of selected DECK experiments are repeated with coupled AOGCM-ISM, where the ice sheet is an interactive component of the AOGCM. Our assessment of the state of existing AOGCMs is that coupled models including an interactive Greenland ice sheet can realistically be expected for CMIP6, however including the Antarctic ice sheet remains challenging (because of the greater complexity of its response to climate forcing, and the issues associated with simulations of the Southern Ocean). It is for these reasons that ISMIP6 heavily relies on standalone ice sheet models driven offline by CMIP6 climate models for projections of sea level.

Existing CMIP exp. used by ISMIP6 (AGCM-AOGCM only, no dynamic ice sheet required)	Standalone ISMIP6 ice sheet model exp. (ISM only)
<ul style="list-style-type: none"> - AMIP simulation (<i>amip</i>) - CMIP6 Historical Simulation (<i>historical</i>) - Pre-Industrial Control (<i>piControl</i>) - 1% yr CO2 to quadrupling CO2 (<i>1pctCo2</i>) - ScenarioMIP SSP5-8.5 up to year 2300 (<i>ssp5-8.5</i>) 	<ul style="list-style-type: none"> - ISM control (<i>piControlforcedism</i>) ** - ISM for last few decades forced by <i>amip</i> (<i>amipforcedism</i>) - ISM for the historical period forced by <i>historical</i> (<i>historicalforcedism</i>) - ISM forced by <i>1pctCo2</i> (<i>1pctCo2forcedism</i>) for quantification of feedback** - ISM for 21st century and up to 2300 sea level forced by ScenarioMIP <i>ssp5-8.5</i> (<i>ssp5-8.5forcedism</i>) * - Other ISMIP6 specific experiments** to explore uncertainty due to ISM.
New ISMIP6 CMIP6 exp. (Coupled AOGCM-ISM)	
<ul style="list-style-type: none"> - Pre-Industrial Control (<i>piControlwithism</i>) ** - 1% yr CO2 to quadrupling CO2 (<i>1pctCo2withism</i>) ** - ScenarioMIP SSP5-8.5 up to year 2300 (<i>ssp5-8.5withism</i>) 	

Table 1: Overview of experimental framework for ISMIP6 (further details on experimental design and motivation are explained in later sections). Name of experiments are indicated in italic. *These types of standalone ensemble ISM experiments were implemented in the European ice2sea and SeaRISE efforts for IPCC-AR5, but using forcing derived from AR4 (See www.ice2sea.eu and http://websrv.cs.umt.edu/isis/index.php/SeaRISE_Assessment) **These types

of experiments, where the ice sheet is an interactive component of the AOGCM, have been recently run as part as the European COMBINE effort (<https://www.combine-project.eu/>) by three modeling groups: IPSL, MIP-I, and DMI.

➤ An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*

The primary goal of ISMIP6 is an analysis of the historical and future estimates of ice sheet contribution to sea level, and associated uncertainty, via evaluation of the ensemble simulations. For evaluating the feedbacks introduced by coupling dynamic ice sheets to AOGCM, we will compare the results of simulations of AOGCM with and without dynamic ice sheet models, and of ice sheet forced by offline with AOGCM and fully coupled to AOGCM.

This goal therefore also requires that three components of the Earth system are evaluated and analyzed by comparing to in situ, airborne and satellite observations:

- 1) The ice sheet dynamics. Flux gates will be defined along grounding lines at the coast, where estimated transports derived from observed surface velocities may be employed. The coupled system allows for an assessment of the total ice sheet contribution to sea level rise, which may be evaluated against GRACE data (available for 2003-present). Model performance may also be assessed with observed changes in surface elevation from ERS-1 and ERS-2 (1992-present) and ICESat (2003-2009), along with ice velocities, ground line and ice front locations.
- 2) The atmosphere and surface conditions over the ice sheets (surface radiative and turbulent fluxes, albedo, temperature, surface mass balance). This component may be divided into two parts: climate forcing that would be generated by an AOGCM and processes at the ice sheet surface (that may or may not be adequately simulated by an AOGCM but will be used in standalone ice-sheet models, such as SMB and albedo evolution). The evaluation of atmospheric state variables including temperature can make use of observations from established automatic weather station networks and surface radiation budget observatories at South Pole and Summit. Surface temperature and albedo may also be evaluated with remote sensing values from AVHRR (1982-present) and MODIS (2000-present). Simulated accumulation may be evaluated at in situ locations along the K-transect for Greenland, and with shallow ice cores distributed across both ice sheets. A comparison with regional climate model output and atmospheric reanalyses is also suggested as a quality test.
- 3) The ocean around the ice sheets (sea surface height, ocean, temperature, ocean induced melting rates, wind stress, hydrographic properties at the margins of the ice sheets to the extent available, sea-ice cover). This component may also be divided into two parts: ocean forcing that would be generated by an AOGCM and processes at the ice sheet boundary (that may or may not be captured within an AOGCM but will also be included in standalone ice-sheet models, such as ice-shelf melt). A key concern will be the validation of ocean thermal forcing of the ice sheets, which is likely to focus on evolving temperature at depth and, in particular, AOGCM simulation of the Southern Ocean.

➤ Proposed timing*

The analysis of atmospheric and oceanic climate over and surrounding the ice sheets from the CMIP5 archive will begin immediately in order to assess the quality and implied change in surface mass balance and temperatures. Analysis of the CMIP6 data would be ongoing and follow the simulation phase of CMIP6.

ISMIP6 started the design of the standalone ice sheet experiments during a workshop in July 2014, therefore further refining these experiments and data preparation would be completed by mid 2015. The sea level projections and quantification of the uncertainty in sea level due to ice sheets would begin mid 2015, and continue in tandem with CMIP6. Analysis of the projection simulations and sensitivity experiments would be ongoing in order to identify the dominant sources of uncertainty.

The runs for the AOGCM-ISM simulations would occur towards the beginning or middle of the CMIP6 cycle, so that the knowledge gained from the effect of dynamic ice sheets and associated feedback can be incorporated in other MIPs.

- For each proposed experiment to be included in CMIP6**
 - the experimental design;
 - the science question and/or gap being addressed with this experiment;
 - possible synergies with other MIPs;
 - potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

As summarized in Figure 1 and Table 1, the experimental design for ISMIP6, consist of three different types of modeling efforts: 1) standard AGCM-AOGCM experiments from the DECK, CMIP6 Historical and ScenarioMIP simulations (therefore all climate models participating in CMIP6 will be included), 2) simulations with standalone ice sheet models, and 3) simulations with coupled AOGCM-ISM when possible.

The detailed specifics for the ISMIP6 experiment design will be provided in a paper for the CMIP6 Special Issue. In summary, the following experimental design is proposed:

- 1) Use of selected standard AGCM & AOGCM CMIP6 experiments over and surrounding ice sheets:
 - *amip*: Allows the evaluation of AGCM climate over ice sheets, in particular surface mass balance (SMB: the combination of precipitation, evaporation and surface runoff).
 - *historical*: Allows the evaluation of AOGCM climate over and surrounding the ice sheets for the CMIP6 historical period (1850-2014).
 - *ssp5-8.5*: standard ScenarioMIP SSP5-8.5 simulation starting from 2015 but continued to year 2300 if possible. The experiment would assess projected changes in SMB with fixed ice sheet extent and topography.
 - *piControl* and *1pctCo2*: Will be used to assess the impacts of introducing dynamic ice sheets in AOGCM. For modeling groups taking part in the AOGCM-ISM experiments, the duration of the experiment need to be the same for both AOGCM and AOGCM-ISM simulations.

Note: we would start with the existing CMIP5 output and repeat the analysis when CMIP6 output is available. The output from these experiments will be used to assess the uncertainty in sea level arising from climate forcing and to drive the standalone ice sheet models.

- 2) Standalone ice sheet models experiments:
 - *piControlforcedism*: ISM control, Constant forcing, needed to evaluate model drift.
 - *amipforcedism*: simulation for the last few decades to understand the well observed record of ice sheet changes. ISM would be driven by SMB anomalies obtained from the standard AMIP DECK simulations, and ice shelf basal melting or temperature anomalies from ocean models.
 - *historicalforcedism*: simulation for the historical period to understand the ice sheet contribution to 20th century GMSLR, forced by outputs obtained from the standard CMIP6 Historical simulation. The results of *amipforcedism* and *historicalforcedism* are likely to differ, and the comparison will provide some insight into the relative importance of biases, climate variability and climate change.
 - *1pctCo2forcedism*: simulation forced by 1% yr CO₂ to quadrupling CO₂ obtained from DECK output: for comparison with the AOGCM-ISM experiment in order to evaluate ice sheet feedback.
 - *ssp5-8.5forcedism*: simulation for the 21st century (and maybe up to the 23rd century depending on ScenarioMIP) for the most realistic ice sheet contribution to sea level projections. ISM would be driven by SMB anomalies (with adjustments for ice sheet elevation change) and ice shelf mass balance or temperatures anomalies derived from the standard SSP5-8.5 ScenarioMIP simulation.

- Additional ISM experiments would be designed to assess the uncertainty in sea level projections due to ice sheet models. These experiments would explore the model biases and uncertainties identified in the ice2sea and SeaRISE efforts, which include ice sheet initialization, poorly known basal conditions and subgrid-scale processes. In addition, ISMIP6 would investigate questions such as “How much excess oceanic heat flux is required to trigger marine ice sheet instability?” to shed light on the potential collapse of the Antarctic ice sheet.

Note: Following the approach taken in the ice2sea and SeaRISE efforts, the anomalies derived from the DECK experiments would be added to the forcing used in the ISM control runs. The AGCM/AOGCM output is in most cases not suitable to directly force standalone ISMs, mainly due to differences in spatial resolution. A downscaling procedure (to be later specified) will be necessary to produce surface mass balance and for ice shelves basal mass balance, used to drive the ISMs. Modeling groups could decide to carry out the experiments for both the Greenland and Antarctic ice sheets, or to focus on one ice sheet.

3) Coupled AOGCM-ISMs experiments (same set up as standard CMIP6 experiments but with evolving ice sheets models (ISM): GCM sends ISM an energy balance based SMB, and ISM sends GCM adjustments to land surface elevation and surface type)

- *piControlwithism*: the pre-industrial control, where the aim is to produce a realistic non-drifting coupled state, and assess systematic model bias. The spin up may require the GCM and ISM to be asynchronously coupled until the system reaches quasi-equilibrium, which would be followed by a multi-hundred years run (500 yrs suggested), in order to capture unforced natural variability.
- *1pctCO2withism*: the 1% per yr CO₂ increase to quadrupling CO₂ over 140 yrs and kept constant at 4xCO₂ for an additional two to four centuries. This experiment, along with *piControlwithism*, are the core experiments that will be used for analysis of coupled ice sheet-climate system. Experiment would be compared to the standard DECK without ice sheets and to the standalone ISM forced by the standard DECK, in order to diagnose the strength of ice sheet-climate feedback and the associated uncertainty in projections resulting from excluding ice sheet models. Length of experiment *1pctCO2*, *1pctCO2withism* and *1pctCO2forcedism* therefore needs to be the same for groups participating in this experiment. It is suggested the experiments are run for a minimum of 350 yrs and up to 500 yrs is encouraged, because results from COMBINE effort indicate that ice sheet model coupled runs start to clearly divert from the uncoupled runs after about 250-300 yrs of simulations.
- *ssp5-8.5withism* scenario for analysis of coupled system and sea level projections from a coupled framework, which can be compared to the standalone ice sheet model projection. Experiment would cover the 21st century and preferably run out to the 23rd century. The set up would follow the set up for the standard SSP5-8.5, which may therefore first require the CMIP6 Historical simulation to be performed too with a coupled AOGCM-ISM setting.

Note: We suggest that the pre-industrial control and 1% yr CO₂ to quadrupling CO₂ experiments are performed first, followed by ScenarioMIP SSP5-8.5. Modeling groups could decide to carry out the experiments for both the Greenland and Antarctic ice sheets, or to focus on one ice sheet. These experiments should only differ from the equivalent standard CMIP AOGCM setting in the manner in which the ice sheet is treated, so that the exploration of feedbacks is not affected by other changes. Feedbacks that we propose to explore include albedo-melt feedback, elevation-SMB feedback, precipitation-sea ice feedback, fresh water (runoff and icebergs calving and submarine melting)- ocean feedback, atmospheric circulation – ocean heat flux feedback (e.g. tip-jets, katabatic winds). This type of coupled experiments have been carried out by 3 modeling centers (IPSL, MPI-M, DMI) and soon MeteoFrance as part of the European COMBINE project. These coupled models (**IPSL**: IPSL – GRISLI; **MPI-M**: MPI-ESM – PISM and MPI-ESM – SICOPOLIS; **DMI**: EC-Earth – PISM, and soon CNRM-GRISLI) were described and evaluated in the COMBINE reports ‘Assessment of performance of AOGCMs coupled to Greenland and Antarctic models’ (http://www.combine-project.eu/fileadmin/user_upload/combine/dels/D4.3.pdf) and ‘Feedbacks of individual components: Cryosphere’ (http://www.combine-project.eu/fileadmin/user_upload/combine/dels/D7.7_v2.pdf, in the second part: ‘Impacts of including an interactive Greenland ice sheet in ESM). Efforts of including

dynamic ice sheets into AOGCMs are also occurring with CESM, GFDL and ModelE for example, so it is expected that about 8-10 groups will be in a position to run such experiments for CMIP6.

The primary goal of these experiments is to improve sea level projections due to changes in the ice sheets, and assessing the uncertainty in these projections due to climate forcing versus that arising from ice sheet models. The secondary goal is to understand how ice sheets affect and are affected by climate. These experiments will thus shed light on the key science questions considered by CMIP6: “How does the Earth system respond to forcing?”, “What are the origin and consequences of systematic model biases”, and “How can we assess future climate change given uncertainty in scenarios?”. These goals directly contribute to the Cryosphere Grand Challenge and the Sea Level Rise Grand Challenge of the Climate and Cryosphere (CliC) project and the World Climate Research Program. Finally, the ISMIP6 sea level projections will be relevant to the Impacts Adaptation and Vulnerability (IAV) community and policy makers.

Possible synergies with other MIPs include:

- High Resolution Model Intercomparison Project (HighResMIP). We will use the results from the high-resolution runs to quantify the impact of increased resolution in our standalone ice sheet suite, and to compare against the results from the DECK runs. Particular processes such as atmospheric blocking will be looked at to understand how well extreme melt-rate events are captured in our runs.
- Coordinated Regional Climate Downscaling Experiment (CORDEX). The CORDEX results will be used against the DECK runs to quantify additional sensitivities not captured in the low-resolution runs or HighResMIP runs. Biases and additional variability in the downscaled CORDEX results will be introduced in the offline ice sheet model runs.
- Land Surface, Snow and Moisture (LS3MIP). One of the objectives of LS3MIP is an evaluation of the current state of the snow cover representation in climate models, which impacts the surface mass balance over the ice sheets. These experiments with land-module may help towards understanding and quantifying the uncertainty in sea level due to surface forcing.
- Scenario Model Intercomparison Project (ScenarioMIP). We have already contacted the members of ScenarioMIP steering committee to indicate that ISMIP6 is interested in an extension of the SSP5-8.5 beyond the planned 2015-2100 timeframe (ideally up to year 2300).
- Observations for Model Intercomparison (Obs4MIP). We would use the observations available on the current database to test how the inclusion of dynamic ice sheets affects the simulations. We would also suggest additional datasets that are pertinent to ice sheets and surface mass balance.
- Reanalysis for Model Intercomparison (Ana4MIP). We would use reanalysis in the assessment of the surface mass balance from AOGCM, and potentially as contemporary forcing for the ice sheets.
- Paleoclimate Modelling Intercomparison Project (PMIP), in particular the proposed Last Glacial Maximum and the PlioMIP experiments. The latter is concerned with investigating how does the Earth System respond in the long term to CO₂ forcing analogous to that of the modern and the significance of CO₂ induced polar amplification for the stability of ice sheets, sea-ice and sea level.

These synergies with other MIPs illustrate potential collaborations within the climate community. It is hoped that other MIPs would be interested in using our simulations to investigate how changes in the ice sheets affect the component of the climate that is their expertise.

➤ If possible, a prioritization of the suggested experiments, including any rationale**

For the coupled AOGCM-ISM, we suggest that the pre-Industrial control and 1% yr CO₂ to quadrupling CO₂ experiments are performed first, followed by the SSP5-8.5 of ScenarioMIP. Our Tiers 1 experiments are thus: *piControlwithism* and *1pctCO2withism*, which allow for an easier evaluation of ice-climate feedback and have already been performed by many modeling groups. Our Tiers 2 experiment, *ssp5-8.5withism*, is however more

relevant to our goal of sea-level rise projections that are in sink with the CMIP6 future climate, and SSP5-8.5 will be the focus of our sea level projection with standalone ice sheet models.

- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**

No objections.

- List of output and process diagnostics for the CMIP DECK/CMIP6 data request**
 - whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;
 - whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);
 - whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.
 - whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;
 - the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.

The current CMIP5 CMOR tables Amon (Monthly Mean Atmospheric Fields), Omon (Monthly Mean Ocean Fields), LImon (Monthly Mean Land Cryosphere Fields), and Olmon (Monthly Mean Ocean Cryosphere Fields) already contains many of the output required to diagnose and intercompare the climate over glaciated land/ice sheets and to derive forcing for the ice sheets. However a few additional variables may be needed to properly derive the forcings for ice sheets and to record outputs from the evolving ice sheets in the coupled AOGCM-ISMs experiments (such as ice elevation change). Table 2 list our initial assessment of the Amon and LImon variables that we plan to use in ISMIP6, or that are missing. Unless otherwise stated, these variables would be on the atmosphere grid and contain monthly output. As one of our first task is to evaluate the existing CMIP5 models output for the DECK experiments that we will be using, we will revisit Table 2 during this effort, and include the oceanic variables. We will complete the CMIP6 data request forms by the deadlines set by CMIP and WIP.

Variable	Units	Existing CMOR variable name or comment if new variable
Near Air Temperature	K	tas in Amon
Snow area fraction	%	snc in LImon
Surface Snow and Ice Sublimation Flux	$\text{kg m}^{-2} \text{s}^{-1}$	sbl in LImon
Surface Rainfall rate	$\text{kg m}^{-2} \text{s}^{-1}$	pr in Amon
Surface snowfall rate	$\text{kg m}^{-2} \text{s}^{-1}$	prsn in Amon

Snow Melt rate	$\text{kg m}^{-2} \text{s}^{-1}$	snm in Lmon or snm in Amon
Latent Heat flux	W m^{-2}	hfls in Amon
Sensible Heat flux	W m^{-2}	hfss in Amon
Downwelling Shortwave over ice sheet	W m^{-2}	rsds in Amon
Upward Shortwave over ice sheet	W m^{-2}	rsus in Amon
Downwelling Longwave over ice sheet	W m^{-2}	rlds in Amon
Upward Longwave over ice sheet	W m^{-2}	rlus in Amon
Calving Flux	$\text{kg m}^{-2} \text{s}^{-1}$	The loss of ice sheet due to iceberg calving. Exist in Omon as ficeberg, would be on Omon grid
New variables that would be added to the Lmon Table, these quantities would sometime origin from the ice sheet grid but be remapped to the atmosphere grid. When there are no ice sheets, these values would be reported as "missing" because zero is a valid value in the melt rates, and can lead to mistaken interpretation by analysts.		
Ice sheet area fraction	%	Fraction of grid cell covered by ice sheets or glaciated land (similar to sci in Olmon)
Ice Sheet Altitude	m	The altitude or surface elevation of the ice sheet in the atmosphere portion of the grid cell.
Surface Temperature of Ice Sheet	K	Similar to tsice in Olmon but over glaciated land
Temperature at the interface between ice sheet and snow	K	Similar to tsnint in Olmon but over glaciated land
Rate of Melt at upper surface of ice sheet or ice shelf	$\text{kg m}^{-2} \text{s}^{-1}$	Similar to tmelt in Olmon but over glaciated land or ice shelf
Rate of Melt at lower surface of ice sheet or ice shelf	$\text{kg m}^{-2} \text{s}^{-1}$	Similar to bmelt in Olmon but under glaciated land or ice shelf
Mass flux of surface meltwater which refreezes within the snow or firn	$\text{kg m}^{-2} \text{s}^{-1}$	the existing standard name is surface_snow_and_ice_refreezing_flux
Surface Runoff	$\text{kg m}^{-2} \text{s}^{-1}$	Similar to mrros in Lmon but over glaciated land or ice shelf

Table 2: Data to be saved on the atmosphere grid (monthly) to capture the glaciated/ice sheet surface realm. Most of these variables already exist in the CMIP5 tables.

For diagnosis and intercomparison of the dynamical ice sheet models within AOGCM (the coupled AOGCM-ISM), the variables in Table 3 would be saved on the dynamical ice sheet native grid or on a regular (5x5km) grid designed for the ice sheets (such as done in the SeaRISE effort). These variables would be recorded for the entire length of the experiments involving ice sheets.

Variable	units	comment
Ice Sheet Altitude	m	The altitude or surface elevation of the ice sheet
Ice Sheet Thickness	m	The mean thickness of ice sheet
Bedrock Altitude	m	The bedrock topography.
Calving Flux	$\text{kg m}^{-2} \text{s}^{-1}$	The loss of ice sheet due to iceberg calving (solid ice discharge).
Individual components of Upper Surface Mass Balance	$\text{kg m}^{-2} \text{s}^{-1}$	surface mass balance = snowfall – sublimation – meltwater runoff
Individual components of Surface Mass Balance beneath ice sheet and ice shelf	$\text{kg m}^{-2} \text{s}^{-1}$	
Column mean ice velocity	m/yr	
Magnitude of ice velocity at the upper surface	m/yr	
Magnitude of ice velocity at the lower surface	m/yr	
Magnitude of basal shear stress	Pa	
Basal Water Content	m	The effective thickness of water beneath the grounded ice sheet
Surface Temperature of Ice Sheet	K	
Basal Temperature of Ice Sheet	K	
Mask	1	Grounded ice sheet, floating ice shelf, ice free land

Table 3: Example of data to be saved on the ice sheet grid (monthly or yearly) to capture the dynamical ice sheet model realm. It would therefore be a new CMOR table.

- Any proposed contributions and recommendations for**
 - model diagnostics and performance metrics for model evaluation;
 - observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;
 - tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).

Model evaluation over the ice sheets will include in situ, airborne, satellite data and reanalysis. Most of these cryospheric data products are not currently available in the obs4MIPs database and we will work closely with obs4MIPs to rectify this. The process is complicated by the need to evaluate both climate forcing over and around the ice sheets (i.e., AOGCM model output) and ice-sheet model response. We plan to have a workshop with data providers and modelers in 2015 to finalize the evaluation plan. The recent IMBIE project (Shepherd et al 2013) provides an excellent example of the ice-sheet observational community work together to provide a reconciled product suitable for testing ice-sheet models.

- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms. **

- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF. **

Some new standard CF names will be needed for ice sheet quantities, and there may be a need for ice sheet-grids to be handled in order to record the fields in Table 3, perhaps by CMOR.

Land Surface, Snow and Soil Moisture (LS3MIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

Land Surface, Snow and Soil moisture MIP (LS3MIP)

- Co-chairs of MIP: Bart van den Hurk (hurkvd@knmi.nl), Gerhard Krinner (krinner@ujf-grenoble.fr), Sonia Seneviratne (sonia.seneviratne@ethz.ch), Chris Derksen (Chris.Derksen@ec.gc.ca), Taikan Oki (taikan@iis.u-tokyo.ac.jp) and Hyungjun Kim (hjkim@rainbow.iis.u-tokyo.ac.jp)
- Members of the Scientific Steering Committee: Martin Best, Paul Dirmeyer, Herve Douville, Richard Essery, Stefan Hagemann, Alex Hall, Randy Koster, Dave Lawrence, Twan van Noije, Helmut Rott, Andrew Slater, Matthew Sturm, Andrea Alessandri, Greg Flato
- Endorsement: CliC and GEWEX
- Link to website: <http://hydro.iis.u-tokyo.ac.jp/GSWP3>, <http://www.iac.ethz.ch/GLACE-CMIP>, and <http://www.climate-cryosphere.org/activities/targeted/esm-snowmip>

Goal of the MIP and brief overview

The goal of the LS3MIP experiment is to provide a comprehensive assessment of land surface-, snow-, and soil moisture-climate feedbacks, and to diagnose systematic biases in the land modules of current ESMs using constrained land-module only experiments. The solid and liquid water stored at the land surface has a large influence on the regional climate, its variability and its predictability, including effects on the energy and carbon cycles. Notably, snow and soil moisture affect surface radiation and flux partitioning properties, moisture storage and land surface memory. They both strongly affect the atmospheric conditions, in particular air temperature, but also large-scale circulation patterns and precipitation. However, models show divergent responses and representations of these feedbacks as well as systematic biases in the underlying processes. LS3MIP will provide the means to quantify the associated uncertainties and to better constrain climate change projections, of particular interest for highly vulnerable regions (densely populated regions, polar regions, agricultural areas, land ecosystems).

A short description of the role of snow and soil moisture in the climate system and of the rationale for the proposed experiments is provided hereafter.

Snow processes and snow-climate feedbacks

Snow cover is an essential component of the Earth System that interacts with the atmosphere and the surfaces it covers (land, ice, sea ice). It is also an important source of (positive) feedbacks within the climate system. A WCRP/CliC Initiative was proposed in 2013 for an ESM-SnowMIP intercomparison programme as a contribution to the WCRP Grand Challenge Cryosphere in a Changing Climate. This initiative builds on the evaluation of the current state of snow cover representation in climate models, which is being broadly addressed by observational and modeling groups across the snow community. It is a core element of the LS3MIP experiment.

It has been shown that CMIP5 models underestimate the observed spring snow cover trend in the Arctic (Derksen and Brown, 2012) and in the Northern Hemisphere (Brutel-Vuilmet et al., 2013). Snow-related climate feedbacks in the climate system arise primarily because of the well-known albedo feedback (e.g. Qu and Hall, 2007) that is also one of the main mechanisms leading to Arctic

Amplification (e.g. Holland and Bitz, 2003). Snow-related biases in climate models may arise through this feedback, but also through the energy sink induced by snow melting in spring and through the strong thermal insulation effect of snow on the underlying soil. Koven et al. (2012) related strong biases in the simulated Northern Hemisphere permafrost extent in CMIP5 models to the representation of snow in these models.

Because of strong snow/atmosphere feedbacks, it is difficult to distinguish and quantify the various potential causes for disagreement in observed versus model snow trends. These causes include: the underestimation of the recent spring warming trend in CMIP5 models (e.g., Brutel-Vuilmet et al., 2013), weaknesses in their representation of snow processes, especially regarding the snow/albedo feedback (Qu and Hall, 2014), a positive pre-melt snow water equivalent (SWE) bias in CMIP5 models across the mid latitudes and the Arctic (Brown and Mote, 2009), increased deposition of light-absorbing impurities on snow which is not accounted for in most models (e.g., Dumont et al., 2014), or a combination of these with other unknown processes.

A better understanding of the links between snow cover and climate is critical to interpret the observed changes in recent years including links to variability in the atmospheric and ocean circulation, and the misrepresentation of polar amplification by climate models in the Arctic. It is a prerequisite for increasing the confidence in the projections of snow cover and its role in the subarctic (boreal) and Arctic climate. This understanding is also necessary for the long-term improvement of the representation of snow in climate models, which will also impact seasonal to interannual prediction of temperature, runoff and soil moisture.

The SnowMIP1 (Etchevers et al., 2002) and SnowMIP2 projects (Essery et al., 2009) evaluated the capacity of snow models of different complexity to simulate the snowpack evolution from local meteorological forcings. These projects were based on the evaluation of stand-alone simulations of snow models over a limited number of instrumented sites (see also <http://www.wcrp-climate.org/index.php/modelling-wgcm-mip-catalogue/57-unifying-themes/modelling-wgcm/catalogue-of-model-intercomparison-projects/276-modelling-wgcm-catalogue-snowmip>). However these pioneering projects did not explore snow-climate interactions, and were limited to the site scale.

LS3MIP will consider both stand-alone snow simulations at the global scale and snow outputs from climate simulations. Dedicated experiments will be designed for evaluating and understanding snow feedbacks within current climate models and assessing snow-related uncertainties in future projections. These will be a key action of the WCRP Grand Challenge Cryosphere in a Changing Climate coordinated by CliC/WCRP.

Soil moisture processes and soil moisture-climate feedbacks

Soil moisture modulates the energy and water balance at the land surface to a large extent (Koster et al., 2004; Seneviratne et al., 2010; van den Hurk et al., 2011). It interacts with vegetation, melting snow, ground water, boundary layer processes, atmospheric moisture, and is a key element for available fresh water resources, heat wave and drought propagation and soil erosion.

The modulating role of soil moisture is eminent at many relevant time scales: diurnal cycles of land surface fluxes, (sub-)seasonal predictability of droughts, floods, and hot extremes, annual cycles governing the water buffer in dry seasons, and shifts in the climatology in response to changing patterns of precipitation and evaporation (e.g. Betts 2004, Ek and Holtslag 2004, Santanello et al. 2009, Koster et al. 2010a,b, Douville et al. 2012, Mueller and Seneviratne 2012, Quesada et al. 2012, Dirmeyer et al. 2013, Miralles et al. 2014, Greve et al. 2014).

An important notion is the difficulty in generating reliable observations of soil moisture and land surface fluxes that can be used as boundary conditions for modeling and predictability studies. Satellite observations, in situ observations, offline model experiments and indirect estimates all have

a potential to generate relevant information, but are largely inconsistent, covering different subdomains of the states, and suffer from methodological flaws. As a consequence, the pioneering work on deriving soil moisture related predictability and regional/global climate responses has been carried out using (ensembles of) modeling experiments. The following studies are particularly relevant in this respect.

The Global Soil Wetness Project (particularly phase 2, GSWP2; Dirmeyer et al., 2006) yielded a 10-year “climatology” (1986-1995) of all land surface states including soil moisture and surface fluxes based on an ensemble of offline land surface models, driven by pseudo-observed climatological forcings. Various follow-up projects to extend the period and applications of this product have taken place or are being planned. *For CMIP6 it is of utmost relevance to document the characteristics of the land surface component of the coupled models under observation-based constrained conditions, and document its main systematic biases.* A third edition of GSWP is being prepared (GSWP3; see <http://hydro.iis.u-tokyo.ac.jp/GSWP3>). Participation by a large subset of the land surface models used in the CMIP6 ensemble allows the generation of a well constrained CMIP6 climatology of land surface characteristics, and provides input to model evaluation and predictability studies. *Therefore, incorporating GSWP3 in the CMIP6 program can be seen as the LMIP of CMIP6, an analogy to AMIP or OMIP. The LMIP simulations will build upon the GSWP3 experiments and were identified, together with OMIP, as possible future DECK experiments at the recent WGCM-18 meeting. In CMIP6, these “proto-DECK” experiments are recommended for Tier1. They will allow an assessment of the representation of soil moisture and snow processes, as well as of other land surface processes (e.g. vegetation) and associated fluxes of water and energy in the CMIP6 land surface models.*

The Global Land Atmosphere Coupling Experiment (GLACE: Phases 1 and 2 on seasonal forecasting (Koster et al. 2004; 2010) and GLACE-CMIP5 on climate change projections (Seneviratne et al., 2013)) provided first assessments of the role of soil moisture for the climate system. The GLACE-1 analysis (Koster et al., 2004) pioneered the identification of regions where soil moisture has a significant impact on the local hydroclimate, based on an ensemble of idealized model simulations. At the seasonal time scale transitional wet-dry climate regions, mostly coinciding with monsoon regions, display an identifiable soil moisture-precipitation coupling. Expanding the GLACE framework at the climate time scale and for regional climate simulations in Europe, Seneviratne et al. (2006) illustrate changes in patterns of coupling strength between present and future climate conditions, showing a shift of the area of strong land-atmosphere interactions from the Mediterranean region to Central and Eastern Europe. More recently, the GLACE-CMIP5 multi-model experiment (Seneviratne et al., 2013) uses this expanded GLACE framework to investigate the role of soil moisture in modifying the regional temperature and precipitation response to a future climate forcing. *The experimental design of the GLACE-CMIP5 study, carried out with a limited CMIP5 ensemble with prescribed SSTs (AGCMs) and vegetation, is used as blueprint for the second set of proposed LS3MIP experiments, described in detail below. The new LS3MIP experiments will allow a full quantification of soil moisture-climate feedbacks in the CMIP6 models and provide reference diagnostics for the evaluation of the CMIP6 ESMs, which will be of key relevance for the application of constraints to reduce uncertainties in projections.*

In addition, LS3MIP will include an assessment of changes in land-based predictability in the CMIP6 models. These experiments build upon the GLACE2 predictability experiment (Koster et al., 2010a), in which the actual temperature and precipitation skill improvement of using observation constrained estimated soil moisture initializations is shown to be much lower than suggested by the coupling strength diagnostics. Limited quality of the initial states, limited predictability and poor representation of essential processes determining the propagation of information through the hydrological cycle in the models all play a role. *An update of the land surface related predictability in state of the art climate models will reveal essential information about the models’ ability to represent the terrestrial hydrological processes, the inherent limitations to predictability, and possible shifts in*

patterns of predictability in response to climate change (Dirmeyer et al., 2013). This will be evaluated in the third branch of LS3MIP experiments.

Both the LMIP (GSWP3) and the soil moisture-based LS3MIP experiments are key action items of the WCRP grand challenges on water availability and climate extremes, which are coordinated by the GEWEX project.

Objectives of LS3MIP

The Land Surface Snow and Soil moisture MIP (LS3MIP) will embrace a small number of multi-model experiments, encompassing simulations driven in offline mode (land-surface only), coupled to the atmosphere (driven by prescribed sea surface temperatures, SSTs), and embedded in fully coupled AOGCMs. The experiments are subdivided in two components, the first one addressing land systematic biases (“LMIP”, building upon the GSWP3 experiment) and the second one addressing land feedbacks in an integrated framework (“LFMIP”, building upon the ESMsnowMIP and GLACE-CMIP blueprints). The LS3MIP experiments address together the following objectives:

- an evaluation of the current state of land processes including surface fluxes, snow cover and soil moisture representation in CMIP6 DECK runs, revealing main *systematic biases and their dependencies* (LMIP-protoDECK)
- a *multi-model estimation* of the long-term terrestrial energy/water/carbon cycles, using the surface modules of CMIP6 models under observation constrained historical (land reanalysis) and projected future (impact assessment) conditions considering land use/land cover changes. (LMIP)
- an assessment of the role of snow and soil moisture feedbacks in the regional response to altered climate forcings, focusing on controls of climate extremes, water availability and high-latitude climate in historical and future scenario runs (addressing Arctic amplification and drought/heatwave characteristics) (LFMIP)
- an assessment of the contribution of land surface processes to the current and future *predictability* of regional temperature/precipitation patterns. (LFMIP)

These objectives respond to each of the three CMIP6 overarching questions: what are regional feedbacks and responses to climate change, what are the systematic biases in the current climate models, and what are the perspectives concerning the generation of predictions and scenarios.

Embedding of LS3MIP within WCRP and CMIP6

As illustrated in Figure 1, LS3MIP is addressing core research questions of the WCRP and is relevant for a large fraction of the WCRP activities. It is initiated by two out of four WCRP core projects (CliC and GEWEX) and directly related to three WCRP Grand Challenges (Cryosphere in a Changing Climate, Changes in Water Availability, and Climate Extremes). The LMIP experiment will provide best estimates of historical changes in snow and soil moisture on global scale, thus allowing the evaluation of changes in freshwater, agricultural drought, and streamflow extremes over continents. The LFMIP experiment is of high relevance for the assessment of key feedbacks and systematic biases of land surfaces processes in coupled mode, and is also addressing two of the main feedback loops over land: The snow-albedo-temperature feedback, involved in Arctic Amplification, and the soil moisture-temperature feedback leading to major changes in temperature extremes. Hence LS3MIP is directly addressing some of the main questions underlying the *Cryosphere in a Changing Climate* and *Changes in Water Availability* Grand Challenges, and will also provide essential insights on temperature and hydrological extremes for the *Climate Extremes* Grand Challenge. In addition, LS3MIP will also allow the exchange of data and knowledge across communities, as snow and soil moisture dynamics are often interrelated (e.g. (Hall et al. 2008) and contribute together to

hydrological variability (e.g. Koster et al. 2010b). LS3MIP will thus constitute a core element within WCRP, binding together several communities that, in fact, address a common physical object: water on land, in its liquid or solid form.

LS3MIP within WCRP Core Projects and Grand Challenges

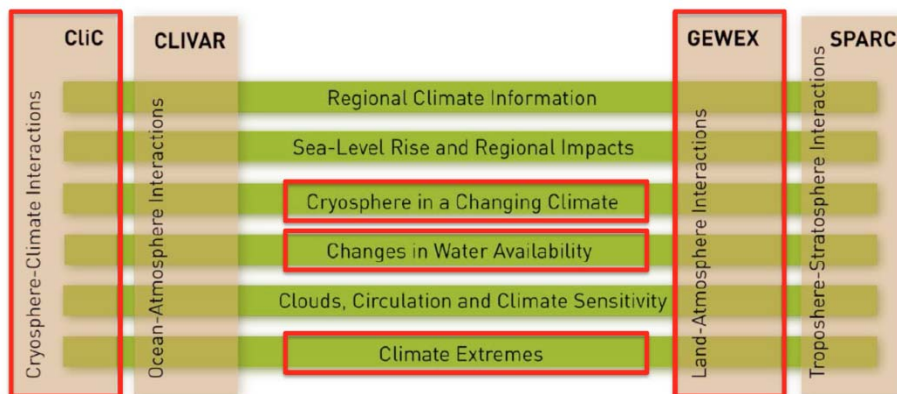


Figure 1: Relevance of LS3MIP for WCRP Core Projects and Grand Challenges

In addition, LS3MIP will provide relevant insights for other research communities within WCRP, such as estimates of freshwater inputs to the oceans (which are relevant for sea-level changes and regional impacts), the assessment of feedbacks shown to strongly modulate regional climate variability and thus relevant for regional climate information, as well as the investigation of land climate feedbacks on large-scale circulation patterns and cloud occurrence. This will thus also imply potential contributions to the other WCRP grand challenges and core projects.

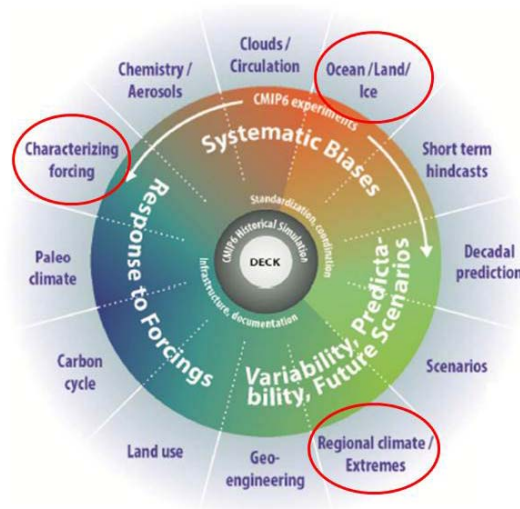


Figure 2: Embedding of LS3MIP within CMIP6

Figure 2 illustrates the embedding of LS3MIP within CMIP6. LS3MIP clearly fills a major gap, by allowing the consideration of land systematic biases and land feedbacks within the CMIP6 framework. In this context, LS3MIP can be seen as part of a larger “LandMIP” series of experiments fully addressing biases, uncertainties, feedbacks and forcings from the land surface (Figure 3), which are complementary to similar experiments for ocean or atmospheric processes. In particular, we note that while LS3MIP focuses on *systematic biases* in land surface processes (LMIP) and on *feedbacks* from the land surface processes on the climate system (LFMIP), the complementary LUMIP experiment (separate proposal) addresses the role of land surface *forcing* on the climate system. The role of vegetation and carbon stores in the climate system is a point of convergence

between LUMIP and LS3MIP. In particular, the LMIP/GSWP3 experiment will serve as land-only reference experiments for both the LS3MIP and LUMIP experiments. In addition, there will also be links to the C4MIP experiment with respect to impacts of snow and soil moisture processes (in particular droughts) on terrestrial carbon exchanges and resulting feedbacks to the climate system.

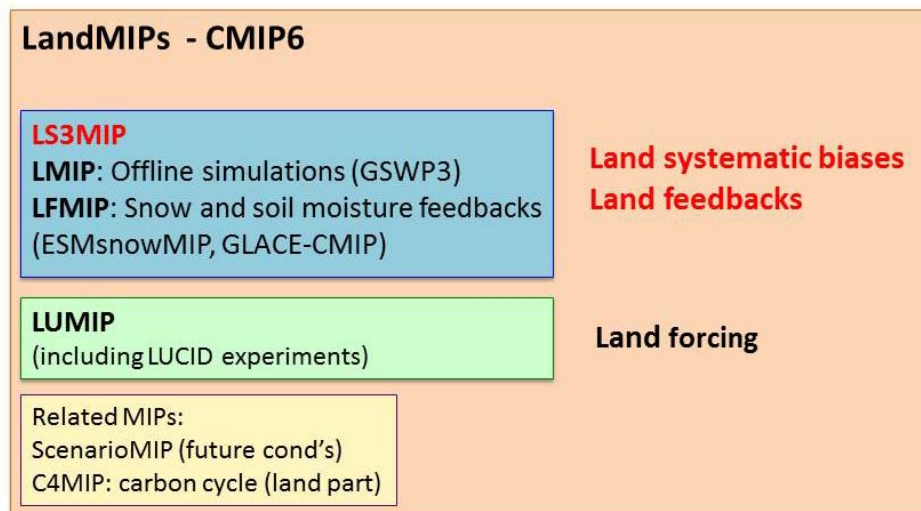


Figure 3: Overview of the embedding of LS3MIP in land-related MIPs. LS3MIP will allow the quantification of land systematic biases and feedbacks induced by snow and soil moisture processes, while LUMIP addresses land forcing on climate.

Overview of the proposed experiments

A number of complementary experiments are proposed as part of LS3MIP (see Figure 4 and Table 1):

(1) Offline land model experiment ("Land offline MIP", LMIP):

In the context of GSWP3 meteorological forcings are made available to drive land modules from climate models in an offline mode. Offline land simulations of land surface states and fluxes allow for the evaluation of trends and variability of snow, soil moisture and land surface fluxes, carbon stores and vegetation states, and climate change impacts. Ancillary data (e.g., land use/cover changes, surface parameters, CO₂ concentration) and documented protocols to spin-up and execute the experiments are currently being compiled.

(1a) Land reanalysis: LMIP-Hist

One set of reference forcing data and a standard bias correction strategy will be provided to drive each land surface model for the historical (1850-2014) simulations. 1d time series of in-situ observational forcing variables from selected reference sites are implanted into the generated forcing data for additional site level validations. Although this historical experiment is not a formal member of the DECK simulations, the WGCM recognized the importance of these offline experiments for the process of model development and benchmarking. The subset (1979-2014) of this historical run, largely analogous to AMIP, constitute Tier 1 of LMIP and is proposed to become part of the DECK in future CMIP exercises. A future implementation into the DECK is foreseen and the LMIP simulations were therefore identified as proto-DECK experiments.

(1b) Climate change impact assessment: LMIP-Fut

The future simulations (2015-2100) constitute Tier 2 of LMIP. In these simulations, the atmospheric output of at least 2 scenarios based on the ScenarioMIP (tentatively, SSP5-8.5 and SSP4-3.7) will be exploited as forcing data with a statistical bias correction method for constant and time varying

conditions of carbon cycle related factors. It focuses on climate change impact assessment (e.g., on water availability and climate extreme) and estimation of the sensitivity of land modules of CMIP6 GCMs to the projected future.

(2) Prescribed land surface states to assess the impact of snow and soil moisture feedbacks (“Land Feedback MIP”, LFMIP):

Here the GLACE-CMIP5 protocol is followed, where apart from the CMIP6 DECK experiments a set of forced experiments is carried out, where land surface states are prescribed from an a priori defined database. In contrast to the earlier experiments coupled AOGCM simulations are anticipated, where the Historical (1980-2014) and future (2015-2100) simulations will be used as reference. For the future a single scenario from the ScenarioMIP will be selected at a later stage. The land surface states that are prescribed may vary across the participating models depending on the model structure, but at least include the water reservoirs (soil moisture, snow mass), but may be extended to other prognostic quantities related to vegetation or temperature.

(2a) Core experiments: 2 experiments are considered to be “core”

- prescribed climatology derived from “present climate” conditions (e.g. 1980-2014), aiming at diagnosing the role of land-atmosphere feedback at the climate time scales
- prescribed climatology using a transient 30-yr running mean, where a comparison to the standard CMIP6 runs allows diagnosing shifts in the regions of strong land-atmosphere coupling, and shifts in potential predictability related to land surface states.

Both simulations cover the historical period and extend to 2100, based on a forcing scenario to be identified at a later stage.

Output in high temporal resolution (daily, as well as sub-daily for some fields and time slices) is planned in order to address the role of land surface-climate feedbacks (including snow and soil moisture feedbacks) on climate extremes on land. These outputs may be generated for shorter time slices only.

A single member of each of these core simulations is considered to be part of the Tier 1 simulations, but multi-member experiments are encouraged (and included in a Tier 2 set of simulations).

(2b) As (2a) for AGCM simulations

The AOGCM simulations from (2a) are duplicated with a prescribed SST configuration (AGCM), and also these simulations are included in the Tier 2 set of LS3MIP experiments.

(2c) Separate effects of soil moisture and snow, and role of additional land parameters and variables

Additional experiments in which only snow, snow albedo or soil moisture is prescribed will be conducted to assess the respective feedbacks in isolation, and have control on possible interactions between snow cover and soil moisture content. At a later stage, also vegetation parameters and variables (e.g. leaf area index) could be considered. These experiments are all part of the Tier 2 batch of LS3MIP.

(2d): As (2a) for fixed land use conditions

In conjunction with the Land Use MIP (LUMIP) a repetition of experiment (2a) under unchanging land cover and land use conditions is planned. This experiment highlights the role of soil moisture in modulating the climate response to land cover and land use. It is a tier 2 set of experiments in LS3MIP.

Apart from the above experiments, particular sensitivity experiments are proposed to isolate the role of individual processes (such as prescribed albedo to address snow-related feedbacks, or

vegetation parameters addressing carbon/water interactions). These all will be Tier 2 experiments, and be designed throughout the runtime of LS3MIP.

(3) Prescribed land surface states derived from pseudo-observations (LFMIP-predictability)

The use of experimental batch (1) (offline land models) to initialize the AOGCM experiments (batch 2) allows a set of predictability experiments in line with the GLACE2 set-up. Here historical runs from 1980 to 2014 are proposed in AOGCM mode, with a prescribed series of ‘reconstructed’ land surface states, either derived from the offline simulations or derived from various observational data sources (such as for SWE or snow albedo, using satellites, reanalysis and land surface model outputs). The predictability assessments include the evaluation of the contribution of snow cover melting and its related feedbacks to the underestimations of recent boreal polar warming by climate models.

Figure 4 and Table 1 summarizes the experimental overview, where experiments focusing on specific processes and the LUMIP configuration (2c and 2d) are not included in this inventory.

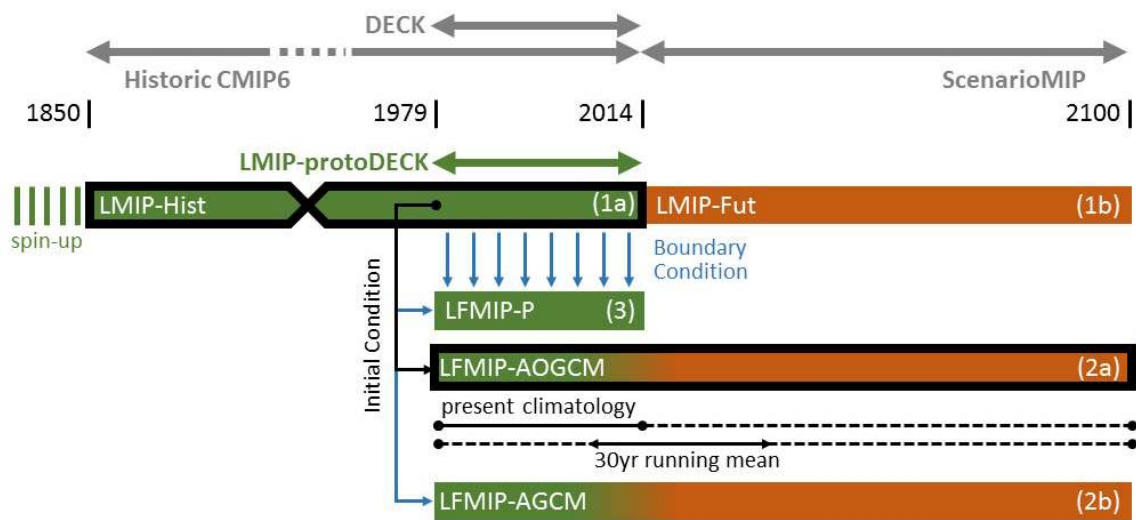


Figure 4: Schematic diagram for the experiment structure of LS3MIP (black-outline for the Tier 1 experiment).

Table 1: Summary of LS3MIP experiments. Details on separate sensitivity studies and selected scenarios have not been included.

Experiment Name	Tier	Experiment Description / Design	Configuration	Start	End	# Years Per Simulation	# Ens	# Total Years	Science Question and/or Gap Being Addressed with this	Possible Synergies with other MIPs
LMIP-Hist	1	Land only simulations	LND	1850	2014	165	2	330	Land reanalysis	LUMIP
LMIP-Fut	2	Land only simulations	LND	2015	2100	86	4	344	Climate trend analysis	ScenarioMIP
LFMIP-Hist-AOGCM	1	Prescribed land conditions 1980-2014 climate	LND-ATM-OC	1980	2100	121	1	121	diagnose land-climate feedback including ocean response	ScenarioMIP
LFMIP-Hist-AOGCM	2	Prescribed land conditions 1980-2014 climate	LND-ATM-OC	1980	2100	121	4	484	diagnose land-climate feedback including ocean response	ScenarioMIP
LFMIP-Hist-AGCM	2	Prescribed land conditions 1980-2014 climate; SSTs prescribed	LND-ATM	1980	2100	121	5	605	diagnose land-climate feedback over land	ScenarioMIP
LFMIP-AOGCM	1	Prescribed land conditions 30yr running mean	LND-ATM-OC	1980	2100	121	1	121	diagnose land-climate feedback including ocean response	ScenarioMIP
LFMIP-AOGCM	2	Prescribed land conditions 30yr running mean	LND-ATM-OC	1980	2100	121	4	484	diagnose land-climate feedback including ocean response	ScenarioMIP
LFMIP-AGCM	2	Prescribed land conditions 30yr running mean; SSTs prescribed	LND-ATM	1980	2100	121	5	605	diagnose land-climate feedback over land	ScenarioMIP
LFMIP-P	2	Initialized pseudo-observations land	LND-ATM-OC	1980	2014	35	10	350	land-related seasonal predictability	Historic runs

Overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments

LS3MIP brings together climate modelers, snow and soil moisture model specialists and experts in local and remotely sensed data of soil moisture and snow properties, mass and extent. This diversity

is reflected in the composition of the steering group of LS3MIP and ensures that the experiment setups, model evaluations and analyses/interpretations of the results are pertinent.

Analyses for snow

Concerning the analysis of climate model runs, large-scale datasets of snow mass (SWE) and snow cover extent (SCE) are particularly relevant for the analysis of the historical simulations in the LS3MIP framework (i.e. the AMIP runs and the historical coupled run). These large-scale, high-quality snow data are available through close links to the Satellite Snow Product Intercomparison and Evaluation Experiment (SnowPEX, <http://calvalportal.ceos.org/projects/snowpex>), via the composition of the steering group. The quality of the representation of these fundamental snow-related variables in the historical simulations (coupled and AMIP) will be evaluated against these datasets. Output from the historical simulations are required to update analyses of the agreement between observations and historical simulations, and determine new projections of the variability and trends in terrestrial snow cover extent and mass (this was examined with CMIP3 and CMIP5 output in studies such as Brown and Mote (2009); Derksen and Brown (2012); Brutel-Vuilmet et al. 2013). These analyses, besides their genuine interest, can also provide clues to the interpretation of general model deficiencies in the representation of boreal and polar climates. The representation of albedo over snow-covered areas in DECK simulations will be analyzed. Multiple satellite-derived datasets are available for the evaluation of simulated albedo, including 16-day MODIS data (2001-present; <http://modis-atmos.gsfc.nasa.gov/ALBEDO/>) and the recently updated twice-daily APP-x product (1982-2011; <http://stratus.ssec.wisc.edu/products/appx/appx.html>). Specific attention will be paid to the role of the models' representation of snow cover fraction in forested and mountainous areas. The DECK simulations will be used to update analyses of observed and simulated snow-albedo feedback, an important diagnostic in determining climate sensitivity to snow cover (Qu and Hall, 2014; Fletcher et al. 2012).

The LS3MIP will be analyzed in concert with the control runs to quantify various climatic effects of snow, including very accurate estimates of snow albedo feedback. For example, the prescribed albedo experiments (simulation set 2c) do not allow the optical properties of vegetation to change in snow-covered areas as the climate warms. However, the prescribed SWE experiments do allow for this effect. The surface albedo change in the Prescribed SWE experiments can be compared to the overall albedo change in the control experiments to quantify the degree to which the surface albedo changes in snow-covered areas are due to vegetation changes, rather than snow changes. These estimates can be used to confirm that snow albedo feedback effects diagnosed from the Prescribed Albedo experiments are not misleading due to vegetation effects. Similarly, the Prescribed albedo experiments contain changes in soil moisture and hydrology due to melting snow. These can be compared to the control experiments to ascertain the degree to which snowmelt influences hydrology independently of its substantial influence on surface absorbed solar radiation. In this way, one can assess the degree to which the Prescribed SWE experiments produce snow effects unrelated to snow albedo feedback.

The geographical focus of the first stage of this project is on the continental snow cover of both hemispheres, both in ice-free areas (Northern Eurasia and North America) and on the large ice sheets (Greenland and Antarctica). In later stages of LS3MIP, the effect of snow on sea ice will be analyzed. Major scientific questions concerning snow on sea ice are related to strong recent trends of Arctic sea ice decline and the potential amplifying effect of earlier snow melt. These questions can be tackled by AGCM runs with a dynamic atmospheric nudging to eliminate biases related to misrepresentation of NH circulation trends (AO, NAO). Some of the modeling groups that have declared interest in participating in the snow-related part of LS3MIP (ESM-SnowMIP) are currently carrying out "proof of concept" simulations using prescribed snow mass (SWE) with an AMIP-type DECK control run; note that similar experiments have already been carried out (e.g., Lawrence and

Slater, 2009; Alexander et al., 2011), demonstrating the feasibility and scientific interest of the proposed experiments.

Analyses for soil moisture

The analyses will focus on 1) systematic biases in offline land simulations (LMIP/GSWP3 simulations) and on 2) the role of soil moisture – climate feedbacks for past and projected changes in land climate conditions.

In the case of systematic land biases, the LMIP/GSWP3 simulations will be evaluated with observations available over the historical time period (e.g. for runoff, storage anomalies, vegetation activity) to assess their degree of realism and typical biases compared to measurements. Uncertainties of current land surface models in the representation of historical variations in land water availability/droughts (due to model parameterizations and/or atmospheric forcings, Sheffield et al. 2012, Trenberth et al. 2013, Greve et al. 2014) as well as systematic biases in water, energy and carbon exchanges between the land and the atmosphere (e.g. Mueller and Seneviratne 2014) will be assessed. These assessments will be used for the evaluation of the offline simulations for future land conditions as well as the coupled experiments.

In the case of soil moisture-climate feedbacks, the focus will be set on the following topics:

1. The quantification of the impact of soil moisture variability for climate variability (trends, decadal variability, interannual anomalies, extremes) on land and its interaction with large-scale drivers (large-scale modes of variability, ocean-climate interactions)
2. The attribution of model disagreement in land temperature, precipitation, runoff vegetation activity, carbon sink to the representation of soil moisture, related processes (plant transpiration and photosynthesis) and feedbacks to the atmosphere
3. The derivation of emergent constraints to reduce uncertainties in projections of mean climate and extremes (hot temperatures, droughts, floods) using observations characterizing the identified soil moisture-climate feedbacks
4. The regional assessment of the relationship between bias in modelled soil moisture/land surface representation and climate response
5. A robust estimate on the geographical patterns of “hot spots” of changes in soil moisture dynamics and their impact on occurrence of droughts, heat waves, irrigation limitations or river discharge anomalies.
6. The assessment of the role of soil moisture for subseasonal to seasonal predictability over land in both present and future climate

Proposed timing

The proposed experiments are continuous model runs duplicating the Historical and ScenarioMIP simulations. AMIP mode runs are foreseen as a Tier 2 set of experiments. The experimental setup requires a reasonable amount of additional coding for reading and prescribing land surface characteristics, while many groups already participated in one of the earlier experiments. It is anticipated that the LS3MIP simulations are carried out after the first set of core CMIP6 experiments (i.e. after 2016 for historical runs and after 2017 or 2018 for ScenarioMIP runs). Stand-alone simulations with the ESMs' land surface modules in uncoupled mode are currently planned in the context of GSWP3, and will initiate early 2015. The evaluation of land surface processes in CMIP6 Historical Simulation experiments will start as soon as historical runs are available.

A 6 month preliminary period will be dedicated to a wide consultation of the climate modeling community aiming at finalizing the detailed experiments design.

Group commitment

The following ESM groups and contact persons (represented in Scientific Steering Committee) will certainly participate in LS3MIP:

- MPI (Stefan Hagemann)
- EC-Earth (Andrea Alessandri)
- BCC (?)
- CESM (David Lawrence)
- CMCC (?)
- CNRM (Hervé Douville)
- GISS (?)
- IPSL (Gerhard Krinner)
- MIROC (Taikan Oki)
- CCCma (Greg Flato)

A list of potentially interested groups includes:

- ACCESS (Andy Pitman?)
- GFDL (Kirsten Findell?)
- UKESM (Martin Best?)
- MRI (?)

References

- Alexander, M.A., R. Thomas, C. Deser, and D. Lawrence, 2009: The atmospheric response to projected terrestrial snow changes in the late 21st century. *J. Clim.*, **23**, 6430–6437, doi: 10.1175/2010JCLI3899.1.
- Betts, A.K., 2004. Understanding hydrometeorology using global models. *B. Am. Meteorol. Soc.* 85 (11), 1673–1688.
- Brown, R. D., and P. W. Mote, 2009: The Response of Northern Hemisphere Snow Cover to a Changing Climate*. *J. Clim.*, **22**, 2124–2145, doi:10.1175/2008JCLI2665.1.
- Brutel-Vuilmet, C., M. Ménégoz, and G. Krinner, 2013: An analysis of present and future seasonal Northern Hemisphere land snow cover simulated by CMIP5 coupled climate models. *The Cryosphere*, **7**, 67–80, doi:10.5194/tc-7-67-2013.
- Derksen, C., and R. Brown, 2012: Spring snow cover extent reductions in the 2008–2012 period exceeding climate model projections. *Geophys. Res. Lett.*, **39**, L19504, doi:10.1029/2012GL053387.
- Dirmeyer, P. A., X. Gao, M. Zhao, Z. Guo, T. Oki, and N. Hanasaki, 2006: GSWP-2: Multimodel Analysis and Implications for Our Perception of the Land Surface. *Bull. Am. Meteorol. Soc.*, **87**, 1381–1397, doi:10.1175/BAMS-87-10-1381.
- , S. Kumar, M. J. Fennessy, E. L. Altshuler, T. DelSole, Z. Guo, B. A. Cash, and D. Straus, 2013: Model Estimates of Land-Driven Predictability in a Changing Climate from CCSM4. *J. Clim.*, **26**, 8495–8512, doi:10.1175/JCLI-D-13-00029.1.
- Douville, H., A. Ribes, B. Decharme, R. Alkama, and J. Sheffield, 2012: Anthropogenic influence on multidecadal changes in reconstructed global evapotranspiration. *Nature Climate Change*, DOI: 10.1038/NCLIMATE1632.
- Dumont, M., and Coauthors, 2014: Contribution of light-absorbing impurities in snow to Greenland's darkening since 2009. *Nat. Geosci.*, **7**, 509–512.
- Ek, M.B., Holtslag, A.A.M., 2004. Influence of soil moisture on boundary layer cloud development. *J. Hydrometeorol.* 5 (1), 86–99.
- Essery, R., and Coauthors, 2009: SNOWMIP2: An Evaluation of Forest Snow Process Simulations. *Bull. Am. Meteorol. Soc.*, **90**, 1120–1135, doi:10.1175/2009BAMS2629.1.

- Fletcher, C. G., S. C. Hardiman, P. J. Kushner, and J. Cohen, 2009: The Dynamical Response to Snow Cover Perturbations in a Large Ensemble of Atmospheric GCM Integrations. *J. Clim.*, **22**, 1208–1222, doi:10.1175/2008JCLI2505.1.
- , H. Zhao, P. J. Kushner, and R. Fernandes, 2012: Using models and satellite observations to evaluate the strength of snow albedo feedback. *J. Geophys. Res. Atmospheres*, **117**, D11117, doi:10.1029/2012JD017724.
- Greve, P., B. Orłowsky, B. Mueller, J. Sheffield, M. Reichstein, and S.I. Seneviratne, 2014: Global assessment of trends in wetting and drying over land. *Nature Geoscience*. Published online, doi: 10.1038/NGEO2247.
- Hall, A., X. Qu, and J.D. Neeling, 2008: Improving predictions of summer climate change in the United States. *Geophys. Res. Lett.*, **35**, L01702, doi:10.1029/2007GL032012.
- Holland, M. M., and C. M. Bitz, 2003: Polar amplification of climate change in coupled models. *Clim. Dyn.*, **21**, 221–232, doi:10.1007/s00382-003-0332-6.
- van den Hurk, B., M. Best, P. Dirmeyer, A. Pitman, J. Polcher, and J. Santanello, 2011: Acceleration of Land Surface Model Development over a Decade of Glass. *Bull. Am. Meteorol. Soc.*, **92**, 1593–1600, doi:10.1175/BAMS-D-11-00007.1.
- Koster, R. D., and Coauthors, 2004: Regions of Strong Coupling Between Soil Moisture and Precipitation. *Science*, **305**, 1138–1140, doi:10.1126/science.1100217.
- , and Coauthors, 2010a: Contribution of land surface initialization to subseasonal forecast skill: First results from a multi-model experiment. *Geophys. Res. Lett.*, **37**, L02402, doi:10.1029/2009GL041677.
- , S.P.P. Mahanama, B. Livneh, D.P. Lettenmaier, and R.H. Reichle, 2010b: Skill in streamflow forecasts derived from large-scale estimates of soil moisture and snow. *Nature Geoscience*, **3**, 613–616.
- Koven, C. D., W. J. Riley, and A. Stern, 2012: Analysis of Permafrost Thermal Dynamics and Response to Climate Change in the CMIP5 Earth System Models. *J. Clim.*, **26**, 1877–1900, doi:10.1175/JCLI-D-12-00228.1.
- Lawrence, D. M. and A.G. Slater, 2009: The contribution of snow trends to future ground climate. *Clim. Dyn.*, **34**, 969–981, doi:10.1007/s00382-009-0537-4.
- Miralles, D.G., A.J. Teuling, C.C. van Heerwaarden, and J. Vila-Guerau de Arellano, 2014: Mega-heatwave temperatures due to combined desiccation and atmospheric heat accumulation. *Nature Geoscience*, published online, DOI: 10.1038/NGEO2141.
- Mueller, B., and S.I. Seneviratne, 2014: Systematic land climate and evapotranspiration biases in CMIP5 simulations. *Geophys. Res. Lett.*, **41** (1-7), doi:10.1002/2013GL058055
- Qu, X., and A. Hall, 2007: What Controls the Strength of Snow-Albedo Feedback? *J. Clim.*, **20**, 3971–3981, doi:10.1175/JCLI4186.1.
- , and —, 2014: On the persistent spread in snow-albedo feedback. *Clim. Dyn.*, **42**, 69–81, doi:10.1007/s00382-013-1774-0.
- Quesada, B., R. Vautard, P. Yiou, M. Hirschi, and S.I. Seneviratne, 2012: Asymmetric European summer heat predictability from wet and dry Southern winter/springs. *Nature Climate Change*, **2**, 736–741, doi:10.1038/nclimate1536.
- Santanello Jr., J.A., Peters-Lidard, C.D., Kumar, S.V., Alonge, C., Tao, W.K., 2009. A modeling and observational framework for diagnosing local land–atmosphere coupling on diurnal time scales. *J. Hydrometeorol.* **10** (3), 577–599.
- Seneviratne, S. I., D. Luthi, M. Litschi, and C. Schar, 2006: Land-atmosphere coupling and climate change in Europe. *Nature*, **443**, 205–209, doi:10.1038/nature05095.
- , T. Corti, E. L. Davin, M. Hirschi, E. B. Jaeger, I. Lehner, B. Orłowsky, and A. J. Teuling, 2010: Investigating soil moisture–climate interactions in a changing climate: A review. *Earth-Sci. Rev.*, **99**, 125–161, doi:10.1016/j.earscirev.2010.02.004.

- , and Coauthors, 2013: Impact of soil moisture-climate feedbacks on CMIP5 projections: First results from the GLACE-CMIP5 experiment. *Geophys. Res. Lett.*, **40**, 2013GL057153, doi:10.1002/grl.50956.
- Sheffield, J., E.F. Wood, and M.L. Roderick, 2012: Little change in global drought over the past 60 years. *Nature*, **491**, 435-438.
- Trenberth, K.E., A. Dai, G. van der Schrier, P.D. Jones, J. Barichivich, K.R. Briffa, and J. Sheffield, 2013: Global warming and changes in drought. *Nature Climate Change*, **4**, 17-22.
- Weedon, G. P., and Coauthors, 2011: Creation of the WATCH Forcing Data and Its Use to Assess Global and Regional Reference Crop Evaporation over Land during the Twentieth Century. *J. Hydrometeorol.*, **12**, 823–848, doi:10.1175/2011JHM1369.1.

Land-Use Model Intercomparison Project (LUMIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

➤ **Name of MIP***

Land-Use Model Intercomparison Project (LUMIP)

➤ **Co-chairs of MIP (including email-addresses)***

George Hurtt (george.hurtt@umd.edu)

David Lawrence (dlawren@ucar.edu)

➤ **Members of the Scientific Steering Committee***

Victor Brovkin (victor.brovkin@mpimet.mpg.de)

Kate Calvin (katherine.calvin@pnnl.gov)

Andrew Jones (adjones@lbl.gov)

Chris Jones (chris.d.jones@metoffice.gov.uk)

Peter Lawrence (lawrence@ucar.edu)

Nathalie de Noblet-Ducoudré (nathalie.de-noblet@lsce.ipsl.fr)

Julia Pongratz (julia.pongratz@mpimet.mpg.de)

Sonia Seneviratne (sonia.seneviratne@ethz.ch)

Elena Shevliakova (elena@princeton.edu)

➤ **Link to website (if available)***

<https://www2.cgd.ucar.edu/research/mips/lumip>

➤ **Goal of the MIP and a brief overview***

Human land-use activities have resulted in large changes to the biogeochemical and biophysical properties of the Earth surface, with resulting implications for climate. In the future, land-use activities are likely to expand and/or intensify further to meet growing demands for food, fiber, and energy. CMIP5 achieved a qualitative scientific advance in studying the effects of land-use on climate, for the first time explicitly accounting for the effects of global gridded land-use changes (past-future) in coupled carbon-climate model projections. Enabling this advance, the first consistent gridded land-use dataset (past-future) was developed, linking historical land-use data, to future projections from Integrated Assessment Models, in a standard format required by climate models. Results indicate that the effects of land-use on climate, while uncertain, are sufficiently large and complex to warrant an expanded activity focused on land-use for CMIP6. Land-use change is an essential forcing of the Earth System, and as such LUMIP is directly relevant and necessary for CMIP6 Question 1: “How does the Earth System respond to forcing?” LUMIP will also play a strong role in addressing the WCRP Grand Challenges, particularly with respect to the “AIMES theme for collaboration: biospheric forcings and feedbacks”. Due to the broad range of effects of land-use change and the major activities proposed, LUMIP is also of cross-cutting relevance to

CMIP6 science questions 2 and 3, and to many of the WCRP Grand Challenges including Climate Extremes, Regional Climate Information, and Water Availability.

The goal of LUMIP is to take the next steps, and enable, coordinate, and ultimately address the most important science questions related to the effects of land-use on climate. The primary science questions of LUMIP are:

- What are the effects of land use and land-use change on climate and biogeochemical cycling (past-future)?
- Are there regional land management strategies with promise to help mitigate and/or adapt to climate change?
- What are the effects of climate change on land-use and land-use change? *

In addressing these questions, LUMIP will also address a range of more detailed science questions to get at process level attribution, uncertainty, data requirements, and other related issues in more depth and sophistication for the community than possible to date. Of particular focus will be the separation and quantification of the effects on climate from fossil fuel emissions and land-use change, biogeochemical from biophysical effects, and land cover change from land management effects.

Three major sets of science activities are envisioned. First, a set of metrics and diagnostic protocol will be developed to quantify model performance, and related sensitivities, with respect to land use. As part of this activity, benchmarking data products will be identified to help constrain models. These metrics will be incorporated into the International Land Model Benchmarking (ILAMB) system. This benchmarking/metrics emphasis in LUMIP dovetails with expanding emphasis in CMIP on metrics.

Second, data standardization efforts will build off the lessons learned and protocols in CMIP5, and work with new historical data, present data, IAMS, and ESMs to produce an enhanced standardized land-use data for CMIP6 model experiments passing the maximum amount of common information between these relevant domains. New output data standardization will also enrich and improve analysis of model experiment results. Particular emphasis is on promoting the archival of subgrid land information in CMIP6. In most land models, physical, ecological, and biogeochemical land state and surface flux variables are calculated separately for several different land surface type or land management ‘tiles’ (e.g., natural and secondary vegetation, crops, pasture, urban, lake, glacier). Frequently, including in the CMIP5 archive, the tile-specific quantities are averaged and only grid-cell mean values are reported. Consequently, a large amount of valuable information is lost with respect to how each surface type responds to climate change and/or direct anthropogenic modifications. LUMIP is developing a proposal outlining the need and protocol for archival for selected key variables on multiple land tiles (see Appendix A for draft proposal).

Third, an efficient model experiment design including both idealized and scenario-based cases has been developed to isolate and quantify land-use effects. These experiments, described in greater detail below, include both idealized and realistic scenario simulations with and without transient land use. The experimental protocol enables integrated analysis of coupled and offline land models (forced with observed meteorology) which will support understanding and assessment of the forced response and climate feedbacks associated with land-use and the relationship of these responses to land and atmosphere model biases.

LUMIP priorities and model experiments have been developed in close consultation with several existing model intercomparison activities and research programs that focus on the role of land use in climate including LUCID, GSWP3, LUC4C, TRENDY, and AgMIP. In addition, discussions have

* Note that experiments to address this question are not included in this proposal because our understanding is that very few Earth System Models have the capability to address this question yet. We maintain this question within LUMIP because it is a high priority land use change science question that LUMIP will promote through individual model efforts until enough models have the capability to do two-way climate-land use interactions.

begun and are ongoing with other proposed CMIP MIPs to ensure that our proposed experiments are complementary and not duplicative. These proposed MIPs include ScenarioMIP, AerChemMIP, C4MIP, LS3MIP, DAMIP, and RFMIP.

➤ **An overview of the proposed experiments***

LUMIP proposes a two phase, tiered, model experiment plan. Phase one, which can start soon, will feature idealized coupled and land-only model experiments designed to improve process understanding and assess how models represent the impact of changes in land use on climate, as well as quantify model sensitivity to potential land cover and land use changes. Phase two will be based on historical land use and realistic scenarios identified by ScenarioMIP and is designed to isolate the role of historical and projected future land-use changes on climate. As there are more possible experiments than are achievable with available resources by all groups, experiments are tiered in order of importance.

Details of the model experiments are included below. The total request includes (all at standard resolution):

- Tier 1: 485 years GCM/ESM; 165 years LND-only
- Tier 2: 380 years GCM/ESM; 1650 years LND-only
- Tier 3: 285 years GCM/ESM; 120 years LND-only

Overview of Phase 1 experiments

Phase 1 consists of two sets of experiments (see Table 1). The first set are idealized deforestation experiments that will enable analysis of impact of biogeophysical and biogeochemical response to land cover change on climate in a controlled and consistent set of simulations. The idealized 1% (or 2%) deforestation experiment is new to the land use change modeling community and is designed to be somewhat analogous/complementary to the 1% CO₂ simulations in the DECK (note that it is not exactly analogous to 1%CO₂ simulations; 1% of initial forest area is removed each year rather than 1% of remaining forest area, yielding a linear decrease in forest area). This idealized deforestation experiment has the advantage that it will be much easier to ensure conformity across models in terms of the land cover change (differences in the representation of realistic land cover changes across different models is a problem that has plagued prior land cover change model intercomparison projects, e.g. LUCID). Two modeling centers are planning test 1%(2%) deforestation simulations in Fall 2014. The regional deforestation simulations are planned for LUCID/LUC4C, who will determine the precise experimental protocol.

The second set of Phase 1 experiments are a series of offline land-only simulations, which will build on the LMIP simulation proposed in LS3MIP. This series of experiments is designed to assess how the specification of land cover change and increasingly comprehensive treatment of land management affects the carbon, water, and energy cycle response to land use change. Only a limited number of models will be able to perform all the experiments, but the experimental design will allow for multiple levels of participation, according to each model’s capabilities. This set of experiments utilizes cutting edge model developments anticipated across several contacted modeling centers and will contribute to the setting of priorities for land use for future CMIPs. Test experiments are planned for late 2014 and early 2015 to finalize the experimental design.

Table 1: Phase 1 experiments.

Process understanding	Idealized experiments designed to assess biogeophysical role of land cover change on climate	
CPL_1%DF	Idealized 1% or 2% per year deforestation, once	1850-????

	global deforest, continue run for 30 years (Tier 1)	
LND_DF, ATM_DF, CPL_DF	Land, atm, cpl simulations with some set of tropical, boreal, or temperate deforestation (defined by LUC4C/LUCID) (Tier 3)	1980-2010
Land cover versus land management change (Tier 2)	Assess relative impact of land cover and incrementally more comprehensive land management change on fluxes of water, energy, and carbon; forced with historical observed climate and projected climate anomalies (1700 to 2014 or 2100?)	
LND_LULCC_AM	All LULCC and All Management (AM) features for each particular model turned on; 1700 start; transient CO ₂ , N-dep, aerosol dep, etc.; This run is same as LMIP-Hist (LS3MIP) if GCM runs include all management capabilities	
LND_LULCC1850	LND_LULCC_AM with land use change starting at 1850 (testing impact of pre-1850 land use)	
LND_noLULCC	LND_no land cover change (Same as Tier 1, LND_noLULCC_hist)	
LND_grasscrop	LCC with 'grassland' crop/pasture; no land management	
LND_gross_vs_net	LND_grasscrop except with net transitions instead of gross	
LND_fire	LND_grasscrop with human fire management	
LND_woodharv	LND_grasscrop or LND_fire with wood harvest	
LND_pasture	LND_grasscrop but with grazing on pastureland	
LND_crop	LND_grasscrop but with crop area utilizing prognostic crop model	
LND_crop-irrig	LND_crop with realistic transient irrigated area	
LND_crop-irrig-fert	LND_crop-irrig with realistic transient fertilization	

* It is still being discussed whether additive or subtractive scheme is preferred for these land only offline experiments.

Overview of Phase 2 experiments

The Phase 2 experiments build off of the CMIP6 Historical and historical LMIP simulations as well as the ScenarioMIP simulations. They will include land-only and coupled historical and future simulations with land use held constant or modified to an alternative land use scenario (Table 2). These simulations will be used to assess the role of land use on climate from the perspective of both the biogeophysical and biogeochemical impacts and will be of interest to the Detection and Attribution MIP. For the projection period, LUMIP plans to include an additional simulation off of both a high and a low radiative forcing scenario with land use from a different RCP-SSP configuration but with all other forcings remaining the same as in the original ScenarioMIP simulation. **The precise scenarios are still to be determined in consultation with ScenarioMIP, C4MIP, and AerChemMIP. LUMIP, ScenarioMIP, C4MIP, AerChemMIP are in ongoing discussions about which set of scenarios is the most mutually beneficial for all groups.** Ideally, the 'alternative' land use scenarios for the high and low radiative forcing cases will be selected to allow assessment of a predominantly greater deforestation or afforestation pathway relative to the original land use projection. Note that these simulations should be considered sensitivity simulations since they will include a set of forcings that are inconsistent with each other (e.g., land use from

SSP3-RCP7 in a simulation that in all other respects is equivalent to SSP3-RCP3.7). See figure 1 for further details of the proposed design.

Table 2: Phase 2 experiments.

Land use change impact on land to atmosphere fluxes of water, energy, carbon (Tier 1)		
LND_noLULCC_hist	Same as LMIP-Hist (LS3MIP) except with land use and land cover held constant at 1850, no human impact	1850-2014
Land use change impact on past and future climate (Tier 1)		
CPL_noLULCC_hist	Same as historical CMIP6 except with land cover/use held constant at 1850, concentration (for DAMIP) and emission driven (Tier 3), no human impact	1850-2014
CPL_landpolicy_future	Additional land use change scenarios with strongly different land use to the control; keep all emissions the same as control scenario, only change land use; emissions driven runs if possible (based on esmssp7 / esmssp2.6; C4MIP)	2015-2100 (3 ens)
CPL_noLULCC_future	Future simulation with same RF scenario with LULCC held at 2014 levels; emissions driven runs if possible (Tier 2) (based on esmssp7 / esmssp2.6; C4MIP)	2015-2100

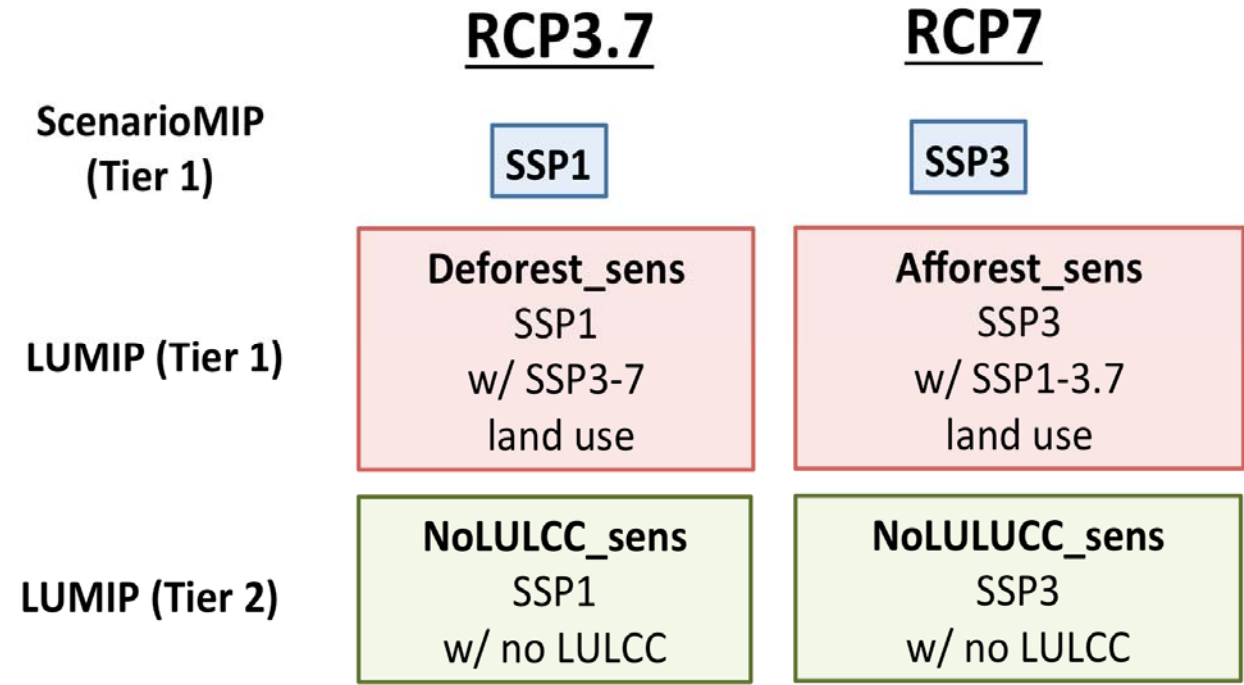


Figure 1: Example set of realistic and sensitivity studies designed to assess potential impact of strongly different land use trajectories on climate outcome.

➤ An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*

The goal is to establish a useful set of model diagnostics that enable a systematic assessment of land use-climate feedbacks and improved attribution of the roles of both land and atmosphere in terms of generating these feedbacks. The need for more systematic assessment of the terrestrial and atmospheric response to land cover change is one of the major conclusions of the LUCID study. Boisier et al. (2012) and de Noblet-Ducoudré et al. (2012) argue that the different land use-climate relationships displayed across the LUCID models highlights the need to improve diagnostics for land surface model evaluation. These analyses need to assess how land surface models respond to a land-cover perturbation in uncoupled (off-line) simulations as well as coupling between land and atmosphere components. One axis of analysis that has previously not been investigated in great detail is how a particular model's regional land-atmosphere coupling strength signature affects how the model simulates the impact of land use change on climate. Here, LUMIP will interface with LS3MIP to investigate the cross-relationship between land-atmosphere coupling strength and land-use change impacts on weather and climate.

In addition, LUMIP will promote the development of biogeophysical and biogeochemical metrics of land use change, based on observations, that will help constrain model dynamics and dovetails with expanding emphasis in CMIP on metrics. Any useful metrics will be integrated into the International Land Model Benchmarking (ILAMB) package that is currently under development. The availability of both land-only and coupled historic simulations enables a much more systematic assessment of the roles of land and atmosphere in the simulated response to land use change.

LUMIP also proposes to develop a set of metrics that quantify a model response to land use across a range of spatial scales and temporal scales that can then be used to quantitatively compare model response across different models, regions, and land management scenarios. For a given variable, say surface air temperature, the diagnostic calculations will be completed for a pair of simulations (offline or coupled) with and without land use change. Across a range of spatial scales, spanning from a single grid cell up to regional (5° by 5° and 10° by 10°) to continental to global, seasonal mean differences between control and land use change simulations will be examined. Differences will be expressed both in terms of seasonal mean differences (and their statistical significance based on student-t tests) and in terms of signal to noise (where 'noise' refers to the natural interannual climate variability simulated in the model). Effects on extremes (e.g. Davin et al. 2014) will receive particular attention.

Analysis could focus on critical regions, such as the intensive agricultural region in the central United States and the deforestation region in the Amazon, telescoping out from point to continental scale for each region. Five primary variables will be considered (net radiation, evapotranspiration, temperature, precipitation, and land carbon stocks) that together characterize the biogeophysical and biogeochemical impacts of land use on climate. The first two variables, net radiation and evapotranspiration (ET) define the surface biogeophysical response to land use change and will be evaluated in both offline and coupled model contexts. The temperature and precipitation response to biogeophysical changes in net radiation and ET will be evaluated in land-atmosphere simulations only. Land carbon stocks can be evaluated in offline and coupled simulations.

There are several axes of analysis that can be performed within this framework that are relevant to assessing land use-climate effects relative to natural variability and greenhouse gas-induced climate change. For instance, by varying the number of years and/or the number of ensemble members included in our analysis, one can establish over what time/spatial scale a land use change signal can be detected. One can also investigate the relative difficulty in isolating a land use-climate signal in transient climate simulations with anthropogenic greenhouse gas forcing versus, for example, timeslice atmosphere-land simulations.

➤ **Proposed timing***

The plans for LUMIP have been developed through conference calls and especially during a series of meetings during the summer of 2014.

2013 August 5-9: Initial concept, Aspen

2013 October 3: Presentation of Initial concept, WGCM Meeting
 2014 Spring: Workshop 1, GLP Meeting
 2014 July 17-18: GEWEX – Biogeophysics
 2014 July 21-22: Hamburg – Biogeochemistry
 2014 July 28-Aug 1: EMF Snowmass Meeting
 2014 August 4-8: AGCI Aspen Joint-MIP Workshop
 2014 September 1: Begin testing of idealized model experiments
 2014 September 15: Initial proposal due to CMIP6 Panel
 2014 October 8-10: Presentation of revised proposal, WGCM Meeting
 2015 January: New prototype land use data/data format released to modeling groups
 2015 Final proposal due to CMIP6 Panel
 2015 GMD paper documenting detailed experimental design
 2015 Model I/O and testing with new prototype land use data
 2015 September: Initiate multi-model idealized Phase 1 experiments
 2016 January: Final land use data made available* (*pending final scenario selection)
 2016 March: Phase 1 experimental delivered to ESGF
 2016 March-September: Phase 1 experiments analysis and papers
 Starting mid-2016: Phase 2 GCM/ESM realistic experiments, contingent on ScenarioMIP schedule
 2018-2019: Model analysis and synthesis

➤ Selected Key References *

Note that this list of references is representative only. Many additional references on land use and land use change impact on climate are available.

- Boisier, J. P., de Noblet-Ducoudré, N., & Ciais, P. (2013). Inferring past land use-induced changes in surface albedo from satellite observations: a useful tool to evaluate model simulations. *Biogeosciences*, 10, 1501-1516.
- Boysen, L. R., Brovkin, V., Arora, V. K., Cadule, P., de Noblet-Ducoudré, N., Kato, E., ... & Gayler, V. (2014). Global and regional effects of land-use change on climate in 21st century simulations with interactive carbon cycle. *Earth System Dynamics Discussion*, 5, 443-472.
- Brovkin, V., Boysen, L., Arora, V. K., Boisier, J. P., Cadule, P., Chini, L., ... & Weiss, M. (2013). Effect of anthropogenic land-use and land-cover changes on climate and land carbon storage in CMIP5 projections for the twenty-first century. *Journal of Climate*, 26(18), 6859-6881.
- Davin, E.L., S.I. Seneviratne, P. Ciais, A. Oliso, and T. Wang, 2014: Preferential cooling of hot extremes from cropland albedo management. *Proc. Natl Acad. Sci.*. Published ahead of print June 23, 2014, doi:10.1073/pnas.1317323111.
- Hurtt, G. C., Chini, L. P., Frohking, S., Betts, R. A., Feddema, J., Fischer, G., ... & Wang, Y. P. (2011). Harmonization of land-use scenarios for the period 1500–2100: 600 years of global gridded annual land-use transitions, wood harvest, and resulting secondary lands. *Climatic Change*, 109(1-2), 117-161.
- Kumar, S., Dirmeyer, P. A., Merwade, V., DelSole, T., Adams, J. M., & Niyogi, D. (2013). Land use/cover change impacts in CMIP5 climate simulations: A new methodology and 21st century challenges. *Journal of Geophysical Research: Atmospheres*, 118(12), 6337-6353.
- Lawrence, P. J., Feddema, J. J., Bonan, G. B., Meehl, G. A., O'Neill, B. C., Oleson, K. W., ... & Thornton, P. E. (2012). Simulating the biogeochemical and biogeophysical impacts of transient land cover change and wood harvest in the Community Climate System Model (CCSM4) from 1850 to 2100. *Journal of Climate*, 25(9), 3071-3095.
- de Noblet-Ducoudré, N., Boisier, J. P., Pitman, A., Bonan, G. B., Brovkin, V., Cruz, F., ... & Voldoire, A. (2012). Determining robust impacts of land-use-induced land cover changes on surface climate over North America and Eurasia: results from the first set of LUCID experiments. *Journal of Climate*, 25(9), 3261-3281.
- Pitman, A. J., de Noblet-Ducoudré, N., Cruz, F. T., Davin, E. L., Bonan, G. B., Brovkin, V., ... & Voldoire, A. (2009). Uncertainties in climate responses to past land cover change: First results from the LUCID intercomparison study. *Geophysical Research Letters*, 36(14).
- Pongratz, J., Reick, C. H., Raddatz, T., Caldeira, K., & Claussen, M. (2011). Past land use decisions have increased mitigation potential of reforestation. *Geophysical Research Letters*, 38(15).

➤ **Appendix A: Proposal for sub-grid archiving of selected land output for CMIP6 (Draft, September 10, 2014)**

Co Task Leads: Elena Shevliakova and David Lawrence

N.B. This draft proposal is not yet complete and is included to provide background information on efforts have begun to redefine output for land variables in CMIP. The intention, once the draft proposal is complete, is to circulate the proposal to the other land MIPs and to major modeling centers for comment. The initial variable list included here is very preliminary and is likely to undergo extensive revision prior to a final variable list that will be proposed to CMIP.

1. Motivation

The majority of CMIP5-class climate models and Earth system models (ESMs) represent land sub-grid spatial heterogeneity by splitting each land grid into sections (i.e. tiles or units) with similar ecological, biogeochemical, and hydrological characteristics. Current land components capture two kinds of sub-grid heterogeneity: 1) hydrological - land surfaces covered by liquid or frozen water (e.g. lakes, wetlands, glaciers) or not (e.g. bare and vegetated surfaces) and 2) land-use and land management induced (e.g. cropland, pastures, urban, natural and secondary, i.e., harvested forests, plantations, abandoned land). Sub-grid tiling applies to both above- and below-ground sections of the land components. Physical, ecological, and biogeochemical land state variables and surface fluxes are calculated separately for each tile. However, frequently, including in the CMIP5 archive, the tile-specific variables were averaged and the grid-cell mean values were reported. Consequently, a large amount of valuable information was lost with respect to how each surface type with different hydrological and land-use properties responds to climate change and/or direct anthropogenic modifications.

In order to better characterize surface climate, its variability and change, we propose to expand the CMOR data convention in order to capture horizontal land sub-grid heterogeneity. In addition to the land-grid cell values, we propose to request a subset of selected variables on multiple land tiles. This reporting and archiving modification will significantly expand the utility of Earth System Model output for scientific analysis and climate change impacts studies.

Each land model has a unique tiling scheme (e.g., CLM, Fig. 1; LM3, Fig. 2), so the archiving protocol needs to be general enough to work for the range of existing model structures.

2. Proposed sub-grid reporting

2.1 Types of tiles

In the context of CMIP6 we propose to report tile-specific information *for a subset of 4 categories* to capture land-use induced surface heterogeneity: (1) Natural and Secondary land types (including bare ground and vegetated wetlands), (2) pasture-land, (3) croplands, and (4) urban. The remaining tiles, such as lakes, rivers and glaciers, will be excluded from the reported tile-specific values. The proposed set of land-use tile reporting units closely corresponds to land-use units to be used in the CMIP6 historical land-use reconstructions and future scenarios. Primary (i.e., natural vegetation never affected by LULCC activity) and secondary vegetation (i.e., natural vegetation that has previously been harvested or establishes on abandoned agricultural lands) are combined because most land models do not yet distinguish between these two land types.

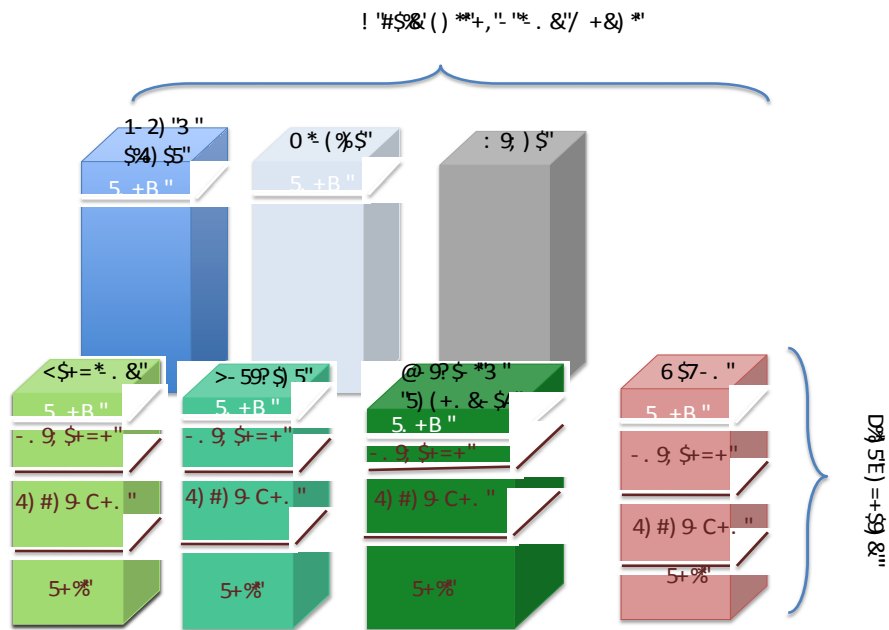


Figure 1. proposed reporting structure for land-tiles

<i>Tile type</i>	<i>Tile Suffix</i>	<i>Comment</i>
<i>Urban and rural settlement</i>	<i>“urb”</i>	
<i>Primary and secondary lands</i>	<i>“psl”</i>	<i>Forest, grasslands, etc or bare</i>
<i>Croplands</i>	<i>“crp”</i>	
<i>Managed pasturelands</i>	<i>“pst”</i>	

For selected key variables, data should be reported for each tile separately, in addition to the grid cell mean. The tiles containing biogeochemical information will report up to four stocks of biogeochemical tracers (e.g. Carbon, Nitrogen) – vegetation, soil, litter, and anthropogenic storage. The latter is used in a subset of land models and reflects the fact that some harvested carbon is not released into the atmosphere immediately, but rather with some time-delay from a year to century (e.g., wood products, food)

2.2 Variables reported by tile

We propose to distinguish 4 tiled variables representing the model tiling structure and how this changes through time, biogeophysical variables, biogeochemical variables, and ???

(A) Sub-grid structure

These variables will report for each tile and will include annual gridded fractional coverage of each tile through time (e.g. under land-use scenario or climate change). If a tile fraction does not change

in time (e.g. if urban area does not change for a particular model), the fractional area should be reported as static. Static files also could include belowground depth of tile(s).

frac_urb
frac_nsl
frac_pst
area_crp

(B) Biogeochemical and ecological variables

Variables to assess effects of LULCC on biogeochemical characteristics and functioning. Tiles resolving biogeochemical cycling and/or representations of vegetation/soils/humans will report. If different vegetation tiles share soil tile, then report same soil tile value for both land use categories.

Biospheric carbon fluxes

gpp_tile – gross primary productivity
npp_tile – net primary productivity
cfire_tile – carbon lost through fire
soil_resp_tile – soil respiration or total respiration (hr???)
nee_tile – net ecosystem exchange

Carbon pools – only instantaneous values at Jan 1, 0Z (instantaneous values requested to enable calculations of carbon cycle closure)

cSoil_tile – carbon mass in soil pool
cVeg_tile – carbon mass in vegetation
cLitter_tile – carbon mass in above and belowground litter pools
cAnthrop_tile – anthropogenic pool (e.g. harvested crop on cropland or grazed carbon on pastures, wood harvest on natural and secondary tiles)

(C) Biogeophysical variables

Energy and hydrological variables to assess effects of LULCC on biophysical characteristics and functioning

all 4 tiles will report (do we need all variables for full energy balance closure?)

tas_tile – near-surface air temperature (2m)
tlsi_tile – surface ‘skin’ temperature
huss_tile – near-surface specific humidity
hfls_tile – latent heat flux (split out ET partitioning?)
hfss_tile – sensible heat flux
hgr_tile – ground heat flux (new CMOR variable)
fah_urb – anthropogenic heat flux (only for urban)
rsus_tile – surface upwelling shortwave (to calculate albedo, downwelling not needed at tile level)
rlus_tile
snd_tile – snow depth (or snow water equivalent, note that in CMIP5 the swe variable is listed as liquid water content of snow layer – should be snow water equivalent)
lai_tile – leaf area index
mrsos_tile – total soil water content in the top 10 cm

(D) LULCC area changes and carbon transfers/fluxes

The variables in this category are specifically requested to permit analysis of how models represent LULCC and how LU processes are represented in the ESMs.

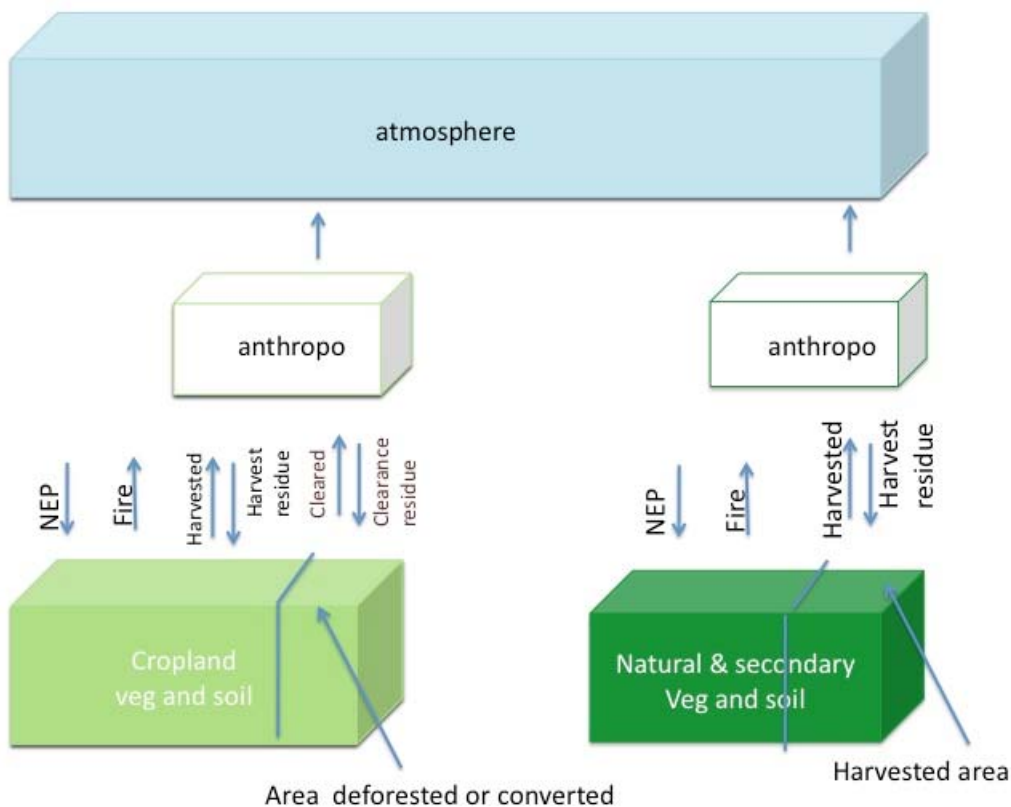
frac_out_tile – annual gross fraction of each tile that was transferred into other land use category

frac_in_tile – annual gross fraction of tile that came from other categories

frac_net_tile – annual net change in the tile fraction

frac_harv_nsl – annual fraction of natural and secondary vegetation tile harvested for wood

We recognize that models have very different implementation of LU processes and only would be able to report a subset of variables



If model has explicit anthropogenic pools

$C_{harv_anthrop_tile}$ – carbon from wood harvest on LU tiles that enters anthropo tile

$C_{cleared_anthrop_tile}$ – cleared carbon on LU tiles (e.g. for deforestation) that enters anthropo tile

$C_{other_anthrop_tile}$ – other LULCC-induced carbon removal that enters anthropo tile (e.g., by grazing or crop harvesting)

$C_{anthrop_atm}$

If models have no explicit anthropogenic pools:

$C_{harv_atm_tile}$ – harvested carbon on LU tiles released to atmosphere

$C_{cleared_atm_tile}$ – cleared carbon on LU tiles (e.g. for deforestation) released to atmosphere

All models:

C_clearance_residue_tile – carbon that is removed from biosphere during clearing or harvesting and is returning into litter or soil (not anthropogenic pool or atmosphere)

C_harv_resid_tile - carbon after harvesting that left as residue into litter or soil

All biospheric fluxes should follow convention of Chapman et al. (2006)

3. Examples of tile-reporting/aggregation from existing models

3.1 CLM

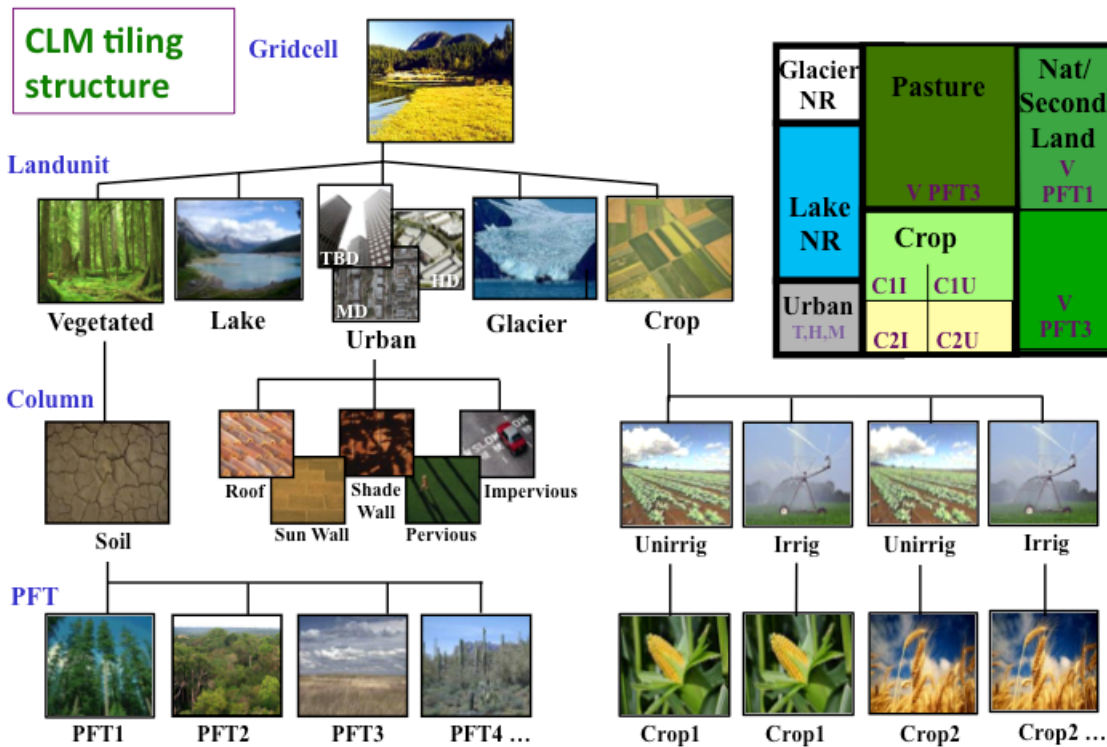


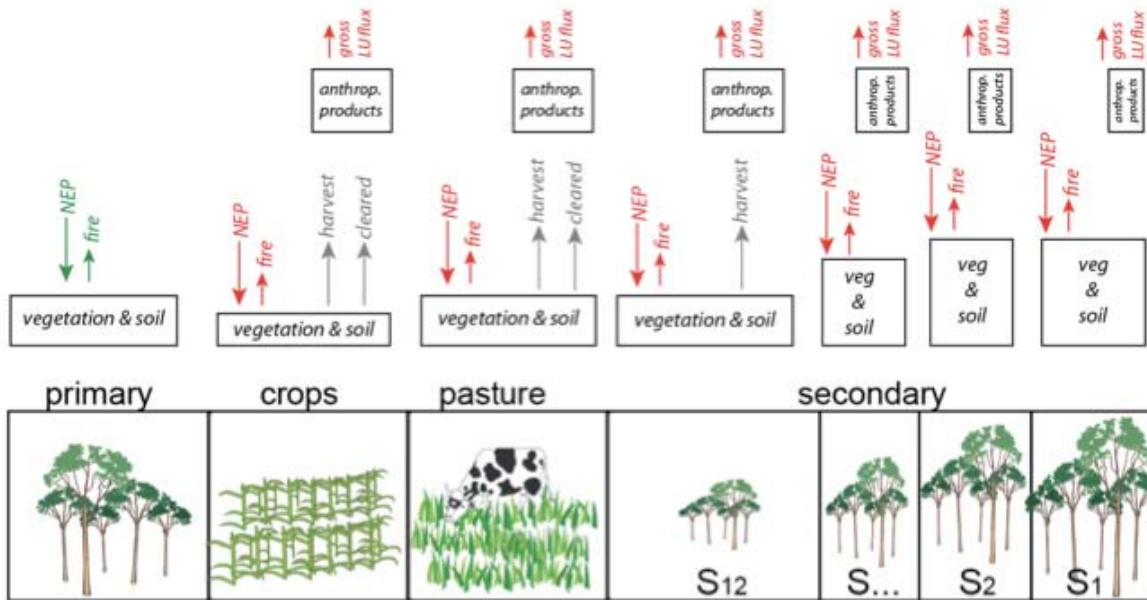
Figure #: CLM tiling structure and possible reporting under proposed subgrid structure.

CLM captures a variety of ecological and hydrological sub-grid characteristics.

In order to meet requirements of the proposed land-tile oriented design, the following aggregation would be required:

- (1) all un-irrigated and irrigated crop fractions would be aggregated
- (2) pasture???
- (3) all vegetated PFTs including bare soil PFT would be aggregated
- (4) tall building, high density, and medium density fractions would be aggregated

3.2 GFDL LM3 example ...



Ocean Carbon Cycle Model Intercomparison Project, Phase 6 (OCMIP6)

Application for CMIP6-Endorsed MIPs

Date: 22 September 2014

No Update received

Proposals from MIPs should include the following information:

- * *Preliminary information used to determine whether a MIP should be endorsed for CMIP6 or not.*
- ** *Information that must be provided later (and before the panel can determine which experiments, if any, will be incorporated in the official CMIP6 suite).*

Name of MIP: OCMIP6 – Ocean Carbon Cycle Model Intercomparison Project, Phase 6

OCMIP6 Chair: James Orr (james.orr@lscce.ipsl.fr)

OCMIP6 Scientific Steering Committee members (to be invited):

Andreas Oschlies, Scott Doney, Corinne Le Quere, Jorge Sarmiento,
Fortunat Joos

Link to website: <http://ocmip5.ipsl.jussieu.fr/OCMIP/>

Goal of OCMIP6 and brief overview

OCMIP is an open international collaboration that aims to improve and accelerate development of global-scale, three-dimensional, ocean biogeochemical models that include the carbon cycle and related biogeochemical and ecosystem components. OCMIP focuses on model evaluation and intercomparison while providing a forum for international discussion and collaboration. OCMIP5 has assessed results from the ocean biogeochemical model components of the earth system models that participated in CMIP5.

References

Bopp, L., Resplandy, L., Orr, J. C., Doney, S. C., Dunne, J. P., Gehlen, M., Halloran, P., Heinze, C., Ilyina, T., Séférian, R., Tjiputra, J., and Vichi, M.: Multiple stressors of ocean ecosystems in the 21st century: projections with CMIP5 models, *Biogeosciences*, 10, 6225-6245, doi:10.5194/bg-10-6225-2013, 2013.

Orr, J. C., V. J. Fabry, O. Aumont, Laurent Bopp, S. C. Doney, R. M. Feely, A. Gnanadesikan, N. Gruber, A. Ishida, F. Joos, R. M. Key, K. Lindsay, E. Maier-Reimer, R. Matear, P. Monfray, A. Mouchet, R. G. Najjar, G.K. Plattner, K. B. Rodgers, C. L. Sabine, J. L. Sarmiento, R. Schlitzer, R. D. Slater, I. J. Totterdell, M.F. Weirig, Y. Yamanaka, and A. Yool. Anthropogenic ocean acidification over the twenty-first century and its impact on calcifying organisms, *Nature*, 437, 681–686.

Overview of the proposed experiments

OCMIP6 will exploit results from the planned CMIP6 experiments. In addition, new OCMIP6 protocols will be developed (1) to run CMIP6's ocean dynamical-biogeochemical models in stand-alone mode, forced by data-based historical forcing (reanalysis data) and (2) to update protocols to evaluate circulation models with passive tracers, namely CFC's and SF6.

Overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments

OCMIP6 will compare results from the ocean biogeochemical components of the CMIP6 earth system models, many having much higher resolution. OCMIP6 will also analyze the analogous forced ocean simulations with the CMIP6 ocean biogeochemical models, focusing in part on how internal variability differs between coupled and forced simulations. Additionally, OCMIP6 will validate the CMIP6 ocean model components by comparing their simulations of 2 passive tracers (CFC and SF6) to a large global observational database. To promote wider collaboration, a single OCMIP6 access interface will allow easy access to a multifaceted database including CMIP6 data, corresponding derived data and forced model results, as well as data from related model intercomparison efforts (e.g., RECCAP and MAREMIP).

Proposed timing: Coincident with CMIP6

Ocean Model Inter-comparison Project (OMIP)

Application for CMIP6-Endorsed MIPs

Date: 28 October 2014

Name of MIP: Ocean Model Inter-comparison Project (OMIP)

Co-chairs of MIP:

Gokhan Danabasoglu, NCAR, US (gokhan@ucar.edu)
Stephen M. Griffies, GFDL/NOAA, US (stephen.griffies@noaa.gov)

Members of the Scientific Steering Committee:

CLIVAR Ocean Model Development Panel (OMDP) and collaborators:

Claus Boning (Germany)
Eric Chassignet (US)
Enrique Curchitser (US)
Helge Drange (Norway)
David Holland (US)
Yoshiki Komuro (Japan)
William Large (US)
Simon Marsland (Australia)
Simona Masina (Italy)
George Nurser (UK)
Andreas Oschlies (Germany)
Anna Pirani (CLIVAR ICPO, Italy)
Anne-Marie Treguier (France)
Mike Winton (US)
Stephen Yeager (US)

Link to website:

The proposed OMIP is based on the Coordinated Ocean-ice Reference Experiments phase II (CORE-II) framework. The current web site is

<http://www.clivar.org/clivar-panels/omdp/core-2>

Goal of the MIP and a brief overview:

The primary goal of the OMIP is to provide a framework for evaluation, understanding, and improvements of ocean components of the earth system models that contribute to CMIPs. The framework describes a protocol (Griffies et al. 2012) for performing global ocean and sea-ice coupled simulations forced with common atmospheric data sets. The OMIP will use the Coordinated Ocean-ice Reference Experiments (CORE) inter-annually varying atmospheric data sets (Large and Yeager 2009), representing the second phase of the CORE, i.e., CORE-II. These data sets cover the 62-year period from 1948-2009. In the

oceanographic community, the CORE-II simulations are usually referred to as hindcast experiments.

In addition to the primary goal stated above, the OMIP experiments have additional benefits and applications. These include their use in: investigation of mechanisms for seasonal, inter-annual, and decadal variability; attribution of ocean-climate events to forced and natural variability; evaluation of robustness of mechanisms across models; and bridging observations and modeling, by complementing ocean reanalysis from data assimilation approaches. They also provide consistent ocean and sea-ice states that can be used for initialization of climate (e.g., decadal) prediction experiments.

To date, CORE-II simulations have been performed worldwide by over twenty modeling groups. The simulations are being analyzed in about ten separate studies, each focusing on a specific aspect of the solutions. These include analysis of mean states in the North Atlantic with a focus on the Atlantic Meridional Overturning Circulation (Danabasoglu et al. 2014) and an assessment of global and regional sea level changes (Griffies et al. 2014). The manuscripts are being published in a Special Issue of Ocean Modelling.

As in the current CORE-II effort, the OMIP will be coordinated by the WCRP Climate Variability and Predictability (CLIVAR) Ocean Model Development Panel (OMDP; formerly Working Group on Ocean Model Development, WGOMD). The CORE atmospheric forcing data sets are collaboratively supported by the National Center for Atmospheric Research (NCAR) and the Geophysical Fluid Dynamics Laboratory (GFDL). All data sets, codes for the bulk flux formulae, technical report, and other support codes along with the release notes are freely available at the above web site.

References:

- Danabasoglu, G., S. G. Yeager, D. Bailey, E. Behrens, M. Bentsen, D. Bi, A. Biastoch, C. Boning, A. Bozec, V. Canuto, C. Cassou, E. Chassignet, A. C. Coward, S. Danilov, N. Diansky, H. Drange, R. Farneti, E. Fernandez, P. G. Fogli, G. Forget, Y. Fujii, S. M. Griffies, A. Gusev, P. Heimbach, A. Howard, T. Jung, M. Kelley, W. G. Large, A. Leboissetier, J. Lu, G. Madec, S. J. Marsland, S. Masina, A. Navarra, A. J. G. Nurser, A. Pirani, D. Salas y Melia, B. L. Samuels, M. Scheinert, D. Sidorenko, A.-M. Treguier, H. Tsujino, P. Uotila, S. Valcke, A. Voldoire, and Q. Wang, 2014: North Atlantic simulations in Coordinated Ocean-ice Reference Experiments phase II (CORE-II). Part I: Mean states. *Ocean Modelling*, **73**, 76-107, doi:10.1016/j.ocemod.2013.10.005.
- Griffies, S. M., M. Winton, B. Samuels, G. Danabasoglu, S. Yeager, S. Marsland, H. Drange, and M. Bentsen, 2012: Datasets and protocol for the CLIVAR WGOMD Coordinated Ocean-ice Reference Experiments (COREs). *WCRP Report No. 21/2012*, pp.21.
- Griffies, S. M., J. Yin, P. J. Durack, P. Goddard, S. C. Bates, E. Behrens, M. Bentsen, D. Bi, A. Biastoch, C. W. Boning, A. Bozec, E. Chassignet, G. Danabasoglu, S. Danilov, C. M. Domingues, H. Drange, R. Farneti, E. Fernandez, R. J. Greatbatch, D. M. Holland, M. Ilicak, W. G. Large, K. Lorabacher, J. Lu, S. J. Marsland, A. Mishra, A. J. G. Nurser, D. Salas y Melia, J. B. Palter, B. L. Samuels, J. Schroter, F. U. Schwarzkopf, D. Sidorenko, A. M. Treguier, Y.-H. Tseng, H. Tsujino, P. Uotila, S. Valcke, A. Voldoire, Q. Wang, M. Winton, and X. Zhang,

2014: An assessment of global and regional sea level for years 1993-2007 in a suite of interannual CORE-II simulations. *Ocean Modelling*, **78**, 35-89, doi: 10.1016/j.ocemod.2014.03.004.

Large, W. G., and S. G. Yeager, 2009: The global climatology of an interannually varying air-sea flux data set. *Climate Dynamics*, **33**, 341-364, doi: 10.1007/s00382-008-0441-3.

An overview of the proposed experiments:

The OMIP consists of only 1 global ocean – sea-ice coupled simulation, run for a minimum of five repeating cycles of the forcing period. With the current 62-year forcing data, the integration length is 310 years. The solutions from the fifth cycle are used for analysis. This simulation is a Tier 1 experiment.

The details of the datasets and the experimental protocol are available in Griffies et al. (2012) and Danabasoglu et al. (2014). Here, we include a very brief summary. The ocean models are initialized using the January-mean potential temperature and salinity climatology from observations and typically from a state of rest. The sea ice models are generally initialized from an existing state taken from another simulation. The surface heat fluxes are determined by the radiative fluxes from CORE-II and turbulent fluxes computed based on the ocean state and CORE-II atmospheric state. It is highly recommended that bulk formulae for the turbulent fluxes follow the ones described in the OMIP protocol in order to facilitate comparisons between the model simulations. There is no restoring term applied to the surface temperature field. In contrast, the surface salinity field is damped to a monthly observational climatology. However, the protocol does not specify a particular recipe for salinity restoring and it is left to the modelers to choose their optimal salinity restoring procedure. Using a unified salinity restoring across all models is not feasible, due to physical sensitivities related to high latitude processes identified in Griffies et al. (2009).

Reference:

Griffies, S. M., A. Biastoch, C. Boning, F. Bryan, G. Danabasoglu, E. P. Chassignet, M. H. England, R. Gerdes, H. Haak, R. W. Hallberg, W. Hazeleger, J. Jungclaus, W. G. Large, G. Madec, A. Pirani, B. L. Samuels, M. Scheinert, A. S. Gupta, C. A. Severijns, H. L. Simmons, A. M. Treguier, M. Winton, S. Yeager, and J. Yin, 2009: Coordinated Ocean-ice Reference Experiments (COREs). *Ocean Modelling*, **26**, 1-46, doi:10.1016/j.ocemod.2008.08.007.

An overview of the proposed evaluation / analysis of the CMIP DECK and CMIP6 experiments:

The CLIVAR OMDP and collaborators have produced an updated version of the CMIP ocean model output request document (Griffies et al. 2009; Griffies et al. 2014). This document presents recommendations for sampling physical ocean fields for CMIP6 and its MIPs, including the OMIP. The goal is to precisely define a suite of ocean model diagnostics related to physical properties and processes within the simulated ocean and associated boundary fluxes.

The broader ocean modeling community is interested in analyzing ocean model output fields from models participating in OMIP. Moreover, we believe that modeling groups

themselves will be keenly interested in evaluating their ocean model simulations as the primary goal of the OMIP is to provide a common framework for evaluation, understanding, and improvements of ocean components of their coupled models.

References:

Griffies, S. M., A. Adcroft, H. Aiki, V. Balaji, M. Bentson, F. Bryan, G. Danabasoglu, S. Denvil, H. Drange, M. England, J. Gregory, R. Hallberg, S. Legg, T. Martin, T. J. McDougall, A. Pirani, G. Schmidt, D. Stevens, K. Taylor, and H. Tsujino, 2009: Sampling physical ocean fields in WCRP CMIP5 simulations. ICPO Publication Series 137, WCRP Informal Report No. 3/2009.

Griffies, S. M., A. J. Adcroft, V. Balaji, G. Danabasoglu, P. J. Durack, P. J. Gleckler, J. M. Gregory, J. P. Krasting, R. J. Stouffer, and K. E. Taylor, 2014: Sampling the physical ocean in CMIP6 simulations. (draft)

Proposed timing:

We envision that initial OMIP experiments will use our existing forcing data sets and protocol. Starting in early 2015, we plan to revisit several aspects of the atmospheric and river runoff data sets, including their extension to year 2014. Thus, we expect to update the OMIP protocol as new developments and extensions become available.

Precipitation Driver and Response Model Intercomparison Project (PDRMIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

Proposals from MIPs should include the following information:

- * Preliminary information used to determine whether a MIP should be endorsed for CMIP6 or not.
- ** Information that must be provided later (and before the panel can determine which experiments, if any, will be incorporated in the official CMIP6 suite).
- Name of MIP* *Precipitation Driver and Response Model Intercomparison Project – PDRMIP*
- Co-chairs of MIP (including email-addresses)* *Gunnar Myhre (gunnar.myhre@cicero.oslo.no) and Piers Forster (P.M.Forster@leeds.ac.uk)*
- Members of the Scientific Steering Committee* *Olivier Boucher, Drew Shindell, Toshihiko Takemura, Francis Zwiers, Slava Kharin, Jean-Francois Lamarque, Apostolos Voulgarakis, Bjørn Samset, Øivind Hodnebrog, Jana Sillmann*

Link to website (if available)* <http://cicero.uio.no/PDRMIP/> (under construction)

- Goal of the MIP and a brief overview* *PDRMIP will compare the precipitation response to various climate drivers, across models. Analyses planned include a better understanding of the drivers' importance for inter-model differences in precipitation changes, energy budget analysis and extremes related to precipitation. An additional result from PDRMIP will be on model quantification of different climate sensitivity from climate drivers, in particular if this is related to spatial location of the drivers as recently suggested.*
- References (if available)* *The main set of simulations have many similarities with single model calculations in Andrews et al. (2010); Kvalevåg et al. (2013)*
- An overview of the proposed experiments* *The main focus of PDRMIP is on changes in precipitation from various drivers of climate change. The proposed experiments includes dedicated simulations with various drivers such as CO₂, solar irradiance changes and different aerosol types to investigate the degree of difference in mean and extreme precipitation between the drivers. The simulations are a combination of slab ocean/full ocean and fixed SST. A subset of experiments will also perturb aerosols regionally, to investigate the precipitation impact of the longitudinal shift in aerosol loading across models.*
 - *The 5 core experiments consist of simulations with doubling of CO₂ concentration, tripling of the CH₄ concentration, changes in the solar constant and two simulations with aerosol one for sulphate and one for black carbon. The additional simulations include changes in sulphate and black carbon over Europe and Asia.*
 - *Ideally equal aerosol distribution should be implemented in the PDRMIP runs. The aerosol distribution will be provided on the required model spatial resolution by the PDRMIP core group. The models can also be run with aerosol emissions if implementation of a fixed aerosol distribution is not possible. These two approaches have already been tested in a few PDRMIP models.*
- Nine modelling groups have confirmed participation in PDRMIP (see more details about the models at our website).
- An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments* *PIcontrol and 4*CO₂*
- Proposed timing* *2014-2018*
- For each proposed experiment to be included in CMIP6**
 - the experimental design;
 - the science question and/or gap being addressed with this experiment;

- possible synergies with other MIPs;
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.
- If possible, a prioritization of the suggested experiments, including any rationale**
- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**
- List of output and process diagnostics for the CMIP DECK/CMIP6 data request**
 - whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;
 - whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);
 - whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.
 - whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;
 - the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.
- Any proposed contributions and recommendations for**
 - model diagnostics and performance metrics for model evaluation;
 - observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;
 - tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).
- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms.**
- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF.**

Andrews, T., Forster, P., Boucher, O., Bellouin, N. and Jones, A.: Precipitation, radiative forcing and global temperature change, *Geophysical Research Letters*, 37, L14701, 2010.

Kvalevåg, M. M., Samset, B. H. and Myhre, G.: Hydrological sensitivity to greenhouse gases and aerosols in a global climate model, *Geophysical Research Letters*, 40(7), 1432-1438, 2013.

Paleoclimate Modeling Intercomparison Project (PMIP)

Application for CMIP6-Endorsed MIPs

Date: 2 December 2014

Please return to CMIP Panel Chair Veronika Eyring (email: Veronika.Eyring@dlr.de)

Proposals from MIPs should include the following information:

* *Preliminary information used to determine whether a MIP should be endorsed for CMIP6 or not.*

** *Information that must be provided later (and before the panel can determine which experiments, if any, will be incorporated in the official CMIP6 suite).*

➤ **Name of MIP***

Paleoclimate Modeling Intercomparison Project
(This will be the fourth phase of PMIP: PMIP4)

➤ **Co-chairs of MIP (including email-addresses)***

- Pascale Braconnot (Pascale.Braconnot@lsce.ipsl.fr)
- Sandy Harrison (s.p.harrison@reading.ac.uk)

➤ **Members of the Scientific Steering Committee***

- Pascale Braconnot / LSCE, France (model and model-data)
- Sandy P. Harrison / University of Reading, UK and Macquarie University, Australia (data and model-data)
- Ayako Abe-Ouchi / AORI, University of Tokyo (ice-sheet and PCMIP)
- Pat Bartlein / University of Oregon, USA (Continental data)
- Alan Haywood / University of Leeds, UK (Mid-pliocene)
- Sylvie Joussaume / LSCE, France
- Johann Jungclaus / MPI-M, Germany (Last millennium)
- Michal Kucera / MARUM, Germany (Ocean data)
- Bette Otto-Bliesner / NCAR, USA (warm climates)
- Gilles Ramstein / LSCE, France (glacial and ice sheet)
- Karl Taylor / PCMDI, USA (Link with CMIP5)
- Paul Valdes / BRIDGE, UK (abrupt changes)

➤ **Link to website (if available)***

<http://pmip3.lsce.ipsl.fr> + <http://pmip.lsce.ipsl.fr> (PMIP1) and <http://pmip2.lsce.ipsl.fr> (PMIP2)

Goal of the MIP and a brief overview*

Since the 1990s, PMIP has developed with the following objectives:

- to evaluate the ability of climate models used for climate prediction in simulating well-documented past climates outside the range of present and recent climate variability
- to understand the mechanisms of these climate changes, in particular the role of the different climate feedbacks

To achieve these goals, PMIP has actively fostered paleo-data syntheses, multi-model analyses, including analyses of relationships between model results from past and future simulations, and model-data comparisons. These have first been focusing on the results from Atmospheric General Circulation Models (PMIP1) and then been extended to coupled Ocean-Atmosphere General Circulation Models and AOGCM including carbon cycle feedbacks, thereby closely following model developments for CMIP (PMIP2 and PMIP3). Three PMIP3 simulations were part of the CMIP5 ensemble of simulations: the last millennium, the mid-Holocene (~6,000 years ago) and the Last Glacial Maximum (~21,000 years ago), hence allowing, for the first time, the rigorous comparison of model results for past and future climates. The rationale for considering these periods was:

- for the Last Glacial Maximum, to evaluate the models on a well-documented climatic extreme, especially in terms of temperatures, and study the role of forcings and feedbacks in establishing this climate;
- for the mid-Holocene, to evaluate and analyse the models on a climate “optimum” for the northern hemisphere, characterized by enhanced monsoons, extra-tropical continental aridity and much warmer summers;
- for the last millennium, to study the mechanisms (natural variability vs impact of solar, volcanic and anthropogenic forcings) of decadal to centennial climate variability and evaluate the models’ performance w.r.t numerous detailed records.

For CMIP6, we propose to include two new warm periods in the PMIP/CMIP set of experiments: the Last Interglacial and the Mid-Pliocene, for which simulations have been performed and significantly contributed to AR5.

PMIP3/CMIP5 and PlioMIP have been very successful in terms of participation and publications. 19 groups have contributed to PMIP3/CMIP5 simulations, 12 groups have taken part in PlioMIP. PMIP3/CMIP5 simulations have been used in more than 40 publications (as of Sept 11th, 2014) and PlioMIP simulations have been the topic of more than 20 publications. PMIP simulations have brought strong contribution to 2 IPCC AR5 chapters: chapter 5 “information from paleoclimate archives” and chapter 9 “evaluation of climate models”.

PMIP simulations specifically address CMIP6 key question on “How does the Earth System respond to forcing” for a variety of forcings and with possible comparisons to data for climates states very different from the current or historical climate. PMIP also addresses question 2 (“What are the origins and consequences of systematic model biases?”) about systematic model biases, with the perspective given by documented climates different from today: PMIP simulations, with comparisons to data, can help assessing whether the biases for present-day are also found for other climate states and whether present-day biases have an impact on the

magnitude of simulated climate changes. Finally, PMIP is relevant for question 3 (“How can we assess future climate changes given climate variability, predictability and uncertainties in scenarios?”), by examining these very questions for documented past climate cases and via the use of the last millennium simulations as reference state for natural variability.

PMIP simulations are being analyzed within the Grand Challenge “Clouds, Circulation and Climate Sensitivity”. They can also provide valuable input for other grand challenges, such as those on the Cryosphere and on Regional Climate Information, with the challenge of paleoclimate modelling at fine scale. Indeed, PMIP model output is increasingly used in “paleo-impact studies”, on biodiversity or on understanding the potential impact of climate and environmental changes on early Humans. Several initiatives have already been proposed along these themes and will be reinforced in the future (e.g. Future Earth “Fast Track Initiatives and Cluster Activities” project “Making better use of the Paleoclimate Modeling Intercomparison Project simulations (MAPS)” led by P. Braconnot, a project concerning WGCM, PAGES, CLIVAR, CLiC and bioDISCOVERY).

The five proposed experiments constitute a reference ensemble for further studies within PMIP: single forcing experiments, transient experiments (testing the models on abrupt climate change and on glacial-interglacial transitions).

➤ References*

PMIP1:

Joussaume, S. and Taylor, K. E., 1995. Status of the Paleoclimate Modeling Intercomparison Project (PMIP), Proceedings of the first international AMIP scientific conference (Monterrey, California, USA, 15-19 May 1995), WCRP report 92, 425-430. Text available at: <https://pmip.lsce.ipsl.fr/publications/overview.html>

PMIP2:

Braconnot, P., B. Otto-Bliesner, S. Harrison, S. Joussaume, J.-Y. Peterchmitt, A. Abe-Ouchi, M. Crucifix, E. Driesschaert, T. Fichefet, C. D. Hewitt, M. Kageyama, A. Kitoh, A. Laîné, M.-F. Loutre, O. Marti, U. Merkel, G. Ramstein, P. Valdes, S. L. Weber, Y. Yu, and Y. Zhao, 2007. Results of PMIP2 coupled simulations of the Mid-Holocene and Last Glacial Maximum – Part 1: experiments and large-scale features. *Climate of the Past*, 3, 261–277, www.clim-past.net/3/261/2007/.

Braconnot, P., B. Otto-Bliesner, S. Harrison, S. Joussaume, J.-Y. Peterchmitt, A. Abe-Ouchi, M. Crucifix, E. Driesschaert, T. Fichefet, C. D. Hewitt, M. Kageyama, A. Kitoh, M.-F. Loutre, O. Marti, U. Merkel, G. Ramstein, P. Valdes, L. Weber, Y. Yu, and Y. Zhao, 2007. Results of PMIP2 coupled simulations of the Mid-Holocene and Last Glacial Maximum – Part 2: feedbacks with emphasis on the location of the ITCZ and mid- and high latitudes heat budget. *Climate of the Past*, 3, 279–296, www.clim-past.net/3/279/2007/.

PMIP2/PMIP3:

Braconnot, P., S. P. Harrison, M. Kageyama, P. J. Bartlein, V. Masson-Delmotte, A. Abe-Ouchi, B. Otto-Bliesner and Y. Zhao, 2012. Evaluation of climate models using palaeoclimatic data, *Nature Climate Change*, DOI: 10.1038/NCLIMATE1456

Schmidt, G.A., J. D. Annan, P. J. Bartlein, B. I. Cook, E. Guilyardi, J. C. Hargreaves, S. P. Harrison, M. Kageyama, A. N. LeGrande, B. Konecky, S. Lovejoy, M. E. Mann, V. Masson-Delmotte, C. Risi, D. Thompson, A. Timmermann, L.-B. Tremblay, and P. Yiou, 2014. Using paleo-climate comparisons to constrain future projections in CMIP5, *Climate of the Past*, 10, 221-250

An overview of the proposed experiments*

The following table summarizes the experiments proposed by PMIP for CMIP6. These experiments all build from the DECK experiments and are part of the core of PMIP simulations (~10), which will themselves constitute a basis for other PMIP experiments (sensitivity analyses, transient simulations starting from the core ones). Within PMIP, each PMIP working group will organize their set of simulations, as PMIP also federates focused MIPs such as PlioMIP on the Pliocene climate, LIGMIP on the Last Interglacial, PAST2K on the last two millennia.

Table 1: summary of proposed experiments. In yellow: PMIP3/CMIP5 experiments. In green: new experiments for CMIP6. The PMIP3/CMIP5 experiment names in the ESFG nomenclature are indicated in italic below each period name.

Period	Purpose	Imposed boundary conditions	# of years
Last millennium <i>(past1000)</i> 850-1850 CE	a) Evaluate the ability of models to capture observed variability on multi-decadal and longer time-scales. b) Determine what fraction of the variability is attributable to “external” forcing and what fraction reflects purely internal variability. c) Provides a longer-term perspective for detection and attribution studies	<ul style="list-style-type: none"> • Solar variations • Volcanic aerosols • Atmospheric concentration of well mixed greenhouse gases • Land use • Orbital parameters 	1000 (after spin-up period)
Mid-Holocene <i>(midHolocene)</i> 6 kyr ago	a) Compare with paleodata the model response to known orbital forcing changes and changes in greenhouse gas concentrations. b) Relationships between changes in mean state and variability	<ul style="list-style-type: none"> • Orbital parameters • Atmospheric concentration of well-mixed greenhouse gases 	≥100 (after spin-up period)
Last Glacial Maximum <i>(lgm)</i> 21 kyr ago	a) Compare with paleodata the model response to ice-age boundary conditions. b) Attempt to provide empirical constraints on global climate sensitivity.	<ul style="list-style-type: none"> • Ice-sheet and land-sea mask • Atmospheric concentration of well-mixed greenhouse gases • Orbital parameters 	≥100 (after spin-up period)
Last Interglacial 128 kyr ago	a) Evaluate climate model for warm period, high sea-level stand b) Impacts of smaller ice-sheets/higher sea-level on climate	<ul style="list-style-type: none"> • Orbital parameters • Atmospheric concentration of well-mixed greenhouse gases 	≥100 (after spin-up period)
Mid-Pliocene Warm Period 3.2 Ma ago	a) How does the Earth System respond in the long term to CO ₂ forcing analogous to that of the modern? b) What is the significance of CO ₂ induced polar amplification for the stability of the ice sheets, sea-ice and sea-level?	<ul style="list-style-type: none"> • Ice-sheet and land-sea mask, topography (smaller ice-sheets) • Atmospheric concentration of well-mixed greenhouse gases • Orbital parameters 	≥100 (after spin-up period)

For all these periods the model to be used is the same as the one used for future climate projections. Therefore depending on the groups the model will be only atmosphere-ocean coupled models or Earth System models. The reference for the analyses will be the CMIP6 pre-industrial simulation. Hereafter, we shortly describe the Mid-Holocene and Last Glacial Maximum, periods which have already been a focus of PMIP since its start and which have been part of the PMIP3-CMIP5 simulations. More details are given below on the Last Millennium (part of PMIP3-CMIP5 as well) and on the two new periods proposed for CMIP6: the Last Interglacial and the Mid-Pliocene Warm Period.

➤ *Mid-Holocene (midHolocene) and Last Glacial Maximum (lgm):*

The mid-Holocene (~6000 years ago) and the Last Glacial Maximum (LGM, ~21000 years ago) constitute the most recent quasi-stable climatic extremes: the mid-Holocene is often described as a warm state, or “climate optimum”, in which dominant features of the global hydrological cycle, such as the North African and Asian monsoon, were amplified; the LGM is a cold extreme in which greenhouse gas concentrations were at their minimum and continental ice-sheet at their maximum size, covering large areas of northern North America and northwestern Eurasia.

These periods have been the focus for paleo-data syntheses since the beginning of the PMIP project and therefore are well documented in terms of temperature, hydrological cycle and land surface type. Some long standing model-data disagreement are echoing preoccupations for future climate change, such as the systematic underestimation of the northward penetration of the African monsoon rainfall onto the continent compared to available records for the Mid-Holocene. The LGM is relevant for studying feedback mechanisms at work in establishing a temperature response as large as (although with an opposite sign) as that predicted for the end of the 21st century. Both periods constitute test cases for our understanding of mechanisms of climate change, such as the interplay between circulation changes and radiation/cloud changes, the respective strengths of feedbacks from different components of the climate system, and for our understanding of the connections between global and regional climate changes.

Compared to the previous phases of PMIP a particular emphasis will be put on the impact of dust on the mean climate and climate feedbacks, as well as on uncertainties in boundary conditions or surface feedbacks related to the vegetation or interactive carbon cycle.

The reference experiments for both the *midHolocene* and *lgm* simulations are the pre-industrial control and it is very interesting to compare those experiments with an idealized experiment designed to study mechanisms of future climate change, such as abrupt4xCO₂. PMIP4 will benefit from idealized experiments proposed by CFMIP, such as AMIPminus4K or abrupt0.5CO₂ which will help comparing feedbacks at work in setting up a cold climate vs. those at work for a warm climate. Similarly, an AMIP experiment with insolation prescribed at a 6ky BP value will be very useful to analyze the strengths of forcings and feedbacks within the climate system and the mechanisms for common/different responses for past and future climates. These sensitivity experiments will be discussed as part of PMIP in the Past to Future working group. They would echo PMIP1 simulations (<http://pmip.lsce.ipsl.fr>), while ESM simulations would echo PMIP2 and PMIP3 (<http://pmip2.lsce.ipsl.fr> and <http://pmip3.lsce.ipsl.fr>) simulations, hence allowing a characterization of the models' evolution in their ability to represent documented large climate changes.

➤ Last Millennium (*past1000*):

The last millennium is the best-documented period of climate change in a multi-century time frame. Climate has varied considerably during the late Holocene and these changes left their traces in history (Medieval Climate Optimum, Little Ice Age). However, the relative magnitude of natural fluctuations due to internal variability of the Earth's climate system and to variations in the external forcings (Sun, orbital, volcanic) and the present global warming, attributed to anthropogenic greenhouse gases, is still under debate. Simulations of the last millennium (LM) therefore directly address the first CMIP6 key scientific question "How does the Earth System respond to forcing". Investigating the response to (mainly) natural forcing under climatic background conditions not too different from today is crucial for an improved understanding of climate variability, circulation, and regional connectivity. LM simulations also allow assessing climate variability on decadal and longer scales and provide information on predictability under forced and unforced conditions. These are crucial for near-term predictions and thus address the third CMIP5 scientific question "How can we assess future climate changes given climate variability, predictability and uncertainties in scenarios". In providing in-depth model evaluation with respect to observations and paleoclimatic reconstructions in particular addressing details of response to forcing, LM simulations serve to "understand origins and consequences of systematic model biases", thus addressing also the second CMIP6 scientific question.

LM will build on DECK experiments, in particular the pre-industrial control simulation as unforced reference and the historical simulations. Moreover, LM provide initial conditions for historical simulations starting in the 19th century that are considered superior to the *piControl* state as it includes integrated information from the forcing history (e.g. large volcanic eruptions in the early 19th century).

Within PMIP, a considerable number of individual researchers and modelling groups is committed to perform LM simulations. The simulations will base on experience gained in PMIP3/CMIP5 where more than a dozen modelling groups participated and a total of 15 LM experiments were stored in the ESGF database. Several studies, partly reflected by entries in the AR5 chapter 5, have highlighted the value of the LM multi-model ensemble. The PMIP3 LM working group (WG Past2K) is closely cooperating with the PAGES initiative PAGES2k promoting regional reconstructions of climate variables and variability modes. Collaborative work has focused on reconstruction-model intercomparison (e.g. Bothe et al., 2013) and assessment of variability modes (e.g. Raible et al., 2014). Integrated assessment of reconstruction and simulations has led to progress in model evaluation and process understanding (e.g. Lehner et al., 2013; Sicre et al., 2013; Jungclaus et al., 2014). WG Past2K will promote future common analyses and workshops bringing together observational and modelling expertise.

For CMIP6 progress is expected owing to new, more comprehensive reconstructions of volcanic forcing (Sigl et al., in preparation), improved models, and an experimental protocol that ensures seamless simulations from the pre-industrial past to the future. Higher-resolution simulations will allow assessing more regional details and processes, e.g. storm-tracks, precipitation.

➤ Last Interglacial:

The Summary for Policymakers for both the IPCC WG1 AR4 and AR5 included statements on the Last Interglacial (LIG):

“There is very high confidence that maximum global mean sea level during the last interglacial period (129,000 to 116,000 years ago) was, for several thousand years, at least 5 m higher than present, and high confidence that it did not exceed 10 m above present. During the last interglacial period, the Greenland ice sheet very likely contributed between 1.4 and 4.3 m to the higher global mean sea level, implying with medium confidence an additional contribution from the Antarctic ice sheet. This change in sea level occurred in the context of different orbital forcing and with high-latitude surface temperature, averaged over several thousand years, at least 2°C warmer than present (high confidence).”

Yet the AR4 and AR5 had no coordinated simulations for the LIG to assess the interplay of polar amplification of temperature, seasonal memory of sea ice, and precipitation/storm track changes on the stability of the Greenland ice sheet and its contribution to the sea level high stand nor the interplay of oceanic and atmospheric temperatures and circulation on the stability of the Antarctic ice sheet. Climate model simulations for the LIG assessed in the AR5, although completed by many modeling groups, varied in their forcings and often were not made with the same model/same resolution as the CMIP5 future projections, thus providing a useful but incomplete means for assessment (Chapter 5; Lunt et al., 2013). Similarly, Greenland ice sheet simulations assessed in the AR5 used offline models with a variety of climate forcing setups, not then allowing feedbacks among the Earth system components (Chapter 5). No simulations were available to assess the Antarctic ice sheet (particularly, the West Antarctic Ice Sheet) contribution to the LIG sea level high stand.

We propose a CMIP6 time-slice experiment for the LIG to determine the interplay of warmer atmospheric and oceanic temperatures, changed precipitation, and changed surface energy balance on ice sheet thermodynamics and dynamics during this period. Still uncertain are how well ice sheet-climate models can predict the stability of the ice sheets and if thresholds may be passed this century. A LIG simulation will be of high societal relevance because of implications for sea level changes as well as sea ice and monsoons. The LIG simulation will also provide an ‘out-of-sample’ evaluation of new features of CMIP6 models: coupled climate-ice sheet models. The LIG is the most suitable of the warm interglacials for a CMIP6 assessment because of the wealth of data including: ice cores providing measurements of well-mixed greenhouse gases, aerosols including dust and sea salt, and stable water isotopes as a proxy for temperature, as well as for Greenland, ice sheet elevation and extent; marine records for ocean temperatures and geotracers that can be interpreted in terms of water masses and overturning strength; speleothems that provide indication of monsoon strength; and terrestrial records that indicate temperature and vegetation. As well, new records are refining our knowledge of sea ice extent, fire, and biodiversity.

The proposed CMIP6 simulation for the LIG is particularly relevant to the WCRP Grand Challenges: Changes in Cryosphere and Regional Sea-level Rise, but also to Regional Climate Information and Clouds, Circulation and Climate Sensitivity because of the large forcings and thus large regional responses as recorded in the data. It addresses well the broad scientific questions: 1. How does the Earth System respond to forcing? and 2. What are the origins and consequences of systematic model biases (especially at high latitudes and relevant to the stability

of the ice sheets)? As part of PMIP, some groups will additionally perform transient coupled ice sheet-climate simulations that will provide rates of change for sea level, including regional sea level if offline GIA models applied, as well as a measure of the capability of these models to initiate the next glacial inception.

➤ **Pliocene warm period**

The Pliocene epoch was the last time in Earth history when atmospheric CO₂ concentrations approached modern values (~400 ppmv) whilst at the same time retaining a near modern continental configuration. The IPCC 5th Assessment report chapter 5 (Masson-Delmotte et al., 2013) states that model–data comparisons for the Pliocene provide high confidence that mean surface temperature was warmer than pre-industrial (Dowsett et al., 2012; Haywood et al., 2013). Global mean sea surface temperatures have been estimated to be +1.7°C above the 1901–1920 mean based on large data syntheses (Lunt et al., 2010; Dowsett et al., 2012). Existing climate model simulations have produced a range of global mean surface air temperature of +1.9°C and +3.6°C relative to the 1901–1920 mean (Haywood et al., 2013). Model simulations have indicated that meridional temperature gradients were reduced (due to high latitude warming), which has significant implications for the stability of polar ice sheets and sea level in the future (e.g. Miller et al. 2012). Compilations of vegetation (Salzmann et al., 2008) have indicated that the global extent of arid deserts decreased and boreal forests replaced tundra, and climate models predict an enhanced hydrological cycle, but with a large inter-model spread (Haywood et al., 2013). The East Asian Summer Monsoon, as well as other monsoon systems, may also have been enhanced (Zhang et al. 2013). Although climate model simulations for the Pliocene were assessed in the AR5, these simulations were not derived from the same model/same resolution as the CMIP5 future projections, thus reducing the communities' ability to assess and compare changes in global and regional Pliocene climates, vis-à-vis similar predictions of future climate change (Haywood et al., 2013).

We propose a CMIP6 time-slice experiment for the Pliocene to understand the long term response of the Earth's climate system to a near modern concentration of atmospheric CO₂ (longer term climate sensitivity or Earth System Sensitivity), and to understand the response of ocean circulation, Arctic sea-ice, modes of climate variability (e.g. ENSO), as well as the global response in the hydrological cycle and regional changes in monsoon systems. A Pliocene simulation will be of high societal relevance because of its potential to inform policy makers on required emission reduction scenarios designed to prevent global annual mean temperatures increase by more than 2 to 3 °C in the long term (beyond 2100 AD).

The proposed CMIP6 simulation for the Pliocene is relevant to two of the WCRP Grand Challenges. This includes Clouds, Circulation and Climate Sensitivity because of the enhanced CO₂ forcing (contemporaneous with modern CO₂ forcing), providing a unique opportunity to examine an equilibrium climate state to a near modern concentration of atmospheric CO₂. The pattern of polar amplification preserved Pliocene climate archives can be compared directly with the latest generation of CMIP models making a valuable contribution towards addressing the potential polar amplification problem. Through the analysis of Pliocene polar amplification in CMIP models, and examining the geological interpretation of a seasonally sea-ice free Arctic Ocean during the Pliocene, our CMIP6 simulation will also address the WCRP Grand Challenge of Changes in the Cryosphere. Whilst uncertainty exists in Pliocene sea level reconstruction,

IPCC AR5 states with high confidence that Pliocene sea-levels were higher than the pre-industrial era, with a number of independent methods indicating a sea-level rise of between 10 and 20 m. This indicates potential long term instability of both the Greenland and Antarctic Ice Sheets (Miller et al. 2012) with CO₂ concentrations at approximately 400 ppmv.

CMIP6 Pliocene experiments will be used within the Pliocene Ice Sheet Model Intercomparison Project in order to better constrain the climatological forcing in ice sheet model simulations for the Pliocene in the future. There is a well-organized and highly active community of Pliocene climate modellers within PMIP, with the Pliocene working group being one of the most successful working groups within PMIP3. The working group is closely associated with the United States Geological Survey (USGS) who has had a highly productive core program focused on Pliocene environmental reconstruction for the last 25 years, and their data has been used to underpin almost all model-data comparisons performed for the Pliocene. Thus, CMIP6 can expect a high degree of continued support and new Pliocene data sets from the USGS for comparison with model outputs.

The experiment will address the broad scientific questions: 1 How does the Earth System respond in the long term to CO₂ forcing analogous to that of the modern? and 2 What is the significance of CO₂ induced polar amplification for the stability of the ice sheets, sea-ice and sea-level?

An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*

midHolocene and lgm: evaluation w.r.t available data (systematic benchmarking, cf. Harrison et al, Climate Dynamics, 2013), both in terms of temperature and hydrological cycle. These evaluations make use of independent climate reconstructions over land and ocean. A specific focus will be put on the link with model biases and model results for future climate. Specific working groups in PMIP have been set up to improve the comparisons with marine data (COMPARE group) and isotopic data (cf. <http://pmip3.lsce.ipsl.fr/>, “working groups” tab). This provides new methodologies and new possibilities for quantitative model assessments.

past1000: In-depth analyses using novel statistical approaches (Sundberg et al., 2012; Moberg et al., 2014; Bothe et al., 2013) and detection/attribution techniques (Schurer et al., 2014).

Process-oriented analyses on variability and changes in circulation modes. Partly supported by dedicated sensitivity studies, e.g. in VolMIP.

LIG: The CMIP6 experiment will analyse the strength of feedbacks at work in the Arctic, and their potential implications for the stability of the Greenland ice sheet. A particular emphasis will be put on the annual redistribution heat by the ocean circulation and the potential role of the transmission of the subsurface warming from North Atlantic to Southern ocean, with implication for basal melting of West Antarctic ice Sheet. High latitude feedbacks from sea-ice, water vapor and cloud will be a focus, as well as the relative changes between the tropical and high latitude water cycle.

Pliocene Warm Period: The CMIP6 experiment will evaluate the ability of models to simulate a recent interval of CO₂-induced global warmth, and assess the response of critical components of the climate system to near modern CO₂ forcing in the long term (sea-ice, modes of variability, monsoons, storm tracks, vegetation). Unlike other warm periods or interglacials the Pliocene retains critical modern boundary conditions such as the continental configuration and astronomical forcing. The signal of change in Pliocene is large and therefore the signal to uncertainty ratio enables model-predicted changes to be attributed with confidence.

Some of these diagnoses and model evaluations will be performed as part of PMIP transverse analyses groups. In particular, the PMIP “Past2Future” working group aims at identifying and understanding relationships between model simulations for past and future climates and at using available paleodata to evaluate the consistency of these relationships. Its work is therefore potentially based on all PMIP simulations together with selected simulations relevant for future climate change.

Proposed timing*

Ideally past1000 should be run before the historical simulations. All other experiments can be run as soon as the reference simulation in DECK is run.

Experimental design of proposed CMIP6 experiments

➤ midHolocene and lgm

(taken from Braconnot et al, Nature Climate Change, 2012) + <http://pmip3.lsce.ipsl.fr>

Mid-Holocene (MH) and Last Glacial Maximum (LGM) simulations are equilibrium experiments, presenting a “snapshot” of climate at a specific time. Table 2 summarises the boundary conditions used for MH and LGM experiments during the various phases of the Palaeoclimate Modelling Intercomparison Project (PMIP). The ultimate external forcing (or driver) of climate is change in incoming solar radiation (insolation) as determined by changes in the Earth’s orbit. These changes can be specified precisely. Due to the slow variations of Earth’s orbital parameters, the seasonal and latitudinal distribution of MH insolation was different from present (1950 C.E), enhancing the magnitude of the seasonal contrast in the Northern Hemisphere by about 60 Wm^{-2} . Insolation forcing at the LGM was very similar to present. When models do not explicitly simulate slow processes such as the build up of ice sheets, concomitant changes in land-sea distribution, or the evolution of atmospheric composition, all of which lead to changes that have to be considered as climate forcings on shorter timescales, then these boundary conditions (hereafter forcings) have to be prescribed in the MH and LGM experiments. As models have evolved in complexity, so the set of forcings that has to be prescribed has also evolved. In the first phase of the Palaeoclimate Modelling Intercomparison Project (PMIP1), the experiments were performed with atmospheric general circulation models and the state of the ocean was prescribed as a forcing. In the second phase of PMIP (PMIP2), some models incorporated vegetation dynamics but vegetation cover and albedo still had to be specified for the coupled ocean-atmosphere general circulation models (OAGCMs). Some processes, such as those associated with the terrestrial and marine carbon cycle, have been ignored in the earlier PMIP experiments, but will be included as interactive components of some of the models used in PMIP3. In all experiments the atmospheric composition is prescribed using results from ice-cores.

The next phase of PMIP will make use of the PMIP3 boundary conditions whenever possible. A major foreseen evolution is related to the interactive computation of the dust cycle in atmospheric models, for which changes in vegetation also have to be taken into account. PMIP2 and PMIP3 recommended the use of either interactive vegetation or prescribed pre-industrial vegetation. For PMIP4, those models which include an interactive representation of the dust cycle will have to account for changes in vegetation. This particular topic will be discussed with the modelling groups.

The *lgm* experiment will be the reference from which sensitivity experiments to uncertainties in boundary conditions will be developed. In particular, the sensitivity to ice sheet reconstructions will be tested within the PMIP working group on LGM ice sheet uncertainties. These CMIP6 and PMIP4 LGM experiments will also be starting points for transient deglaciation experiments (coordinated by the working group on the deglaciation).

Table 2 : Evolution of the boundary conditions prescribed in the different phases of the PMIP project. Boundary conditions that remain the same between different sets of simulations are highlighted in yellow; blue highlighting shows boundary conditions that are not included in a given set of experiments. More details on the protocols used in PMIP3 can be found on the PMIP3 web site (see <http://pmip3.lsce.ipsl.fr/>), which also provides links to the webpages detailing the protocols used in PMIP1 and PMIP2. Note that in the MH experiment the CO₂ concentration is the pre-industrial one. CO₂ctrl refers to the CO₂ concentration of the present-day control simulation.

	PMIP1	PMIP2	PMIP3
Mid Holocene (6000 years BP)*			
*In this experiment ice-sheet, coastline, solar constant and aerosols are prescribed as in the PI simulation.			
Insolation	eccentricity = 0.018682 obliquity = 24.105° perihelion-180° = 0.87°	eccentricity = 0.018682 obliquity = 24.105° perihelion-180° = 0.87°	eccentricity = 0.018682 obliquity = 24.105° perihelion-180° = 0.87°
Trace gases	CO ₂ = 280 ppm or 280/345* CO ₂ ctrl CH ₄ = 650 ppb N ₂ O = 270 ppb CFC = 0 O ₃ = not considered	CO ₂ = 280 ppm CH ₄ = 650 ppb N ₂ O = 270 ppb CFC = 0 O ₃ = not considered	CO ₂ = 280 ppm CH ₄ = 650 ppb N ₂ O = 270 ppb CFC = 0 O ₃ = same as in CMIP5 PI
Vegetation and land surface	Prescribed to be the same as modern vegetation	Either prescribed to be the same as modern vegetation or computed using a dynamical vegetation module	Computed using a dynamical vegetation module, Or prescribed as in PI, with phenology computed for models with active carbon cycle or prescribed from data
Carbon cycle	Not considered	Not considered	Interactive, with atmospheric concentration prescribed and ocean and land carbon fluxes diagnosed as recommended in CMIP5
Last Glacial Maximum (21000 years BP) *			
* In this experiment solar constant and aerosols are prescribed as in the PI simulations.			
Insolation	eccentricity = 0.018994 obliquity = 22.949° perihelion-180° = 114.42°	eccentricity = 0.018994 obliquity = 22.949° perihelion-180° = 114.42°	eccentricity = 0.018994 obliquity = 22.949° perihelion-180° = 114.42°
Trace gases	CO ₂ = 200 ppm or (200/280) * CO ₂ ctrl CH ₄ = 350 ppb N ₂ O = 190 ppb CFC = 0 O ₃ = same as in PI	CO ₂ = 185 ppm CH ₄ = 350 ppb N ₂ O = 200 ppb CFC = 0 O ₃ = same as in PI	CO ₂ = 185 ppm CH ₄ = 350 ppb N ₂ O = 200 ppb CFC = 0 O ₃ = same as in PI
Ocean	SST prescribed from CLIMAP (1981) Or SST computed using a slab ocean model	3D Ocean model and sea-ice	3D ocean model and sea-ice

Ice sheet	ICE-4G (Peltier et al, 1994)	ICE-5G (Peltier et al, 2004)	PMIP3 Blended ice sheet
Land-sea mask	-105 m sea level	Prescribed following Peltier (2004) land-sea mask -120 m	Prescribed from the blended ice-sheet land-sea mask. Sea-level change consistent with the change in land-sea mask.
Freshwater		Excess LGM freshwater added to the ocean in 3 different regions	Excess LGM freshwater added to the ocean in 3 different regions
Ice sheet ice streams	Not considered	Not considered	Not considered
River runoff	Not considered	As in CTRL or river pathway modified	As in PI or river pathway modifier according to PMIP protocol
Mean ocean salinity	Not considered	Not considered	+1 PSU everywhere
Carbon cycle	Not considered	Not considered	Interactive, with atmospheric concentration prescribed and ocean and land carbon fluxes diagnosed as recommended in CMIP5 For PCMIP: fully interactive with atmospheric concentration computed by the model

➤ Last Millennium

Updated PMIP3 protocol (<http://pmip3.lsce.ipsl.fr>) based on Schmidt et al (2011, 2012):

Schmidt, G. A. et al. (2011). Climate forcing reconstructions for use in PMIP simulations of the last millennium (v1.0), *Geosci. Model Dev.*, 4, 33–45.

Schmidt, G. A. et al. (2012). Climate forcing reconstructions for use in PMIP simulations of the Last Millennium (v1.1), *Geosci. Model Dev.*, 5, 185–191

Transient simulations 850-1849 followed by historical experiments, set of boundary conditions for solar, volcanic, land-cover-change, greenhouse gases to be blended with those for historical (1850-2010) simulations. The continuity between the past1000 and historical scenarios has to be improved and fully discussed within CMIP6.

➤ Last Interglacial

Based on the protocol discussed within PMIP3.

For CMIP6, we propose to perform a simulation for the 128ky BP time slice - large orbital forcing (Figure 1), large responses.

- Orbital parameters set to 128ka.
- Greenhouse gas concentrations well-known from ice cores [CO_2 275ppm; CH_4 709ppb; N_2O 266 ppb].
- modern geography, ice sheets, and vegetation;
- Initialize from CMIP6 pre-industrial DECK simulation;
- length: ≥ 100 years after spinup.

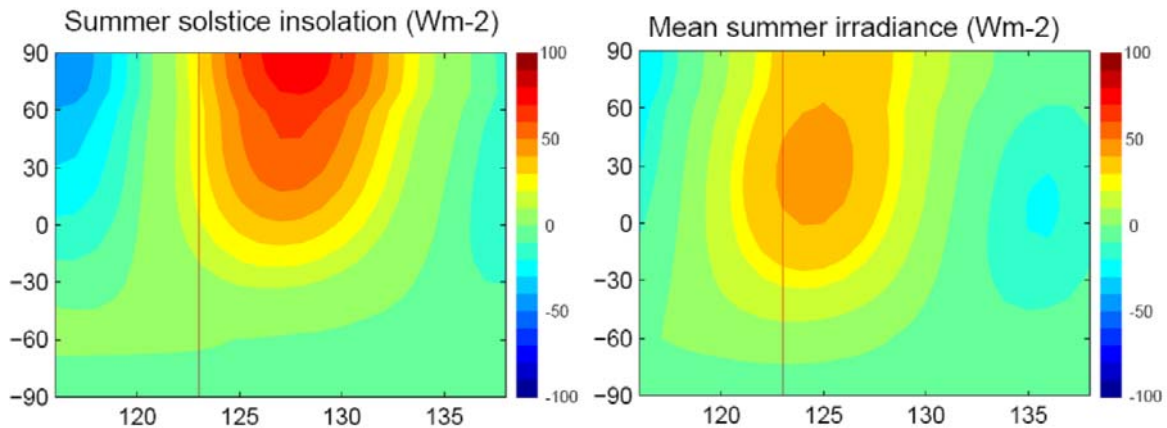


Figure 1. Anomalies of summer solstice insolation (left) and mean summer irradiance (right) as compared to present. Left axes are latitude and bottom axes are in thousands of years before present.

This simulation will constitute a reference for PMIP LIG simulations: other snapshots within the last interglacial (125, 122 ky BP), transient simulations for the whole interglacial. This will also be a target period for testing AOGCMs coupled with polar ice-sheet models, as proposed in ISMIP. Discussions are on-going with ISMIP are on-going in order to coordinate a LIG experiment with them.

➤ Pliocene warmth

Time slice equilibrium climate experiment modifying CO_2 (to 400 ppmv), topography, ice sheet extent and running with dynamic vegetation.

Updated from PlioMIP experiment 2 (Haywood et al, GMD, 4, 571-577, 2011), under discussion for minimum changes in boundary conditions w.r.t. pre-industrial.

As for the other proposed CMIP6 experiments, this Pliocene experiment is the basis for a full range of experiments coordinated within PMIP by the PlioMIP working group. In particular, the sensitivity of the results to insolation, ice sheet configuration and other boundary conditions will be investigated.

Science question and/or gap addressed with the PMIP experiments

Cf. introduction and summary excel table.

New foci for analyses will be:

- Forced vs. internal variability, putting in context climate changes in the industrial historical period
- Clouds/Circulation: WCRP Grand challenge Initiative on Leveraging the past record (<http://www.wcrp-climate.org/index.php/gc-clouds-circulation-activities/gc4-clouds-initiatives/116-gc-clouds-initiative4>)
- Analyses of cryospheric feedbacks under natural forcings (transient simulations over the last millennium put in perspective the recent changes e.g. in Arctic Sea ice, coupling between ice-sheets and climate (*lgm*, LIG, Pliocene)
- Regional climate and decadal predictions:
 - Improved assessment of decadal to centennial variability as carrier of near-term prediction potential (*past1000*, *midHolocene*, *lgm*).
 - Regional assessment of response to natural forcing and interaction with variability modes and teleconnections (all experiments)
- Assessment of extremes under natural forcing, e.g. volcanoes. Natural variations in droughts in connection with paleo-reconstructions (*past1000*). Analyses of mechanisms of mega-droughts (*midHolocene*). → link with Grand challenges on extremes and on water availability.

Possible synergies with other MIPs

PMIP simulations can serve to interact with other MIPs on the following themes:

- CF-MIP (cloud feedbacks): dedicated common idealized sensitivity experiments to be run in aquaplanet set up: AMIP simulations with SSTs minus 4K, abrupt0.5xCO₂, abrupt solar perturbation experiments, to be co-analysed in CF-MIP and PMIP.
- OCEAN/SEA_ICE: Mutual assessment of the role of the ocean in low-frequency variability, e.g. multi-decadal changes in ocean heat content or heat transport. Provide initial conditions for the ocean including long-term forcing history.
- CARBON CYCLE (C4MIP): Assessment of carbon-cycle evolution and feedbacks between sub-components of the Earth System. Evaluation of paleo reconstructions of carbon storage.
- LAND USE: Links should be reinforced for better connecting past1000 to historical simulations. Useful for analysis of past1000 simulations, for biophysical as well as carbon cycle aspects.
- VolMIP (volcanic forcing): analysis of specific volcanic events very useful for critical analysis of past1000 simulations. VolMIP would systematically assess uncertainties in the climate response to volcanic forcing, whereas LM simulations describe the climate response to volcanic forcing in long transient simulations where related uncertainties are due to chosen input data for volcanic forcing: mutual assessment of forced response.
- DETECTION/ATTRIBUTION: long millennium simulations can be very useful for this topic.

Potential benefits of the experiments

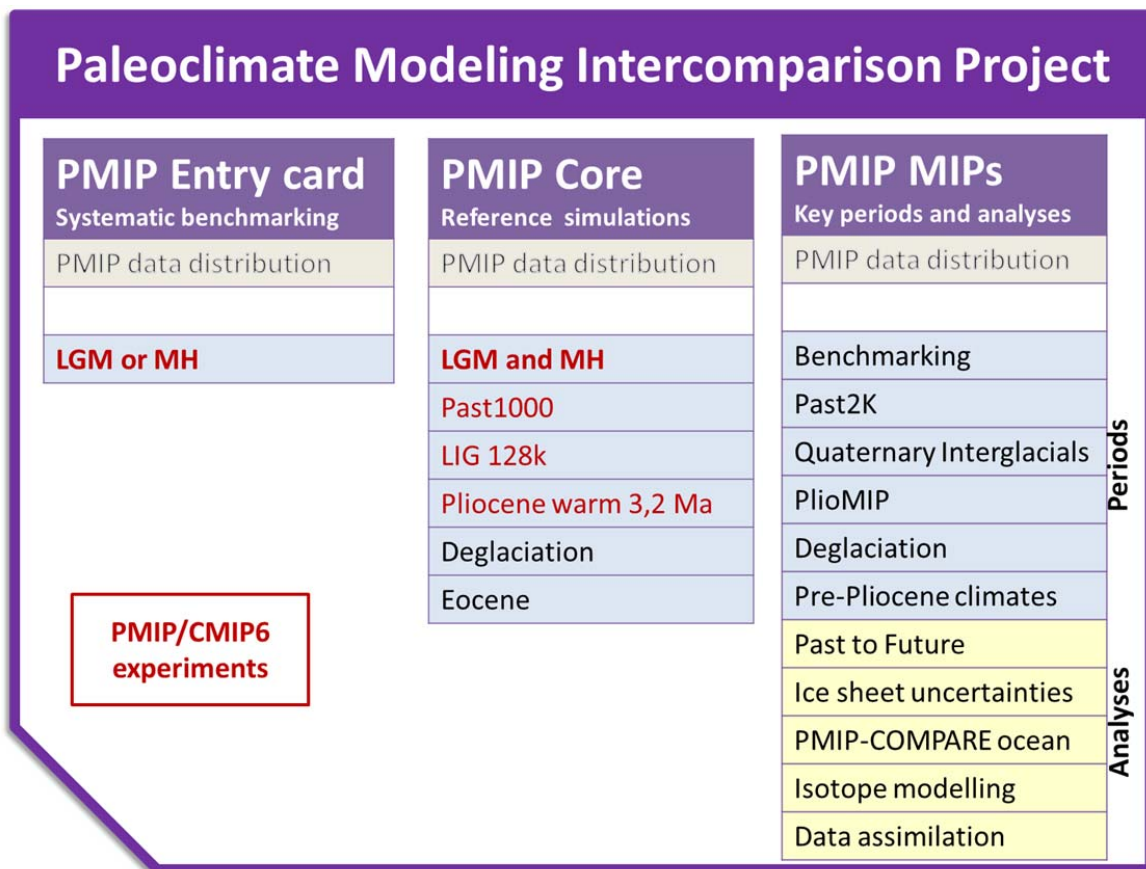
Potential benefits of the experiments to

- (A) climate modeling community
 - Improved assessment of forced response and forced vs. internal variability
 - Improved knowledge on which processes are important in the forced response to natural forcing (e.g. ozone changes owing to solar radiation changes for the past1000 experiment)
- (B) Integrated Assessment Modelling (IAM) community,
- (C) Impacts Adaptation and Vulnerability (IAV) community
 - All experiments:
 - Identification of thresholds for ecosystems and water availability under different climate conditions
 - Improved assessment of natural variability including extreme events under pre-industrial boundary conditions. Identification of regions where, under natural forcing, changes, changes lead to specific vulnerability (e.g. regional sea-level)
- , and (D) policy makers.
 - Quantification of magnitude and speed of a range of past climatic changes compared to the natural variability and recent and future climatic trends. Impact of these changes on water availability and ecosystems.

Prioritization of the proposed experiments

- If possible, a prioritization of the suggested experiments, including any rationale**

Each proposed PMIP experiment for CMIP6 can be run independently, because they focus on different time periods. The *midHolocene* and *lgm* experiments have been the focus of PMIP since its start and allow for an evaluation of new model versions since the first atmosphere-only GCMs in PMIP1. We therefore require one of these two simulations to be performed as an entry card to CMIP6-PMIP4 experiments. All five PMIP experiments proposed for CMIP6 have equal priority, each experiment being the core of a set of sensitivity experiments to be run within PMIP. The organization of the PMIP experiments w.r.t CMIP6 is given in the figure below.



Model output

- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**

PMIP (all experiments): no objections

- List of output and process diagnostics for the CMIP DECK/CMIP6 data request**
 - whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;

PMIP (all experiments): same set of CMOR variables as historical/scenario (possibly reduced set of high-frequency output owing to length of experiment), some simulations with COSP simulator (subset of years)

- whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);

PMIP all-experiments: diagnostics needed for ESM (i.e all components of the ESM + forcings + feedback analyses) + tracers and isotopes when available (list to be established)

- whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.

PMIP all periods: subset of variables for driving regional climate models, ice-sheet models (ISMIP) or ecological models (land surface variables) or dust models.

PMIP *past1000* and *midHolocene*: subset of variables for investigating extreme events or variability

- whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;

PMIP: same as for CMOR variables from historical/scenario;

- the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.

See previous PMIP requests (CMIP5 or PMIP3 ESGF): same set of CMOR variables as historical/scenario (possibly reduced set of high-frequency output owing to length of experiment)

- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms.**

Needs to discussed with all MIPs

- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF.**

PMIP benefits from two entries on ESGF: via the CMIP5 Project for PMIP3-CMIP5 experiments or the PMIP3 project for other PMIP3 experiments or for groups which do not take part in CMIP5. It would be very convenient to still be able to search through both (or indeed multi-MIP) data bases on the same system, as can be done now.

Proposed contributions for model diagnostics and evaluation

- Any proposed contributions and recommendations for**

- model diagnostics and performance metrics for model evaluation;

past1000: use diagnostics that have been defined for DECK historical/scenario simulations. In addition to integrated quantities such as hemispheric temperature averages, *past1000* experiments will increasingly be analysed w.r.t. circulation regimes, extreme events etc.

midHolocene, *lgm*: PMIP specific diagnostics have been developed for benchmarking. A working group is dedicated to this topic. cf. Harrison et al (2014), *Climate Dynamics*, 43, 671–688

- observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;

past1000 simulations will benefit from observations to be extended to the early 19th century. *past1000* simulations will be compared, mutually analysed with paleo reconstructions, most importantly the growing set of PAGES2K reconstructions that are available through Paleodata data bases.

PMIP data syntheses for *midHolocene* and *lgm* (<http://pmip3.lsce.ipsl.fr/synth/>)

new syntheses will be available for characterizing high resolution variability during the Holocene (paleoVar PMIP working group)

- tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).

Common analyses scripts are being discussed within PMIP.

For *past1000*: in the framework of the PMIP working group Past2K advanced statistical analyses and evaluation tools have been developed (e.g. Bothe et al., 2013; Moberg et al., 2014).

Expression of interest from modelling groups

On Nov 28th 2014, PMIP has received the expression of interest from 9 modelling groups for the LIG experiment, 10 modelling groups for the *midHolocene*, *lgm* and *past1000* experiments and 11 modelling groups for the Pliocene Warm Period experiment.

Radiative Forcing Model Intercomparison Project (RFMIP)

Application for CMIP6-Endorsed MIPs

Date: 2 December 2014

➤ Name of MIP*

Radiative Forcing Model Intercomparison Project (RFMIP)

➤ Co-chairs of MIP (including email-addresses)*

Robert Pincus, University of Colorado, US; Robert.Pincus@colorado.edu

Piers Forster, University of Leeds, UK; P.M.Forster@leeds.ac.uk

Bjorn Stevens, Max Planck Institute for Meteorology, Germany; Bjorn.Stevens@mpimet.mpg.de

➤ Members of the Scientific Steering Committee*

Viktor Brovkin, Max Planck Institute for Meteorology, Germany (representing LUMIP)

Gunnar Myhre, CICERO, Norway (representing AerChemMIP)

Hideo Shiogama, National Institute for Environmental Studies, Japan (representing DAMIP)

Karl Taylor, Program for Climate Model Diagnosis and Intercomparison, US (general expertise)

Jean-Louis Dufresne, LMD/IPSL, France (representing IPSL)

James Manners, UK Met Office, UK (representing UKMO)

Miho Sekiguchi, Tokyo University of Marine Science and Technology, Japan (representing MIROC)

➤ Link to website (if available)*

<http://www.wcrp-climate.org/modelling-wgcm-mip-catalogue/modelling-wgcm-mips/418-wgcm-rfmip>,

though this is out of date

➤ Goal of the MIP and a brief overview*

RFMIP aims to understand the radiative forcing to which models are subject. The project will assess the accuracy of instantaneous radiative forcing calculations for greenhouse gases and aerosols in each model by comparing these to reference calculations across a range of states representative of present-day, past, and future climates. We will increase the accuracy and spatial detail with which effective radiative forcing is known for each model and for each DECK or other experiment by requesting and analyzing matching simulations designed for this purpose, carefully diagnosing the degree to which the diversity in effective radiative forcing is due to variations in rapid adjustments, radiative forcing and climatological base state. We will close the circle by requesting historical-to-near-future simulations in which anthropogenic aerosol optical and cloud-active properties are tightly controlled, allowing us to determine which aspects of the observed historical record consistently emerge and so can be attributed to aerosol forcing.

The project is aligned with CMIP6 criteria as follows:

The MIP addresses at least one of the key science questions of CMIP6

One of the guiding questions for CMIP6 centers on models' response to forcing; it is not possible to answer this question without the ability to quantify the forcing precisely. RFMIP is central to the

question of response to forcing and relevant for the other two CMIP questions. Accurate diagnosis of forcing and its errors in CMIP models is key to understanding the spread of models response across simulations and, for example, understanding any possible model bias. Interpreting projections also requires an understanding of the forcings to which models are subject and understanding the degree to which the treatment of radiative transfer introduces errors is central to evaluating a key source of model biases.

The MIP follows CMIP standards in terms of experimental design, data format and documentation

Model integration requests for calculations of effective radiative forcing follow CMIP protocols.

Integrations with prescribed aerosol properties may require small code changes to existing models but will largely follow a protocol that has been already implemented and tested by a number of modeling centers through the Easy Aerosol project (<http://www.wcrp-climate.org/index.php/gc-clouds-circulation-activities/gc4-clouds-initiatives/368-gc-clouds-initiative3-easy-aerosol>).

The data request for the assessment of aerosol optical properties is unusual in requesting spectrally-detailed information at a few snapshots, but otherwise is a simple request for diagnostic information.

Requests for offline radiative transfer calculations will build on the “site” infrastructure. Reference line-by-line calculations will be performed by RFMIP, formatted to comply with CMIP conventions, and made available to participating modeling centres on the ESG.

A sufficient number of modeling groups have agreed to participate in the MIP

These requests were developed through preliminary email discussions and a dedicated workshop (Hamburg, Sept, 2014). Input has been solicited from climate modeling groups (HadGEM, GFDL, CCCM, CanESM, ECHAM, IPSL, NorESM, MIROC) and the radiative forcing community. The second two phases of RFMIP have been developed in consultation with the EC-Earth and CNRM communities.

Eight modeling groups have agreed to contribute integrations needed to estimate effective radiative forcing (NCAR, GFDL, CCCma, MIROC, MPI, MOHC, NCC, GISS).

An EU proposal has been submitted that would fund five modeling centers (MPI, IPSL, UKMO, KNMI, CNRM) to participate in the prescribed-aerosol simulations. The prototype Easy Aerosol project has attracted further participation by CMA, NCAR, GFDL, and the NorESM consortium.

The following modeling groups have agreed to participate in off-line radiative transfer calculations: GFDL, UKMO, CCCma, the MIROC consortium, RRTMG. The last contribution comes from the developers of a widely-used radiation parameterization used in variants of CAM/CESM and by the ECWMF. Combined with models derived from the UKMO Hadley Centre this list constitutes all but a few modeling centers.

The MIP builds on the shared CMIP DECK experiments

The project will precisely quantify forcing in 4xCO₂ and historical integrations, It may lead to decreased spread in forcing by uncovering errors in radiative transfer and will, by design, narrow the spread in forcing for one set of experiments.

A commitment to contribute to the creation of the CMIP6 data request and to analyze the data

Data requests have been drafted. The US DoE has funded two of the coordinators (Pincus, Forster) to perform part of the analyses. The same proposal will support reference radiative transfer calculations (Mlawer, Collins, Ramaswamy). An EU proposal supporting the participation of five modeling centers in the prescribed-aerosols component has been submitted (Stevens, Forster). Several groups have committed to the analyses (CICERO/UiO, Leeds, UKMO, GFDL, Colorado, Berkley, PCMDI, MPI, LMD/CNRS, GISS).

A commitment to identify observations needed for model evaluation and improved process understanding, and to contribute directly or indirectly to making such datasets available as part of obs4MIPs

Reference calculations by line-by-line models are the radiative forcing equivalent of observations because they have been extensively tested against well-calibrated, spectrally detailed observations. These reference calculations will be made available on the ESG so models can run ongoing tests on their radiative transfer.

Observations of clear-sky flux will constrain radiative forcing for the prescribed aerosol runs of RFMIP and are already distributed through the ESG.

The proposed experiment is of central importance to CMIP6

Without quantifying forcings the community will not be able to understand the spread of model responses.

The proposed experiment has been run at least by two modeling groups already

Fixed SST simulations were performed by nine modeling groups in CMIP5.

The Easy Aerosol methodology has been tested by eight modeling centers (MPI, UKMO, CAM, NCAR, GFDL, IPSL, CICERO).

A pilot assessment of radiative parameterization accuracy using offline calculations under $4\times\text{CO}_2$ conditions attracted 6 GCM codes (representing roughly 16 modeling centers) and 2 reference codes. RFMIP differs from this prototype only in scale.

The proposed experiment is useful in a multi-model context and to a number of climate researchers.

Our focus is on understanding the multi-model spread in response, especially the degree to which this is due to spread in forcing as opposed to spread in sensitivity. This is by definition essential for establishing origins of systematic biases. We expect a very wide community to take full advantage of the data we request and the reference calculations we make. We have support from other proposed MIPs (e.g. AerChemMIP, LUMIP, PDRMIP, VolMIP, ScenarioMIP).

A commitment to scientifically analyze, evaluate and exploit the proposed experiment.

The PIs and their groups are very active publishers of CMIP and related model results. We expect the results of the project to make important contributions to the refereed literature.

➤ References (if available)*

Collins, W. D., and Coauthors, 2006: Radiative forcing by well-mixed greenhouse gases: Estimates from climate models in the Intergovernmental Panel on Climate Change (IPCC) Fourth Assessment Report (AR4). *J. Geophys. Res.*, **111**, D14317, doi:10.1029/2005JD006713.

Shindell, D. T., and Coauthors, 2013: Radiative forcing in the ACCMIP historical and future climate simulations. *Atmos. Chem. Phys.*, **13**, 2939–2974, doi:10.5194/acp-13-2939-2013.

Forster, P. M., T. Andrews, P. Good, J. M. Gregory, L. S. Jackson, and M. Zelinka, 2013: Evaluating adjusted forcing and model spread for historical and future scenarios in the CMIP5 generation of climate models. *J. Geophys. Res.*, **118**, 1139–1150, doi:10.1002/jgrd.50174.

➤ An overview of the proposed experiments*

RFMIP-IRF-AER: We will request *detailed diagnostic information* on the vertically- and spectrally-resolved optical properties of aerosols and the surface, along with detailed information about atmospheric physical and chemical state, for snapshots at present-day and pre-industrial conditions. These will be paired with requests for model computations of clear-sky instantaneous radiative forcing. These will allow the diversity in relationships between burden and optical properties to be quantified and will provide data required for accuracy assessments.

RFMIP-IRF-GHG: We will request calculations of vertically-resolved broadband-integrated longwave and shortwave fluxes made with *off-line radiative transfer models* identical to the model's online version, using specified atmospheric states (distribution of temperature and humidity) and surface properties over many profiles. A series of such calculations will assess the accuracy of radiative transfer approximations for gases under various conditions.

RFMIP-ERF: We will request a series of 30-year *uncoupled (atmosphere+land only) simulations with prescribed sea-surface temperatures* and sea-ice concentrations from the preindustrial AOGCM control simulation of the model. Each simulation is matched to an existing simulation from the DECK or some other MIP and will enable the accurate diagnosis of effective radiative forcing for this experiment. Direct radiative forcing and rapid adjustments can be analyzed using newly developed kernel methodologies. Longer runs with fixed sea-surface temperature, paired with the proposed DAMIP AOGCM historical and RCP8.5 simulations, will allow us to calculate transient radiative forcings.

RFMIP-Historical: We will request *small ensembles of “historical+RCP8.5” coupled runs* from 1850 to 2020 in which spectrally-dependent anthropogenic aerosol optical properties and cloud interactions are directly prescribed. Three types of simulations are being requested: historical simulations with all forcings, historical simulations with only natural forcings (here the DAMIP simulations will be used) and, at lower priority, historical simulations with only non-GHG forcing. The time-dependent spatially-patterned prescription will be based on a hybrid model-observation climatology. Such an approach has already been developed and implemented within the MPI-ESM. A non-stationarity factor will account for day-to-day variations. These simulations will be used to assess which aspects of the historical period emerge robustly from ensembles in which aerosol direct and indirect effects are constrained.

➤ An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*

We will perform no direct analysis of DECK or other experiments in isolation.

The assessment of radiative parameterization accuracy (RFMIP-GHG, RFMIP-AER) is applicable to most DECK experiments, but especially the AMIP, coupled historical, and 4xCO₂ experiments.

Thirty year fixed SST/sea-ice simulations (RFMIP-ERF) will allow us to determine accurate regional variations of effective radiative forcing, rapid adjustments and direct radiative forcing for the DECK's coupled historical, and 4xCO₂ experiments, and for simulations requested by the Detection and

Attribution MIP (anthropogenic, well-mixed greenhouse gases, aerosols+ozone, natural) and Land Use MIP (land use).

Simulations with prescribed anthropogenic aerosol optical properties (RFMIP-Historical) will provide a useful complement to historical coupled simulations present in the DECK and will complement activities under the Detection and Attribution MIP. They will be used to evaluate the hypothesis that a present day aerosol forcing stronger than -1 W/m^2 is incompatible with the temperature record prior to 1950.

➤ **Proposed timing***

Requests for simulations required to diagnose effective radiative forcing, including specification of the diagnostic fields are required, are available now.

Detailed requests (CMOR tables) for aerosol diagnostic outputs have been drafted.

Atmospheric specifications for off-line radiation calculations will be available by summer 2015.

Prescriptions of aerosol optical properties to be used in specified-aerosol simulations will be available by the end of 2015.

➤ For each proposed experiment to be included in CMIP6**

- the experimental design;
- the science question and/or gap being addressed with this experiment;
- possible synergies with other MIPs;
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

➤ If possible, a prioritization of the suggested experiments, including any rationale**

Aerosol diagnostic requests for RFMIP-IRF-AER

These requests will allow for complete characterization of the diversity in aerosol properties that contribute to the diversity in aerosol direct radiative forcing. They will also provide the basis for the sampling of aerosol optical properties needed to assess radiative transfer errors in the project's next stage.

TIER 1

These requests are unusual in that they require three-dimensional fields at a few instants, and that the spectral detail is requested.

Snapshots of aerosol optical properties (optical thickness, single-scattering albedo, asymmetry parameter) on the model grid (including the spectral grid), along with surface conditions (spectral albedo, temperature), top-of-atmosphere insolation, and fields of temperature, pressure, and water vapor content or relative humidity.

1. Present-day (2015)
2. Pre-industrial (1850)

Offline radiative transfer calculations for RFMIP-IRF-GHG

Offline radiative transfer calculations are intended to identify the degree to which errors in radiative transfer contribute to diversity in estimates of instantaneous radiative forcing, and how this error depends on the background state of the atmosphere and on the forcing itself.

These calculations require an investment in infrastructure (the creation or adaptation of an off-line radiative transfer code traceable to the on-line version) but are a very small computational burden.

TIER 1

These calculations provide baseline error estimates at present-day (PD, 2015) and pre-industrial (PD, 1850) conditions, for the forcing used to assess equilibrium climate sensitivity ($4xCO_2$) and under conditions used to assess rapid adjustments (PD+4K). We request the following clear sky calculations for a set of prescribed atmospheric conditions.

1. PD greenhouse gas concentrations with PD atmospheres (aerosol-free)
2. PI greenhouse gas concentrations with PD atmospheres (aerosol-free)
3. $4 \times$ PI CO_2 , other gases at PD concentrations, in PD atmospheres (aerosol-free)
4. PD greenhouse gas concentrations in PD+4K conditions, assuming constant relative humidity and using a vertical shift transform to map surface warming to atmospheric structure (aerosol-free)
5. “Future” combining simultaneous increases in CO_2 , temperature, and humidity (aerosol free)

TIER 2a

These calculations explore the accuracy of CO_2 forcing across a range of concentrations relevant to past values (relevant for Last Glacial Maximum calculations for PMIP) and future concentrations. All use aerosol-free clear-sky PD atmospheric conditions with CO_2 concentrations drawn from {0.25, 0.5, 2, 3, 8} times the PI value.

TIER 2b

These calculations explore the error in radiative forcing estimates from different well-mixed greenhouse gases. Calculations are for aerosol-free clear-sky conditions using PD atmospheres and greenhouse gas concentrations with one (or all) set to its PI value:

1. CH_4
2. N_2O
3. CO
4. HC
5. O_3
6. All

Model integrations to diagnose effective radiative forcing for RFMIP-ERF

Model integration requests are designed to give consistent radiative forcings in both emission-based and prescribed-concentration frameworks and we recommend that CMIP6 standardizes to this methodology for computing radiative forcings across all MIPs and that other MIPs link to RFMIP as far as possible. RFMIP concentrates on understanding forcing changes from preindustrial (PI, 1850) to present day (PD, 2015).

The protocol for time-slice experiments is a single 30-year uncoupled simulation in which sea surface temperature (SST) and sea ice distributions are specified and vegetation may interact. Sea ice and anthropogenic forcing agents are specified at PI values unless noted. Tests of this method demonstrate that 30-year integrations constrain broad regional patterns of forcing to better than 0.1 Wm^{-2} .

TIER 1

These integrations will allow us to quantify the radiative forcing at present day. We request the following integrations:

- a) 30 year PI control with monthly averaged fixed SST and sea-ice climatology from the PI AOGCM control integration. To be used as control for all other integrations
- b) As a) with $4x\text{CO}_2$
- c) As a) with PD ANTHROPOGENIC forcings
- d) As a) with PD WMGHGs (+indirect effects in emission based models)
- e) As a) with PD AEROSOLS and OZONE (linking to additional forcing estimates from AerChemMIP)
- f) As a) with PD LAND USE (surface albedo/roughness, transpiration). Change vegetation but fix GHG concentrations and preindustrial levels. (linking to LUMIP)

TIER2a

Non-linearity of aerosol cloud interactions are particularly important for understanding historical forcing evolution (Carslaw et al., 2013). These two experiments are designed to quantify this across the models.

- g) As a) with PD AEROSOLS and OZONE x 0.1
- h) As a) with PD AEROSOLS and OZONE x 2.0

TIER2b

Forcings at periods other than present are important for understanding aspects of historical and future change. Different applications will likely require different time-slices. It is also difficult to evaluate transitory volcanic and solar forcing using a time-slice methodology of Tier 1. Therefore we propose to evaluate forcings over the full 1850-2100 period, concentrating on natural forcings not evaluated in Tier 1.

Transient forcings from CMIP5 models relied on a crude two-step residual method to diagnose globally averaged ERF estimates (Forster and Taylor, 2006; Forster et al., 2013). This method assumed that a CO_2 -based climate sensitivity was applicable across scenarios and the method had significant errors due to noise. The two-step method is a possible fall-back for CMIP6.

The two-step method can be improved on by using transient climate model integrations with fixed SST and sea ice (PD ctrl). Preliminary work with HadGEM2 indicates that three ensemble members with this method would be needed to constrain interannual variation on global forcing to better than 0.1 Wm^{-2} . Further work is needed to verify the number of accuracy of different ensemble sizes.

We will request small ensembles of 1850-2100 fixed SST and sea ice climate integrations (including interactive vegetation). These requests are matches of DAMIP AOGCM requests to have pairs of consistent forcing and response integrations, maximizing the science benefit.

- i) Historical + RCP8.5 (all forcing integration) with forcings added to PI control from a)
- j) Natural (solar and volcanoes) forcings added to a), only to 2015
- k) Aerosols (and their indirect effects) added to a),
- l) WMGHG changes added to a)

TIER 3

Transient WMGHG forcing can likely be inferred from tier 1 results. However, a transient AOGCM WMGHG run is proposed by DAMIP and it is useful to check its forcing in a few models.

Model integrations with specified aerosol properties for RFMIP-Historical

Here we request integrations in which aerosol optical properties including cloud-radiation interactions are prescribed. The prescription is expected to be a single analytically-described spatial pattern with time-constant spectral variation (building on the Easy Aerosols experience) and time-varying strength. Our goal is to determine which aspects of the historical record robustly emerge from ensembles in which the radiative forcing by anthropogenic aerosols is tightly constrained.

TIER 1

A 4-member ensemble of coupled simulations from 1850-2020 (merging the historical period and RCP 8.5 or an appropriate scenario for GHGs for the 2015-2020 period) with all forcings including specified aerosols.

TIER 2

A set of 4-ensembles to be used in detection and attribution experiments

1. Hist-Nat, 1850-2020 (joint with DAMIP)
2. Hist-Aer 1850-2020 (with prescribed aerosol)
3. Corresponding prescribed SST ERF runs (see part III) to diagnose aerosol forcing

TIER 3

1. AMIP, 1980-2020 (with prescribed aerosol)
2. Single AMIP, 1980-2020, nudged to observed winds (with prescribed aerosol)

- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**
- List of output and process diagnostics for the CMIP DECK/CMIP6 data request***
 - whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;
 - whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);
 - whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.

- whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;
- the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.
- Any proposed contributions and recommendations for**
 - model diagnostics and performance metrics for model evaluation;
 - observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;
 - tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).
- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms.**

Aerosol diagnostics will require the addition of a model-dependent spectral dimension and the addition of roughly fifteen variables describing the spectrally-dependent characteristics of aerosols, surface properties, and top-of-atmosphere solar insolation.
- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF.**

Scenario Model Intercomparison Project (ScenarioMIP)

Application for CMIP6-Endorsed MIPs

Date: 2 December 2014

Proposals from MIPs should include the following information:

- * *preliminary information that will be used to determine whether a MIP should be endorsed for CMIP6 or not.*
- ** *information that must be provided later (and before the panel can determine which experiments, if any, will be incorporated in the official CMIP6 suite).*

➤ Name of MIP*

ScenarioMIP

➤ Co-chairs of MIP (including email-addresses)*

Brian O'Neill (boneill@ucar.edu), Claudia Tebaldi (tebaldi@ucar.edu), Detlef van Vuuren (detlef.vanvuuren@pbl.nl)

➤ Members of the Scientific Steering Committee*

Veronika Eyring (DLR, Germany), Pierre Friedlingstein (U of Exeter, UK); George Hurtt (U of Maryland, USA); Reto Knutti (ETH, Switzerland); Jean-Francois Lamarque (NCAR, USA); Jason Lowe (MetOffice, UK); Jerry Meehl (NCAR, USA); Richard Moss (Joint Global Change Research Institute, USA); Ben Sanderson (NCAR, USA)

➤ Link to website (if available)*

<https://www2.cgd.ucar.edu/research/mips/scenario-mip>

➤ Goal of the MIP and a brief overview*

Overall objectives

The goal of ScenarioMIP is to simulate future climate outcomes based on alternative plausible future scenarios in order to:

- (1) Facilitate integrated research leading to a better understanding not only of the physical consequences of these scenarios on the climate system, but also of the climate impact on societies, including considerations of mitigation of and adaptation to climate change. ScenarioMIP will be the main provider of new climate information for plausible future scenarios that will facilitate integrated research across multiple communities including the (1) climate science, (2) integrated assessment modeling (IAM) and mitigation, and (3) impacts, adaptation and vulnerability (IAV) communities.
- (2) Provide a basis for addressing targeted science questions regarding the climate effects of particular aspects of forcing relevant to scenario-based research, e.g., the effect of different assumptions in near-term climate forcings (NTCFs, namely tropospheric aerosols, ozone and methane) and land use on climate change and impacts.
- (3) Provide a basis for various international efforts that target improved methods to quantify projection uncertainties based on multi model ensembles, taking into account model performance, model dependence and observational uncertainty. This builds on the DECK experiments and the CMIP6 Historical Simulation and allows for the quantification of uncertainties on different timescales.

The first objective on “integration” is considered to be the highest priority for the following reasons:

- Scenarios for integration serve a large scientific audience, underpinning hundreds of scenario-based studies addressing a wide variety of scientific questions regarding physical climate changes, mitigation, impacts, and adaptation. Having common climate and socioeconomic scenarios serves as a critical means to hold key factors constant across a wide variety of studies, allowing synthetic conclusions to be drawn that would not be possible from a variety of uncoordinated studies.
- Climate simulations based on such broad-use scenarios are critical elements of the new scenario process established at Nordwijkerhout in 2007 (Moss et al., 2007; 2010); without climate simulations to support integrated studies that draw on both climate and societal futures, the scenario process cannot function. CMIP5 simulations will continue to underpin this process through 2020, and CMIP6 scenarios are seen as a critical continuation of that contribution.
- Scenarios for integration serve as a key means for connecting assessments in IPCC WGs 1, 2, and 3, as well as in the Synthesis Report. WG1 assesses the climate implications of scenarios; WG2 assesses the impact consequences of those same scenarios; WG3 assesses the mitigation required to achieve those same scenario outcomes.
- Scenarios for integration provide information on alternative climate and societal futures that thus does not need to be generated independently for each individual study.

Because targeted questions regarding the climate effects of land use, aerosols, and radiative forcing overshoot pathways are also very important to scenario-based research, a set of variants of the scenarios proposed here are being proposed in other MIPs (see below) to address these targeted questions. Thus the scenarios in ScenarioMIP serve not only the function of integration across research communities, but also serve as anchoring scenarios from which variants are designed to address targeted questions in AerChemMIP, C4MIP, DAMIP, GeoMIP, ISMIP6, LUMIP, and RFMIP.

Background

A ScenarioMIP Scientific Steering Committee (SSC) was formed following the October 2013 WGCM17 meeting as an outcome of earlier discussions among the IAM, IAV and climate modeling communities at the annual meeting of the integrated assessment and impacts communities in Snowmass, CO, in July 2013, and the AGCI session on CMIP6 in Aspen, CO, in August 2013. The ScenarioMIP SSC systematically investigated a number of issues that could substantially influence the experimental design, including the possibility of statistically sampling climate model-scenario combinations, the potential for pattern scaling or other statistical emulators of climate model output to meet some of the demand for scenario-based climate information,¹ and the differences between scenarios (in terms of global average forcing or temperature change) that is required to produce climate outcomes that are significantly different at the grid-cell level. Conclusions of these investigations were that a sparse statistical sampling approach to design was unworkable, that pattern scaling has not yet been demonstrated to be able to reliably replace the need for climate model simulations to generate information for impact studies (although it might play a limited role for some applications), and that scenario differences of at least 0.3 C (approximately 0.75 W/m²) are likely necessary to generate statistically significant differences in local climate outcomes over a substantial fraction of the surface.

Informed by these conclusions, the a proposal was prepared by the SSC after close interaction with the IAM and IAV communities through follow-up meetings at Snowmass and Aspen in summer 2014, discussions with representatives of IAM groups producing candidate scenarios for CMIP6, and

¹ A three-day workshop on pattern scaling was organized by ScenarioMIP co-chairs and others in April 2014 to address this question, see <https://www2.image.ucar.edu/event/PS2014>.

discussions with key individuals in other relevant research communities, including through the International Committee On New Integrated Climate change assessment Scenarios (ICONICS) and the WCRP-IPCC WG1 meeting in Bern, Switzerland, in September 2014. That proposal was submitted to and discussed at the October 2014 WGCM18 meeting, and has since been revised to reflect feedback at that meeting, additional coordination with other MIP proposals, and feedback from a presentation of the proposal at the annual meeting of the Integrated Assessment Modeling Consortium (IAMC) in November.

➤ References (if available)*

➤ An overview of the proposed experiments*

Scientific questions

The scientific questions addressed by the experiments proposed in ScenarioMIP fall under two of the three broad questions of interest to CMIP6, and also address the CMIP6 themes based on the WCRP grand challenges:

1. **How does the Earth system respond to forcing?** Scenarios for integration and targeted scenarios will address variants of this question as follows:
 - How does the Earth system respond to forcing pathways relevant to IAM and IAV research and to policy considerations?
 - What is the uncertainty in global and regional climate change due to *plausible* variations in future **land use** and **NTCFs** emissions, and how does it compare to multi-model uncertainty in the response to a given forcing pathway?
 - How much do plausible alternative shapes of forcing pathways (e.g. **overshoot**) matter to climate change outcomes, and therefore to questions about mitigation, impacts, and adaptation?
 - What is the uncertainty in global and regional climate as a result of model uncertainty (as opposed to scenario variations), and how can this be estimated from a model ensemble of opportunity without a specific design to sample uncertainty?
 - Can emergent constraints (i.e., statistical relationships between features of current and projected future climate that emerge from considering the multi-model ensemble as a whole) be used to recalibrate the ensemble and to reduce the uncertainty in the response to a given scenario of future forcing?
 - In which part of the Earth System, and when, are such constraints expected to emerge, how do they trace back to modelled processes, are those processes adequately represented, and how can this information be used to improve models, point to critical observations and monitoring programs, and link process understanding, detection and attribution, projections, and uncertainty quantification?
2. **How can we assess future climate changes given climate variability, climate predictability, and uncertainties in scenarios?**
 - How can we assess future climate changes for forcing pathways spanning a range of uncertainties in global and regional forcing relevant to IAM and IAV research, as well as to policy?
3. How will plausible future forcing pathways affect **climate extremes, global and regional climate information, regional sea level rise, water availability, and biospheric feedbacks**, and how will these affects influence mitigation and adaptation possibilities?

Overview of proposed design and caveats

Here we describe the principal features of the experimental design. We refer to this proposed design as “preliminary” because, as described below, we plan to finalize it over the period December 2014 – March 2015, and submit a final proposal at that time. Given the importance of climate runs based on these scenarios to multiple research communities, we believe it is important to seek additional feedback and interaction on these choices. Further interactions, the gathering of additional information and the final scenario selection have been made part of the experimental design, as described below.

The proposed preliminary experimental design consists of six scenarios grouped into two tiers by priority. Considered together, these scenarios produce climate information for two types of scenarios:

(1) “SSP-based RCPs”: new versions of the RCPs that are based on the Shared Socioeconomic Pathways (SSPs; O’Neill et al., 2014; van Vuuren et al., 2014) and new IAM model simulations derived from them. The SSPs are new societal development pathways that have been developed as part of the parallel process. In contrast, the existing RCPs were derived from SRES scenarios.

(2) “Gap scenarios”: new forcing pathways not covered by the RCPs that include new unmitigated SSP baselines or new mitigation pathways.

PRELIMINARY PROPOSAL

Forcing category	Type of Scenario	Forcing in 2100 ¹ (W/m ²)	SSP ²	Short name	Use by other MIPs ³
Tier 1 ⁴					
High	SSP-based RCP	8.5	5	SSP5-8.5	C ⁴ MIP, DAMIP, GeoMIP, ISMIP6, RFMIP
Medium	SSP-based RCP	6.0	1, 2, 4, 5	SSP1/2-6.0	GeoMIP
Low	SSP-based RCP	2.6	1, 4, 2, 5	SSP1-2.6	AerChemMIP, LUMIP
Tier 2					
High	Gap: Baseline	7.0	3, 2	SSP3-7	AerChemMIP, LUMIP
Medium	SSP-based RCP	4.5	2, 5, 1, 3	SSP2/5-4.5	
Low	Gap: Mitigation	3.7	4, 1, 2, 5	SSP4-3.7	
Tier 1 scenario, 9-member ensemble ⁵	SSP-based RCP	8.5	5	SSP5-8.5	
Overshoot ⁶	Gap: Mitigation	TBD	TBD	TBD	C ⁴ MIP
Tier 1 scenario(s), long-term extension ⁷	SSP-based RCP	TBD	TBD	TBD	ISMIP6, C ⁴ MIP

Notes

1 The leading candidate forcing level is shown in bold. While the forcing category (first column) for each scenario is fixed, the specific forcing level may change by switching scenarios between tiers, depending on final design choices.

2 The leading candidate SSP for each scenario is shown in bold, with other options listed next. For the two “Medium” forcing scenarios, two leading candidates have been identified.

3 Current plans by other MIPs to use ScenarioMIP scenarios either directly or as a basis for a variant to be run as part of their own design are indicated here, but should be considered tentative.

4 We recommend that the Tier 1 runs be considered entry requirements for models participating in ScenarioMIP and that models be encouraged to run as many additional scenarios as possible, guided by this prioritization. If models are going to run only one scenario, we request they run the Tier 1 High scenario.

5 We request that models run 9 or more additional initial condition ensemble members for the SSP5-8.5 scenario (if not 10, then as many as possible). However, these additional ensemble members would be considered Tier 2 runs (i.e., not required runs for participation in ScenarioMIP).

6 An overshoot scenario may be carried out by another MIP (e.g., C⁴MIP or GeoMIP) but otherwise it would be included in the ScenarioMIP design.

7 Long-term extensions to 2300 will be included in the design; the number and type of extension remains to be determined.

More specific justifications for these scenarios are provided in the appendix. Overall, the proposal has the following general features:

- A small number of scenarios (3 in Tier 1) required for any model participating in this MIP, with model runs of additional scientific value in Tier 2.
- Each Tier contains scenarios at high, medium, and low forcing levels in 2100. Tier 1 contains new versions of RCPs based on SSPs. Tier 2 contains high and low “gap” scenarios for forcing levels of interest that are not covered by the RCPs, plus an additional SSP-based RCP at a relevant medium forcing level.

- Tier 2 also contains scenarios that build on Tier 1 scenarios, including additional ensemble members and long-term extensions, as well as an overshoot scenario that would preferably be carried out as part of another MIP but otherwise will be carried out as part of ScenarioMIP.
- The new versions of the RCPs will continue to support scenario-based IAM and IAV research into the mid-2020s. These new RCPs will be based not only on new (CMIP6) climate models, but also on updated forcing pathways generated by new IAM model runs based on the SSPs.
- The two new “gap” scenarios complement the existing RCPs: an SSP baseline (i.e., no mitigation) useful to impact assessments that falls between RCPs 6.0 and 8.5, and a mitigation scenario (3.7 W/m²) of high interest to mitigation policy discussions that falls between RCPs 2.6 and 4.5.
- Scenarios that can anchor experiments in a number of other MIPs (see below) to investigate targeted questions, including for example the influence of land use, aerosols and other NTCFs, and overshoot on climate outcomes; carbon cycle feedbacks; and ice sheet-climate interactions.

We consider the basic structure of the proposal fixed: two tiers, with high, medium and low forcing pathways in each. However, we leave open the possibility that the final experimental design may switch corresponding forcing pathways between tiers, based on the needs of other proposed MIPs and further feedback from research communities. Also, as noted in the table, the specific SSP on which each forcing pathway is based will be finalized once IAM scenarios have been completed in early 2015.

We have outlined below (in the section on timing) the interactions that will occur between December 2014 and March 2015, when we aim to finalize this proposal. Critical open questions to be discussed during that period that will affect the final design choices include:

- The nature of the final versions of the IAM scenarios (changes from current versions might influence design decisions). Information available January 2015.
- Information from IAV community on need for updated climate model scenarios.
- Design of other MIPs (AerChemMIP, C4MIP, DAMIP, GeoMIP, ISMIP6, LUMIP, RFMIP) that could influence ScenarioMIP choices.

Emissions- vs concentration-driven simulations

We recommend that the scenarios specified in the ScenarioMIP design be run as concentration-driven experiments. Such scenarios are more consistent with the “integration” role that these scenarios will play in the broader research community. The conceptual framework for scenario-based research is based on investigating the implications of alternative climate futures. These climate futures will be more similar (for a given scenario) in concentration-driven runs than in emissions-driven runs (given uncertainties in the carbon cycle), and therefore will better serve this purpose of the overall scenario framework.

Concentration driven scenarios still represent uncertainty in the carbon cycle and in climate-carbon cycle feedbacks, through their influence on the anthropogenic carbon emissions allowable for a given concentration pathway. Indeed, ESM results indicating the uncertainty across models in allowable emissions will be a very important outcome for the IAM community. We recognize that concentration-driven scenarios do not allow for assessing amplification effects of feedbacks (in which climate change influences the carbon cycle, producing more emissions and more climate change, and further influencing the carbon cycle, etc.). However, amplification could be investigated in other C⁴MIP simulations, including in the overshoot scenario (see below).

Ensemble size

It is important for scenario-based research to represent the influence of internal variability on climate outcomes. To accommodate this need, while also economizing on model runs, we request that models

run multiple initial condition ensemble members for one scenario, but not for others, based on the assumption that variability estimated for one scenario can be applied to outcomes for others. Currently, our leading candidate for a multiple-ensemble member scenario is the Tier 1 High forcing scenario, although it is possible this could change as a result of further interactions with other MIPs that may wish to draw on this ensemble for their own experiments. We request that models run 9 or more additional ensemble members (if not 9, then as many as possible). These additional ensemble members would be considered Tier 2 scenarios (i.e., not required model runs for participation in ScenarioMIP). For all other scenarios, only a single ensemble member is requested.

Long-term extensions

There is strong interest from the climate and impacts communities in long-term extensions of scenarios beyond 2100. We have so far had direct expressions of interest from some other MIPs (ISMIP6, C⁴MIP; see below) and from representatives of the impacts community. A number of options for the selection and design of such scenarios have been discussed, including extensions assuming zero emissions, constant emissions, constant concentrations, idealized overshoot, and plausible extrapolations. We anticipate including one or more extensions in the final ScenarioMIP proposal, after further discussion with the climate modeling, IAM, and IAV communities on the most useful design.

CMIP5 vs CMIP6 models

For multiple research communities it will be useful to evaluate the difference in climate outcomes for plausible future scenarios that is due to a new generation of climate models, rather than to new scenarios of emissions and land use. For example, such an evaluation is valuable in order to determine whether CMIP5 and CMIP6 results could be used together in research on impacts and adaptation (and how), or whether IAM and IAV researchers should abandon CMIP5 runs in favor of CMIP6 runs when they become available. It is not part of the ScenarioMIP design being proposed to CMIP6 to carry out simulations that would inform this evaluation. However, we believe it would be interesting to the community if at least a few climate modeling teams investigated this question. Possible approaches include running the new SSP-based scenarios with the previous (CMIP5) generation of models, running the previous (RCP) scenarios using the new (CMIP6) generation of models, or carrying out relevant analyses with climate model emulators.

Connections to other MIPs

The ScenarioMIP design is intended to provide a basis for targeted scenarios to be run in other MIPs in order to address specific questions regarding the sensitivity of climate change outcomes to particular aspects of these scenarios, especially land use and emissions of NTCFs. In addition, we are pursuing coordination with other MIPs on overshoot scenarios, including both their climate and carbon cycle consequences. We describe here current plans for coordinated experiments. A summary of the scenarios within the ScenarioMIP design that are currently part of plans for other MIPs is provided in the experimental design table above.

Aerosols and Chemistry MIP (AerChemMIP)

AerChemMIP plans to design experiments with similar overall goals as LUMIP, but directed at the sensitivity of climate to near term climate forcers. One possibility that has been discussed is to use the same approach and the same two scenarios from ScenarioMIP as LUMIP plans to do, but switching assumptions about NTCFs rather than land use change. An alternative possibility is to use the SSP3-baseline from ScenarioMIP as a starting point and devise high and low air pollutant variants of this scenario by either assuming no pollution controls, or maximum feasible reductions in air pollutants.

Coupled Climate Carbon Cycle MIP (C⁴MIP)

ScenarioMIP will coordinate with C⁴MIP on targeted scenarios regarding concentration vs emission driven simulations. While the ScenarioMIP protocol will recommend concentration-driven

simulations (see above), C⁴MIP/Tier 1 will recommend emission-driven simulations for the SSP5-8.5 in order to explore the implications of carbon cycle feedbacks on projected climate change. As mentioned before, C⁴MIP also has an interest in extensions of scenario beyond 2100 (e.g. up to 2300 as in CMIP5) in order to investigate climate change impacts on Earth System components that operate on longer time scales (vegetation, permafrost, oceanic circulation and carbon export, etc.) In addition, there is a possibility for C⁴MIP/Tier 2 to include an overshoot scenario given the importance of carbon cycle responses to the implications of overshoot for mitigation. C⁴MIP has expressed some interest in running an emissions-driven overshoot scenario that could be compared to one of the ScenarioMIP scenarios that does not have substantial overshoot; the leading candidate at the moment is an SSP-based RCP-2.6 scenario. Overshoot scenarios are also of potential interest to GeoMIP given that geoengineering may be an option for avoiding overshoot.

Detection and Attribution MIP (DAMIP)

DAMIP plans to use the high scenario in Tier 1 (currently identified as SSP5-8.5) as an anchoring scenario to which individual forcing simulations extended to the end of the century will be compared. These experiments are aimed at distinguishing the climate effects of different forcings and facilitating the identification of observational constraints and their use in future projections. Additionally, DAMIP needs to choose one of the scenarios in ScenarioMIP as input for the extension of their historical single forcing experiments to 2020. The choice would be driven by two considerations: it would be useful for DAMIP to have multiple ensemble members available for comparison; it would be preferable to choose the scenario whose aerosol forcings are most in line with historical/observed forcings in the last years. Both of these issues will be clarified in the near-future stages of the proposals.

Geoengineering MIP (GeoMIP)

GeoMIP has proposed several experiments that will use two scenarios from ScenarioMIP as a basis from which geoengineering measures would be implemented. Forcing pathways from other ScenarioMIP scenarios would serve as targets for those measures. In particular, SSP5-8.5 would be used as a basis for four experiments: using geoengineering to reduce forcing to the SSPx-6.0 (G6Sulfur and G6Solar experiments) or SSPx-2.6 (G6Sulfur_SSP1-2.6) forcing pathways, investigating the effect of cirrus cloud thinning (G7Cirrus experiment), and investigating the effect of fixed levels or stratospheric aerosol injections (GeoFixed10, 20, 50). In addition, SSPx-6.0 would be used as a basis for a stratospheric aerosol injection experiment (G4SSA).

Ice Sheet MIP (ISMIP6)

ISMIP will be proposing two types of experiments that will draw on long-term extensions of a scenario from ScenarioMIP in order to investigate ice sheet response and ice-climate interactions on centennial timescales. In particular, an extension of SSP5-8.5 to 2300 would be used to provide climate model output for offline (uncoupled) ice sheet simulations, and to provide emissions/concentrations for fully coupled ice sheet-climate model experiments. ISMIP has indicated no strong preference for any particular type of extension (i.e., plausible extrapolation, constant concentration, etc.).

Land Use MIP (LUMIP)

LUMIP plans to design experiments that use two scenarios from ScenarioMIP as a basis for testing sensitivity to land use change. These two scenarios would differ both in forcing levels and in land use change. Tentatively, these two scenarios could be the SSP3-baseline and the SSP1-2.6. These two scenarios span a range of approximately 4.5 W/m² (7.0 vs 2.6 W/m² in 2100), and likely will differ substantially in land use change, with substantial deforestation in the SSP3-baseline and net afforestation in SSP1-2.6. An early suggestion for the LUMIP experimental design called for re-running both scenarios with land use change switched between them, and all else the same. However, other variants are under consideration.

Radiative Forcing MIP (RFMIP)

RFMIP has plans to estimate radiative forcing in different models for a plausible future scenario, preferably a high forcing pathway. At the moment the candidate is SSP5-8.5, whose forcings would

be applied to current day fixed SSTs in the idealized setting of the RFMIP experiments.. RFMIP interest in anchoring their experiment to SSP5-8.5 in ScenarioMIP or another high-forcing scenario would increase the importance of a large number of models running such a scenario.

Vulnerability, Impacts, Adaptation (VIA) Advisory Group

Researchers examining the consequences of climate change and potential adaptations are a key user group of CMIP outputs and products. ScenarioMIP will establish a close link with the impact community through the VIA Advisory Board and other relevant groups to facilitate integrated research that leads to a better understanding not only of the physical consequences of these scenarios on the climate system, but also of the climate impact on societies. In particular ScenarioMIP will link with the VIA Advisory Board to ensure that the designing and prioritizing of the scenarios will serve the VIA community and that the climate model output from the scenarios allows for sector-specific indices being derived (e.g., heat damage degree days for ecosystems, consecutive dry days for agriculture and water resources).

➤ An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*

➤ Proposed timing*

Consistent with the timeline developed at the August 2014 Aspen meeting on MIPs for scenarios, land use, and aerosols, we envision the following timing for ScenarioMIP:

December 2014 – February 2015 Interaction with IAM, IAV and climate modeling communities on the details of ScenarioMIP design as described above. To include:

- (1) solicited feedback from the WCRP Working Group on Regional Climate; IPCC Task Group on Data and Scenario Support for Impact and Climate Analysis (TGICA); CMIP6 VIA Advisory Group; International Committee On New Integrated Climate change assessment Scenarios (ICONICS); Integrated Assessment Modeling Consortium (IAMC) Scientific Working Group on Scenarios; informed users from the climate policy community;
- (2) Continuing interactions with other proposed MIPs (see list above) to coordinate experimental designs;
- (3) Continuing interactions with IAM groups producing SSP-based scenarios;
- (4) Continuing interactions with CMIP6 modelling groups;
- (5) Consideration of comments and recommendations for ScenarioMIP from the review process organized by WGCM and the CMIP Panel as part of the MIP endorsement process.

31 March 2015	Submission of final ScenarioMIP design to CMIP panel
Summer 2015	Submit paper on ScenarioMIP design to the CMIP6 Special Issue
April 2015 – October 2016	Specification of future emissions and land use scenarios from IAMs, harmonization with historical emissions/land use, specification of future atmospheric concentrations
October 2016	Provision of IAM scenario information to ESMs; ScenarioMIP ESM runs begin

- For each proposed experiment to be included in CMIP6**
- the experimental design,
 - the science question and/or gap being addressed with this experiment,

- possible synergies with other MIPs,
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

See accompanying worksheet.

- If possible, a prioritization of the suggested experiments, including any rationale**

See above and accompanying worksheet.

- List of output and process diagnostics for the CMIP DECK/CMIP6 data request**
 - Please indicate whether the standard output archived from the *CMIP DECK* experiments needs to be complemented by additional diagnostics to be useful for the MIP or whether the output is only suggested for the additional experiments,
 - Some output might only be relevant if certain components or diagnostic tools are used interactively (e.g. carbon cycle, chemistry, simulators); please indicate clearly if this is the case,
 - If the data request is large, please indicate the importance of the various data to be archived via a tiered listing.
- Proposed contributions and recommendations for**
 - model diagnostics and performance metrics for model evaluation,
 - observations/reanalysis that could be used to evaluate the proposed experiments. Status in obs4MIPs/ana4MIPs (in database / to be included).
 - when possible tools, code or scripts for model evaluation in open source languages (e.g., python, NCL, R).
- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms.**
- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF.**

References

Meehl, G.A., Moss, R., Taylor, K.E., Eyring, V., Stouffer, R.J., Bony, S. and B. Stevens (2014) Climate Model Intercomparisons: Preparing for the Next Phase. *EOS Transactions of the American Geophysical Union* 95 (9): 77-78.

Moss, R. H. et al. (2008) Towards New Scenarios for Analysis of Emissions, Climate Change, Impacts, and Response Strategies (IPCC Expert Meeting Report, IPCC, Geneva) .

Moss, R. H. et al. (2010). The next generation of scenarios for climate change research and assessment. *Nature* 463: 747-756. doi:10.1038/nature08823

O'Neill, B., Kriegler, E., Riahi, K., Ebi, K.L., Hallegatte, S., Carter, T.R., Mathur, R. and D.P. van Vuuren (2014) A new scenario framework for climate change research: the concept of shared socioeconomic pathways. *Climatic Change, Special Issue, Nakicenovic N, Lempert R, Janetos A (eds) A Framework for the Development of New Socioeconomic Scenarios for Climate Change Research*. DOI 10.1007/s10584-013-0906-1

van Vuuren DP, Kriegler E, O'Neill BC, Ebi KL, Riahi K, Carter TR, Edmonds J, Hallegatte S, Kram T, Mathur R, Winkler H (2014) A new scenario framework for Climate Change Research: scenario matrix architecture. *Climatic Change, Special Issue, Nakicenovic N, Lempert R, Janetos A (eds) A Framework for the Development of New Socioeconomic Scenarios for Climate Change Research*. DOI 10.1007/s10584-013-0906-1

Appendix: Motivations for specific scenario experiments

Tier 1:

High forcing, SSP-based RCP (SSP5-baseline, approximately equivalent to SSP5-8.5)

- Provides updated climate information based on CMIP6 climate models and new forcing pathways similar to RCP8.5 to support continued scenario-based research in the IAM and IAV communities.
- Represents the highest forcing pathway in the new set of IAM scenarios, and is therefore useful to represent a high-end climate change scenario.
- Current preferred SSP: SSP5
 - o SSP5 is currently the only one that produces forcing of 8.5 W/m² by 2100.
- Other SSP options
 - o SSP3: A variant of SSP3, in which economic growth and emissions were higher than in the current version, might also reach 8.5 W/m² and could be an alternative.

Medium forcing, SSP-based RCP (SSPX-6.0)

- Provides updated climate information based on CMIP6 climate models and new forcing pathways for RCP6.0 to support continued scenario-based research in the IAM and IAV communities.
- Current preferred SSP: SSP1 or SSP2
 - o SSP1: Would allow together with SSP5-8.5 to span the range of baseline scenarios. The full set would thus allow comparisons of the baseline climate range with lower mitigation/climate policy scenarios. In addition, it can be combined with SSP1-2.6 for differential impact analyses.
 - o SSP2: Middle of the road development path with a medium forcing pathway (between 7.0/8.5 on the high end, and 4.5/2.6 on the low end) that represents the level of forcing roughly consistent with the current level of climate policy ambition. It could therefore serve as reference case to measure the benefit of strengthening climate policy.
- Other SSP options:
 - o SSP5 and SSP 4

Low forcing, SSP-based RCP (SSPX-2.6)

- Provides updated climate information based on CMIP6 climate models and new forcing pathways for RCP2.6 to support continued scenario-based research in the IAM and IAV communities.
- RCP2.6 is the only RCP that produces a >50% chance of remaining below 2 C of global average temperature increase, an agreed upon climate policy goal.
- IAM studies indicate that RCP2.6 remains technically feasible, even if emissions mitigation is not undertaken until after 2020.
- Current preferred SSP: SSP1
 - o Provides an optimistic scenario in terms of both climate outcomes and societal development, useful as a point of contrast for other scenarios (for example, it may be of use to LUMIP and AerChemMIP in their targeted experiments)
- Other SSP options:
 - o SSP2, 4, and 5 have all been discussed as possibilities (SSP3 may not be feasible)

Tier 2:

High forcing, SSP baseline “gap” scenario (SSPX-baseline)

- Provides climate information for SSP baseline scenarios that currently lack it. Runs based on any one of the three SSP baselines (2, 3, 4) that fall in the gap between RCP6.0 and RCP8.5 (reaching about 7.0 W/m²) would likely produce climate information that would correspond at least approximately to all three of them. The lack of climate information for these scenarios prevents the study of the impacts associated with them. In particular, many impact studies are interested in quantifying “avoided impacts,” requiring evaluation of impacts in an unmitigated baseline scenario and comparing them to the reduction in impacts achieved by mitigating to a lower forcing pathway.
- Combined with a low scenario, it would provide simulations that can test for emergent constraints on future climate model projections.
- Current preferred SSP: SSP3
 - o provides high forcing scenario with pessimistic development pathway to contrast with optimistic (in terms of development) SSP5 pathway, which is useful for impact studies investigating the sensitivity of impacts to alternative development pathways
 - o preliminary versions have high aerosol emissions and substantial land use change (high deforestation), so that this scenario is a good candidate as an anchor for AerChemMIP and LUMIP variants investigating the sensitivity of climate and air quality to changes in NTCFs and land use. (Preliminary plans for LUMIP and AerChemMIP make use of this scenario, see below. However, note that other proposals for anchoring scenarios involve RCP4.5.)
- Other SSP options
 - o SSP2: intermediate scenario

Medium forcing, SSP-based RCP (SSPX-4.5)

- Provides updated climate information based on CMIP6 climate models and new forcing pathways for RCP4.5 to support continued scenario-based research in the IAM and IAV communities.
- Planned to be used for downscaling by CORDEX
- Current preferred SSPs: SSP2 or SSP5
 - o SSP2: middle of the road development pathway paired with middle of the road climate outcome (which is consistent with the rationale of selecting SSP3/5 for high forcing levels and SSP1/4 for the low forcing levels). If SSP2 is selected for 4.5, then SSP1 would be the preferred option for the 6.0 W/m² scenario, since it would allow for differential impact assessment based on SSP1 (6→2.6 W/m²)
 - o SSP5: Allows comparison with SSP5 baseline and would be a good baseline for geoengineering experiments, which may be relevant for a RCP4.5 pathway. If SSP5 is selected for this forcing level, then SSP2 is proposed for the 6.0 W/m² scenario.
- Other SSP options:
 - o SSP3: allow for comparison with SSP3 baseline
 - o SSP1: allows impacts to be compared with SSP1-2.6 within same development pathway

Low forcing, mitigation gap scenario (SSPX-3.7)

- Provides climate information for a mitigation scenario that currently lacks it. There is substantial mitigation policy interest in scenarios that reach 3.7 W/m² by 2100, since mitigation costs differ substantially between RCP4.5 and RCP2.6. Climate model simulations would allow for impacts of a 3.7 scenario to be compared to those occurring in the 4.5 or 2.6 scenarios, to evaluate relative costs and benefits of these scenarios.
- IAMs have already produced scenarios that achieve 3.7 W/m² based on each of the SSPs, so no additional modeling would be necessary.
- Current preferred SSP: SSP4
 - o Chosen to balance climate model runs across SSPs
- Other SSP options:
 - o SSP2: As a middle of the road with a slightly less optimistic climate outcome, it would provide a contrast with SSP1-2.6, which has the most optimistic development pathway and climate outcome.
 - o SSP5: Would be a good candidate for a geoengineering and/or overshoot scenario.

Solar Model Intercomparison Project (SolarMIP)

Application for CMIP6-Endorsed MIPs

Date: 1 December 2014

Proposals from MIPs should include the following information:

* Preliminary information used to determine whether a MIP should be endorsed for CMIP6 or not.

➤ Name of MIP* *SolarMIP*

➤ Co-chairs of MIP (including email-addresses)*

Katja Matthes (kmatthes@geomar.de)

Bernd Funke (bernd@iaa.es)

➤ Proposed members of the Scientific Steering Committee*

Dan Marsh (US)

Drew Shindell (US)

Lon Hood (US)

Rémi Thiéblemont (Germany)

Hauke Schmidt (Germany)

Klairie Tourpali (Greece)

Stergios Misios (Greece)

Adam Scaife (UK)

Lesley Gray (UK)

Dann Mitchell (UK)

Amanda Maycock (UK)

➤ Link to website (if available)*

TBC, SPARC/SOLARIS-HEPPA website <http://www.solarisgeomar.de>

➤ Goal of the MIP and a brief overview*

Goal of the MIP

The purpose of SolarMIP is to: 1) understand and quantify contributions of solar forcing to past and future regional climate variability with respect to other natural and anthropogenic forcings, 2) estimate the impact of uncertainties in solar forcing for atmospheric radiation and chemistry in the past, 3) assess the role of a possible future grand solar minimum for future regional climate change. SolarMIP addresses the question as to how the Earth system responds to solar forcing and tries to detect systematic model biases with respect to the radiation scheme, the ozone field, and the background climatology. This requires not only specially designed model experiments but also the need for additional model output of the CMIP6 experiments. Questions that need to be addressed to understand solar signals in climate model simulations are: 1) What is the role of solar induced ozone signals (prescribed or interactively calculated)?, 2) What is the role of the spectral resolution in the radiation schemes?, 3) What is the role of the background climatology? Another important question with respect to the recent global warming hiatus is to assess the role of a new grand solar minimum for future climate change. We will work in close collaboration with other MIPs (DAMIP, CFMIP, DCP, AerChemMIP, DynVarMIP) to limit the number of extra model experiments and extra output for the modeling groups.

Overview

The importance of solar forcing in particular for regional climate variability is becoming increasingly evident (Gray et al., 2010; Seppälä et al., 2014). Together with volcanic activity, solar variability could be an important external source of natural climate variations, superimposed on the anthropogenic global warming. Because of its prominent 11-year cycle, solar variability offers a degree of predictability for regional climate variability. If the Sun's effect on climate is substantial, foreseeable fluctuations in solar output could help reduce the uncertainty of future regional climate predictions on decadal time scales.

However there are still uncertainties in the atmospheric solar signal and its transfer mechanism(s). Proposed transfer mechanisms include changes in total and spectral solar irradiance as well as in solar-driven energetic particles. Recent work in addition suggests a lagged response in the North Atlantic European region due to atmosphere-ocean coupling (Gray et al., 2013, Scaife et al., 2013). Recent modeling efforts have made progress in defining the pre-requisites to simulate solar influence on regional climate more realistically but the lessons learned from CMIP5 show that a more systematic analysis of climate models within CMIP6 is required to better understand the differences in model responses to solar forcing (Mitchell et al., 2014; Misios et al., 2014; Hood et al., 2014). In particular the role of solar induced ozone changes and the need to prescribe spectrally resolved solar irradiance variations and therefore the need for a suitable resolution in the model's radiation scheme is becoming increasingly evident. The proposed solar only experiment will facilitate the unambiguous solar signal detection in particular to separate clearly solar and volcanic effects and hence will allow a more systematic analysis of the differences in model responses.

To address the uncertainty in spectral solar irradiance forcing (Ermolli et al., 2013), we propose to run an experiment with the so-far standard NRLSSI dataset in order to compare with earlier CMIP results.

Recent studies have investigated a strong future solar forcing change and show only small impact on a global scale. However, a systematic assessment of the regional impacts of a more realistic future solar forcing is still to be done. For example, on a regional scale a future grand solar minimum could potentially reduce the Arctic amplification significantly (Chiodo et al., 2014). We propose to investigate the sensitivity of future regional climate to secular variations of the solar background by an experiment with a more realistic Maunder/Dalton-type future solar forcing.

➤ References (if available)*

Chiodo, G., G. Garcia-Herrera, N. Calvo, J.M. Vaquero, and J.A. Anel, The impact of a future solar minimum under a climate change scenario, under revision Nature Climate Change, 2014.

Ermolli, I., Matthes, K., Dudok de Wit, T., Krivova, N. A., Tourpali, K., Weber, M., Unruh, Y. C., Gray, L., Langematz, U., Pilewskie, P., Rozanov, E., Schmutz, W., Shapiro, A., Solanki, S. K., Thuillier, G. und Woods, T. N. (2013) Recent variability of the solar spectral irradiance and its impact on climate modelling Atmospheric Chemistry and Physics, 13. pp. 3945-3977. DOI [10.5194/acp-13-3945-2013](https://doi.org/10.5194/acp-13-3945-2013).

Gray, L.J., J. Beer, M. Geller, J.D. Haigh, M. Lockwood, K. Matthes, U. Cubasch, D. Fleitmann, G. Harrison, L. Hood, J. Luterbacher, G. A. Meehl, D. Shindell, B. van Geel, and W. White, 2010: Solar Influences on Climate, Rev. Geophys., 48, RG4001, doi:10.1029/2009RG000282.

Gray, L.J. et al., A lagged response to the 11 year solar cycle in observed winter Atlantic/European weather patterns, *J. Geophys. Res.*, **118**, doi:10.1002/2013JD020062, 2013.

Hood, L., S. Misios, D. Mitchel, L.J. Gray, K. Tourpali, K. Matthes, H. Schmidt, G. Chiodo, R. Thiéblemont, E. Rozanov, D. Shindell, A. Krivolutsky (2014), *Solar Signals in CMIP-5 Simulations: The Ozone Response*, to be submitted to *Q. J. Roy. Met. Soc.*

Misios, S., D. Mitchel, L.J. Gray, K. Tourpali, K. Matthes, L. Hood, H. Schmidt, G. Chiodo, R. Thiéblemont, E. Rozanov, D. Shindell, A. Krivolutsky (2014), *Solar Signals in CMIP-5 Simulations: Effects of Atmosphere-Ocean Coupling*, to be submitted to *Q. J. Roy. Met. Soc.*

Mitchell, D., S. Misios, L.J. Gray, K. Tourpali, K. Matthes, L. Hood, H. Schmidt, G. Chiodo, R. Thiéblemont, E. Rozanov, D. Shindell, A. Krivolutsky (2014), *Solar Signals in CMIP-5 Simulations: The Stratospheric Pathway*, submitted to *Q. J. Roy. Met. Soc.*

Scaife, A. A., Ineson, S., Knight, J., R., Gray, L. J., Kodera, K. & Smith, D. M., A mechanism for lagged North Atlantic climate response to solar variability, *Geophys. Res. Lett.*, **40**, 1-6 doi:10.1002/grl.50099 195, 2013.

Seppälä, A., K. Matthes, C. Randall, and I. Mironova (2014), *What is the solar influence on climate? – Overview of activities during CAWSES-II, special issue of Progress in Earth and Planetary Science (PEPS)*, accepted.

- An overview of the proposed experiments*

See attached spreadsheet.

- An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*

The SolarMIP and related CMIP6 experiments (DECK-AMIP, CMIP6 historical simulation, as well as DCP A Component and CFMIP abrupt solar forcing experiments, if selected) will be evaluated and analyzed under the following aspects:

- *Detection of (lagged) solar signals in the stratosphere, the Earth's surface and the ocean on global and regional scale.*
- *Systematic assessment of signals in dependence of model class (climate models with and without interactive chemistry, high-top vs. low-top, resolution of radiation scheme).*

- Proposed timing*

As soon as the solar forcing and the ozone concentration database (with a consistent solar signal! Will be coordinated with AerChemMIP and CCMI) is ready and the DECK experiments have been finished, the additional proposed experiments could be started.

- For each proposed experiment to be included in CMIP6**

1. "Solaronly" Experiment (Tier 1)

- the experimental design;

Historical simulation with solar forcing only (preindustrial GHG+ODS, 1850 level). Identical to the DAMIP-histNAT experiment, but volcanic forcing switched off. 3 Ensembles and extension to 2020 to be consistent with DAMIP and DCP.

- the science question and/or gap being addressed with this experiment;
Assess solar-only effects, separate solar and volcanic effects, additional output to clarify importance of radiation, ozone (interactive or prescribed) and stratospheric dynamics for solar signals.
- possible synergies with other MIPs;
 - *DAMIP will make use of Solaronly to attribute observed changes to contributions from solar forcing.*
 - *To assess solar cycle contribution to prediction, SolarMIP runs could be used to estimate solar forcing spread and compare with Component A of DCPD (for which additional output is desirable, see below)*
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

Primarily A, to a lesser extent B and D.

2. “futureSolarMin” Experiment (Tier 1)

- the experimental design;
Future simulation (based on RCP4.5&RCP8.5) using solar forcing running into a new Dalton/Maunder Minimum type.
- the science question and/or gap being addressed with this experiment;
Assess impact of future grand solar minimum for climate change under increasing GHG & ODS concentrations.
- possible synergies with other MIPs;
 - *CFMIP: the “futureSolarMin” experiment is complementary to the proposed CFMIP abrupt solar forcing change experiment. A joint evaluation of both experiments is envisaged.*
 - *DCPD: futureSolarMin experiment will be used to assess solar cycle contributions to decadal predictions and will be compared with Component A of DCPD*
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

Primarily A, to a lesser extent B, C, and D.

3. “NRLSens” Experiment (Tier 2)

- the experimental design;
Sensitivity experiment based on in the CMIP6 Historical Simulation, but using a different spectral solar irradiance (NRLSSI) forcing.
- the science question and/or gap being addressed with this experiment;

Assess uncertainty in solar SSI forcing and generate reference for earlier CMIP5 experiments (all using NRLSSI).

- possible synergies with other MIPs;
 - *NRLSens will provide solar signal uncertainty estimates relevant for DAMIP.*
- potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

Primarily A, to a lesser extent B and D.

- If possible, a prioritization of the suggested experiments, including any rationale**

Prioritization in the order as described above.

- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**

no objections

- List of output and process diagnostics for the CMIP DECK/CMIP6 data request**

1. *Zonal mean shortwave (SW) and longwave (LW) heating rates (as requested by DynVar)*
2. *2D or 3D ozone fields (prescribed or interactively calculated).*
3. *O₂ and O₃ photolysis rates from climate models with interactive chemistry (may be already requested by AerChemMIP?).*
4. *O_x as well as O_x total production and loss rates (may be already requested by AerChemMIP?).*
5. *TEM diagnostics (monthly mean v*, w*, EPflux divergence) indices (already requested by DynVar)*
6. *daily zonal mean temperatures and zonal wind*
7. *3D geopotential height at least at 10hPa level for the detection of sudden stratospheric warmings (SSWs) and the calculation of NAM and SAM*

- whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;

Request for collecting the above mentioned additional output for DECK-AMIP & CMIP6 historical simulations. This output might be also requested for some runs in DAMIP, CFMIP (abrupt solar forcing change experiments), and DCPP (Component A experiments) in order to add the analysis to the experiments requested in SolarMIP.

- whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);

- whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.

Variables 1-4 are required to disentangle the different pathways of direct effects of solar forcing on the atmosphere. Variables 5-7 will be used to investigate the indirect dynamical responses to the solar forcing.

- whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;

Data can be regridded.

- the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.

N/A

➤ Any proposed contributions and recommendations for**

- model diagnostics and performance metrics for model evaluation;

Request for detailed information on spectral resolution in SW radiative heating rate and photolysis calculations as well as respective parameterizations used (to understand the differences in heating rates and production rates better than in CMIP5) (may be already requested by AerChemMIP?)

- observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;

None for the proposed SolarMIP experiments. Available reanalysis satellite data for evaluation of solar signals in DECK AMIP and CMIP6 historical simulations.

- tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R).

N/A

➤ Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms.**

None

➤ Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF.**

None

Model Intercomparison Project on the climatic response to Volcanic forcing (VolMIP)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

Davide Zanchettin^{1*}, Claudia Timmreck², Myriam Khodri³, Alan Robock⁴, Gabi Hegerl⁵, Anja Schmidt⁶, Matthew Toohey⁷, Francesco S. R. Pausat⁸, Benjamin Black⁹, Oliver Bothe¹⁰, Jason M. English¹¹, Edwin Gerber¹², Hans F. Graf¹³, Allegra N. LeGrande¹⁴, Graham Mann⁶, Timothy Osborn¹⁵, Steven J. Phipps¹⁶, Christoph C. Raible¹⁷, Björn Stevens², Didier Swingedouw¹⁸, Kostas Tsigaridis^{14,19}, Qiong Zhang⁸

1	University of Venice, Italy
2	Max-Planck-Institute for Meteorology, Hamburg, Germany
3	IRD/IPSL/Laboratoire d'Océanographie et du Climat, France
4	Department of Environmental Sciences, Rutgers University, New Brunswick, USA
5	GeoScience, U. Edinburgh, UK
6	School of Earth and Environment, University of Leeds, UK
7	GEOMAR Helmholtz Centre for Ocean Research Kiel, Germany
8	Department of Meteorology (MISU), Stockholm, Sweden
9	University of California , Berkley, U.S.A
10	HZG, Helmholtz Center Geesthacht, Geesthacht, Germany
11	Laboratory for Atmospheric and Space Physics University of Colorado, Boulder, USA
12	Courant Institute of Mathematical Sciences, New York University
13	University of Cambridge, UK
14	NASA GISS, Columbia University, USA
15	Climatic Research Unit, School of Environmental Sciences, University of East Anglia, UK
16	University of New South Wales, Sydney, Australia
17	Universität Bern, Swiss
18	Université de Bordeaux, France
19	Center for Climate Systems Research, Columbia University

* To whom the correspondence should be sent: University of Venice, Dept. of Environmental Sciences, Informatics and Statistics, Calle Larga Santa Marta, Dorsoduro 2137, Venice, Italy (davide.zanchettin@unive.it)

VolMIP

Name of MIP:

Model Intercomparison Project on the climatic response to Volcanic forcing (**VolMIP**)

Co-chairs of MIP (including email addresses):

Davide Zanchettin (davide.zanchettin@unive.it)

Claudia Timmreck (claudia.timmreck@mpimet.mpg.de)

Myriam Khodri (myriam.khodri@locean-ipsl.upmc.fr)

Members of the Scientific Steering Committee:

Gabi Hegerl (gabi.hegerl@ed.ac.uk)

Alan Robock (robocock@envsci.rutgers.edu)

Anja Schmidt (A.Schmidt@leeds.ac.uk)

Matt Toohey (mtoohey@geomar.de)

Edwin Gerber (gerber@cims.nyu.edu)

Link to website (if available):

WCRP webpage:

<http://www.wcrp-climate.org/index.php/modelling-wgcm-mip-catalogue/modelling-wgcm-mips/505-modelling-wgcm-volmip>

Official webpage: under construction

Goal of the MIP and a brief overview

VolMIP is central to the three broad CMIP questions:

- How does the Earth system respond to external forcing?
- What are the origins and consequences of systematic model biases?
- How can we assess future climate changes given climate variability, predictability and uncertainties in scenarios?

VolMIP is motivated by the large uncertainties regarding the climatic responses to strong volcanic eruptions identified in CMIP5 simulations with respect to, e.g., the Northern Hemisphere's winter response (e.g., Driscoll et al., 2012, Charlton-Perez et al., 2013) and the response of the oceanic thermohaline circulation (Ding et al., 2014), and by the apparent mismatch between simulated and reconstructed post-eruption surface cooling for volcanic eruptions during the last millennium (Mann et al., 2012, 2013; Anchukaitis et al., 2012; D'Arrigo et al., 2013; Schurer et al., 2013). Therefore, VolMIP will assess to what extent responses of the coupled ocean-atmosphere system to strong volcanic forcing are robustly simulated across state-of-the-art coupled climate models and identify the causes that limit robust simulated behavior, especially differences in their treatment of physical processes.

VolMIP is closely linked to the WCRP Grand Challenge on:

- "Clouds, circulation and climate sensitivity," in particular through improved characterization of volcanic forcing and improved understanding of how the hydrological cycle and the large-scale circulation respond to volcanic forcing. VolMIP further contributes to the initiative on leveraging the past record through planned experiments describing the climate response to historical eruptions that are not (or not sufficiently) covered by CMIP6-DECK or other MIPs. VolMIP will contribute towards more reliable models through improved understanding of how model biases affect the response to volcanic forcing.
- "Climate extremes" and "Regional climate information," in particular through a more systematic assessment of regional climate variability – and associated predictability and prediction - during periods

of strong volcanic forcing at both intraseasonal-to-seasonal (e.g., post-eruption Northern Hemisphere's winter warming) and interannual-to-decadal (e.g., post-eruption delayed winter warming) time scales.

- “Water Availability,” in particular through the assessment of how strong volcanic eruptions affect the monsoon systems and the occurrence of extensive and prolonged droughts.

VolMIP addresses specific questions related to:

- The apparent mismatch between simulated and reconstructed post-eruption surface cooling for volcanic eruptions during the last millennium (Mann et al., 2012; Anchukaitis et al., 2012; D’Arrigo et al., 2013; Schurer et al., 2013). A possible reason for the mismatch are the large uncertainties in the volcanic forcing for eruptions that occurred during the pre-instrumental period and for which no direct observations are available. Therefore, VolMIP will provide new consensus forcing input data and related coupled climate simulations for some of the major volcanic eruptions that occurred during the pre-industrial period of the last millennium. Forcing data will be in the form of best estimates with uncertainties or of a range of estimates if a best estimate is not feasible with the given uncertainties.
- The mismatch between observed and modeled seasonal to interannual dynamical responses to volcanic eruptions during the instrumental period. Observations suggest that volcanic eruptions are followed by an anomalously strong Northern Hemisphere’s winter polar vortex, and significant positive anomalies in the North Atlantic Oscillation and Northern Annular Mode, but CMIP5 models do not robustly reproduce this behavior (e.g., Driscoll et al., 2012, Charlton-Perez et al., 2013). Observed volcanic events are, however, few and of limited magnitude, and their associated dynamical climate response is very noisy (e.g., Hegerl et al., 2011). The short-term dynamical response is now known to be sensitive to the particular structure of the applied forcing (Toohey et al., 2014). Using carefully constructed forcing fields and a sufficient number of realizations, VolMIP will investigate the inter-model robustness of the short-term dynamical response to volcanic forcing, and elucidate the mechanisms through which volcanic forcing leads to changes in surface dynamics.
- The large uncertainties in the interannual and decadal dynamical climatic responses to strong historical volcanic eruptions. As described above, coupled climate simulations produce a considerable range of atmospheric and oceanic dynamical responses to volcanic forcing, which likely depend on various aspects of model formulation, on the simulated background internal climate variability (e.g., Zanchettin et al, 2013), and also on eruption details including magnitude, latitude and season (e.g., Timmreck, 2012). VolMIP will help to identify the origins and consequences of systematic model biases affecting the dynamical climate response to volcanic forcing and to clarify how regional responses to volcanic forcing are affected by the background climate state, especially the phase of dominant modes of internal climate variability. As a consequence, VolMIP will improve our confidence in the attribution and dynamical interpretation of reconstructed post-eruption regional features and provide insights into regional climate predictability during periods of strong volcanic forcing.
- The large uncertainties in the multidecadal and longer-term climate repercussions of prolonged periods of strong volcanic activity (e.g., Miller et al., 2012; Schleussner and Feulner, 2013; Zanchettin et al., 2013). VolMIP proposes an experiment describing the climate response to the close succession of strong volcanic eruptions that affected the early 19th century, whose long-term repercussions may be relevant for the initialization of CMIP6-Nucleus *historical* simulations.

In summary, VolMIP will contribute towards advancing our understanding of the dominant mechanisms behind simulated post-eruption climate evolution, but also more generally of climate dynamics and decadal variability. Volcanic eruptions offer the opportunity to assess the climate system’s dynamical response to changes in radiative forcing, a major uncertainty in future climate projections. Careful sampling of initial climate conditions and the possibility to consider volcanic eruptions of different strengths (e.g., Fröhlicher et al., 2012; Muthers et al., 2014a,b; Zanchettin et al., 2014b) will allow a better understanding of the relative role of internal and externally-forced climate variability during periods of strong volcanic activity, hence improving the evaluation of climate models and enhancing our ability to accurately simulate past, as well as future, climates.

For these purposes, VolMIP defines a common protocol to improve comparability of results across different Earth system models and coupled general circulation models, and accordingly subjects them to the same set of idealized volcanic perturbations under similar background climate conditions (Zanchettin et al., in prep, 2014a).

VolMIP experiments will be designed based on a twofold strategy.

- A first set of experiments is designed to systematically investigate inter-model differences in the long-term (up to the decadal time scale) dynamical climate response to idealized volcanic eruptions that are characterized by a high signal-to-noise ratio in the response of global-average surface temperature. The main goal of these experiments is to assess the signal propagation pathways of volcanic perturbations within the simulated climates, the associated determinant processes and their representation across models.
- A second set of experiments will be used to systematically investigate inter-model differences in the short-term dynamical response to volcanic eruptions characterized by a low signal-to-noise ratio in the response of global-average surface temperature. The main goal of these experiments is to quantify the uncertainty in the short-term climate response to a 1991 Pinatubo-like eruption and discriminate the parts that are due to internal variability and to model characteristics. The proposed set of experiments will include idealized sensitivity experiments designed to determine the different contributions to such uncertainty that are due to the direct radiative (i.e., surface cooling) and to the dynamical (i.e., stratospheric warming) response.

Generation of forcing input data for both types of experiments is an integral part of VolMIP. Some of the participating modeling groups are currently testing the proposed methodologies through coordinated activities within VolMIP and in cooperation with other MIPs.

An overview of the proposed experiments

An overview of the proposed experiments is provided in Tables 1, 2 and 3, where they are summarized according to their prioritization. VolMIP experiments are divided into two main branches: long-term volcanic forcing experiments and short-term volcanic forcing experiments.

Long-term volcanic forcing experiments

Experiments based on coupled climate simulations to assess inter-model differences in the climate response to *very strong* volcanic eruptions up to the decadal time scale.

- *VolLongS100EQ*: This Tier 1 experiment is designed to realistically reproduce the radiative forcing resulting from the 1815 eruption of Mt. Tambora, Indonesia. The experiment will not account for the actual climate conditions when the real event occurred (e.g., presence and strength of additional forcing factors). Instead, the experiment is designed to span very different initial climate states to systematically assess uncertainties in the post-eruption behavior that are related to background climate conditions.
- *VolLongS100HL*: An additional, non-mandatory experiment which applies the same approach as *VolLongS100EQ* and extends the investigation to the most relevant historical high-latitude volcanic eruption (1783-1784 Laki, Iceland). The unique eruption style (large SO₂ mass releases: 100 Tg SO₂, and close temporal spacing: 5 active phases within 5 months) will substantially contribute to outstanding questions about the magnitude of the climatic impact of high-latitude eruptions. Results of this experiment may have implications for sulfate aerosol geo-engineering.
- *VolLongC19th*: A “volcanic cluster” experiment to investigate the climate response to a close succession of strong volcanic eruptions. The proposed experiment is designed to realistically reproduce the volcanic forcing generated by the early 19th century volcanic cluster (including the 1809 eruption of unknown location and the 1815 Tambora and 1835 Cosigüina eruptions). The early 19th century is the coldest period in the past 500 years (Cole-Dai et al., 2009) and therefore of special interest for multidecadal variability. In addition long-term repercussions may be relevant for the initialization of CMIP6-Nucleus *historical* simulations.

Short-term volcanic forcing experiments

Experiments based on coupled climate simulations to assess uncertainty and inter-model differences in the seasonal-to-interannual climate response to volcanic eruptions characterized by a rather low signal-to-noise ratio in the response of global-average surface temperature.

- *VolShort20EQfull*: This Tier 1 experiment uses the same volcanic forcing recommended for the 1991 Pinatubo eruption which is used in the CMIP6-Nucleus *historical* simulation, but produces a large ensemble of short-term simulations in order to accurately estimate simulated responses to volcanic forcing which may be small compared to the internal variability.
- *VolShort20EQsurf/strat*: Additional non-mandatory simulations, which are aimed at investigating the mechanism(s) connecting volcanic forcing and short-term climate anomalies. Specifically, these experiments will aim to disentangle dynamical responses to the two primary thermodynamic consequences of aerosol forcing: stratospheric heating and surface cooling.
- *VolShort20EQslab*: Non-mandatory slab-ocean experiment, which is proposed to clarify the role of coupled atmosphere-ocean processes (most prominently linked to the El Niño-Southern Oscillation) in determining the dynamical response.
- *VolShort20EQini*: Non-mandatory experiment to address the impact of volcanic forcing on seasonal and decadal climate predictability and predictions. The experiment will address the climate implication of a future Pinatubo-like eruption.

Experimental set-up:

Length of integration

- *LongS*: for each simulation: at least 20 years (mandatory), but preferably longer (30-40 years) to cover the multi-decadal oceanic response;
- *LongC*: at least 50 years to cover the multi-decadal oceanic response and to assess stationarity of post-cluster climate;
- *Short*: for each simulation: 3 years, since the experiment focuses on the short-term responses;
- *Short.ini*: 10 years for each initialized run (hindcast, forecast).

Initial conditions:

- *LongS*: predefined states describing different states of dominant modes of variability (see “ensemble size”) sampled from an unperturbed control integration, under common constant boundary forcing across the different models (*PiControl* simulations from DECK). The VolMIP experiments should maintain the same constant boundary forcing as the control integration, except for the volcanic forcing;
- *LongC*: as *LongS*, but inclusion of background volcanic forcing and a dedicated spin-up procedure for this experiment are currently under discussion to account for possible implications of volcanic forcing on ocean heat content in long transient simulations (e.g., Gregory, 2010);
- *Short*:
- *Short.ini*: initialized on 1 January 2014.

Ensemble size:

- *LongS*: should be large to systematically account for the range of variability depicted by the dominant processes influencing interannual and decadal climate variability. VolMIP will accordingly identify a set of desired initial conditions. Nine simulations are planned for the Tier 1 experiment, which would allow spanning warm/cold/neutral and strong/weak/neutral states of El Niño-Southern Oscillation (ENSO) and of the Atlantic Meridional Overturning Circulation (AMOC), respectively;
- *LongC*: at least an ensemble of 3 simulations;
- *Short*: same rationale as for *LongS*, but further taking into account additional phenomena primarily contributing to internal atmospheric variability, such as the Quasi Biennial Oscillation (QBO), the characteristics of the polar vortex and the North Atlantic Oscillation (NAO). A core of 25 simulations is requested for the Tier 1 experiment, but a larger ensemble size is recommended;
- *Short.ini*: at least 5-member ensembles, but preferably 10-member ensembles.

Forcing input:

Forcing data should be consistent across the participating models for all events included in the protocol. Therefore, VolMIP will provide a self-consistent set of forcing parameters that can be used by all models, in order to ensure the best possible consistency between models in the resulting radiative forcing. Depending on the number of participating coupled climate models including modules for interactive stratospheric chemistry and aerosols microphysics, VolMIP may pose an additional focus on the simulated climatic response to given SO₂ emissions beyond the proposed CMIP6 simulations. In this stage, VolMIP will benefit from global aerosol model studies conducted within the framework of the Stratospheric Sulfur and its Role in Climate (SSiRC) initiative.

- *Long:* The forcing input data will be in the form of aerosol optical properties (e.g., aerosol optical depth, effective radius, single scattering albedo, asymmetry factor), which will allow the applied forcing in the different models to be constrained. Coupled climate models including modules for stratospheric chemistry and aerosol microphysics will be selected and used to generate the forcing input. Ongoing coordinated activities mainly involving MPI-M and IPSL are currently devoted to testing the methodology. If ad-hoc forcing inputs cannot be generated for an event through the proposed methodology, VolMIP will indicate reference forcing data sets to be used that are already available to the community.
- *Short:* The mandatory Tier1 experiment will use the volcanic forcing for the 1991 Pinatubo eruption which is recommended for the CMIP-Nucleus *historical* simulation (assumed: Sage_4λ¹). The additional mechanistic forcing experiments that are aimed at dissecting the contributions from direct radiative and dynamical responses will make use of prescribed surface radiative flux anomalies and of heating rates in the stratosphere. To generate such input data, specific diagnostics from the Tier-1 experiments are required (if these are not made available, the VolMIP protocol will provide reference input data to the community).

The observation-based volcanic-forcing to be used in the CMIP *historical* and VolMIP *VolShort20EQfull* experiments contains information about the real-world structure of the stratospheric circulation at the time of the eruptions, which does not necessarily match the states of individual free-running model realizations. To further investigate the impact of the forcing structure on the dynamical response, VolMIP will support the development of an idealized Pinatubo volcanic forcing dataset, where the spatial structure of the forcing is much more uniform than observation-based forcings. This work shares parallels with the WCRP Grand Challenge initiative “Easy Aerosol”, and we envision cooperation in the future months between the two groups. Additional dedicated sensitivity experiments will be carried out by individual model Centers to contribute to this activity.

An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments

VolMIP experiments will provide context to CMIP6-DECK (AMIP) and -Nucleus simulations where volcanic forcing is among the dominant sources of climate variability and uncertainty.

¹VolMIP experiments will be designed in a way that any recommended CMIP6 volcanic data set is applicable. The indications provided so far by the CMIP6 panel about the recommend volcanic forcing data for the CMIP6-Nucleus experiments are not definitive (email V. Eyring, 27.11.2014). It is assumed that the recommended volcanic forcing dataset for the CMIP6 *historical* simulations is based on the SAGE_4λ dataset (Arfeuille et al., 2013), since Larry Thomason is the designed responsible for volcanic forcing (page 10 ofCMIP6FinalDesign_WGCMMeeting_141110_Sent.pdf).

Proposed timing

2014 November	High-latitude volcano workshop in Stockholm: definition of pre-studies on high-latitude volcanic eruptions
2014 November	Revised version submitted to CMIP6 panel
2014-2015	Experimental design phase and definition of consensus volcanic forcing input
2015 January	Experiment and variable list sent to CMIP6 panel
2015 February	MiKlip/SPECS workshop in Offenbach. Experimental set-up for volcanic prediction runs (DCPP, VolMIP)
2015 April	VolMIP splinter meeting at Tambora conference in Bern (Switzerland)
2015	GMD Paper documenting detailed experimental design
2015 -2016	Work on idealized volcanic forcing fields
2016	Execution of Tier1 experiments
2017- 2019	Execution of Tier2 (Tier3) experiments
2017	Public sharing and analysis of model output

Possible synergies with other MIPs:

VolMIP is closely linked to and will co-operate with the following ongoing modeling activities and MIPs:

- **PMIP** (<https://pmip3.lscce.ipsl.fr/>) – PMIP and VolMIP provide complementary perspectives on one of the most important and less understood factors affecting climate variability during the last millennium. VolMIP systematically assesses uncertainties in the climatic response to volcanic forcing associated with initial conditions and structural model differences. In contrast, the PMIP last-millennium experiments, i.e., the *past1000* simulations, describe the climatic response to volcanic forcing in long transient simulations where related uncertainties are due to the reconstruction of past volcanic forcing, the implementation of volcanic forcing within the models, initial conditions, the presence and strength of additional forcings, and structural model differences. VolMIP and PMIP are expected to tighten cooperation in the upcoming months to strengthen the synergies between the two MIPs.
- **GeoMIP** (<http://climate.envsci.rutgers.edu/GeoMIP/>) – GeoMIP and VolMIP share interest on the climatic effects of massive stratospheric aerosol loadings. The closest association between proposed experiments is between VolMIP *Long* and GeoMIP G6sulfate simulations.
- **RFMIP** (Radiative Forcing MIP) – Precise quantification of the forcing to which models are subject is central for both RFMIP and VolMIP. RFMIP has encouraged other MIPs to standardize as far as possible to the RFMIP methodology for computing radiative forcings. RFMIP has planned transient volcanic and solar forcing experiments with fixed preindustrial SST to diagnose volcanic and solar effective forcing, instantaneous forcing and adjustments, which seems to be complementary to the *Short* experiments for VolMIP.
- **DAMIP** (Detection and Attribution MIP) – DAMIP and VolMIP share the common interest of assessing the relevance of volcanic forcing over the historical past. In particular, VolMIP can address the substantial uncertainty associated with the effects of volcanism on the historical periods. DAMIP's

histALL, histNAT, histVLC and histALL_aerconc can provide context to the *Short* set of VolMIP simulations, since they include the 1991 Pinatubo eruption within transient climate situations.

- **DCPP** (Decadal climate prediction panel) - VolMIP and DCPD are closely working together on the impact of future volcanic eruptions on seasonal and decadal predictions, with a common experiment. The proposed VolMIP's *Short* experiment including 1991 Pinatubo-like volcanic forcing in decadal prediction runs (*Short20EQini*) and the DCPD experiment C2.1 are identical and will be jointly prepared/discussed in a meeting planned for February 2015 in Offenbach (Germany).
- **SPARC DYNVAR** (<http://www.sparcdynvar.org/>) – The SPARC DynVar group aims to assess the impact of uncertainty in atmospheric dynamics on climate projections and is therefore deeply involved in the setup and analysis of VolMIP's *Short* experiments.
- VolMIP is closely linked to with the ongoing modeling activities within **SPARC-SSiRC** (<http://www.sparc-ssirc.org/>). The Stratospheric Sulfur and its Role in Climate Initiative (SSiRC) model intercomparison uses global aerosol models to understand the radiative forcing of stratospheric aerosols (background, volcanic) and to assess related parameter uncertainties. The SSiRC study "Pinatubo Emulation in Multiple models" (PoEMs) will inter-compare and evaluate Pinatubo perturbation to stratospheric aerosol properties and radiative forcings across AGCMs with prognostic stratospheric aerosol modules.

Potential benefits of the experiment to (A) climate modeling community, (B) Integrated Assessment Modelling (IAM) community, (C) Impacts Adaptation and Vulnerability (IAV) community, and (D) policy makers.

- A. VolMIP will contribute towards identifying the causes that limit robust simulated behavior under strong volcanic forcing conditions. Uncertainty in simulated estimates of clear-sky radiative forcing is largest around strong volcanic eruptions, which poses VolMIP at the core of CMIP6. VolMIP will also clarify more general aspects of the dynamical climatic response to strong external forcing, especially differences in the models' treatment of physical processes. VolMIP will further evaluate the possibility of robustly identifying key climate feedbacks in coupled climate simulations following well-observed eruptions (e.g., Soden et al., 2002), and assess the role of model biases for simulations-observations discrepancies.
- B. VolMIP will contribute towards advancing our understanding of the dominant mechanisms behind simulated post-eruption climate evolution, but also more generally of climate dynamics, decadal variability and of past transitions between different multi-centennial climate states, such as the transition between the so-called Medieval Climate Anomaly and Little Ice Age. Careful and systematic sampling of initial climate conditions and consideration of volcanic eruptions of different strength will help in better understanding the relative role of internal and externally-forced climate variability during periods of strong volcanic activity, hence improving the evaluation of climate models and advancing our understanding of past climates.
- C. VolMIP will identify regions that are most robustly significantly affected by strong volcanic eruptions, and it will provide a framework for assessing the immediate as well as decadal climate repercussions of future volcanic events.
- D. VolMIP will contribute towards advancing our understanding of the relative role of internal and volcanically-forced climate variability, therefore providing relevant information to policy makers concerning how the latter may contribute to the spread of future climate scenarios (where volcanic forcing is presently not accounted for).

All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.

No objection

List of output and process diagnostics for the CMIP DECK/CMIP6 data request:

VolMIP output is planned to be converted into the standard format using the CMOR package, following the same criteria adopted for *past1000* and *historical* simulations. Additional output is needed for *Short* experiments, in particular for the DYNVAR diagnostic tool, which includes key diagnostics of parameterized and resolved wave forcings, radiative and latent heating rates. A daily temporal resolution of output data for the stratosphere is desirable.

References

- Anchukaitis K, Breitenmoser P, Briffa K, Buchwal A, Büntgen U, Cook E, D'Arrigo R, Esper J, Evans M, Frank D, Grudd H, Gunnarson B, Hughes M, Kirdyanov A, Körner C, Krusic P, Luckman B, Melvin T, Salzer M, Shashkin A, Timmreck C, Vaganov E, Wilson R. (2012) Tree-rings and volcanic cooling. *Nature Geoscience*, 5: 836-837 doi:10.1038/ngeo1645
- Arfeuille, F., B. P. Luo, P. Heckendorn, D. Weisenstein, J. X. Sheng, E. Rozanov, M. Schraner, S. Brönnimann, L. W. Thomason, and T. Peter (2013), Uncertainties in modelling the stratospheric warming following Mt. Pinatubo eruption, *Atmos. Chem. Phys.*, 13, 11221-11234, doi:10.5194/acp-13-11221-2013, 2013
- Berdahl, M., and A. Robock (2013) Northern Hemispheric cryosphere response to volcanic eruptions in the Paleoclimate Modeling Intercomparison Project 3 last millennium simulations, *J. Geophys. Res. Atmos.*, 118, 12,359–12,370, doi:10.1002/2013JD019914
- Cole-Dai J, D. Ferris, A. Lanciki, J. Savarino, M. Baroni, MH Thiemens (2009) Cold decade (AD 1810–1819) caused by Tambora (1815) and another (1809) stratospheric volcanic eruption, *Geophys. Res. Lett.*, 36, L22703 doi:10.1029/2009GL04088.
- Driscoll, S., A. Bozzo, L. J. Gray, A. Robock, and G. Stenchikov (2012) Coupled Model Intercomparison Project 5 (CMIP5) simulations of climate following volcanic eruptions, *J. Geophys. Res.*, 117, D17105, doi:10.1029/2012JD017607
- D'Arrigo, R., Wilson, R., & Anchukaitis, K. J. (2013) Volcanic cooling signal in tree ring temperature records for the past millennium. *Journal of Geophysical Research: Atmospheres*, 118(16), 9000-9010
- Ding, Y., J. A. Carton, G. A. Chepurin, G. Stenchikov, A. Robock, L. T. Sentman, and J. P. Krasting (2014) Ocean response to volcanic eruptions in Coupled Model Intercomparison Project 5 (CMIP5) simulations. *J. Geophys. Res.*, 119, 5622-5637, doi:10.1002/2013JC009780.
- Driscoll, S., Bozzo, A., Gray, L. J., Robock, A., & Stenchikov, G. (2012) Coupled Model Intercomparison Project 5 (CMIP5) simulations of climate following volcanic eruptions. *Journal of Geophysical Research: Atmospheres*, 117, D17105, doi:10.1029/2012JD017607.
- Froelicher, T. L., F. Joos, C. C. Raible, J. L. Sarmiento (2013) Atmospheric CO₂ response to volcanic eruptions: the role of ENSO, season, and variability. *Global Biogeochemical Cycles*, 27, 239-251
- Gregory, J. M. (2010) Long-term effect of volcanic forcing on ocean heat content. *Geophys. Res. Lett.*, 37, L22701, doi:10.1029/2010GL045507
- Hegerl, G., J. Luterbacher, F. González-Rouco, S. F. B. Tett, T. Crowley and E. Xoplaki (2011) Influence of human and natural forcing on European seasonal temperatures. *Nat. Geosc.* 4:99-103, doi:10.1038/NGEO1057
- Mann, M.E., Fuentes, J.D., Rutherford, S. (2012) Underestimation of volcanic cooling in tree-ring based reconstructions of hemispheric temperatures. *Nature Geosciences*, doi 10.1038/ngeo1394
- Mann, M. E., Rutherford, S., Schurer, A., Tett, S. F., & Fuentes, J. D. (2013) Discrepancies between the modeled and proxy-reconstructed response to volcanic forcing over the past millennium: Implications and possible mechanisms. *Journal of Geophysical Research: Atmospheres*, 118(14), 7617-7627
- Mignot, J., M. Khodri, C. Frankignoul, and J. Servonnat (2011), Volcanic impact on the Atlantic Ocean over the last millennium, *Clim. Past*, 7, 1439–1455, doi:10.5194/cp-7-1439-2011

- Miller, G. H., Geirsdóttir, Á., Zhong, Y., Larsen, D. J., Otto-Bliesner, B. L., Holland, M. M., Bailey, D. A., Refsnider, K. A., Lehman, S. J., Southon, J. R., Anderson, C., Björnsson, H., and Thordarson, T. (2012) Abrupt onset of the Little Ice Age triggered by volcanism and sustained by sea-ice/ocean feedbacks, *Geophys. Res. Lett.*, 39, L02708, doi:10.1029/2011GL050168
- Muthers, S., J. G. Anet, E. Rozanov, C. C. Raible, T. Peter, A. Stenke, A. Shapiro, J. Beer, F. Steinhilber, S. Broennimann, F. Arfeuille, Y. Brugnara, and W. Schmutz (2014a) Sensitivity of the winter warming pattern following tropical volcanic eruptions to the background ozone climatology, *Journal of Geophysical Research*, 119, 1340-1355. DOI:10.1002/2013JD020138
- Muthers, S., F. Arfeuille, and C. C. Raible (2014b) Dynamical and chemical ozone perturbations after volcanic eruptions: Role of the climate state and the strength of the eruption. *Journal of Geophysical Research*, submitted
- Schurer, A., Hegerl, G.C., Mann, M., Tett, S.F.B., Phipps, S (2013) Separating forced from chaotic variability over the last millennium. *J Climate*, doi:10.1175/JCLI-D-12-00826.1
- Schleussner, C. F. and Feulner, G. (2013) A volcanically triggered regime shift in the subpolar North Atlantic Ocean as a possible origin of the Little Ice Age, *Clim. Past*, 9, 1321–1330, doi:10.5194/cp-9-1321-2013
- Soden, B. J., R. T. Wetherald, G. L. Stenchikov, and A. Robock (2002) Global cooling after the eruption of Mount Pinatubo: A test of climate feedback by water vapor. *Science* 296(5568): 727-730, doi:10.1126/science.296.5568.727
- Timmreck C. (2012) Modeling the climatic effects of volcanic eruptions, invited review paper *Wiley Interdisciplinary Reviews: Climate Change*, doi: 10.1002/wcc.192
- Toohey M, K. Krüger, M. Bittner, C. Timmreck, H. Schmidt (2014) The impact of volcanic aerosol on the Northern Hemisphere stratospheric polar vortex: mechanisms and sensitivity to forcing structure, *Atmos. Chem. Phys. Discuss.*, 14, 16777-16819, doi:10.5194/acpd-14-16777-2014, ACP accepted
- Zanchettin, D., C. Timmreck, H.-F. Graf, A. Rubino, S. Lorenz, K. Lohmann, K. Krueger, and J. H. Jungclaus (2012) Bi-decadal variability excited in the coupled ocean–atmosphere system by strong tropical volcanic eruptions. *Clim. Dyn.*, 39:1-2, 419-444, doi:10.1007/s00382-011-1167-1
- Zanchettin, D., O. Bothe, H. F. Graf, S. J. Lorenz, J. Luterbacher, C. Timmreck and J. H. Jungclaus (2013) Background conditions influence the decadal climate response to strong volcanic eruptions, *J. Geophys. Res. Atmos.*, 118, doi:10.1002/jgrd.50229
- Zanchettin, D., et al. (2014a) VolMIP - Model Intercomparison Project on the climate response to volcanic forcing. In preparation
- Zanchettin, D., O. Bothe, C. Timmreck, J. Bader, A. Beitsch, H.-F. Graf, D. Notz and J. H. Jungclaus (2014b) Inter-hemispheric asymmetry in the sea-ice response to volcanic forcing simulated by MPI-ESM (COSMOS-Mill). *Earth Syst. Dynam.*, 5, 223–242, doi:10.5194/esd-5-223-2014

Table 1 – Tier 1 VolMIP experiments

<u>Name</u>	<u>Description</u>	<u>Start year</u>	<u>Configuration</u>	<u>Ens. Size</u>	<u>Years per simulation (minimum)</u>	<u>Total years</u>	<u>Connection with other MIPs</u>	<u>Gaps of knowledge being addressed with this experiment</u>
VolLongS100EQ	Idealized equatorial eruption corresponding to an initial emission of 100 Tg of SO ₂ . This eruption has a magnitude roughly corresponding to the 1815 Tambora eruption, the largest historical tropical eruption, which was linked to the so-called “year without a summer” in 1816	PID (from <i>PiControl</i>)	AOGCM/ESM	9	20	180	PMIP	Uncertainty in the climate response to strong volcanic eruptions, with focus on coupled ocean -atmosphere feedbacks and interannual to decadal global as well as regional responses. The mismatch between reconstructed and simulated climate responses to historical strong volcanic eruptions, with focus on the role of simulated background internal climate variability.
VolShort20EQfull	1991 Pinatubo forcing as used in the CMIP6 <i>historical</i> simulations. Requires special diagnostics of parameterized and resolved wave forcings, radiative and latent heating rates. A large number of ensemble members is required to address internal atmospheric variability	PID	AOGCM/ESM	25	3	75	DYNVAR DCPP	Uncertainty in the climate response to strong volcanic eruptions with focus on short-term response. Robustness of volcanic imprints on Northern Hemisphere’s winter climate and of associated dynamics.

Vol = Volcano, Long = long-term simulation, Short = short-term simulation, S = Single (XXX = approx. amount of Tg of SO₂ release), C = Cluster (XXX = approx. period of the cluster), HL = high latitude, EQ = equator, full = full-forcing simulation, surf = short-wave forcing only, strato = stratospheric thermal (long-wave) forcing only, slab = slab ocean simulation, ini = simulation initialized for decadal prediction

Table 2 – Tier 2 VolMIP experiments

<u>Name</u>	<u>Description</u>	<u>Start year</u>	<u>Configuration</u>	<u>Ens. Size</u>	<u>Years per simulation</u>	<u>Total years</u>	<u>Connection with other MIPs</u>	<u>Gaps of knowledge being addressed with this experiment</u>
VolLongS100HL	Idealized high-latitude (60°N) eruption emitting 100 Tg of SO ₂ over five months. The eruption's strength and length roughly correspond to that of the 1783-84 Laki eruption.	PID	AOGCM/ESM	9	20	180	PMIP, GeoMIP	Uncertainty in climate response to strong high-latitude volcanic eruptions (focus on coupled ocean-atmosphere). Laki has a unique eruption style (large SO ₂ mass releases occurred at short temporal intervals). Outstanding questions about the magnitude of the climatic impact of high-latitude eruptions.
VolLongC19thC	Early 19th century cluster of strong tropical volcanic eruptions, including the 1809 event of unknown location, and the 1815 Tambora and 1835 Cosigüina eruptions.	PID (integration starts on year 1809)	AOGCM/ESM	3	50	150	PMIP, GeoMIP	Uncertainty in the multi-decadal climate response to strong volcanic eruptions (focus on long-term climatic implications). Contribution of volcanic forcing to the climate of the early 19th century, the coldest period in the past 500 years. Discrepancies between simulated and reconstructed climates of the early 19th century.
VolShort20EQsurf	As VolShort20EQfull, but with prescribed surface cooling patterns or net surface flux changes	PID	AOGCM/ESM	25	3	75	DYNVAR DCPP	Mechanism(s) underlying the dynamical atmospheric response to large volcanic eruptions, in particular in Northern Hemisphere's winters. The experiment considers only the effect of volcanically induced surface cooling. Complimentary experiment to VolShort20EQstrat.
VolShort20EQstrat	As VolShort20EQfull, but with prescribed aerosol heating in the stratosphere	PID	AOGCM/ESM	25	3	75	DYNVAR DCPP	Mechanism(s) underlying the dynamical atmospheric response to large volcanic eruptions, in particular in Northern Hemisphere's winter. The experiment considers only the effect of volcanically-induced stratospheric heating. Complimentary experiment to VolShort20EQstrat.

Vol = Volcano, Long = long-term simulation, Short = short-term simulation, S = Single (XXX = approx. amount of Tg of SO₂ release), C = Cluster (XXX = approx. period of the cluster), HL = high latitude, EQ = equator, full = full-forcing simulation, surf = short-wave forcing only, strato = stratospheric thermal (long-wave) forcing only, slab = slab ocean simulation, ini = simulation initialized for decadal prediction

Table 3 – Tier 3 VolMIP experiments

<u>Name</u>	<u>Description</u>	<u>Start year</u>	<u>Configuration</u>	<u>Ens. Size</u>	<u>Years per simulation</u>	<u>Total years</u>	<u>Connection with other MIPs</u>	<u>Gaps of knowledge being addressed with this experiment</u>
VolShort20EQslab	As VolShort20EQfull, but with a slab ocean	PID	AOGCM/ESM	25	3	75	ENSOMIP DCPP	Effects of volcanic eruptions on ENSO dynamics.
VolShort20EQini/ DCPP C2.1	As VolShort20EQfull, but as decadal prediction runs joint experiment with DCPP	PID	AOGCM/ESM	10(5)	10		DCPP	Influence of large volcanic eruptions in future climate. Influence of large volcanic eruptions on seasonal and decadal climate predictability

Vol = Volcano, Long = long-term simulation, Short = short-term simulation, S = Single (XXX = approx. amount of Tg of SO₂ release), C = Cluster (XXX = approx. period of the cluster), HL = high latitude, EQ = equator, full = full-forcing simulation, surf = short-wave forcing only, strato = stratospheric thermal (long-wave) forcing only, slab = slab ocean simulation, ini = simulation initialized for decadal prediction

WCRP COORDINATED REGIONAL DOWNSCALING EXPERIMENT (CORDEX)

Application for CMIP6-Endorsed MIPs

Date: 29 November 2014

- Name of MIP: Coordinated Regional Downscaling Experiment (CORDEX)
- Co-chairs of MIP (including email-addresses)
 - Filippo Giorgi <giorgi@ictp.it> and William Gutowski <gutowski@iastate.edu>
- Members of the Scientific Steering Committee
 - Isabelle Anguelovski <Isabelle.Anguelovski@uab.cat>
 - Hyung-Suk Kang <hyunskang@korea.kr>
 - R. Krishnan <krish@tropmet.res.in>
 - Chris Lennard <lennard@csag.uct.ac.za>
 - Grigory Nikulin <grigory.nikulin@smhi.se>
 - Silvina Solman <solman@cima.fcen.uba.ar>
 - Tannecia Stephenson <tannysyd@yahoo.com>
 - Bertrand Timbal <B.Timbal@bom.gov.au>
 - Fredolin Tangang <ftangang@gmail.com>
 - WCRP liaison: Michel Rixen <mrixen@wmo.int>
- Link to website (if available): <http://wcrp-cordex.ipsl.jussieu.fr/>
- Goal of the MIP and a brief overview:
 - CORDEX has a set of six goals:
 1. To produce quality-control intercomparable data sets of information based on regional climate downscaling (RCD) for the recent historical past and 21st century projections, covering the majority of populated land regions on the globe, aimed at improving understanding of regional to local climate change information and related uncertainties
 2. To build a common set of domains and simulation protocols for dynamical and statistical downscaling activities and define a standard set of variables, frequency and format for output and archival at a number of CORDEX data centers and for distribution via the ESGF infrastructure
 3. To coordinate RCD activities for the defined domains forced by analyses of observations (currently ERA-Interim) aimed at providing a benchmark framework for model evaluation and assessment.
 4. To develop Regional Analysis and Evaluation Teams to evaluate the ensemble of RCD simulations, develop a suitable set of common and region-specific evaluation metrics, collect suitable observational data to evaluate high-resolution RCD output, design experiments to investigate the added-value of RCD methods and the role of regional forcings (e.g. land-use, aerosols) and provide recommendations for future regional priorities in climate research
 5. To engage the broad RCD community in its activities and discussions
 6. To support and inform climate impact assessment and adaptation groups interested in utilizing CORDEX RCD results in their research.

The RCD information samples uncertainties in Regional Climate Change associated with varying forcing GCM simulations and greenhouse gas concentration scenarios, natural climate variability and different downscaling methods. The CORDEX downscaling activities base themselves as much as possible on the latest sets of GCM climate simulations. For example the CORDEX Phase I RCM

experiments were based on driving GCMs participating to CMIP5, which was an invaluable resource for the design and implementation of CORDEX.

More generally, RCD techniques, including both dynamical and statistical approaches, are being increasingly used to provide higher-resolution climate information than is available directly from contemporary global climate models. The techniques available, their applications, and the community using them are broad and varied, and this is a growing area. These techniques, and the results they produce must be applied appropriately and their strengths and weaknesses need to be understood. This requires a better evaluation and quantification of the performance of the different techniques for application to specific problems, along with an understanding of uncertainties underlying regional climate projections. Building on experience gained in the global modelling community, a coordinated, international effort to objectively assess and intercompare various RCD techniques provides a means to evaluate their performance, to illustrate benefits and shortcomings of different approaches, to produce multi-model, multi-method based information and to provide a more solid scientific basis for impact assessments and other uses of downscaled climate information.

The WCRP views regional downscaling as both an important research topic and an opportunity to engage a broader community of climate scientists in its activities. The Coordinated Regional Climate Downscaling Experiment (CORDEX) has served as a catalyst for achieving this goal.

➤ References:

Many papers have been published using simulations in the CORDEX framework; some are listed at <http://wcrp-cordex.ipsl.jussieu.fr/index.php/cordex-peer-review-publications>. Giorgi et al. (2009, "Addressing climate information needs at the regional level: the CORDEX framework", *WMO Bulletin*, **58**, 175-183) and Jones et al. (2011, "The Coordinated Regional Downscaling EXperiment CORDEX, an international downscaling link to CMIP5." *CLIVAR Exchanges*, 16, 34-40) give a brief overview of initial program plans. General updates appear in the *WCRP CORDEX Newsletter* (<http://wcrp-cordex.ipsl.jussieu.fr/index.php/cordex-newsletters>).

➤ An overview of the proposed experiments:

The anticipated CORDEX experiments are downscaling activities that will use CMIP DECK, CMIP6 Historical Simulation and ScenarioMIP output to provide input conditions for both statistical and dynamical downscaling under the CORDEX framework. CORDEX has a general framework of specified regions, resolutions and simulation periods that all regional CORDEX activities adhere to. Specific details of downscaling experiments are a function of plans generated by groups participating in each of the CORDEX regions. In particular, for each region a matrix of GCM-RCD experiments is designed based on the need to cover as much as possible different dimensions of the uncertainty space (different scenarios, GCMs, RCD models and techniques). The dimension of this matrix depends on the participation of groups in the different regional domain activities.

An optimal design of GCM-RCM matrices requires the availability of a broad range of driving GCM data (6 hourly meteorological fields), spanning a high-end, mid level and low-end GHG emission scenario, and all or at least a large portion of GCMs participating in CMIP6. For the initial stages of the CORDEX activities, the focus will be on historical climate simulations for the 20th century and projections for 21st century, implying that data would be needed minimally for the period 1950-2100 (but ideally 1900-2100). Therefore, as for CMIP5, 6-hourly forcing data from one realization of each contributing GCM is a minimal requirement.

CORDEX activities provide a unique opportunity to deliver a full range of the uncertainties attached with regional climate change projections by creating GCM-RCD matrices. It is therefore important that the uncertainties attached to the human activities in the 21st century are encapsulated; multiple scenarios will allow us to evaluate some of the uncertainty due to human choices and are therefore an important additional request should they become available as part of the CMIP6 simulations. In addition, multiple

realizations from some GCMs would allow us to explore also another dimension of the uncertainty space, GCM/RCM internal variability.

- An overview of the proposed evaluation/analysis of the CMIP DECK, CMIP6 Historical Simulation and ScenarioMIP experiments:

CORDEX experiments would use output from

- 1) 30 years of the pre-industrial simulation (CMIP DECK)
- 2) 1950-2014 from the historical climate simulation (CMIP6 Historical Simulation)
- 3) 2015-2100 from the transient scenario climate simulation that uses RCP8.5 and 4.5 for one realization of future projection (ScenarioMIP)

We request RCP4.5 output, even though it is part of ScenarioMIP Tier 2, for continuity with CMIP5-based downscaling that used RCP 4.5. Although one realization is requested providing output from more realizations and more scenarios (from ScenarioMIP) is very welcome.

The downscaling activities will contribute to answering all three of the key questions for CMIP6 through regional simulations with different climate forcings (key question 1), evaluation of physical processes affecting added value and biases in the downscaled results (key question 2) and characterization of the impact of unforced variability, both internally generated and via ensemble boundary conditions, on the ratio of regional climate change signals versus the noise of unforced variability (key question 3).

The downscaling activities will contribute primarily to the WCRP grand challenges of regional climate information and climate extremes. Some of the downscaling will include evaluation of regional feedbacks associated with land-use change and aerosols, along with regional rendition of GCM responses to different climatic forcings.

Downscaled results using CMIP output will be evaluated for their ability to provide added value to the CMIP simulations. This will occur in three ways:

- 1) Analysis during the historical period (1950-2014) will indicate where and when the downscaling provides regional detail of physical behavior that agrees better with observations than the driving GCM output and provides robust additional fine scale climate information.
- 2) Analysis of downscaled projections (2015-2100) will assess where and when the downscaling provides regional detail of physical behavior that exceeds noise levels of unforced internal variability.
- 3) Analysis of downscaled CMIP DECK simulations for the pre-industrial control compared to the transient forcing case will determine potential regional climate-change detection. Should additional CMIP6 simulations occur that specify changes in just one of the major forcings (e.g., solar output, greenhouse gases, volcanic aerosols), then item 3) above would include additional downscaling of those runs with an eye toward regional attribution.

- Proposed timing:

At least some regional modeling groups will be poised to use CMIP DECK, CMIP6 Historical Simulation and ScenarioMIP output suitable for RCM boundary conditions as it becomes available.

The statistical downscaling program under CORDEX is in development. However, some participants in the program have been using CMIP5 output and should be ready to use appropriate CMIP DECK, CMIP6 Historical Simulation and ScenarioMIP output as it becomes available.

- For each proposed experiment to be included in CMIP6: N/A

- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale:

CORDEX is preparing a Memorandum of Understanding for output produced by CORDEX modelers that will follow the availability terms of CMIP output.

➤ List of output and process diagnostics for the CMIP DECK/Historical/ScenarioMIP data request:

CORDEX requests output from the targeted CMIP DECK, CMIP6 Historical Simulation and ScenarioMIP simulations sufficient to allow the downscaling activities and associated analyses listed above:

- 1) Output sufficient for dynamical and empirical statistical downscaling (transient climate-change simulation)
- 2) Output from multiple realizations of the same GCM for both the pre-industrial and transient climate-change simulations, to bring unforced variability into downscaling boundary conditions
- 3) Output that could allow regional detection and attribution work. This would entail boundary conditions from pre-industrial control runs (CMIP DECK) and runs with changes in only one climate forcing (if part of CMIP6).

Output variables needed from CMIP DECK, CMIP6 Historical Simulation and ScenarioMIP runs:

- Preferred output period: 1951-2100 for transient climate change (RCP8.5 and RCP4.5 for 2015-2100); 30 years of pre-industrial control.
- For dynamical downscaling:
 - 6-hourly instantaneous surface pressure
 - 6-hourly instantaneous three-dimensional fields at model levels of temperature, atmospheric specific humidity, zonal wind and meridional wind

We suggest saving these variables to files with the same time period (e.g., 6 months, one year), to ensure uniform time periods covered for a GCM's files for all variables and to avoid very large files (many Gb) that are awkward to handle.

- For statistical downscaling, in addition to the 6-hourly three-dimensional fields listed above, values for integrated quantities will be required:

- maximum daily surface (2m) temperature
 - minimum daily surface (2m) temperature
 - daily surface temperature (2m)
 - daily surface dewpoint temperature (2m)
 - daily zonal wind (10m)
 - daily meridional wind (10m)
 - daily precipitation
 - daily vertical atmospheric column of water (or precipitable water)
 - monthly sea surface temperature
- Supplementary variables that are desirable:
- daily soil moisture (vertically integrated)
 - daily snow density
 - daily snow albedo
 - daily low and medium cloud cover
 - 6-hourly instantaneous geopotential height at 850, 700 and 500 hPa

➤ Any proposed contributions and recommendations for observations

Assessments of added value will seek fine resolution (25-50 km or less) observational datasets. The obs4MIPs and ana4MIPs efforts are potentially useful and there is already some CORDEX interaction with obs4MIPs. For some regions, fine resolution observational datasets are being sought in all CORDEX regions, especially those that could support evaluation of higher resolution CORDEX runs. CORDEX will help with efforts to make new datasets accessible in standardized formats via the ESGF infrastructure.

- Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms: NONE

- Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF: NONE

DynVar – Diagnostic MIP

Dynamics and Variability of the Stratosphere - Troposphere System

Co-Chairs:

Edwin Gerber (gerber@cims.nyu.edu)

Elisa Manzini (elisa.manzini@mpimet.mpg.de)

Members of the Scientific Steering Committee:

Amy Butler (amy.butler@noaa.gov)

Natalia Calvo (nataliac@fis.ucm.es)

Andrew Charlton-Perez (a.j.charlton-perez@reading.ac.uk)

Marco Giorgetta (marco.giorgetta@mpimet.mpg.de)

Adam Scaife (Adam.scaife@metoffice.gov.uk)

Tiffany Shaw (tas2163@columbia.edu)

Shingo Watanabe (wnabe@jamstec.go.jp)

Website: <http://www.sparcdynvar.org/>

Goal of the MIP and a brief overview

DynVar focuses on the interactions between atmospheric variability, dynamics and climate change, with a particular emphasis on the two-way coupling between the troposphere and the stratosphere. The key questions addressed by the activity are:

- How do dynamical processes contribute to persistent model biases in the mean state and variability of the atmosphere, including biases in the position, strength, and statistics of blocking events, storm tracks and the stratospheric polar vortex?
- How does the stratosphere affect climate variability at intra-seasonal, inter-annual and decadal time scales?
- What is the role of dynamics in shaping the atmospheric circulation response to anthropogenic forcings (e.g. global warming, ozone depletion) and how do dynamical processes contribute to uncertainty in future climate projections?

An overview of the proposed experiment

Rather than proposing new experiments, we are *requesting additional output*, critical for understanding the role of atmospheric dynamics in both present and past climate, and future climate projections. *Without this output, we will not be able to fully assess the dynamics of mass, momentum, and heat transport - essential ingredients in projected circulation changes - nor take advantage of the increasingly accurate representation of the stratosphere in coupled climate models.* Our rationale is that by simply extending the standard output relative to that in CMIP5, there is potential for significantly expanding our research capabilities in atmospheric dynamics.

An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments

Understanding circulation changes in the atmosphere, particularly of the mid-latitude storm tracks, has been identified by the World Climate Research Programme (WCRP)

as one of the grand challenges in climate research. Changes in the storm tracks are significantly coupled with lower atmosphere processes such as surface temperature gradients and moisture availability (e.g. Booth et al. 2013) as well as with processes in the stratosphere, from natural variability on synoptic to intraseasonal timescales (e.g. Baldwin and Dunkerton 2001) to the response to changes in stratospheric ozone (e.g. Son et al. 2008) and other anthropogenic forcings (e.g. Scaife et al. 2012). The storm tracks depend critically on the transport of momentum, heat and chemical constituents throughout the whole atmosphere. Both resolved (primarily Rossby) and parameterized (gravity) waves play the key roles in these transports, and it is important that the standard output of the DECK experiments, the CMIP6 Historical Simulation and (in principle) any MIP experiment allow proper diagnosis of these wave fluxes.

The lack of output is particularly acute in the stratosphere, where daily means of standard variables (e.g., zonal and meridional winds, geopotential height and temperature) and parameterized gravity wave forcings (a key driver of the circulation) were not well documented in CMIP5, and resolved waves could at best be coarsely assessed, given the importance of the vertical structure to momentum and mass transport. As detailed by Hardiman et al. (2013), the stratospheric community had to rely on direct collaboration to obtain necessary diagnostics to assess the Brewer-Dobson circulation, the first order circulation of mass and momentum in the stratosphere. Daily means of standard variables in both the troposphere and stratosphere would expand our ability to assess the synoptic dynamics of the atmosphere.

Investigation of the impact of solar variability and volcanic eruptions on climate also relies heavily on atmospheric wave forcing diagnostics, as well as radiative heating rates (particularly in the short wave). By extending our request to the energy budget and including diagnostics such as diabatic heating from cloud-precipitation processes, research on the links between moist processes and atmospheric dynamics will be enabled as well. The interplay between moist processes and circulation is central to the WCRP Grand Challenge on Clouds, Circulation and Climate Sensitivity (Bony et al. submitted to Nature Geoscience, 2014).

The CMIP5 saw a significant upward expansion of models with a more fully resolved stratosphere (e.g. Gerber et al. 2012), and several multi-model studies have investigated the role of the stratosphere in present climate and in projections of future climate (e.g., Anstey et al. 2013; Charlton-Perez et al. 2013; Gerber and Son, 2014; Hardiman et al. 2013; Lott et al. 2014; Manzini et al. 2014; Min and Son 2013; Shaw et al. 2014; Wilcox and Charlton-Perez 2013) in addition to many other single model studies. These studies document a growing interest in the role of middle and upper atmosphere in climate, research that would take full advantage of these diagnostics.

Key science questions of CMIP6: DynVar primarily addresses CMIP6 key science questions on the origin and consequences on systematic models biases in the context of atmospheric dynamics and on the storm track theme of the Clouds, Circulation and Climate Sensitivity Grand Challenge, by further enabling and stimulating research on atmospheric dynamics and storm tracks with CMIP models. We envision as well contributions to the questions on how the Earth System responds to forcing, assessments of future climate changes, and on the Grand Challenges on Regional

Climate Information, Climate Extremes and on the Biospheric Forcings and Feedbacks theme.

Synergy with other MIPs: We envision analyses of the atmospheric circulation with the DECK experiments at the highest priority. Availability of dynamically oriented diagnostics within the DECK and for the CMIP6 Historical Simulation will also provide the benchmark for any other MIP. In addition, we envision fruitful potential collaborations with the following proposed MIPs: AerChemMIP, DAMIP, DCP, ENSOMIP, SolarMIP and VOLMIP.

List of output and process diagnostics for the CMIP DECK/CMIP6 data request

We stress the need of archiving standard variables (e.g. zonal and meridional winds, temperature, and geopotential height) as daily means in the troposphere and stratosphere. We expect that the location and total number of vertical pressure levels for daily mean fields will be discussed during the definition of the standard output.

We request archival of the Transformed Eulerian Mean (TEM) atmospheric circulation, which allows diagnosis of resolved wave driving and transport, and of parameterized atmospheric gravity wave driving. These diagnostics are also widely used in the analysis of chemistry climate models (e.g. CCMVal and CCMI, here AerChemMIP). The TEM diagnostics are particularly sensitive to vertical resolution and model formulation (Hardiman et al. 2010), and so ideally computed following the model's dynamical core assumptions and on the native grid of the model, before coarsened for archival. In addition, we request the archival of heating rates. *Note that the requested diagnostics are 2-D fields (zonal means) on an atmospheric grid defined by latitudes and pressure levels. We are targeting both daily and monthly diagnostics.*

List of proposed variables:

long name	units	comment
residual northward wind	ms ⁻¹	Transformed Eulerian Mean diagnostic calculated from high frequency (6hr or shorter time intervals) atmospheric fields. Reference: Andrews et al (1987): Middle Atmospheric Dynamics. Academic Press.
residual upward wind	ms ⁻¹	Transformed Eulerian Mean diagnostic calculated from high frequency (6hr or shorter time intervals) atmospheric fields. Reference: Andrews et al (1987): Middle Atmospheric Dynamics. Academic Press.
residual mean mass stream function	kgs ⁻¹	Transformed Eulerian Mean diagnostic calculated from high frequency (6hr or shorter time intervals) atmospheric fields. Reference: Andrews et al (1987): Middle Atmospheric Dynamics. Academic Press.

northward EP-flux	Nm^{-1}	Transformed Eulerian Mean diagnostic calculated from high frequency (6hr or shorter time intervals) atmospheric fields. Reference: Andrews et al (1987): Middle Atmospheric Dynamics. Academic Press.
upward EP-flux	Nm^{-1}	Transformed Eulerian Mean diagnostic calculated from high frequency (6hr or shorter time intervals) atmospheric fields. Reference: Andrews et al (1987): Middle Atmospheric Dynamics. Academic Press.
EP-flux divergence	$\text{ms}^{-1}\text{d}^{-1}$	Transformed Eulerian Mean diagnostic calculated from high frequency (6hr or shorter time intervals) atmospheric fields. Reference: Andrews et al (1987): Middle Atmospheric Dynamics. Academic Press.
u-tendency by residual northward wind advection	$\text{ms}^{-1}\text{d}^{-1}$	Transformed Eulerian Mean diagnostic calculated from high frequency (6hr or shorter time intervals) atmospheric fields. Reference: Andrews et al (1987): Middle Atmospheric Dynamics. Academic Press.
u-tendency by residual upward wind advection	$\text{ms}^{-1}\text{d}^{-1}$	Transformed Eulerian Mean diagnostic calculated from high frequency (6hr or shorter time intervals) atmospheric fields. Reference: Andrews et al (1987): Middle Atmospheric Dynamics. Academic Press.
u-tendency by orographic gravity waves	$\text{ms}^{-1}\text{d}^{-1}$	Zonal mean of eastward wind tendency by orographic gravity wave parameterization
v-tendency by orographic gravity waves	$\text{ms}^{-1}\text{d}^{-1}$	Zonal mean of northward wind tendency by orographic gravity wave parameterization
u-tendency by non-orographic gravity waves	$\text{ms}^{-1}\text{d}^{-1}$	Zonal mean of eastward wind tendency by non-orographic gravity wave parameterization
v-tendency by non-orographic gravity waves	$\text{ms}^{-1}\text{d}^{-1}$	Zonal mean of northward wind tendency by non-orographic gravity wave parameterization
mean age of air	years	Zonal mean of mean age of air
longwave heating rate	Kd^{-1}	Zonal mean of heating from longwave radiation
shortwave heating rate	Kd^{-1}	Zonal mean of heating from shortwave radiation
latent heating rate	Kd^{-1}	Zonal mean of heating from cloud and precipitation processes

References

- Anstey, J. A. and Coauthors (2013), Multi-model analysis of Northern Hemisphere winter blocking: Model biases and the role of resolution, *J. Geophys. Res. Atmos.*, 118, 3956–3971, doi: 10.1002/jgrd.50231
- Baldwin, M. P., and T. J. Dunkerton (2001), Stratospheric harbingers of anomalous weather regimes. *Science*, 294, 581–584.
- Booth, J. F., S. Wang, L. Polvani (2013), Midlatitude storms in a moister world: lessons from idealized baroclinic life cycle experiments. *Climate Dynamics* 41, 787–802, doi: 10.1007/s00382-012-1472-3
- Charlton-Perez, A. J. and Coauthors (2013), On the lack of stratospheric dynamical variability in low-top versions of the CMIP5 models, *J. Geophys. Res. Atmos.*, 118, 2494–2505, doi: 10.1002/jgrd.50125
- Gerber, E. P. and Coauthors (2012), Assessing and Understanding the Impact of Stratospheric Dynamics and Variability on the Earth System, *Bull. Amer. Meteor. Soc.*, 93, 845-859, doi: 10.1175/BAMS-D-11-00145.1
- Gerber, E. P. and S.-W. Son, (2014), Quantifying the Summertime Response of the Austral Jet Stream and Hadley Cell to Stratospheric Ozone and Greenhouse Gases. *J. Climate*, 27, 5538-5559, doi: 10.1175/JCLI-D-13-00539.1
- Hardiman, S. C., N. Butchart, N. Calvo (2013), The morphology of the Brewer-Dobson circulation and its response to climate change in CMIP5 simulations, *Q. J. R. Meteorol. Soc.*, doi: 10.1002/qj.2258
- Hardiman, S.C. et al. (2010), Using Different Formulations of the Transformed Eulerian Mean Equations and Eliassen–Palm Diagnostics in General Circulation Models, *J. Atmos. Sci.*, 67, 1983-1995. DOI: 10.1175/2010JAS3355.1
- Lott, F. and Coauthors (2014), Kelvin and Rossby-gravity wave packets in the lower stratosphere of some high-top CMIP5 models, *J. Geophys. Res. Atmos.*, 119, 2156–2173, doi: 10.1002/2013JD020797
- Manzini, E. and Coauthors (2014), Northern winter climate change: Assessment of uncertainty in CMIP5 projections related to stratosphere-troposphere coupling, *J. Geophys. Res. Atmos.*, 119, doi: 10.1002/2013JD021403
- Min, S.-K. and S.-W. Son (2013), Multi-model attribution of the Southern Hemisphere Hadley cell widening: major role of ozone depletion, *J. Geophys. Res. Atmos.*, 118, 3007-3015.
- Scaife, A. A. and Coauthors (2012) Climate change projections and stratosphere-troposphere interaction. *Climate Dyn.* doi: 10.1007/s00382-011-1080-7
- Shaw, T. A., J. Perlwitz, O. Weiner (2014), Troposphere-stratosphere coupling: Links to North Atlantic weather and climate, including their representation in CMIP5 models. *J. Geophys. Res.*, 10.1002/2013JD021191
- Son, S.-W. and Coauthors (2008) The impact of stratospheric ozone recovery on the Southern Hemisphere westerly jet. *Science*, 320, 1486–1489.
- Wilcox, L. and A. Charlton-Perez (2013), Final warming of the Southern Hemisphere polar vortex in high- and low-top CMIP5 models. *J. Geophys. Res. Atmos.*, 118, doi: 10.1002/jgrd.50254

Global Dynamical Downscaling Experiment (GDDEX)

Application for CMIP6-Endorsed MIPs

Revised Date: 3 December 2014

Now a Diagnostic MIP (i.e., no proposed experiments rather requesting that certain output is archived)

➤ Name of MIP*:

Global Dynamical Downscaling Experiment (GDDEX)

➤ Co-chairs of MIP (including email-addresses)*:

Kei Yoshimura (kei@ori.u-tokyo.ac.jp, Atmosphere and Ocean Research Institute, the University of Tokyo, Japan), Hans von Storch (hvonstorch@web.de, Institute of Coastal Research, Helmholtz Center Geesthacht, Geesthacht, Germany)

➤ Members of the Scientific Steering Committee*:

Hyungjun Kim, Martina Schubert-Frisius, Frauke Feser, Izuru Takayabu, Song-You Hong, Suryun Ham, Eun-Chul Chang, Tomohito Yamada

➤ Link to website (if available)*:

N/A

➤ Goal of the MIP and a brief overview*:

Earth System Models have been more and more sophisticated and complicated, as results, more and more expensive. However the computing speed and resources has not been improved so much as expected. We have had dynamically downscaling projects such as CORDEX for the needs of spatially high resolution simulations. In CORDEX, multiple Regional Climate Models simulations have been compiled in several regional domains to provide 'added regional information' contributing to one of WCRP's Grand Challenges. However, it resulted in a considerable domain overlapping globally. This does not only cause the redundant resources globally but also controversies in the inconsistent lateral boundary conditions and model ensemble set between the domains.

The goal of the GDDEX project is to provide a seamless suite of hi-res global climate keeping the consistency with the CMIP6, and to investigate the response of the Earth System to the forcing (CMIP6 key scientific question 1) in a higher spatiotemporal resolution (e.g., local extreme events). The outcome is also useful for the land surface communities (e.g., GLASS/GEWEX) and the impact assessment studies (e.g., ISI-MIP) in their offline simulations. As results, GDDEX would contribute WCRP Grand Challenges, particularly for Climate Extremes and Regional Climate Information.

In GDDEX, we propose two sets of experiments, i.e., Atmosphere-Forced Experiments and SST-Forced Experiments. The former uses the spectral nudging concept (von Storch et al., 2000); large scale (over 1000 km horizontally) kinetic atmospheric wave is constantly enforced to create smaller scale waves. It was first performed by Yoshimura and Kanamitsu (2008), and added values in spatiotemporally detailed information are demonstrated by Feser et al. (2011), Kim and Hong (2012), and Chang et al. (2014). This type will act as a magnifying glass of low resolution Earth System Models' climates but for the whole globe. This type of experiment directly contributes the investigation on the response of the Earth System to the forcing (key question 1), and it partly contributes to determine the significance of the future climate change signal with regards to the current climate variability with regional details and contrasts (key question 3).

The latter, SST-Forced Experiments, uses only SST fields generated from low resolution Earth System Models. There is model-dependent bias for the atmospheric modes, such as annular modes, so that this experiment relaxes the constraint to freely create such atmospheric modes by AGCM. This type of experiments is of course very popular (e.g., represented by AMIP; Mizuta et al., 2012), but in this project, we will use ESM-derived SST and compare with the Atmosphere-Forced Run, i.e., with and without the large scale kinetic atmospheric wave, to quantify the impacts by the SST-driven or Atmospheric-driven forcings on spatiotemporally detailed climate changes. By comparing the

SST-Forced Experiments and Atmosphere-Forced Experiments, we expect to partly contribute the question of origin and consequences of systematic model biases (key question 2).

We anticipate participation of AGCMs not only those participating CMIP6 as a part of coupled model system (e.g., ECHAM6), but also those not participating CMIP6 and individually developed as a stand-alone atmospheric model (e.g., Scripps Institution of Oceanography GSM).

➤ **References (if available)*:**

- *von Storch, H., H. Langenberg and F. Feser, 2000: A spectral nudging technique for dynamical downscaling purposes. Mon. Wea. Rev. 128: 3664-3673*
- *Yoshimura, K., and M. Kanamitsu, 2008: Dynamical global downscaling of global reanalysis. Mon. Wea. Rev. 136: 2983-2998*
- *Feser, F., B. Rockel, H. von Storch, J. Winterfeldt, and M. Zahn, 2011: Regional Climate Models add Value to Global Model Data: A Review and selected Examples. Bull. Amer. Meteor. Soc. 92: 1181–1192*
- *Kim, J.-E. and S.-Y. Hong, 2012: A global atmospheric analysis dataset downscaled from the NCEP-DOE Reanalysis, J. Climate, 25: 2527–2534*
- *Chang, E.-C., S.-W. Yeh, S.-Y. Hong, J.-E. Kim, R. Wu, and K. Yoshimura, 2014: Study on the changes in the East Asian precipitation in the mid-1990s using a high-resolution global downscaled atmospheric data set, J. Geophys. Res. Atmos., 119, doi:10.1002/2013JD020903.*
- *Mizuta, R, et al., 2012: Climate simulations using MRI-AGCM3.2 with 20-km grid. J. Meteor. Soc. Japan, 90A: 233–258.*
- *Yoshimura, K. and M. Kanamitsu, 2009: Specification of external forcing for regional model integrations, Mon. Wea. Rev., 137, 1409–1421.*
- *Compo, G.P. et al., 2011: The Twentieth Century Reanalysis Project. Quarterly J. Roy. Meteorol. Soc., 137, 1-28. DOI: 10.1002/qj.776.*

➤ **An overview of the proposed experiments*:**

We anticipate multiple AGCMs to participate in GDDEX by using output data of CMIP6 Historical Simulation and CMIP6 ScenarioMIP Simulations as input data for initial and boundary conditions. Each participant, who doesn't have to be a model-developing body, runs the experiments under the GDDEX framework (briefly described below). GDDEX is dynamical downscaling activity not for specific regions, but for the whole globe at once. GDDEX has an objective to compare global models in terms of dynamical downscaling capability. That is one of the unique objectives.

The Atmosphere-Forced Experiments will be mainly achieved by the global spectral nudging techniques using multiple AGCMs. With typical spectral nudging technique, horizontally larger scale waves than 1000 km of wind fields are enforced with appropriate nudging coefficients. The nudging coefficients may vary in vertical levels. Those atmospheric boundary data will be given from CMIP6 models in vertical levels and at 6-hourly interval. In case these data is not fully available in the CMIP6 data archive, data at least 3 layers and daily interval can be usable (Yoshimura and Kanamitsu, 2009). Ocean data will be given for only surface (AMIP-like). The SST-Forced Experiments will be additionally done with all the same settings as Atmosphere-Forced Experiments but without the atmospheric boundary data and the spectral nudging. Target horizontal resolution is 20 km globally and common to all experiments (but it is flexible with individual models).

○ *Historical Reanalysis Global Downscaling Experiment (HR-GDEX)*

Purpose: *Generate a reference historical climate fields for HC-GDEX and FC-GDEX experiments.*

Time-span: *Preferably 1950-2014 but at least including the last part of 20th century (1980-2010) up to data availability*

Input Data source: *20th Century Reanalysis (Compo et al., 2011) in 2 degree for 1871-2012*

Remarks:

This experiment will contribute to the GSWP3 (a GLASS/GEWEX endorsed project lead by Hyungjun Kim) and LMIP-Hist experiment of LS3MIP which is another proposed satellite MIP of CMIP6 (co-chaired by Hyungjun Kim).

- *Historical CMIP6 Global Downscaling Experiment (HC-GDEX)*
Purpose: Generate 'added-values' on the historical CMIP6 Historical simulations
Time-span: Preferably, 1950-2014 but the core stream can be determined (e.g., 1980-2010) to reduce required resources.
Input Data source: Selected GCM projections of CMIP6 Historical experiments
Remarks: Long-term representation of the atmospheric system responding to the forcing is evaluated by comparing with the HighResMIP experiments.
- *Future CMIP6 Global Downscaling Experiment (FC-GDEX)*
Purpose: Generate 'added-values' on the CMIP6 ScenarioMIP simulations
Time-span: Preferably, 2015-2100 but the core stream can be determined (e.g., 2020-2050 and 2070-2100) to reduce the size of experiment.
Input Data source: CMIP6 ScenarioMIP experiments for the same model subset used in HC-GDEX experiment with selected (at least to) scenarios
Remarks: Resolution dependent model bias and its propagation in the system will be investigated. Downscaled surface fields will contribute to the GSWP3 and LMIP-Fut experiment of LS3MIP.

➤ An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*:

GDDEX would use output from CMIP6 experiments as follows:

1. *CMIP6 Historical Simulation (1950-2014)*
2. *CMIP6 ScenarioMIP Simulations (2015-2100)*

How many scenarios are requested from ScenarioMIP simulations has not been fixed yet, but higher end (RCP8.5) and lower end (RCP2.6) would be appropriate.

We would contribute to answering the three scientific questions of CMIP6, by providing a seamless suite of hi-res global climate keeping the consistency with the CMIP6, particularly by investigating the response of the Earth System to the forcing in a higher spatiotemporal resolution (key question 1). In addition to that, we expect to contribute to elucidate the origin and consequence of systematic model biases by comparing the SST-Forced experiment and Atmosphere-Forced experiment (key question 2), and to figure our regional details and contrasts of the significance of climate change signals (key question 3).

➤ Proposed timing*:

- *2014: Preparation. Call for participants.*
- *2014-2016: HR-GDEX completed.*
- *2016-2019: HC-GDEX and FC-GDEX completed.*

➤ For each proposed experiment to be included in CMIP6**:

N/A

➤ If possible, a prioritization of the suggested experiments, including any rationale**

None

➤ All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**:

No objection

➤ List of output and process diagnostics for the CMIP DECK/CMIP6 data request**

- whether the variable should be collected for all CMIP6 experiments, or only some specified subset and whether the output is needed from the entire length of each experiment or some shorter period or periods;

Variables classified as 6hrLev in CMOR, which are global 3-dimensional instantaneous atmospheric states (U, V, T, q) and surface pressure (Ps), and SST for every 6-hour through the entire periods of Nucleus CMIP6 (1850-2014) and ScenarioMIP.

- o whether the output might only be relevant if certain components or diagnostic tools are used interactively (e.g. interactive carbon cycle or atmospheric chemistry, or only if the COSP simulator has been installed);
No
- o whether this variable is of interest to downstream users (such as impacts researchers, WG2 users) or whether its principal purpose is for understanding and analysis of the climate system itself. Be as specific as possible in identifying why the variable is needed.
The requested variables are not of interest to downstream users, but our high resolution products will be used by downstream users as similar as regional downscaling projects, e.g., CORDEX.
- o whether the variables can be regridded to a common grid, or whether there is essential information that would be compromised by doing this;
Regridded data is OK (either in horizontal and vertical. We have method to recover the compromised information.), however, global domain is crucially required (any regional trimming should not be done).
- o the relative importance of the various variables requested (indicated by a tiered listing) is required if the data request is large.
Among above-mentioned variables (U, V, T, q, Ps, SST), q is less important.

➤ Any proposed contributions and recommendations for**

- o model diagnostics and performance metrics for model evaluation;
None
- o observations/reanalysis data products that could be used to evaluate the proposed experiments. Indicate whether these are available in the obs4MIPs/ana4MIPs database or if there are plans to include them;
None
- o tools, code or scripts for model benchmarking and evaluation in open source languages (e.g., python, NCL, R)..:
N/A

➤ Any proposed changes from CMIP5 in NetCDF metadata (controlled vocabularies), file names, and data archive (ESGF) search terms. **:

Probably none.

➤ Explanation of any proposed changes (relative to CMIP5) that will be required in CF, CMOR, and/or ESGF. **:

Probably none.

SEA-ICE MODEL INTERCOMPARISON PROJECT (SIMIP)

CO-CHAIRS OF MIP (INCLUDING EMAIL-ADDRESSES) *

Alexandra Jahn (NCAR, US, ajahn@ucar.edu)

Dirk Notz (Max Planck Institute for Meteorology, Germany, dirk.notz@mpimet.mpg.de)

Members of the Scientific Steering Committee *

Marika Holland (NCAR, US)

Elizabeth Hunke (Los Alamos National Laboratory, US)

Francois Massonet (Université catholique de Louvain, Belgium)

Julienne Stroeve (NSIDC, US)

Bruno Tremblay (McGill University, Canada)

Martin Vancoppenolle (Laboratoire d'Océanographie et du Climat, France)

LINK TO WEBSITE (IF AVAILABLE) *

Not available yet.

GOAL OF THE MIP AND A BRIEF OVERVIEW *

Defines variables that are necessary to analyze sea-ice evolution in any CMIP6 experiment

This purely diagnostic MIP defines a list of variables that capture the evolution of sea ice in any experiment carried out as part of CMIP6. Given the importance of sea ice both as a driver and as an indicator of climatic changes, the analysis of the changing sea-ice cover in CMIP6 experiments provides insight into the time-integrated evolution of the climate system. To obtain all necessary information for such analysis for any given CMIP6 experiment is the overarching goal of this MIP.

To achieve this aim, we propose a list of those variables that are required to close the three budgets that govern the evolution of sea ice and its impact on the Earth's climate system. These are the conservation of heat, the momentum balance and tracer conservation. In addition, we provide a list of variables that allow for the high frequency analysis of the sea-ice state itself. We aim for the best possible compromise of output frequency and necessity of high-resolution sampling for closing the budgets. To achieve this aim, we group the variables according to their priorities, with the variables of the highest priority being necessary for a basic analysis of the sea-ice evolution in any CMIP6 experiment. By making sure that budgets can be closed, the analysis of sea ice in CMIP6 simulations has the potential to focus on processes rather than only on the sea ice state, leading to improved understanding of the biases in sea ice and the fidelity of projections of sea ice.

AN OVERVIEW OF THE PROPOSED EXPERIMENTS *

None

We do not propose any sea-ice specific experiments. Instead, we clearly define a list of variables that allow any scientist to analyze the sea-ice state in any experiment that is carried out as part of CMIP6. The list of variables is accompanied by guidance as to how a standardized analysis of sea-ice evolution can be carried out that will allow for the straight-forward comparison of sea-ice evolution across different MIPs.

AN OVERVIEW OF THE PROPOSED EVALUATION/ANALYSIS OF THE CMIP DECK AND CMIP6 EXPERIMENTS *

Variables defining sea-ice state and external forcing

The variables that are proposed in this MIP can be divided into those that determine the sea-ice state and those that determine the external forcing that changes this sea-ice state. State variables include standard output such as sea-ice area fraction and thickness, but also more advanced variables with lower priority, such as melt-pond coverage and information on the ice-thickness distribution for those models that have such information available.

Regarding the forcing, the proposed variables both on the atmospheric and on the oceanic side allow for a closure of the main budgets. Hence, they include a description of all heat fluxes that affect the ice, the momentum forcing and the transport of tracers into and out of the ice. Also these forcing variables are split into routine output and more advanced measures.

PROPOSED TIMING *

In parallel with all CMIP6 experiments

At least the standard variables to which we assign priority 1 in this MIP should be saved from any experiment carried out as part of CMIP6.

FOR EACH PROPOSED EXPERIMENT TO BE INCLUDED IN CMIP6 **

We do not propose individual experiments. We therefore here only summarize why the variables we propose should be saved as output from any CMIP6 experiment.

THE EXPERIMENTAL DESIGN

The variables of this MIP should be saved independent of the experimental design

THE SCIENCE QUESTION AND/OR GAP BEING ADDRESSED WITH THIS EXPERIMENT;

The variables of this MIP will allow scientists to answer, for example, the following science questions:

- How sensitive is the sea-ice cover to changes in the external forcing?

- Are these changes primarily driven by changes in the atmosphere or in the ocean?
- What causes biases in the simulation of the sea-ice state?
- How much do simulations of the Earth's climate profit from improvements in the sea-ice model component?
- What's the internal variability of the Earth's sea-ice cover?
- How predictable is the sea-ice cover on time scales ranging from daily to decadal?
- What are the most pressing needs for observations?

POSSIBLE SYNERGIES WITH OTHER MIPs;

Sea ice integrates changes in the atmospheric and oceanic forcing

The variables that we propose will allow the users of any MIP to analyse and to understand the temporal evolution of the sea-ice cover in their simulations. Since the sea-ice state reflects changes in the climate system of the Earth on decadal time scales, changes in the sea-ice cover usually provide direct insight into Earth-System response to climate changes on time scales that are between the atmospheric and the oceanic response time. Such analysis is hence helpful in understanding the temporal evolution of changes in these other compartments of the Earth System.

POTENTIAL BENEFITS OF THE EXPERIMENT TO (A) CLIMATE MODELING COMMUNITY, (B) INTEGRATED ASSESSMENT MODELLING (IAM) COMMUNITY, (C) IMPACTS ADAPTATION AND VULNERABILITY (IAV) COMMUNITY, AND (D) POLICY MAKERS.

- (A) Sea ice is both an integrator and a driver of changes in the climate system. The SIMIP protocol will allow the climate modeling community to understand and to compare the underlying budgets and to hence quantify the role of sea ice for any given experimental setup.
- (B) Changes in the polar sea-ice cover are currently among the most directly observable ones in the Earth's climate system. By allowing for a better understanding of the ongoing changes through the protocol defined here, the IAM community will be able to better estimate the reliability of their models during a period of already observed, significant changes in a specific climate variable.
- (C) Since changes in the sea-ice cover are already ongoing, understanding these changes through the protocol defined here and assessing the reliability of modeled changes allows for a direct assessment of the quality of IAV models.
- (D) Changes in sea ice are one of the most direct measures of ongoing changes in the Earth's climate system. Earth System Models' capability to simulate these changes is a key aspect to underpin the models' credibility for policy makers. Hence, understanding any mismatch between models and observations will be central for understanding the robustness of these simulations for policy decisions. Such understanding will be possible for the modeled sea-ice cover through the protocol that we suggest here.

LIST OF OUTPUT AND PROCESS DIAGNOSTICS FOR THE CMIP DECK/CMIP6 DATA REQUEST **

WHETHER THE VARIABLE SHOULD BE COLLECTED FOR ALL CMIP6 EXPERIMENTS, OR ONLY SOME SPECIFIED SUBSET AND WHETHER THE OUTPUT IS NEEDED FROM THE ENTIRE LENGTH OF EACH EXPERIMENT OR SOME SHORTER PERIOD OR PERIODS;

The variables should be collected for all CMIP6 experiments when possible, ranked by priority if the full set cannot be provided. In addition to the normal monthly data for most variables, we will define certain short periods where daily data should be saved for all variables, in order to allow for a more in-depth analysis.

WHETHER THE OUTPUT MIGHT ONLY BE RELEVANT IF CERTAIN COMPONENTS OR DIAGNOSTIC TOOLS ARE USED INTERACTIVELY (E.G. INTERACTIVE CARBON CYCLE OR ATMOSPHERIC CHEMISTRY, OR ONLY IF THE COSP SIMULATOR HAS BEEN INSTALLED);

The variables of priority 1 are always relevant. Some variables of lower priority (ice-thickness distribution, tracer transport, melt-pond coverage) are only relevant if respective model components are being used

WHETHER THIS VARIABLE IS OF INTEREST TO DOWNSTREAM USERS (SUCH AS IMPACTS RESEARCHERS, WG2 USERS) OR WHETHER ITS PRINCIPAL PURPOSE IS FOR UNDERSTANDING AND ANALYSIS OF THE CLIMATE SYSTEM ITSELF. BE AS SPECIFIC AS POSSIBLE IN IDENTIFYING WHY THE VARIABLE IS NEEDED.

The principle purpose of the variable request is to analyze the climate system.

WHETHER THE VARIABLES CAN BE REGRIDDED TO A COMMON GRID, OR WHETHER THERE IS ESSENTIAL INFORMATION THAT WOULD BE COMPROMISED BY DOING THIS;

Variables should not be regridDED, since this will not allow an a posteriori closure of any budget. However, we provide a short list of very basic variables that could additionally be provided on a common grid or as integrated quantity (i.e., hemispheric sea ice extent, area, and volume) to simplify a superficial analysis of model output

THE RELATIVE IMPORTANCE OF THE VARIOUS VARIABLES REQUESTED (INDICATED BY A TIERED LISTING) IS REQUIRED IF THE DATA REQUEST IS LARGE.

We group the variables in 3 priority levels.

ANY PROPOSED CONTRIBUTIONS AND RECOMMENDATIONS FOR **

MODEL DIAGNOSTICS AND PERFORMANCE METRICS FOR MODEL EVALUATION;

OBSERVATIONS/REANALYSIS DATA PRODUCTS THAT COULD BE USED TO EVALUATE THE PROPOSED EXPERIMENTS. INDICATE WHETHER THESE ARE AVAILABLE IN THE OBS4MIPS/ANA4MIPS DATABASE OR IF THERE ARE PLANS TO INCLUDE THEM;

TOOLS, CODE OR SCRIPTS FOR MODEL BENCHMARKING AND EVALUATION IN OPEN SOURCE LANGUAGES (E.G., PYTHON, NCL, R).

These usually depend on the overarching experiments that save sea-ice variables as part of their output. For the last few decades, satellite observations with varying accuracy are available of sea-ice drift, sea-ice concentration, sea-ice thickness, sea-ice age, sea-ice area, sea-ice volume and sea-ice extent. We will provide guidance to using these products in our variable description.

ANY PROPOSED CHANGES FROM CMIP5 IN NETCDF METADATA (CONTROLLED VOCABULARIES), FILE NAMES, AND DATA ARCHIVE (ESGF) SEARCH TERMS

Updated list of variables, saved in netCDF4 format.

EXPLANATION OF ANY PROPOSED CHANGES (RELATIVE TO CMIP5) THAT WILL BE REQUIRED IN CF, CMOR, AND/OR ESGF

Updated list of variables, saved in netCDF4 format

Vulnerability, Impacts and Adaptation Advisory Board for CMIP6 (VIAAB)

Application for CMIP6-Endorsed MIPs

Date: 22 September 2014

Diagnostic MIP (i.e., no proposed experiments rather requesting that certain output is archived and/or contributing to the evaluation)

This proposal is still under development. The revised version will include other aspects of VIA/CMIP links. If you are interested in updates, please contact the co-chairs.



- Name of Proposed Activity*: VIA Advisory Board
- Co-chairs of MIP (including email-addresses)*: Cynthia Rosenzweig (crr2@columbia.edu) and international co-chair
- Board Coordinator: Alex Ruane (alexander.c.ruane@nasa.gov)
- Proposed Members of the Scientific Steering Committee*: Leaders of major Vulnerability, Impacts and Adaptation Sectors (potentially including: Jean Palutikof, Rob Swart, Dennis Lettenmaier, Dennis Ojima, Jerry Melillo, Almut Arneth, Shari Kovats, John Porter, John Shellenhuber/Katja Frieler, Nigel Arnell, Tim Carter, Linda Mearns, Martin Parry).
- Link to website (if available)*: <http://www.unep.org/provia/HOME/tabid/55173/Default.aspx>
- Goal of the Activity and a brief overview*:

To help form a more coherent interaction between the climate modelers in CMIP6 and the IAV community and help to design CMIP6-endorsed MIPs and online analysis capabilities that would enhance the benefit of CMIP simulations, we propose the creation of a Vulnerability, Impact, and Adaptation (VIA) Advisory Board for CMIP6 under the auspices of the Programme of Research on Climate Change Vulnerability, Impacts, and Adaptation (PROVIA). The VIA Advisory Board would include leaders of the established IAV projects, including AgMIP, WaterMIP, ISI-MIP, etc., as well as senior scientists in impacts sectors who have established credibility within their impacts community as well as in the wider climate change community. PROVIA is recognized within the World Climate Programme as an interface between researchers, stakeholders, and decision-makers within the VIA community and can play the vital role of representing the perspectives of this highly diverse, transdisciplinary community.

The VIA Advisory Board would not propose new experiments, but would serve as a

Diagnostic MIP for planning and evaluation of existing CMIP6 experiments. The VIA Advisory Board members would survey their respective communities (e.g., Agriculture, Urban, Biomes, Forestry, Oceans/Fisheries, Coastal, Water Resources, Health, Economics, Energy, Infrastructure/Transportation), coordinate activities, and provide comprehensive feedback for CMIP6 to consider in designing and prioritizing scenarios and metrics for analysis and benchmarking that would be relevant for VIA. This Advisory Board would also continue the interactions already underway that bring the VIA perspective into the development of new CMIP6-endorsed MIPs, and advise on better integration of the VIA community model processes and results into the Earth Systems Grid where CMIP6 outputs are archived. The VIA Advisory Board would help also raise awareness and establish best practices within the VIA community.

Background on PROVIA: The Global Programme of Research on Climate Change Vulnerability, Impacts and Adaptation (PROVIA) represents an interface between the research community and decision makers and other stakeholders to improve policy-relevant research on vulnerability, impacts and adaptation (VIA), allowing scientists to coordinate and facilitate the dissemination and practical application of their research. PROVIA helps international community of practice share practical experiences and research findings by improving the availability and accessibility of knowledge to the people that need it most. PROVIA aims to do so together with collaborative partners, knowledge networks, and the larger VIA community, by identifying research needs and gaps, helping scientific community to mobilize and communicate the growing knowledge-based on VIA so that governments and other main stakeholders are able to solicit scientific knowledge into their decision making processes.

- An overview of the proposed experiments*: The VIA Advisory Board will provide input in particular on the development and use of the representative concentration pathways and shared socioeconomic pathways (RCPs and SSPs, e.g., interacting with ScenarioMIP) which are the basis for most state-of-the-art projections of climate changes utilized in VIA assessments.
- An overview of the proposed evaluation/analysis of the CMIP DECK and CMIP6 experiments*: The Board will help guide the development of online metrics and visualizations that will appeal to the VIA community or researchers, stakeholders, and decision-makers. These include sector-specific indices (e.g., heat damage degree days for ecosystems, consecutive dry days for agriculture and water resources) and requirements for documentation and online guidance that will facilitate the use of CMIP6 products by the lay public.
- Proposed timing*: The VIA Advisory Board will be convened before March, 2015, with members serving two-year terms and rotating chairs to ensure new perspectives.

The VIA Advisory Board would not propose new experiments, but would serve as a Diagnostic activity for planning and evaluation of existing CMIP6 experiments. We have thus adjusted the below sections to better illustrate the design and outcomes of this activity.

For each proposed activity to be included in CMIP6**

- the activity design: The VIA Advisory Board will include leaders from major impact sectors, each of whom will have a mandate to coordinate with other experts within their sector to provide community-based guidance from their sector that can be integrated at the VIA Advisory Board level and then presented to CMIP6.
 - the science question and/or gap being addressed with this activity: The VIA Advisory Board will provide inputs from the VIA community on experiment and data design for CMIP6, guidelines for best practices in the use of CMIP6 outputs in VIA assessments, and advice on the development of online metrics and visualizations for the Earth Systems Grid.
 - possible synergies with other MIPs: The VIA Advisory Board will provide VIA perspective to MIPs with societal implications, for example including the development of RCPs and SSPs with ScenarioMIP, the use of ecosystem and agricultural models in conjunction with LUMIP, the health impacts of pollution policies in AerChemMIP, the role of water resource management in LandMIP.
 - potential benefits of the activity to
 - (A) climate modeling community: The VIA Advisory Board will improve the relevance of climate model outputs to society through the development of more creative, robust, and efficient applications of GCM outputs. The Board will also facilitate dissemination of important scientific findings and caveats that need to be recognized in the design and communication of climate impact assessments.
 - (B) Integrated Assessment Modelling (IAM) community: The VIA Advisory Board will provide important feedback on the implications of various policies and economic trajectories projected by the IAM community, potentially leading to shifts in the magnitude of feedbacks or the extent of plausible outcomes (e.g. land use, water resource availability, agricultural prices).
 - (C) Impacts Adaptation and Vulnerability (IAV) community: The VIA Advisory Board will dramatically increase the level of communication between CMIP and the IAV community with mutual benefits. In particular, the Board will facilitate the IAV community's access to, and understanding of, key climate model outputs for societal applications. The Board will also help build ties between the IPCC Assessment Report Working Group 2 with Working Groups 1 and 3.
 - (D) policy makers: The VIA Advisory Board will help CMIP6 incorporate the experience of the VIA community interactions with policy makers around the world, leading to online metrics tailored toward policy makers and a greater translation of climate model output toward social outcomes that are at the heart of policy maker interests.
- If possible, a prioritization of the suggested experiments, including any rationale**:
- The VIA Advisory Board will be most interested in the RCP experiments that form the basis for projections out to 2100, decadal prediction runs out to 2035 that inform a large number of decision-makers acting in the time frame where the climate change signal does not substantially differentiate itself from climate variability, and the RCP8.5 DECK experiment that will help link CMIP5 IAV findings with those that result from CMIP6 experiments.

- All model output archived by CMIP6-Endorsed MIPs is expected to be made available under the same terms as CMIP output. Most modeling groups currently release their CMIP data for unrestricted use. If you object to open access to the output from your experiments, please explain the rationale.**: We support open access.
- The Activity addresses at least one of the key science questions of CMIP6: The VIA Advisory Board would facilitate efforts to address all three key science questions of CMIP6. The VIA community would be better able to determine how the Earth System (in particular the impacted elements relevant to society) will respond to forcing, how model biases potentially influence decision-making in impacted sectors, and how climate variability, predictability, and uncertainty may be handled in preparing climate change adaptation and mitigation strategies that benefit impacted sectors.
 - A sufficient number of modeling groups have agreed to participate in the MIP: There are a large number of impacts modeling groups across various sectors that will use CMIP6 outputs and potentially guidance from the VIA Advisory Board. These include AgMIP, WaterMIP, ISI-MIP, and other community projects as well as smaller modeling groups not necessarily attached to large projects.
 - The MIP builds on the shared *CMIP DECK* experiments: The MIP will build in particular on the historical 20th Century simulations and the RCP8.5 simulations in the *DECK*.
 - A commitment to contribute to the creation of the CMIP6 data request and to analyze the data: The VIA Advisory Board will work with CMIP6 to help identify and create metrics and visualizations of relevance to the VIA community.
 - A commitment to identify observations needed for model evaluation and improved process understanding, and to contribute directly or indirectly to making such datasets available as part of obs4MIPs: The VIA Advisory Board will collect information about the observational datasets utilized by various VIA sectors and encourage the addition of those datasets to obs4MIPs.
 - The proposed activity is of central importance to CMIP6: The VIA Advisory Board will enhance the relevance of CMIP6 to society through all impact sectors.
 - The proposed activity has been run at least by two modeling groups already: The VIA Advisory Board will survey the best practices of impacts modelers from groups like AgMIP, ISI-MIP, WATERMIP, and others with a long history of VIA contributions.
 - The proposed activity is useful in a multi-model context and to a number of climate researchers: The VIA Advisory Board will encourage the use of multi-model ensembles both in the driving climate data and in the impacts models utilized. The Board will also expedite the transfer of knowledge and practices from the climate modeling community's long use of ensemble approaches into the VIA community which has only emphasized ensemble approaches in recent years.
 - A commitment to scientifically analyze, evaluate and exploit the proposed experiment: The VIA Advisory Board will enable a large number of researchers, stakeholders, decision-makers, and policy-makers to better integrate climate information into climate impact assessments across a number of sectors, with results also feeding back into the design and implications of climate modeling experiments.